



32 **Main**

33 Wildfires represent a major ecosystem disturbance and aerosol emission source, affecting the  
34 global carbon budget, climate, and human life<sup>1,2</sup>. Short-term influences of wildfires include  
35 damaged infrastructures, degraded air quality, and nutrient redistribution caused by the emission  
36 of smoke aerosols<sup>3,4</sup>. In situ and remote sensing measurements have also suggested the presence  
37 of mineral dust in smoke plumes<sup>3,5</sup>, caused by pyroconvective updraft from nearby burning<sup>6</sup>.  
38 Longer-term influences of wildfires primarily involve vegetation disturbances and resultant  
39 changes in ecosystem, hydroclimate, and geomorphology<sup>7,8</sup>. Among the natural consequences of  
40 the destroyed vegetation, especially the short species such as grasses and shrubs<sup>8</sup>, is the  
41 expansion of bare ground that is particularly susceptible to wind erosion<sup>9</sup>—the detachment of  
42 soil particles from the ground, and dust storms in an extreme condition<sup>10</sup>. The intensity of wind  
43 erosion and resultant dust emission depends on wind friction velocity<sup>11</sup>, vegetation structure<sup>12</sup>,  
44 and soil properties<sup>10</sup>. In addition to the clearance of vegetation and biocrusts cover, several  
45 additional features of wildfires may exacerbate the occurrence of post-fire dust storms. First, the  
46 fire-induced reduction in vegetation leads to an expanded vegetation canopy gap and reduced  
47 vegetation height, aerodynamically intensifying the severity of wind erosion<sup>12</sup>. Second, large  
48 wildfires are often associated with climate-driven, dry fuels and accompanying dry soils<sup>13</sup>, which  
49 favor dust emission. Moreover, fires may alter the physical and chemical properties of soils and  
50 disrupt the wet-bonding forces<sup>14</sup>, thereby further promoting the occurrence of wind erosion from  
51 these burned landscapes.

52

53 In situ observations and modeling studies have confirmed dust emission from post-fire  
54 landscapes mostly in North America<sup>5,15–17</sup>; yet, post-fire dust emission has not yet been globally  
55 examined using observational data. To fill this knowledge gap, the current study aims to (1)  
56 identify global hotspots of post-fire dust emission from a suite of satellite observations, (2) test  
57 the hypothesized driving mechanisms of dust emission after wildfires with observational data  
58 sets, and (3) diagnose the observed recent trends in the intensity and duration of post-fire dust  
59 emission.

60

61 Here we analyze a spectrum of satellite measurements of active fires, aerosol abundance and  
62 characteristics, vegetation cover, and soil moisture, as well as reanalysis wind for the period

63 2003–2020. This collection of global observational data sets well captures the dust emissions  
64 after wildfires during the 2010 and 2012 burning seasons in western United States, as reported by  
65 previous in situ observations and modeling studies<sup>15,16</sup> (**Extended Data Fig. 1**). Based on these  
66 observational datasets, we first identify large fire events with more than 20 active fires occurring  
67 within each 10 km pixel in consecutive seven day, and then search for significant vegetation  
68 reduction and accompanying enhanced dust load during the subsequent 60 days since the end of  
69 each large fire event (see Methods). To demonstrate the capability of currently applied satellite  
70 measurements in characterizing post-fire dust events, we first show extreme dust emission from  
71 the burned areas during the 2019–2020 Australian bushfires. Statistical assessments of post-fire  
72 dust events across the entire globe are shown afterwards.

73

#### 74 **Dust emission after 2019–2020 Australian bushfires**

75 Following the long-lasting drought conditions<sup>18</sup>, a series of large wildfires burned a historic  
76 186,000 km<sup>2</sup> across eastern Australia during the 2019–2020 bushfire season<sup>19</sup>. Satellite  
77 observations indicate a record-breaking low vegetation cover [represented by extremely low  
78 Enhanced Vegetation Index (EVI), a semi-quantitative measurement for the amount of  
79 vegetation, **Figure 1c**] and high dust concentration [represented by extremely high Dust Optical  
80 Depth (DOD), an approximate measure of columnar dust mass, **Figure 1d** and Supplementary  
81 **Fig. 1**] across the burned regions during December 2019 to February 2020. Indeed, the 2019–  
82 2020 bushfire season receives more than doubled dust loading, compared with an average year  
83 during the past two decades for this region<sup>20</sup>. Such a massive amount of dust particles is mainly  
84 emitted from the burned regions that witness the most severe vegetation damage. For example,  
85 the savannahs (around 27.8°S, 152.3°E) to the west of Brisbane experience a substantial  
86 reduction in vegetation cover (EVI drops from about 0.25 to 0.17, compared with a long-term  
87 average of 0.35) after the extensive fires during November 7–13, 2019 (**Figure 1b**). Following  
88 the persistent vegetation disturbances and abnormally dry soils, extreme DOD episodes are  
89 observed in December 2019 (**Figure 1a, b**). These post-fire dust episodes in eastern Australia are  
90 also captured by spaceborne lidar (Cloud-Aerosol Lidar with Orthogonal Polarization, CALIOP)  
91 measurements of total backscatter and depolarization ratio, as well as derived aerosol type  
92 information (**Extended Data Fig. 2**).

93

## 94 **Global occurrence of post-fire dust emission**

95 Based on the analyzed satellite measurements of fire, vegetation, and aerosols during the 18  
96 years of 2003–2020 (see Methods), we identify 151,727 large wildfire events with more than 20  
97 active fires detected in a  $0.1^\circ \times 0.1^\circ$  pixel during consecutive seven days. Among the analyzed  
98 large fire events, 91% and 54% are followed by significant EVI reduction and consequent dust  
99 events, respectively, during the subsequent 60 days. These 87,400 post-fire dust emission events  
100 are distributed across 36,386  $0.1^\circ$  pixels in the fire-prone regions of tropical savannahs in Africa,  
101 South America, and northern and eastern Australia, shrublands in western Australia, grasslands  
102 and croplands in central Asia, and various landscapes in western North America (**Figure 2a**).  
103 Among different landscapes, global savannahs contribute 66% of the currently identified large  
104 fire events, 59% of consequent significant EVI reduction, and 51% of post-fire dust events  
105 (**Figure 2b**).

106

107 The post-fire dust events typically last for 1–25 days and the maximum DOD ranges between 0.2  
108 to 1.8 (**Figure 2c**); the intensity and duration of post-fire dust events vary by land cover type  
109 (**Figure 2d**). The post-fire maximum DODs are on average 351% (192%–578%, 10<sup>th</sup>–90<sup>th</sup>  
110 percentiles of the relative increment among all pixels) larger than the local average DOD. Global  
111 savannahs see the most intensive (average DOD of 0.62, and 10<sup>th</sup>–90<sup>th</sup> percentiles of 0.21–1.56)  
112 and long-lasting (median duration of 3.5 days, and 10<sup>th</sup>–90<sup>th</sup> percentiles of 1–10 days) post-fire  
113 dust events, compared with other landscapes (**Figure 2d**). Regionally, the most severe and long-  
114 lasting post-fire dust events are observed over savannahs in West Africa and tropical Africa to  
115 the north and south of the Congo rainforest, where maximum DOD after large fires reaches 1.8,  
116 about three folds of the local 95<sup>th</sup> percentile of monthly DOD and close to that over the global  
117 leading dust sources, such as the Bodélé Depression in Chad<sup>21,22</sup> (**Extended Data Fig. 3b**).  
118 Moreover, the moderate-to-high DODs above 0.5 are widely seen over the normally dust-free  
119 regions, such as the boreal regions in North America and eastern Asia and mid-to-high latitudes  
120 in the Southern Hemispheric (**Figure 2c**). Indeed, among the 36,386 pixels where post-fire dust  
121 emissions are identified here, only 8% have ever experienced a DOD exceeding 0.2 that cannot  
122 be attributed to antecedent large wildfires during the study years.

123

124 Post-fire dust events occur episodically during weeks after large fires, due to the long-lasting,  
125 pre-fire dry soils and post-fire vegetation disturbances (**Extended Data Fig. 4**). The currently  
126 examined large wildfires cause a reduction in vegetation, represented by a shift in the probability  
127 distribution towards lower EVI during the first week after burning; this vegetation disturbance  
128 typically lasts for several months, accompanied by persistent dry soils (**Extended Data Figure**  
129 **4b, c**). Corresponding to the anomalies in vegetation and soil moisture, the probability  
130 distribution of DOD shifts toward higher values, with the probability of DOD exceeding the  
131 long-term 90<sup>th</sup> percentile between 21% to 40% during the first eight weeks after large wildfires  
132 (**Extended Data Figure 4a**). Unlike the long-lasting EVI and soil moisture anomalies, elevated  
133 surface wind speed mainly occurs during the extensive wildfire events and decays afterwards  
134 (**Extended Data Fig. 4d**).

