

1 **A multi-scale model analysis of ozone formation in the Bangkok Metropolitan** 2 **Region, Thailand**

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10 **Abstract**

11 Over the last three decades, Thailand's rapid industrialization and urbanization have led to
12 an impact on urban air quality. A majority of the country's development has occurred within and
13 around Bangkok (BKK), the capital city of Thailand, and the Bangkok Metropolitan Region
14 (BMR). Since 1995, the BMR has experienced air quality degradation, in particular, enhanced
15 ozone (O₃) due to a combination of the local increase in emissions from accelerated growth in
16 automotive and industrial activities, local meteorology including strong solar radiation, high
17 temperature and high humidity, and potential long-range effects of regional transport from China.
18 To investigate the O₃ formation in the BMR due to the effects of long-range transport and local
19 meteorology feedbacks, we perform a multi-scale simulation with the Weather Research and
20 Forecasting model with Chemistry (WRF-Chem) during the O₃ season (January to March), 2010;
21 since O₃ mixing ratio exceedances in the BMR occur primarily during this period. The results in
22 this study indicate the significance of China's emission reductions on the regional-scale and the
23 local-scale pollutions, as far as the BMR region and southern Thailand. Applying China's oxide
24 of nitrogen (NO_x)-only emission controls, generally, enhance the domain-wide monthly-average
25 peroxyacetyl nitrate (PAN) and O₃ in the regional scale, in the order of ~1 to 7% and ~1 to 5%,
26 respectively, while those in the local scale are ~0.2 to 6% and ~0.1 to 5% compared with the
27 baseline simulation. However, the increases in PAN and O₃ are mitigated by 40% China's Volatile
28 Organic Compound (VOC) reduction along with 40% NO_x reduction. The results, supported by an
29 indicator analysis, suggest that northern and eastern China, northern and central Thailand and the
30 BMR, are likely VOC-limited during the O₃ season. Since the BMR is VOC-limited regime,
31 controlling anthropogenic VOC emissions will show more benefit to control O₃ than controlling
32 NO_x-only emissions. Other factors that influence on O₃ levels in the BMR are biogenic VOC
33 emissions from the Tenasserim range and land- and sea-breeze circulations that recirculate and
34 disperse pollutants along the coastal areas.

35 **Keywords:** Long-range transport; Surface ozone; WRF-Chem; VOC-limited; NO_x-limited

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37

38 1. Introduction

39 The Bangkok Metropolitan Region (BMR) has experienced air quality degradation over
40 the past few decades, due to rapid economic growth, and an increase in pollutant emissions (Bucher
41 et al., 2011; Chueinta et al., 2000; Oanh et al., 2006; Oanh and Zhang, 2004; Ruchirawat et al.,
42 2002; Uttamang et al., 2018). These enhanced emissions are harmful to human health, and the
43 environment (Bucher et al., 2011; Oanh et al., 2006; Ruchirawat et al., 2002; Uttamang et al.,
44 2018). Pollution Control Department (PCD), Thailand (2012) reported that, in Thailand, annual
45 average mixing ratios of O₃ during 2005 to 2012 (17 years) had been increasing continuously due
46 to long-range transport and increases in volatile organic compounds (VOC) emissions. During
47 2012, more than 80 % of the monitoring stations located in Thailand reported exceeded O₃ mixing
48 ratios which very high mixing ratios of O₃ were observed in Bangkok and its vicinities (PCD,
49 2012). Our previous study on air quality in the BMR (Uttamang et al., 2018) showed that during
50 2010 to 2014, among other gaseous criteria pollutants, ozone (O₃) was the only gaseous species
51 whose mixing ratios frequently exceeded Thailand National Ambient Air Quality standard
52 (NAAQs) for O₃ (Thailand NAAQs for hourly O₃ is 100 ppb [PCD, 2018]). Moreover, the annual
53 average O₃ mixing ratios had been increasing during 2010 to 2014 (~16.5 ppb in 2010 and ~20.0
54 ppb in 2014). The hourly O₃ exceedances in the BMR were frequently observed during the dry
55 season (local summer starts from February to May and local winter starts from October to
56 February), especially in January to March when Northeast monsoon winds were the predominant
57 wind direction. Those values were less observed during the transitional period between wet and
58 dry seasons (i.e. during May) and rarely occurred during wet season (May to October) (Uttamang
59 et al., 2018). Annual average mixing ratios of CO, SO₂, O₃ and NO₂ during 2003 to 2014 in the
60 BMR are shown in table S1, supplement material.