135

136 The occurrence and duration of post-fire dust emission largely depends on the temporal and  
137 spatial extent of precedent wildfires (**Figure 3**). As the extent of wildfires increases, the posterior  
138 vegetation cover lowers, weakening its physical protecting and aerodynamical sheltering effects  
139 on soils. Consequently, the probability of less vegetation, more dusty situations elevate after the  
140 occurrence of large wildfires, and further enhances with the increment of precedent fire counts  
141 (**Figure 3a**). Specifically, the probability of extremely low monthly EVI (lower than long-term  
142 10<sup>th</sup> percentile) increases from 39% after a moderate fire event (with the occurrence of 21–30  
143 fires) to 78% after a severe fire event (with more than 100 fires). The intensified precedent fires  
144 are also accompanied by generally lower soil moisture (**Extended Data Fig 5a**), either due to  
145 large-scale climate variations that favor dryer conditions for burning or due to fire-induced soil  
146 moisture depletion, which further favors dust emission. As a result, the probability of extremely  
147 high monthly DOD (exceeding the long-term 90<sup>th</sup> percentile) increases from 28% after a  
148 moderate fire event to 84% after a severe fire event (**Figure 3a**); and the median duration of  
149 post-fire dust events increases from one day (10<sup>th</sup> – 90<sup>th</sup> percentiles of 1–6 days) after a moderate  
150 fire event to 10 days (10<sup>th</sup> – 90<sup>th</sup> percentiles of 8–16 days) after a severe fire event (**Figure 3b**).  
151 Corresponding to this dependence of post-fire dust emission on the extent of burning, the  
152 seasonal peak of post-fire dust emission occurs simultaneously with or shortly after the seasonal  
153 peak of active fires across the majority of global post-fire dust emission hotspots (**Extended**  
154 **Data Fig. 6**).

155

156 The long-lasting (>5 days) post-fire dust events are mainly present over all land cover types  
157 when post-fire EVI falls below 0.20 (**Figure 3b and Extended Data Fig. 7**), a typical EVI value  
158 over the active arid and semi-arid dust sources, such as the Sahara Desert (**Extended Data Fig.**  
159 **3a**). This relationship between post-fire EVI and post-fire dust event duration could serve as the  
160 basis for early warning of extreme post-fire dust activity. Noticeably, a large portion of the long-  
161 lasting post-fire dust events occur after a moderate fire event (**Figure 3b**). These long-lasting  
162 post-fire dust events after moderate fires are mostly observed over savannahs, mainly in Africa  
163 (**Extended Data Fig. 7c**), where transported dust from nearby dust sources<sup>21</sup> (e.g. Sahara, Sahel,  
164 and Middle East) are likely mixed with locally emitted dust from burned areas, thereby  
165 obscuring the accurate duration of these post-fire dust events.

166

167 The occurrence and intensity of post-fire dust events is also modulated by pre-fire drought  
168 conditions, as represented by soil moisture anomalies before the occurrence of fires (**Extended**  
169 **Data Figs. 8 and 9**). Over the semi-arid regions included in this analysis, drought conditions are  
170 favorable for dust emission from the sub-grid bare-soil areas, as reflected by the probability  
171 distribution of DOD towards higher values during dry periods even without fires, compared with  
172 wet periods (**Extended Data Fig. 8**). This difference in DOD probability distribution between  
173 relatively wet and dry periods partially diminishes after fires, indicating the primary role of fires  
174 on determining the occurrence and severity of post-fire dust events (**Extended Data Fig. 8**).  
175 Nevertheless, pre-fire drought conditions favor elevated occurrence and intensity of post-fire  
176 dust events (**Extended Data Fig. 9**).

177

### 178 **Recent trends in post-fire dust emissions**

179 While forest ecoregions have experienced a positive trend of 87.6 cases per decade in post-fire  
180 dust emission events (**Figure 4a**), other ecoregions have exhibited minimal changes in  
181 occurrence. However, the duration of these events has been increasing significantly (all p-values  
182 < 0.05, based on the Mann-Kendall monotonic trend test) over all landscapes, with a positive  
183 trend of 0.82, 0.54, 0.28, 1.05, and 0.80 days per decade over forests, shrublands, savannahs,  
184 grasslands, and croplands during 2003–2020 (**Figure 4b**). Indeed, the most long-lasting dust  
185 emission events are widely seen in either 2019 or 2020 over 4,699 pixels out of the 36,386

186 examined pixels (**Figure 4c**), such as those during the 2019–2020 Australian bushfire season  
187 (**Figure 1**) and 2020 western United States extreme fire season (**Extended Data Fig. 10**). This  
188 recent elongation of post-fire dust emission is attributed here to the increased extent of wildfires,  
189 as indicated by the positive trends in active fire counts (**Figure 4b**). For the forest, shrubland,  
190 savannah, grassland, and cropland pixels examined here, the averaged fire counts per 0.1° pixel  
191 per event increase by 3.50, 0.56, 0.37, 1.12, and 0.24 per decade, respectively, with all p-values  
192 < 0.05 except for the croplands, according to the Mann-Kendall test. Meanwhile, the analyzed  
193 forest, grassland, and cropland pixels exhibit moderately significant reduction ( $p < 0.1$ , based on  
194 the Mann-Kendall test) in pre-fire soil moisture during 2003 to 2020.

195

## 196 **Discussion**

197 Our findings have direct implications on the ecological and societal impacts of intensifying  
198 droughts and wildfires over certain landscapes. In addition to the instantaneous societal  
199 disruptions and health risks, drought and resultant wildfires set the stage for dust storms even  
200 weeks after burning. Our findings are further supported by recent in-situ observations<sup>23</sup> reporting  
201 that wildfires reduce soil biocrusts by 50%, which may also enhance dust emissions<sup>24</sup>. These  
202 post-fire dust storms could be as intensive as those observed in arid to semi-arid lands (**Figure 2**  
203 and **Extended Data Fig. 3b**) and cause similar infrastructure damages and air quality declines<sup>25</sup>.  
204 Compared with the dryland dust storms, the post-fire dust storms may cause even larger  
205 socioeconomic and health impacts, due to their closer location to populated areas and possible  
206 mixing of harmful combustion residuals into the post-fire dust storms. The emitted soil particles  
207 from these disturbed lands may enter the global dust cycle, altering the radiation budget<sup>26</sup>, cloud  
208 and precipitation patterns<sup>27</sup>, as well as oceanic<sup>3</sup> and terrestrial biogeochemistry<sup>28</sup>. For example,  
209 dust particles from Australia are key suppliers of iron, a bio-essential trace metal, to the iron-  
210 limited ecosystems of Southern Ocean. The recently uncovered widespread phytoplankton  
211 blooms from December 2019 to March 2020 in the Southern Ocean downwind of Australia<sup>29</sup>  
212 could be a result of the post-fire dust emission triggered by the 2019-2020 Australian bushfires  
213 (**Figure 1**, **Extended Data Fig. 2**, and **Supplementary Fig. 1**). On the other hand, the high-  
214 latitude post-fire dust emissions (**Figure 2a, c**) may provide an additional source for light-  
215 absorbing aerosols, beyond transported dust and smoke, that may accelerate snow darkening and  
216 melting with warming<sup>30</sup>. Furthermore, the currently identified role of drought and fires on post-

217 fire dust emission (**Figure 3, Extended Data Figs. 8 and 9**) will potentially become more  
218 interactive and complicated in the upcoming decades, given the complex response in drought<sup>31</sup>  
219 and wildfires<sup>32–35</sup> to anthropogenic activity and global warming. Overall, our study calls for  
220 adaptation and/or mitigation strategies for this compound drought-fire-dust hazard, which is  
221 likely to become more frequent and severe with ongoing environmental change.

222

223 The uncertainty of our study mainly derives from the quality of the currently analyzed satellite  
224 data, especially retrieval difficulties. First, transported dust provides inevitable uncertainty for  
225 quantifying the intensity and duration of post-fire dust emissions, particularly over the African  
226 savannahs that are close to dry-land dust hotspots. Although we focus on coarse-mode dust  
227 optical depth as the metric for dust abundance to minimize the impacts of transported, smaller-  
228 sized dust particles on our results (see Methods), the retrieval of coarse-mode dust optical depth  
229 involves errors too<sup>36,37</sup>. A sensitivity test that addresses possibly non-local–originated high  
230 DODs suggests that transported dust may cause an overestimation of post-fire dust duration by  
231 0–3 days (10<sup>th</sup> – 90<sup>th</sup> percentiles of all pixels from both regions) and maximum post-fire DOD by  
232 0–7% in North Africa (**Supplementary Fig. 2**) and Australia (**Supplementary Fig. 3**). Second,  
233 deposited combustion ash from the burned vegetation could be lifted by strong winds along with  
234 mineral dust. Ash (typically 0 – 50 mm thick above surface) consists of mineral materials and  
235 charred organic components with a wide range of particle size, shape, and optical properties that  
236 partially overlap with dust<sup>38</sup>; therefore, the currently identified post-fire dust events, especially  
237 those shortly after burning, may contain a mixture of co-emitted dust and deposited ash. Here we  
238 perform a qualitative test that assumes pure dust emission only occurs after the first day of post-  
239 fire high DOD. This test confirms similar peak intensity of post-fire pure dust emissions versus  
240 potential dust-ash-mixture emissions (**Supplementary Fig. 4**). Nevertheless, a dust-ash-mixture  
241 storm can cause similar socioeconomic and health problems as a pure dust storm. With  
242 development of higher-quality satellite and ground observations, especially the hyperspectral and  
243 mineralogical information of dust and ash emitted from specific geographic locations, we will  
244 continue to quantify regional and global post-fire dust emissions. Based on this post-fire dust  
245 emission inventory, future observational and modeling studies should characterize the dynamical  
246 and optical properties of post-fire dust emissions, quantify their climatic impacts, and compare  
247 with regular dust emissions from dry lands.



248

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257

258 **Author contributions**

259 YY conceived the study, analysed the data, and wrote the manuscript with contribution from PG.  
260 PG retrieved MODIS DOD and FOD data from MODIS Deep Blue aerosol products.

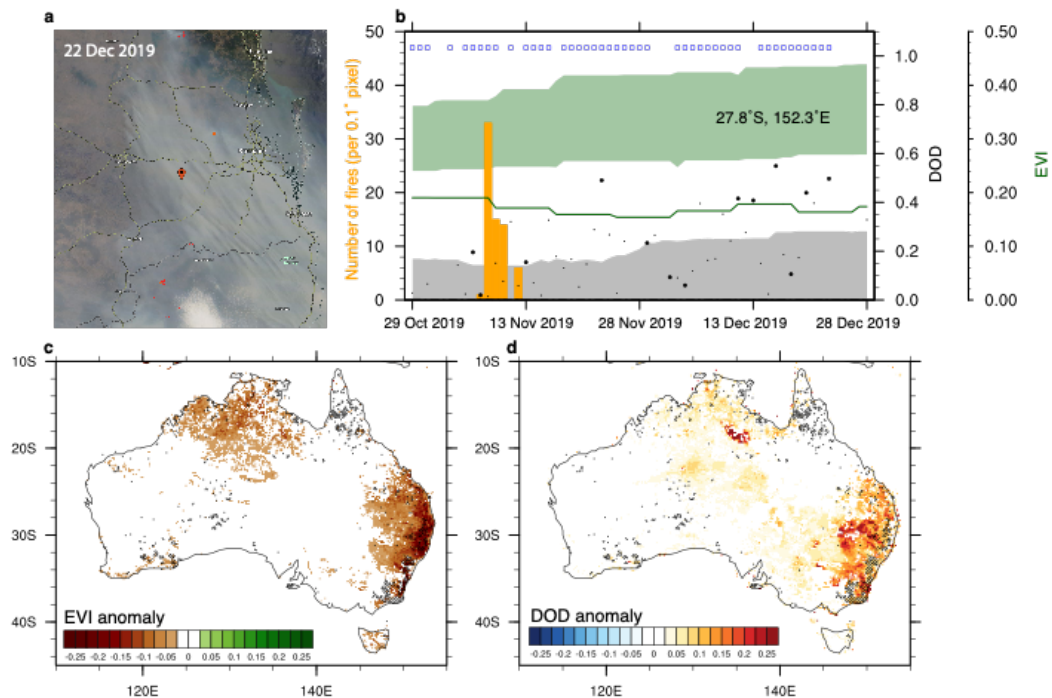
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262 **Competing interests**

263 The authors declare that they have no conflict of interest.

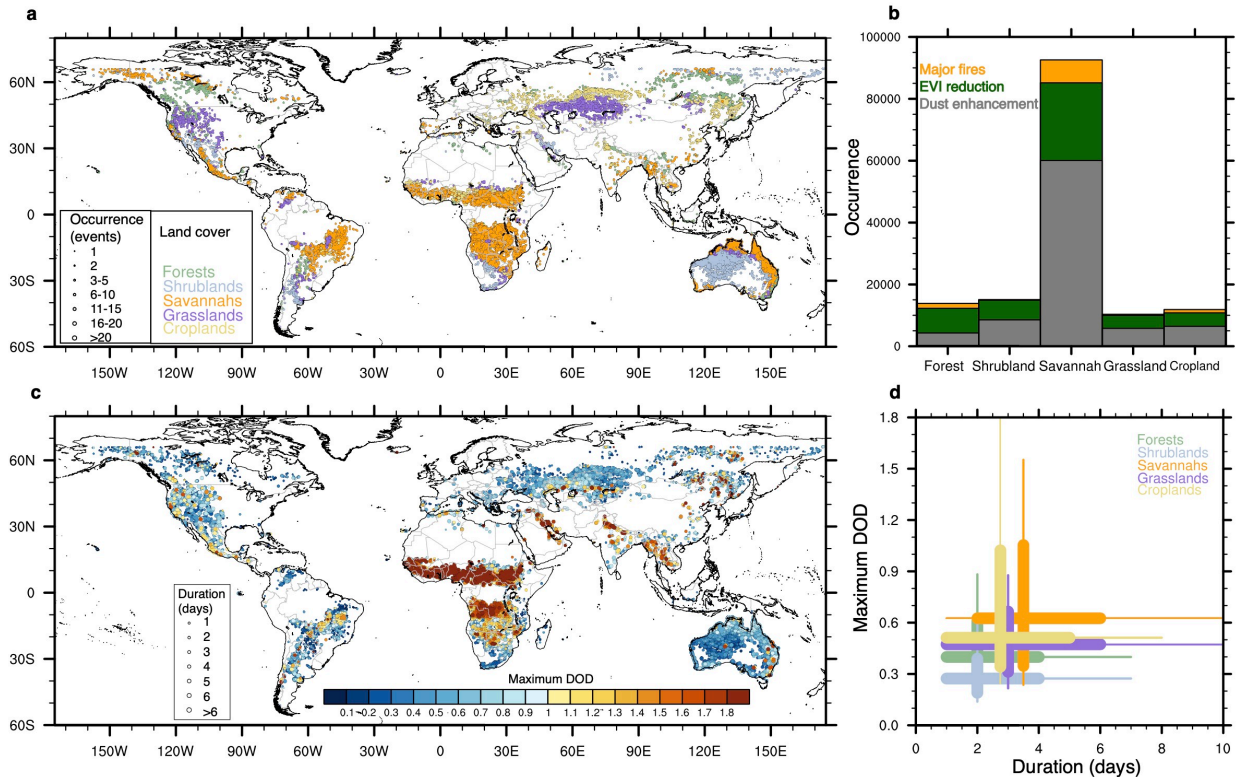
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265 **Figures**



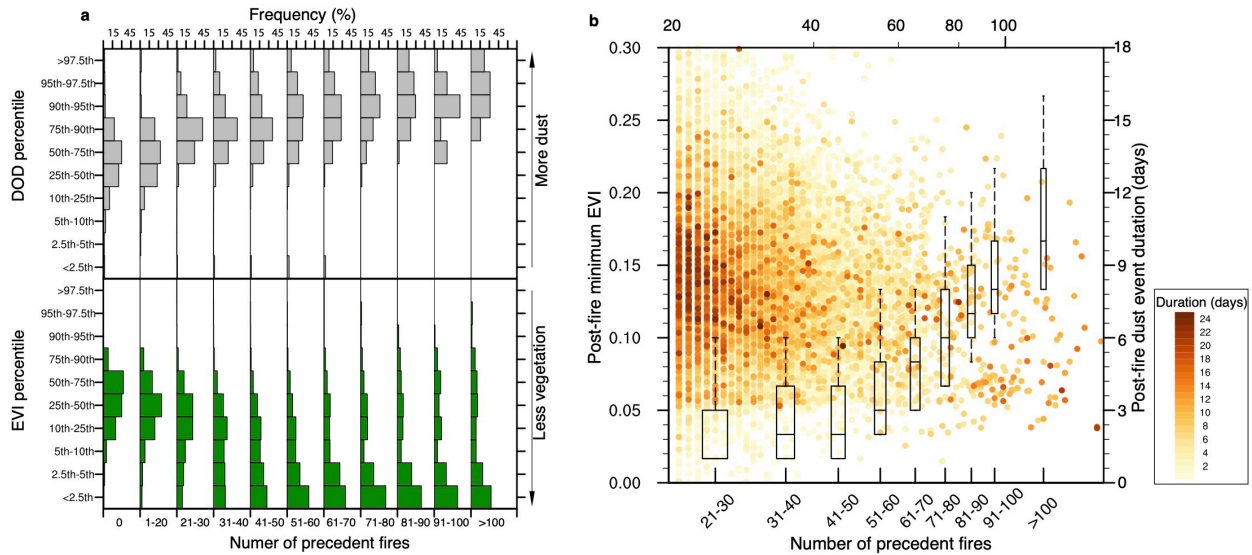
266  
 267 **Figure 1 Extreme dust activity associated with vegetation disturbances caused by the 2019-**  
 268 **2020 Australian bushfires. a.** True color image of aerosol plumes originating from the active  
 269 fires and burned areas in southeastern Australia on December 22, 2019, captured by the  
 270 Moderate Resolution Imaging Spectroradiometer (MODIS) instrument onboard the Terra  
 271 satellite that overpasses in local morning time. The red dots indicate active fires detected by  
 272 MODIS onboard both the Terra and Aqua satellites, both day and night. **b.** Time series of active  
 273 fire count (orange bars, referring to the left y-axis), enhanced vegetation index (EVI, green line,  
 274 referring to rightmost y-axis), and daily maximum dust optical depth (DOD, black dots, referring  
 275 to the inter right y-axis) within  $\pm 0.05^\circ$  of 27.8°S, 152.3°E (location indicated in panel **a**). The  
 276 blue squares indicate dates with abnormally dry soil (below long-term 10<sup>th</sup> percentile). The green  
 277 and grey shadings represent the long-term 10<sup>th</sup>–90<sup>th</sup> percentiles in the daily EVI and DOD,  
 278 respectively. The large and small black dots represent time series of DOD; large dots indicate  
 279 situations with relatively small amount of biomass burning aerosols, represented by below-  
 280 average coincident fine-mode optical depth (FOD). **c.** Anomaly in EVI during December 2019 to  
 281 February 2020, compared to the long-term average during December to February of 2000–2020.