61 East Asia's pollutant emissions play an important role in regional and global air quality. In
62 2010, emissions of oxide of nitrogen (NO_x = NO + NO₂), sulfur dioxide (SO₂), coarse mode
63 particulate matter (PM₁₀), fine mode particulate matter (PM_{2.5}), and non-methane volatile organic
64 compounds (NMVOC) over East Asia were 29.7 Mt, 29.5 Mt, 16.8 Mt, 12.5 Mt, and 25.9 Mt,
65 respectively (Wang et al., 2014b). Compared to the total emissions in East Asia, China contributes
66 ~ 82 to 88 % of NO_x, ~ 94 % of SO₂ and ~ 84 to 88 % of PM₁₀ and PM_{2.5} (Wang et al., 2014b).
67 Several emission control strategies have been suggested to reduce current and future pollution in
68 China (Guo et al., 2014; Liang et al., 2017; Shao et al., 2009; Sun et al., 2018; Xue et al., 2014);
69 however, the reduction targets for major air pollutants are not clear (Zhang et al., 2017, 2014). A
70 10% NO_x emission reduction from 2010 levels was proposed as the national reduction target
71 during China's 12th Five Year Plan (FYP) (2011 to 2015) (Wang et al., 2014a). By applying this
72 strategy, the hourly maximum O₃ mixing ratios in China were reduced ~ 2 % (Wang et al., 2014a).
73 Different emission control strategies were proposed in Shanghai, Jiangsu and Zhejiang. The
74 proposed anthropogenic NO_x emission reductions during the 12th FYP were by 17.5 %, 17.5 %
75 and 18 %, respectively (Zhang et al., 2017). According to the ambient air quality guidelines set by
76 the World Health Organization (WHO), China needs to reduce 60 % of SO₂ emissions along with
77 40 % of NO_x and volatile organic compounds (VOC), and 50 % of PM₁₀ emissions from 2005, in
78 order to attain the guidelines (Wang and Hao, 2012).

79 The effects of long-range transport of air pollutants originating in China on ambient air

80 pollution in Asian countries have been reported in several studies. Kim et al. (2009) reported that
81 the increases in PM₁₀, PM_{2.5} and ionic components of particulate matters over a remote area in
82 Hong Kong was influenced by air masses originating from northern and eastern China. Lee et al.
83 (2013) and Oh et al. (2015) reported that high PM₁₀ levels over Seoul, Korea were caused, in part,
84 from a combination of transboundary pollutants from northern and eastern China and local
85 meteorology (high-pressure system and atmospheric circulation) over Korea. Cuesta et al. (2018)
86 reported that O₃ pollution plumes transported by an anticyclonic circulation from the North China
87 Plain to northern China, Korea and Japan, resulted in elevated O₃ levels over Japan.

88 In this paper, we examined O₃ behavior and regional processes that contribute to elevated
89 O₃ levels in the BMR; in particular, the role of NO_x and VOC emissions in China and the
90 subsequent transport of trace gases and aerosols to the BMR as elevated O₃ mixing ratios are
91 frequently observed during dry seasons when there is a predominant Northeast monsoon wind
92 direction. The effect of China's emission control strategies on O₃ levels in the BMR was
93 investigated using the Weather Research and Forecasting model version 3.9.1 coupled with
94 chemistry (WRF-Chem). We first evaluated the performance of the WRF model to simulate
95 meteorology and O₃ mixing ratios in the BMR. Then we examined the impacts of NO_x and VOCs
96 emission reductions in China on transboundary O₃ precursors transport, and the resulting effect on
97 O₃ levels in the BMR region.

98 **2. Model configuration and simulation design**

99 **2.1 WRF-Chem model configuration**

100 WRF-Chem is a fully coupled online chemical and meteorological model (Grell et al.,
101 2005), which has been widely used for regional-scale air quality studies (Guo et al., 2016; Jose et
102 al., 2017; Karagulian et al., 2019; Liua et al., 2016; Podrascanin, 2019; Sharma et al., 2017;
103 Tessum et al., 2015; Thompson et al., 2008; Yahya et al., 2017).