282 **d.** Anomaly in DOD during December 2019 to February 2020, compared to the long-term  
 283 average during December to February of 2000–2020. In **c** and **d**, only the EVI anomalies and  
 284 DOD anomalies more extreme than the long-term 10<sup>th</sup>–90<sup>th</sup> percentiles are shown in color. In **c**  
 285 and **d**, the stitches and slashes indicate 0.1° pixels with more than 30 and 100, respectively,  
 286 active fires during the December 2019 to February 2020 Australian bushfire season. Satellite  
 287 image from NASA Earth Observatory. Figure created using NCL<sup>39</sup>.

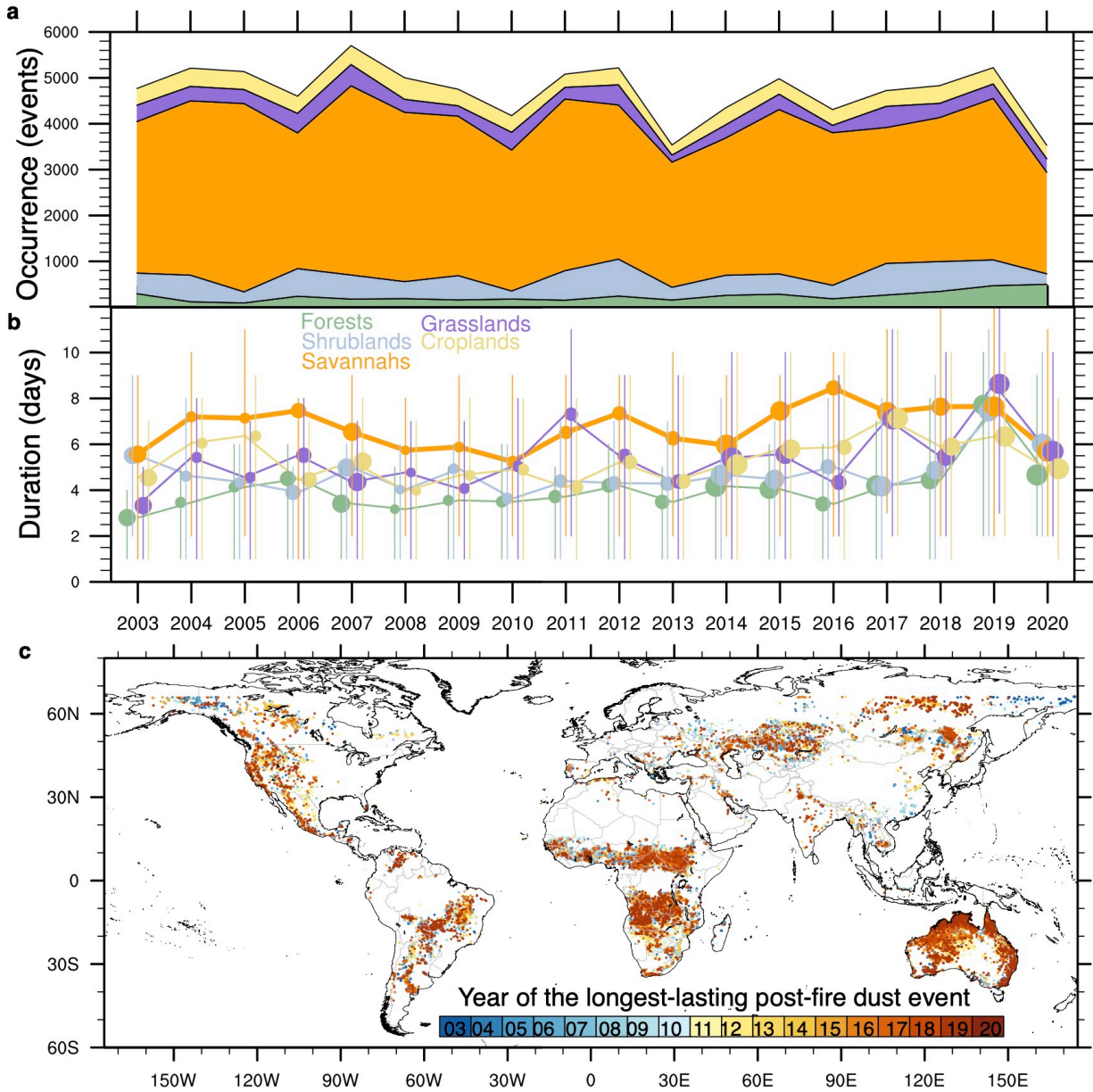


288  
 289 **Figure 2 Global distribution of post-fire dust events.** **a.** Occurrence of post-fire dust events  
 290 during 2003–2020 indicated by the size of dots, with color representing the dominant land cover  
 291 type. The land cover types are identified by the Moderate Resolution Imaging Spectroradiometer  
 292 (MODIS) Terra+ Aqua Combined Land Cover product following the International Geosphere  
 293 Biosphere Programme (IGBP) scheme for the period 2003–2020. **b.** Total occurrence of large  
 294 wildfire events (orange bars), significant Enhanced Vegetation Index (EVI) reduction (green  
 295 bars), and dust emission (grey bars) by land cover type (see Methods). **c.** Maximum Dust Optical  
 296 Depth (DOD, color of dots), representing the columnar dust loading associated with the most  
 297 intensive post-fire dust emission, and mean duration (days, size of dots) of post-fire dust events.  
 298 **d.** Boxplot (thin lines: 10<sup>th</sup> – 90<sup>th</sup> percentiles, thick lines: 25<sup>th</sup> – 75<sup>th</sup> percentiles, intersection:

299 median of both metrics) of maximum DOD and duration of post-fire dust events, by land cover  
 300 type.



301  
 302 **Figure 3 Severity of post-fire dust events regulated by the extent of precedent wildfires and**  
 303 **vegetation disturbance. a.** Probability distribution of post-fire, 30-day average (top) DOD and  
 304 (bottom) EVI as a function of number of precedent fires. The probability distribution is  
 305 represented by the frequency (%) of DOD and EVI below the long-term 2.5<sup>th</sup> percentile, between  
 306 the 2.5<sup>th</sup>–5<sup>th</sup>, 5<sup>th</sup>–10<sup>th</sup>, 10<sup>th</sup>–25<sup>th</sup>, 25<sup>th</sup>–50<sup>th</sup>, 50<sup>th</sup>–75<sup>th</sup>, 75<sup>th</sup>–90<sup>th</sup>, 90<sup>th</sup>–95<sup>th</sup>, 95<sup>th</sup>–97.5<sup>th</sup> percentiles,  
 307 and above the 97.5<sup>th</sup> percentile. **b.** Scatterplot of post-fire minimum EVI (referring to the left y-  
 308 axis) and number of precedent fires. Each dot in **b** corresponds to a post-fire dust event, with the  
 309 color representing the duration (days) of this event. The boxes in **b** indicate the 10<sup>th</sup>, 25<sup>th</sup>, 50<sup>th</sup>,  
 310 75<sup>th</sup>, and 90<sup>th</sup> percentiles of post-fire dust event duration (days, referring to the right y-axis) with  
 311 precedent fires ranging between 21–30, 31–40, 41–50, 51–60, 61–70, 71–80, 81–90, 91–100, and  
 312 above 100 during the burning period.



313  
 314 **Figure 4** Observed evolution of the occurrence and duration of post-fire dust events during  
 315 **2003–2020.** **a–b.** Time series of the **a.** occurrence and **b.** mean duration (days) of post-fire dust  
 316 events for each dominant land cover type during 2003–2020. In **b.**, the sizes of dots are  
 317 proportional to the average fire counts per large wildfire event for each land cover type and year.  
 318 The thicknesses of the lines are proportional to the total occurrence of post-fire dust emission for  
 319 each land cover type during 2003–2020. The vertical lines indicate the 10<sup>th</sup> to 90<sup>th</sup> percentiles of  
 320 duration among all post-fire dust events for each land cover type and year. **c.** Year of the longest-  
 321 duration post-fire dust event.

## 322 **Methods**

### 323 **Satellite-based measurements of wildfire intensity**

324 To assess the intensity of wildfires, we analyze the active fires reported by the Moderate  
325 Resolution Imaging Spectroradiometer (MODIS) onboard the polar-orbiting Terra and Aqua  
326 satellites. The collection 6.1 MODIS active fire product detects fires in 1-km pixels that burn at  
327 the time of overpass under relatively cloud-free conditions using a contextual algorithm. The  
328 detection algorithm uses native (i.e., unprojected swath) 4-, 11-, and 12- $\mu\text{m}$  brightness  
329 temperatures derived from the corresponding 1-km MODIS channels, and, for daytime  
330 observations, 0.65-, 0.86-, and 2.1- $\mu\text{m}$  reflectance, aggregated to 1-km spatial resolution<sup>40</sup>. Daily  
331 wildfire intensity is examined here as the total active fire counts from both the day-time and  
332 night-time MODIS measurements.

333

334 The relatively coarse-resolution satellite measurements of active fires at about 1 km resolution  
335 may miss nearly half of the burned area in Africa detected by higher resolution satellite  
336 measurements (about 20 m resolution) in a given year<sup>41</sup>. The underrepresentation of small fires  
337 may lead to underestimation of small fire-induced dust emission.

338

### 339 **Satellite-based measurements of dust and other aerosols**

340 Dust Optical Depth (DOD) is a column-integration of extinction coefficient by mineral particles.  
341 The current study examines DOD from MODIS onboard the Terra and Aqua satellites and the  
342 non-spherical aerosol optical depth (nsAOD) from the Multiangle Imaging SpectroRadiometer  
343 (MISR) instrument<sup>42</sup> on Terra, during 2003-2020.

344

345 MODIS DOD represents the optical depth of absorbing, coarse-mode aerosols that are often dust  
346 over bare ground or sparsely vegetated regions. Following Pu et al. (2020)<sup>43</sup>, daily MODIS DOD  
347 is retrieved from collection 6.1, level 2 MODIS Deep Blue aerosol products<sup>44,45</sup>, including  
348 aerosol optical depth (AOD), single-scattering albedo ( $\omega$ ), and the Ångström exponent ( $\alpha$ ). All  
349 the daily variables are first interpolated to a  $0.1^\circ \times 0.1^\circ$  grid using the algorithm described by  
350 Ginoux et al. (2010)<sup>46</sup>. To account for dust's absorption of solar radiation and separate dust from  
351 scattering aerosols, such as sea salt, we require the single-scattering albedo at 470 nm to be less  
352 than 0.99 for the retrieval of DOD. Based on the size distribution of dust towards the coarse



353 range and to separate it from fine particles, DOD is retrieved as a continuous function of AOD  
354 and Ångström exponent:

$$355 \text{ DOD} = \text{AOD} \times (0.98 - 0.5089\alpha + 0.051\alpha^2). \quad (1)$$

356

357 This retrieval of DOD is on the basis of Ångström exponent's sensitivity to particle size, with  
358 smaller values of Ångström exponent indicating larger particles<sup>47</sup>, and the previously established  
359 relationship between Ångström exponent and fine-mode AOD<sup>48</sup>. Details about the retrieval  
360 process and estimated errors are summarized by Pu and Ginoux (2018)<sup>37</sup>. MODIS DOD products  
361 have been widely used for the identification and characterization of dust sources<sup>21,49,50</sup>, as well as  
362 examination of variations in regional and global dustiness<sup>37,43,51,52</sup>.

363

364 Following the recommendation from Baddock et al. (2016)<sup>50</sup> and previous applications of  
365 MODIS DOD<sup>37,43,51,52</sup>, here we use DOD with a low-quality flag of QA = 1, under the  
366 assumptions that 1) dust sources are better detected using DOD with a low-quality flag, and 2)  
367 retrieved aerosol products are poorly flagged over dust source regions. For example, when the  
368 standard deviation of AOD between 10×10 pixels is greater than 0.18, the retrieval algorithm  
369 considers the scene as cloudy although intense dust plumes over dust sources could easily reach  
370 such value<sup>50</sup>. Indeed, the comparison with the Aerosol Robotic Network (AERONET) DOD  
371 shows drastically decreased sample size but minimally enhanced consistency for any land cover  
372 type after applying a higher quality flag (**Supplemental Figs. 5–6**).

373

374 The residual of total aerosol optical depth after subtracting DOD, namely fine-mode optical  
375 depth (FOD), represents the abundance of other aerosol species, which is primarily smoke  
376 aerosols over active fires. In short, DOD and FOD represent the atmospheric abundance of  
377 particles that are relatively coarse and fine, respectively, mainly reflecting dust and carbonaceous  
378 aerosols, over burned and active burning areas. Overall, the optimal spatial and temporal  
379 coverage of MODIS aerosol products with over 20 years' record warrant its application for  
380 studying the day-to-day variations and environmental drivers of global aerosol loads.

381

382 It should be noted that limited by the spatial resolution of MODIS aerosol products, our study is  
383 conducted at 0.1° latitude and longitude resolution (about 10 km near the equator), which may

384 not be ideal for accurate representation of fire, dust, and environmental structures over complex  
385 landscapes.

386  
387 Benefiting from its multiangle observations, MISR data can be used to directly retrieve AOD and  
388 particle properties<sup>42</sup>. In the current study, Version 23, Level 2, daily MISR 550-nm nonspherical  
389 AOD (nsAOD) at 4.4-km resolution<sup>53</sup> is compared with MODIS DOD. The MISR nonspherical  
390 AOD fraction is often referred to as “fraction of total AOD due to dust”, as dust is the primary  
391 nonspherical aerosol particle in the atmosphere, especially over sparsely vegetated regions<sup>54</sup>. The  
392 MISR nsAOD has been used to examine variations in dustiness in North Africa and the Middle  
393 East<sup>22,55–57</sup>. Similar to our use of MODIS DOD with a low-quality flag, here we analyze the raw  
394 MISR nsAOD retrieval without quality filtering. MISR nsAOD data is also interpolated to a  $0.1^\circ$   
395  $\times 0.1^\circ$  grid using the algorithm described by Ginoux et al. (2010)<sup>46</sup>. Due to its relatively narrow  
396 swath of  $\sim 380$  km, MISR samples the study region about every 2-16 days. The sparse sampling  
397 of MISR limits its application in understanding day-to-day variations of aerosols, such as in our  
398 current study; but MISR’s capability at distinguishing dust particles from other aerosol species  
399 provides useful benchmark for evaluating other satellite-based approximate measurements of  
400 dust abundance. The correlation between temporally (both onboard the Terra satellite) and  
401 spatially (within  $0.1^\circ$  pixels) collocated MISR nsAOD and MODIS DOD measurements during  
402 2019 to 2020 suggests a generally high consistency between the two measurements of dust mass  
403 loading (**Supplementary Fig. 7**). Correlation exceeding 0.7 is widely seen over the identified  
404 hotspot regions for post-fire dust emission (e.g. in **Figure 2**), such as tropical savannahs in  
405 Africa, South America, and northern and eastern Australia, shrublands in western Australia,  
406 grasslands and croplands in central Asia, and various landscapes in western North America  
407 (**Supplementary Fig. 7**), ensuring the reliability of MODIS DOD in the current analysis.

408  
409 The Version 3, level 2 (cloud screened and quality assured), sub-daily AERONET coarse-mode  
410 AOD (DOD) and fine-mode AOD (FOD) at 500 nm obtained from the 205 sun photometers  
411 across the globe<sup>58</sup> and retrieved by the Spectral Deconvolution Algorithm (SDA)<sup>59</sup> are analyzed  
412 here to evaluate the accuracy of spatially and temporally collocated MODIS DOD and FOD,  
413 especially for various land cover types. Here a “collocated observation” is identified when there  
414 is available MODIS DOD and FOD over the  $0.1^\circ$  grid covering the AERONET site within  $\pm 0.5$



415 hour of the corresponding AERONET site observation. This definition results in a total of 64,390  
416 collocated observations between AERONET and MODIS. The comparison shows acceptable  
417 consistency between MODIS and AERONET in both DOD (**Supplementary Fig. 5**) and FOD  
418 (**Supplementary Fig. 8**) across the major landscapes. Highest correlation and lowest root-mean  
419 square error (RMSE) between MODIS and AERONET DOD is seen over forests and savannahs,  
420 respectively (**Supplementary Fig. 5**). Highest correlation and lowest RMSE between MODIS  
421 and AERONET FOD are both seen over savannahs (**Supplementary Fig. 8**). Higher accuracy of  
422 MODIS DOD and FOD is particularly present for high DOD and FOD (**Supplementary Figs. 5,**  
423 **8**).

424  
425 Total attenuated backscatter and depolarization ratio at 532 nm, as well as aerosol subtype  
426 information from the Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) Version 4.20  
427 (V4) Level 2 aerosol products are analyzed to confirm episodic dust emission from post-fire  
428 landscapes in eastern Australian during the 2019–2020 bushfire season. The spaceborne lidar  
429 instruments, such as CALIOP aboard the Cloud Aerosol Lidar and Infrared Pathfinder Satellite  
430 Observation (CALIPSO) spacecraft<sup>60</sup>, are able to provide vertical structure of aerosol and clouds.  
431 Despite CALIOP's limited spatial coverage (with a diameter of 70 m on Earth surface), the  
432 vertical distribution of aerosol abundance (reflected by total attenuated backscatter), aerosol  
433 shape (reflected by depolarization ratio, with a typical value of 0.2–0.3 for dust<sup>61</sup>), and aerosol  
434 subtypes (identified from altitude, location, surface type, estimated particulate depolarization  
435 ratio, integrated attenuated backscatter<sup>62</sup>) is particularly useful for identifying post-fire dust  
436 emission during selected events. Data for four nighttime overpasses on November 5, 2019,  
437 December 22, 2019, January 4, 2020, and January 22, 2020 are shown in **Extended Data Fig. 2**.  
438 These overpasses capture near-surface dust-smoke mixtures during active burning and pure dust  
439 after burning over the burned land in eastern Australia.

#### 440 441 **Satellite-based measurements of vegetation cover**

442 Collection 6, MODIS Enhanced Vegetation Index (EVI) derived from atmospherically corrected  
443 reflectance in the red, near-infrared, and blue wavebands<sup>63</sup> is analyzed here for vegetation  
444 disturbances caused by wildfires. EVI and another vegetation index, the Normalized Difference  
445 Vegetation Index (NDVI), effectively characterize the global range of vegetation states and

446 processes and have been successfully applied in various ecosystem, climate, and natural  
447 resources management studies<sup>64,65</sup>. Compared with NDVI, EVI minimizes canopy-soil variations  
448 and improves sensitivity over dense vegetation conditions. The VI's use a MODIS-specific  
449 compositing method based on product quality assurance metrics to remove low quality pixels.  
450 From the remaining good quality VI values, a constrained view angle approach then selects a  
451 pixel to represent the compositing period (from the two highest NDVI values it selects the pixel  
452 that is closest-to-nadir). Benefiting from the MODIS sensors aboard both Terra and Aqua  
453 satellites, here we analyze 16-day EVI composite eight days apart from both satellites, thereby  
454 obtaining a higher temporal resolution (8-day) product by combining both data records.  
455 Corresponding to the spatial resolution of analyzed aerosol data, the original 1-km EVI data is  
456 also interpolated to a  $0.1^\circ \times 0.1^\circ$  grid using the algorithm described by Ginoux et al. (2010)<sup>46</sup>.