104 Figure 1 illustrates a triple-nested domain using in this study and 36- (the outermost domain
105 [d01]), 12- (the second domain [d02]) and 4-km (the innermost domain [d03]) horizontal
106 resolutions, which results in 98 × 150, 88 × 106 and 97 × 97 east-west and north-south grid points,
107 respectively. The model vertical resolution consists of 33 sigma-pressure vertical layers. D01
108 covers Thailand and most part of the region of China, where d02 and d03 scale down to the region
109 of Thailand and the BMR, respectively.

110 Table 1 summarizes the physics and chemistry options used in this study. Here we use the
111 Thompson scheme for grid-scale microphysics (Thompson et al., 2008), the Grell-Freitas cumulus
112 parameterization (Grell and Freitas, 2014), the Rapid Radiative Transfer Model (RRTMG) for the
113 longwave and shortwave radiation (Iacono et al., 2008), the Yonsei University planetary boundary
114 layer (YSU-PBL) scheme (Hong et al., 2006), the MM5 similarity surface layer scheme (Fairall et
115 al., 2003), and the Noah Land-Surface model (Noah-LSM) (Chen and Dudhia, 2001) with the
116 united States Geological Survey (USGS) landuse dataset.

117 The Regional Acid Deposition Model version 2 (RADM2) couples with Model Aerosol
118 Dynamics Model for Europe/Secondary Organic Aerosol Model (MADE/SORGAM) is used to
119 simulate the gas-phase mechanism, and aerosol physics and chemistry, respectively. The RADM2

120 is one of the most computationally efficient mechanisms to calculate O₃ and photochemical
121 reactions of gaseous species (Zimmermann and Poppe, 1996). Three particle modes (Aitken,
122 accumulation and coarse mode) and three major particle dynamics (nucleation, condensational
123 growth and coagulation) are simulated using the MADE/SORGRAM aerosol module by assuming
124 the particle size distribution is log-normal (Yang et al., 2018). The Madronich F-TUV is selected
125 for photolysis scheme. The chemistry is configured to run with dry deposition of gas species and
126 aerosols, subgrid convective wet scavenging and aqueous chemistry, vertical turbulent mixing, and
127 direct and indirect feedbacks from the aerosols to the radiation schemes.

128 The total simulation period is 18 December 2009 to 31 March 2010 (which includes an O₃
129 episode which occurred during 5 to 6 March 2010). The first two weeks (18 to 31 December 2009)
130 used as model spin-up (excluded from model analysis), and a 10-day meteorological re-
131 initialization strategy with continuous chemistry.

132 **2.2 Input data**

133 The meteorological simulation is driven by the National Centers for Environmental
134 Prediction-Final Operational Global Analysis data (NCEP-FNL) with 1° × 1° spatial resolution
135 prepared operationally every six hours. Amnuaylojaroen et al. (2014) studied the effects of
136 different emission inventories including the Reanalysis of the TROpospheric chemical
137 composition (RETRO), the Intercontinental Chemical Transport Experiment-Phase B (INTEX-B),
138 the MACCity emissions (adapted from the Monitoring Atmospheric Composition and Climate and
139 megacity Zoom for the Environment projects), the Southeast Asia Composition, Cloud, Climate
140 Coupling Regional Study (SEAC4RS) emissions, and a combination of MACCity and SEAC4RS
141 emissions on simulated O₃ by using WRF-Chem in Southeast Asia. Their results showed that the
142 difference among those emission inventories might reach ~ 30%, and the WRF-Chem model
143 generated only a slight variability of O₃ mixing ratios (~8 %). Lamarque et al. (2010) reported
144 that the global total amount of anthropogenic emissions from RETRO was similar to those from
145 the Emissions Database for Global Atmospheric Research emission inventory (EDGAR),
146 eventhough specific sectors or regions in those inventories were different.

147 In this study, monthly average Emissions Database for Global Atmospheric Research-
148 Hemispheric Transport of Air Pollution (EDGAR-HTAP) for the year 2010 with the finest
149 horizontal resolution (0.1° × 0.1° spatial resolution) is selected (Janssens-Maenhout et al., 2012)
150 (http://edgar.jrc.ec.europa.eu/htap_v2/). This inventory contains CO, NO_x, SO₂, NMVOC,
151 ammonia (NH₃), methane (CH₄), PM₁₀, PM_{2.5}, organic compound (OC) and black carbon (BC)
152 emissions from power, industrial, residential, agriculture, and transport (ground, air and shipping)
153 sectors. The EDGAR-HTAP emissions were spatially processed for the specific trace gas and
154 aerosol emission species needed for RADM2-MADE/SORGAM using the PREP-CHEM-Sources
155 preprocessor, version 1.5 (Oliveira et al., 2016), and then were subsequently processed to be
156 model-ready using the WRF-Chem “convert_emiss” program. Biogenic emissions data are
157 calculated inline in WRF-Chem using the Model of Emissions of Gases and Aerosols from Nature
158 version 2 (MEGAN) (Guenther et al., 2006).