457

#### 458 **Identification of dust emission after wildfires**

459 The identification of post-fire dust events follows these steps: (1) We screen the daily  $0.1^\circ$  active  
460 fire count data to identify the location and time of the occurrence of large fires, namely more  
461 than 20 fires in consecutive seven days. Each location-time combination is defined as a large  
462 wildfire event. We choose 20 fires as the threshold for identifying large fires because of the  
463 minor change in the probability distributions of EVI and DOD after the occurrence of 1-20 fires  
464 during the antecedent burning week (**Figure 3a**). (2) Among these large wildfire events, we  
465 identify those with significant EVI reduction, i.e. any 8-day EVI that falls below the 10<sup>th</sup>  
466 percentile of the long-term spread, during the subsequent 60 days after the end of burning. The  
467 long-term spread for a specific  $0.1^\circ$  pixel on a specific date of year is obtained by aggregating all  
468 16-day EVI measurements from both MODIS-Terra and MODIS-Aqua within  $\pm 15$  days of that  
469 date for that pixel during 2003-2020. (3) Among those large wildfire events that trigger  
470 significant EVI reduction, we then examine DOD and FOD measurements during the 60 days  
471 after the end of burning. A day with significant DOD increase, i.e., exceeding the 90<sup>th</sup> percentile  
472 of long-term spread, from either morning or afternoon measurements and concurrent moderate-  
473 to-low FOD, i.e. below the 50<sup>th</sup> percentile of long-term spread, from both the morning and  
474 afternoon measurements is defined as a post-fire dust emission day. The reason for excluding  
475 high FOD situations is to minimize contamination of flying ash and smoke aerosols over active  
476 fires on the DOD signal, as flying ash and smoke aerosols are mostly smaller particles compared

477 with dust<sup>66,67</sup>. The long-term spread of DOD and FOD for a specific date of year and 0.1° pixel  
478 are aggregated from both MODIS-Terra and MODIS-Aqua within ±15 days of that date for that  
479 pixel during 2003-2020. The duration of a post-fire dust event is defined as the total number of  
480 post-fire dust emission days within 60 days of the end of burning. The maximum DOD of a post-  
481 fire dust event is defined as the maximum DOD among the post-fire dust emission days.

482  
483 Note that the identification of post-fire dust events depends on the thresholds for significant EVI  
484 reduction and DOD increase. Among the analyzed 151,727 large fire events (52% of all week-  
485 pixel combinations that experienced at least one active fire), 91% and 78% are followed by any  
486 8-day EVI falling below long-term 10<sup>th</sup> percentile and 5<sup>th</sup> percentile, respectively, during the  
487 subsequent 60 days; 54% and 37% are followed by any daily DOD reaching long-term 90<sup>th</sup> (with  
488 EVI below 10<sup>th</sup> percentile) and 95<sup>th</sup> (with EVI below 5<sup>th</sup> percentile) percentiles, respectively,  
489 during the subsequent 60 days.

490

#### 491 **Statistics of post-fire dust events for different land cover types**

492 Number of large wildfire events, those with significant EVI reduction, and with post-fire dust  
493 emission events are reported for different land cover types, along with the distribution of the  
494 duration and maximum DOD of post-fire dust emissions. The land cover data is obtained from  
495 the collection 6 MODIS Terra+Aqua Combined Land Cover product<sup>68</sup>. We examine the primary  
496 land cover scheme that identifies 17 classes defined by the International Geosphere-Biosphere  
497 Programme (IGBP), including 11 natural vegetation classes, three human-altered classes, and  
498 three non-vegetated classes. The yearly land cover fraction data for the 17 classes originally at  
499 0.05° latitude and longitude are first regridded to the same 0.1° grid as the aerosol data, following  
500 Ginoux et al. (2010). All global 0.1° pixels are then grouped to five dominant land cover types,  
501 namely forests (Evergreen needleleaf forests, Evergreen broadleaf forests, Deciduous needleleaf  
502 forests, Deciduous broadleaf forests, and mixed forests), shrublands (closed and open  
503 shrublands), savannahs (woody savannahs and savannahs), grasslands, and croplands (croplands  
504 and cropland/natural vegetation mosaics), using the 18-year average land cover fractions. Note  
505 that the uncertainty of the MODIS land cover data, such as the unrealistic savannahs in the  
506 boreal region (**Figure 2a**), may complicate the land cover-specific interpretation of current  
507 results<sup>69</sup>.

508

509 **Soil moisture and wind observations**

510 To explore additional drivers of dust emission after wildfires, we analyze daily soil moisture at  
511 0.1° spatial resolution from European Space Agency (ESA) Climate Change Initiative (CCI) soil  
512 moisture data<sup>70</sup> and hourly 10-m wind speed at 0.1° spatial resolution from European Centre for  
513 Medium-Range Weather Forecasts (ECMWF) Reanalysis v5 (ERA5)<sup>71</sup>. The ESA CCI soil  
514 moisture combines various single-sensor active and passive microwave soil moisture products  
515 into three harmonized products: a merged ACTIVE, a merged PASSIVE, and a COMBINED  
516 active + passive microwave product. Here we analyze the version 06.1 break-adjusted  
517 COMBINED daily soil moisture for the top layer<sup>72</sup>. This product involves several algorithm  
518 updates and represents the most accurate ESA CCI global soil moisture data from 1978-2020.  
519 ERA5 is the latest reanalysis from ECMWF covering the period of 1950 to near real time and  
520 assimilates various observations in the upper air and near surface. Regional evaluation of ERA5  
521 hourly 10-m wind speed show vastly improved accuracy of ERA5 wind data compared with its  
522 older version ERA-Interim, but relatively large discrepancy with in situ observations remains  
523 over complex terrains<sup>73,74</sup>. In this study, we focus on daily maximum 10-m wind speed, which is  
524 directly related to dust emission<sup>75</sup> and obtained from the original hourly ERA5 reanalysis data.  
525 We obtain collocated daily soil moisture and daily maximum wind speed with observed dust at  
526 the nearest pixel of their original grid to the corresponding location of dust pixel.

527

528 **Understanding post-fire dust emissions through examination of aerosols, EVI, soil moisture,  
529 and wind speed responses to large wildfires**

530 To assess the temporal evolution of environmental responses to large wildfires, weekly averaged  
531 EVI, soil moisture, and aerosol, as well as weekly maximum wind speed, are examined in their  
532 probability distribution, i.e. the frequency of these variables falling below the long-term 2.5<sup>th</sup>  
533 percentile, between the 2.5<sup>th</sup>–5<sup>th</sup>, 5<sup>th</sup>–10<sup>th</sup>, 10<sup>th</sup>–25<sup>th</sup>, 25<sup>th</sup>–50<sup>th</sup>, 50<sup>th</sup>–75<sup>th</sup>, 75<sup>th</sup>–90<sup>th</sup>, 90<sup>th</sup>–95<sup>th</sup>,  
534 95<sup>th</sup>–97.5<sup>th</sup> percentiles, and above the 97.5<sup>th</sup> percentile. For example, a frequency of certain  
535 variable falling below the long-term 2.5<sup>th</sup> percentile that exceeds 2.5% indicates a higher  
536 probability of extremely low value of this variable after the occurrence of large fires. The long-  
537 term percentiles of these variables for a specific week of year and 0.1° are aggregated from that  
538 week and ±1 week of the year for that pixel during 2003-2020.

539

540 To assess the response of aerosols, EVI, soil moisture, and wind speed to intensifying precedent  
541 fires, we examine 30-day average DOD, 30-day average and minimum EVI and 30-day  
542 minimum soil moisture, as well as and 30-day upper decile wind speed, either in their probability  
543 distribution or actual value, as a function of active fires during the antecedent week. In the  
544 probability distribution analysis, the long-term percentiles of each variable at  $0.1^\circ$  for a specific  
545 30-day period of a year is constructed from 30-day running averages during  $\pm 30$  days,  
546 respectively, of the center date of the year. Trend analysis of the occurrence and intensity of  
547 post-fire dust events, active fire accounts, and soil moisture involves the application of the  
548 Mann-Kendall nonparametric test for monotonic trend and Theil-Sen robust estimate of linear  
549 trend<sup>76</sup>. A monotonic upward (downward) trend means that the variable consistently increases  
550 (decreases) through time, but the trend may or may not be linear; thus the Mann-Kendall test is a  
551 more general approach than linear regression-based tests for identifying any upward or  
552 downward trend.

553

#### 554 **Data Availability**

555 The datasets for conducting the analysis presented here are all publicly available, including: the  
556 MODIS Collection 6 Active Fire Detections (MCD14ML) acquired from  
557 NASA Fire Information for Research Management System (<https://earthdata.nasa.gov/firms>); the  
558 MODIS Deep Blue aerosol products acquired from the Level-1 and Atmosphere Archive and  
559 Distribution System (LAADS) Distributed Active Archive Center (DAAC)  
560 (<https://ladsweb.modaps.eosdis.nasa.gov/>); the MISR aerosol products acquired from the NASA  
561 Langley Research Center Atmospheric Science Data Center ([https://l0dup05.larc.nasa.gov/cgi-  
562 bin/MISR/main.cgi](https://l0dup05.larc.nasa.gov/cgi-bin/MISR/main.cgi)); the AERONET coarse-mode aerosol optical depth data downloaded  
563 from <https://aeronet.gsfc.nasa.gov>; the ESACCI soil moisture data download from [https://www.esa-  
564 soilmoisture-cci.org/node/238](https://www.esa-soilmoisture-cci.org/node/238); the ERA-5 hourly climate data provided by ECMWF  
565 (<https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era5>); the MODIS MCD12Q1v006  
566 Landcover Type 1 product (<https://lpdaac.usgs.gov/products/mcd12q1v006/>); and the MODIS L3  
567 EVI (MOD13C1 and MYD13C1) from DAAC (<https://lpdaac.usgs.gov/products/mod13c1v006/>).  
568 We generate a list of all identified dust emission cases following large fires available at  
569 <https://doi.org/10.6084/m9.figshare.20648055><sup>77</sup>.

570

## 571 **Code Availability**

572 The code to carry out the current analyses is available from the corresponding author upon  
573 request.

574

## 575 **References**

- 576 1. Bowman, D. M. J. S. *et al.* Fire in the earth system. *Science (80-. )*. **324**, 481–484 (2009).
- 577 2. Bowman, D. M. J. S. *et al.* Human exposure and sensitivity to globally extreme wildfire  
578 events. *Nat. Ecol. Evol.* **1**, 1–6 (2017).
- 579 3. Hamilton, D. S. *et al.* Earth, Wind, Fire, and Pollution: Aerosol Nutrient Sources and  
580 Impacts on Ocean Biogeochemistry. *Ann. Rev. Mar. Sci.* **14**, 303–330 (2022).
- 581 4. Barkley, A. E. *et al.* African biomass burning is a substantial source of phosphorus  
582 deposition to the Amazon, Tropical Atlantic Ocean, and Southern Ocean. *Proc. Natl. Acad.  
583 Sci. U. S. A.* **116**, 16216–16221 (2019).
- 584 5. Schlosser, J. S. *et al.* Analysis of aerosol composition data for western United States  
585 wildfires between 2005 and 2015: Dust emissions, chloride depletion, and most enhanced  
586 aerosol constituents. *J. Geophys. Res. Atmos.* **122**, 8951–8966 (2017).
- 587 6. Wagner, R., Schepanski, K. & Klose, M. The Dust Emission Potential of Agricultural-  
588 Like Fires — Theoretical Estimates From Two Conceptually Different Dust Emission  
589 Parameterizations. *J. Geophys. Res. Atmos.* **126**, e2020JD034355 (2017).
- 590 7. Ichoku, C. *et al.* Biomass burning, land-cover change, and the hydrological cycle in  
591 Northern sub-Saharan Africa. *Environ. Res. Lett.* **11**, (2016).
- 592 8. Bowman, D. M. J. S. *et al.* Vegetation fires in the Anthropocene. *Nat. Rev. Earth Environ.*  
593 **1**, 500–515 (2020).
- 594 9. Duniway, M. C. *et al.* Wind erosion and dust from US drylands: a review of causes,  
595 consequences, and solutions in a changing world. *Ecosphere* **10**, (2019).
- 596 10. Okin, G. S., Gillette, D. A. & Herrick, J. E. Multi-scale controls on and consequences of  
597 aeolian processes in landscape change in arid and semi-arid environments. *J. Arid Environ.*  
598 **65**, 253–275 (2006).
- 599 11. Raupach, M. R. Drag and drag partition on rough surfaces. *Boundary-Layer Meteorol.* **60**,  
600 375–395 (1992).

- 601 12. Webb, N. P. *et al.* Vegetation Canopy Gap Size and Height: Critical Indicators for Wind  
602 Erosion Monitoring and Management. *Rangel. Ecol. Manag.* **76**, 78–83 (2021).
- 603 13. Ellis, T. M., Bowman, D. M. J. S., Jain, P., Flannigan, M. D. & Williamson, G. J. Global  
604 increase in wildfire risk due to climate-driven declines in fuel moisture. *Glob. Chang. Biol.*  
605 **28**, 1544–1559 (2022).
- 606 14. Ravi, S. *et al.* Aeolian processes and the biosphere. *Rev. Geophys.* **49**, 1–45 (2011).
- 607 15. Wagenbrenner, N. S., Germino, M. J., Lamb, B. K., Robichaud, P. R. & Foltz, R. B. Wind  
608 erosion from a sagebrush steppe burned by wildfire: Measurements of PM10 and total  
609 horizontal sediment flux. *Aeolian Res.* **10**, 25–36 (2013).
- 610 16. Wagenbrenner, N. S. A large source of dust missing in Particulate Matter emission  
611 inventories? Wind erosion of post-fire landscapes. *Elem Sci Anth* **5**, (2017).
- 612 17. Jeanneau, A. C., Ostendorf, B. & Herrmann, T. Relative spatial differences in sediment  
613 transport in fire-affected agricultural landscapes: A field study. *Aeolian Res.* **39**, 13–22  
614 (2019).
- 615 18. Deb, P. *et al.* Causes of the Widespread 2019–2020 Australian Bushfire Season. *Earth's*  
616 *Futur.* **8**, (2020).
- 617 19. Nogrady, B. & Nicky, B. the Climate Link To Australia's Fires. *Nature* **577**, (2020).
- 618 20. Yu, Y. & Ginoux, P. Assessing the contribution of the ENSO and MJO to Australian dust  
619 activity based on satellite- And ground-based observations. *Atmos. Chem. Phys.* **21**, 8511–  
620 8530 (2021).
- 621 21. Ginoux, P., Prospero, J. M., Gill, T. E., Hsu, N. C. & Zhao, M. Global-scale attribution of  
622 anthropogenic and natural dust sources and their emission rates based on MODIS Deep  
623 Blue aerosol products. *Rev. Geophys.* **50**, RG3005 (2012).
- 624 22. Yu, Y., Kalashnikova, O. V., Garay, M. J., Lee, H. & Notaro, M. Identification and  
625 Characterization of Dust Source Regions Across North Africa and the Middle East Using  
626 MISR Satellite Observations. *Geophys. Res. Lett.* **45**, 6690–6701 (2018).
- 627 23. Brianne, P., Rebecca, H. & David, L. The fate of biological soil crusts after fire: A meta-  
628 analysis. *Glob. Ecol. Conserv.* **24**, e01380 (2020).
- 629 24. Rodriguez-Caballero, E. *et al.* Global cycling and climate effects of aeolian dust  
630 controlled by biological soil crusts. *Nat. Geosci.* (2022). doi:10.1038/s41561-022-00942-1
- 631 25. Goudie, A. S. & Middleton, N. J. *Desert dust in the global system.* (Springer, 2006).

- 632 26. Ginoux, P. Atmospheric chemistry: Warming or cooling dust? *Nature Geoscience* **10**,  
633 246–247 (2017).
- 634 27. DeMott, P. J. *et al.* Predicting global atmospheric ice nuclei distributions and their impacts  
635 on climate. *Proc. Natl. Acad. Sci.* **107**, 11217–11222 (2010).
- 636 28. Yu, H. *et al.* The fertilizing role of African dust in the Amazon rainforest: A first  
637 multiyear assessment based on data from Cloud-Aerosol Lidar and Infrared Pathfinder  
638 Satellite Observations. *Geophys. Res. Lett.* **42**, 1984–1991 (2015).
- 639 29. Tang, W. *et al.* Widespread phytoplankton blooms triggered by 2019–2020 Australian  
640 wildfires. *Nature* **597**, 370–375 (2021).
- 641 30. Sarangi, C. *et al.* Dust dominates high-altitude snow darkening and melt over high-  
642 mountain Asia. *Nat. Clim. Chang.* **10**, 1045–1051 (2020).
- 643 31. Cook, B. I. *et al.* Twenty-First Century Drought Projections in the CMIP6 Forcing  
644 Scenarios. *Earth's Futur.* **8**, 1–20 (2020).
- 645 32. Zheng, B. *et al.* Increasing forest fire emissions despite the decline in global burned area.  
646 *Sci. Adv.* **7**, (2021).
- 647 33. Abatzoglou, J. T. & Williams, A. P. Impact of anthropogenic climate change on wildfire  
648 across western US forests. *Proc. Natl. Acad. Sci. U. S. A.* **113**, 11770–11775 (2016).
- 649 34. Abram, N. J. *et al.* Connections of climate change and variability to large and extreme  
650 forest fires in southeast Australia. *Commun. Earth Environ.* **2**, (2021).
- 651 35. Yu, Y. *et al.* Machine learning–based observation-constrained projections reveal elevated  
652 global socioeconomic risks from wildfire. *Nat. Commun.* **13**, 1–11 (2022).
- 653 36. Pu, B. & Ginoux, P. How reliable are CMIP5 models in simulating dust optical depth ?  
654 *Atmos. Chem. Phys.* **18**, 12491–12510 (2018).
- 655 37. Pu, B. & Ginoux, P. Climatic factors contributing to long-term variations in surface fine  
656 dust concentration in the United States. *Atmos. Chem. Phys.* **18**, 4201–4215 (2018).
- 657 38. Bodí, M. B. *et al.* Wildland fire ash: Production, composition and eco-hydro-geomorphic  
658 effects. *Earth-Science Rev.* **130**, 103–127 (2014).
- 659 39. UCAR/NCAR/CISL/TDD. The NCAR Command Language (Version 6.6.2) [Software].  
660 (2019). doi:<http://dx.doi.org/10.5065/D6WD3XH5>
- 661 40. Giglio, L., Schroeder, W. & Justice, C. O. The collection 6 MODIS active fire detection  
662 algorithm and fire products. *Remote Sens. Environ.* **178**, 31–41 (2016).



- 663 41. Ramo, R. *et al.* African burned area and fire carbon emissions are strongly impacted by  
664 small fires undetected by coarse resolution satellite data. *Proc. Natl. Acad. Sci. U. S. A.*  
665 **118**, 1–7 (2021).
- 666 42. Diner, D. J. *et al.* Multi-angle Imaging SpectroRadiometer (MISR) instrument description  
667 and experiment overview. in *IEEE Transactions on Geoscience and Remote Sensing* **36**,  
668 1072–1087 (1998).
- 669 43. Pu, B. *et al.* Retrieving the global distribution of the threshold of wind erosion from  
670 satellite data and implementing it into the Geophysical Fluid Dynamics Laboratory land-  
671 atmosphere model (GFDL AM4.0/LM4.0). *Atmos. Chem. Phys.* **20**, 55–81 (2020).
- 672 44. Sayer, A. M., Hsu, N. C., Bettenhausen, C. & Jeong, M. J. Validation and uncertainty  
673 estimates for MODIS Collection 6 ‘deep Blue’ aerosol data. *J. Geophys. Res. Atmos.* **118**,  
674 7864–7872 (2013).
- 675 45. Hsu, N. C. *et al.* Enhanced Deep Blue aerosol retrieval algorithm: The second generation.  
676 *J. Geophys. Res. Atmos.* **118**, 9296–9315 (2013).
- 677 46. Ginoux, P., Garbuzov, D. & Hsu, N. C. Identification of anthropogenic and natural dust  
678 sources using moderate resolution imaging spectroradiometer (MODIS) deep blue level 2  
679 data. *J. Geophys. Res.* **115**, 1–10 (2010).
- 680 47. Eck, T. F. *et al.* Wavelength dependence of the optical depth of biomass burning, urban,  
681 and desert dust aerosols. *J. Geophys. Res. Atmos.* **104**, 31333–31349 (1999).
- 682 48. Anderson, T. L. *et al.* Testing the MODIS satellite retrieval of aerosol fine-mode fraction.  
683 *J. Geophys. Res.* **110**, 1–16 (2005).
- 684 49. Baddock, M. C., Bullard, J. E. & Bryant, R. G. Dust source identification using MODIS:  
685 A comparison of techniques applied to the Lake Eyre Basin, Australia. *Remote Sens.*  
686 *Environ.* **113**, 1511–1528 (2009).
- 687 50. Baddock, M. C., Ginoux, P., Bullard, J. E. & Gill, T. E. Do MODIS-defined dust sources  
688 have a geomorphological signature? *Geophys. Res. Lett.* **43**, 2606–2613 (2016).
- 689 51. Pu, B. & Ginoux, P. Projection of American dustiness in the late 21st century due to  
690 climate change. *Sci. Rep.* **7**, 1–10 (2017).
- 691 52. Pu, B., Ginoux, P., Kapnick, S. B. & Yang, X. Seasonal Prediction Potential for  
692 Springtime Dustiness in the United States. *Geophys. Res. Lett.* **46**, 9163–9173 (2019).
- 693 53. Garay, M. J. *et al.* Introducing the 4.4 km spatial resolution Multi-Angle Imaging

- 694 SpectroRadiometer (MISR) aerosol product. *Atmos. Meas. Tech.* **13**, 593–628 (2020).
- 695 54. Kalashnikova, O. V., Kahn, R., Sokolik, I. N. & Li, W.-H. Ability of multiangle remote  
696 sensing observations to identify and distinguish mineral dust types: Optical models and  
697 retrievals of optically thick plumes. *J. Geophys. Res.* **110**, D18S14 (2005).
- 698 55. Yu, Y. *et al.* Assessing temporal and spatial variations in atmospheric dust over Saudi  
699 Arabia through satellite, radiometric, and station data. *J. Geophys. Res. Atmos.* **118**,  
700 13253–13264 (2013).
- 701 56. Yu, Y., Notaro, M., Kalashnikova, O. V. & Garay, M. J. Climatology of summer Shamal  
702 wind in the Middle East. *J. Geophys. Res. Atmos.* **121**, 289–305 (2016).
- 703 57. Yu, Y. *et al.* Disproving the Bodélé Depression as the Primary Source of Dust Fertilizing  
704 the Amazon Rainforest. *Geophys. Res. Lett.* **47**, 1–12 (2020).
- 705 58. Giles, D. M. *et al.* Advancements in the Aerosol Robotic Network (AERONET) Version 3  
706 database - Automated near-real-time quality control algorithm with improved cloud  
707 screening for Sun photometer aerosol optical depth (AOD) measurements. *Atmos. Meas.*  
708 *Tech.* **12**, 169–209 (2019).
- 709 59. O’Neill, N. T., Eck, T. F., Smirnov, A., Holben, B. N. & Thulasiraman, S. Spectral  
710 discrimination of coarse and fine mode optical depth. *J. Geophys. Res. Atmos.* **108**, 1–15  
711 (2003).
- 712 60. Winker, D. M. *et al.* Overview of the CALIPSO mission and CALIOP data processing  
713 algorithms. *J. Atmos. Ocean. Technol.* **26**, 2310–2323 (2009).
- 714 61. Esselborn, M. *et al.* Spatial distribution and optical properties of Saharan dust observed by  
715 airborne high spectral resolution lidar during SAMUM 2006. *Tellus, Ser. B Chem. Phys.*  
716 *Meteorol.* **61**, 131–143 (2009).
- 717 62. Kim, M. H. *et al.* The CALIPSO version 4 automated aerosol classification and lidar ratio  
718 selection algorithm. *Atmos. Meas. Tech.* **11**, 6107–6135 (2018).
- 719 63. Didan, K., Munoz, A. B., Solano, R. & Huete, A. *MODIS Vegetation Index User ’s Guide*  
720 *(Collection 6)*. (2015).
- 721 64. Seddon, A. W. R., Macias-Fauria, M., Long, P. R., Benz, D. & Willis, K. J. Sensitivity of  
722 global terrestrial ecosystems to climate variability. *Nature* **531**, 229–232 (2016).
- 723 65. Saleska, S. R. *et al.* Dry-season greening of Amazon forests. *Nature* **531**, E4–E5 (2016).
- 724 66. Remer, L. A., Kaufman, Y. J., Holben, B. N., Thompson, A. M. & McNamara, D.

- 725 Biomass burning aerosol size distribution and modeled optical properties. *J. Geophys. Res.*  
726 *Atmos.* **103**, 31879–31891 (1998).
- 727 67. Tegen, I. & Lacis, A. A. Modeling of particle size distribution and its influence on the  
728 radiative properties of mineral dust aerosol. *J. Geophys. Res. Atmos.* **101**, 19237–19244  
729 (1996).
- 730 68. Friedl, M. A. & Sulla-Menashe, D. *User Guide to Collection 6 MODIS Land Cover*  
731 *(MCD12Q1 and MCD12C1) Product.* **6**, (2018).
- 732 69. Sulla-Menashe, D., Gray, J. M., Abercrombie, S. P. & Friedl, M. A. Hierarchical mapping  
733 of annual global land cover 2001 to present: The MODIS Collection 6 Land Cover  
734 product. *Remote Sens. Environ.* **222**, 183–194 (2019).
- 735 70. Dorigo, W. *et al.* ESA CCI Soil Moisture for improved Earth system understanding: State-  
736 of-the art and future directions. *Remote Sens. Environ.* **203**, 185–215 (2017).
- 737 71. Hersbach, H. *et al.* The ERA5 global reanalysis. *Q. J. R. Meteorol. Soc.* **146**, 1999–2049  
738 (2020).
- 739 72. Preimesberger, W., Scanlon, T., Su, C.-H., Gruber, A. & Dorigo, W. Homogenization of  
740 Structural Breaks in the Global ESA CCI Soil Moisture Multisatellite Climate Data  
741 Record. *IEEE Trans. Geosci. Remote Sens.* **59**, 2845–2862 (2021).
- 742 73. Minola, L. *et al.* Near-surface mean and gust wind speeds in ERA5 across Sweden:  
743 towards an improved gust parametrization. *Clim. Dyn.* **55**, 887–907 (2020).
- 744 74. Molina, M. O., Gutiérrez, C. & Sánchez, E. Comparison of ERA5 surface wind speed  
745 climatologies over Europe with observations from the HadISD dataset. *Int. J. Climatol.* **41**,  
746 4864–4878 (2021).
- 747 75. Klose, M. *et al.* Mineral dust cycle in the Multiscale Online Nonhydrostatic Atmosphere  
748 Chemistry model (MONARCH) version 2.0. *Geosci. Model Dev.* **14**, 6403–6444 (2021).
- 749 76. Mondal, A., Kundu, S. & Mukhopadhyay, A. Rainfall Trend Analysis By Mann-Kendall  
750 Test: a Case Study of North-Eastern Part of Cuttack District, Orissa. *Int. J. Geol. Earth*  
751 *Environ. Sci.* **2**, 2277–208170 (2012).
- 752 77. Yu, Y. & Ginoux, P. Dust emission following large wildfires. *Figshare* (2022).  
753 doi:10.6084/m9.figshare.20648055.v2  
754