159 Chemical initial and boundary conditions (ICs/BCs) are prepared using the Model for
160 Ozone and Related Chemical Tracers, version 4, which is driven by meteorological fields from the

161 Goddard Earth Observing System Model, Version 5 (MOZART-4/GEOS-5) with $1.9^\circ \times 2.5^\circ$
 162 horizontal resolution, and 56 vertical layers (Emmons et al., 2010).

163 **2.3 Simulation design**

164 A baseline simulation is set following the model configurations in section 2.1. The results
 165 from the baseline simulation are used for model evaluation. Long-range transport of O₃ and its
 166 precursors (NO_x and VOC) on O₃ behavior in the BMR is examined by performing a sensitivity
 167 analysis of NO_x and VOC reductions in China's emissions, since China is a major contributor of
 168 pollutants in Asia. Four emission control strategies for China including 10 %, 20 %, and 40 % NO_x
 169 emission reductions, and 40 % NO_x emission reduction along with 40 % VOC emission reduction
 170 over China's region are examined for the sensitivity analysis (Table 2). The 10 % NO_x emission
 171 reduction is set as Strategy 1, regarding the national reduction target during China's 12th FYP for
 172 NO_x. The other two NO_x-only control strategies (20 % and 40 % NO_x emission reductions) are set
 173 as Strategy 2 and 3, respectively, to examine the responses of O₃ and its precursors in the BMR
 174 due to different China's NO_x emission reductions. The 40 % NO_x emission along with 40 % VOC
 175 emission reduction proposed by Wang and Hao (2012) is set as Strategy 4, in order to investigate
 176 implications for incorporating VOC emission reduction strategies in China.

177 **2.4 Model evaluation protocol**

178 Several discrete statistics, including the correlation coefficient (r), mean bias (MB), root
 179 mean square error (RMSE), and normalized mean bias (NMB) are used to evaluate the model
 180 baseline simulation for temperature (T), wind speed (WS), relative humidity (RH), pressure (P),
 181 and O₃ mixing ratio during daytime hours (10:00 to 17:00 local time [LT]). This time window was
 182 selected to avoid pollution accumulation by surface inversions (Pochanart et al., 2001; Uttamang
 183 et al., 2018). Hourly meteorology and daily maximum 8-hour-average O₃ mixing ratios (Max 8-h
 184 O₃) from 16 Pollution Control Department (PCD) Thailand monitoring stations located in and
 185 around the BMR region (Fig. 2) were compared with the simulated variables from the lowest model
 186 layer of d03. Details of the PCD monitoring stations equipment are provided in (Uttamang et al.,
 187 2018). A mean normalized bias error (MNBE) was applied instead of NMB for the O₃ mixing
 188 ratios evaluation, as the U.S. EPA suggests a MNBE criteria of ± 10 to 15 % for regulatory
 189 modeling applications (Žabkar et al., 2013).

$$190 \quad \text{MNBE} = \frac{1}{N} \sum \left(\frac{M_i - O_i}{O_i} \right) \times 100 \%$$

191 Where M_i = modeled value i , O_i = observed value i , N = number of paired observed-
 192 modeled values.

193 The comparison between the simulations and the observations are provided using geo-
 194 reference information (latitude and longitude) of the monitoring stations.

195

196 3. Results of model evaluation

197 3.1 Model evaluation during the 2010 dry season

198 Table 3 summarizes the average January to March 2010 model evaluations of meteorology
 199 and Max 8-h O₃ mixing ratios. The comparison between hourly simulated T, RH, P and WS in d03
 200 with the observations from 16 monitoring stations show that the model accurately captures the
 201 variations of T ($\bar{r} = \sim 0.8$), RH ($\bar{r} = \sim 0.7$) and P ($\bar{r} = \sim 0.8$), but less accurately captures the
 202 variations of WS ($\bar{r} = \sim 0.2$). Overall, the model tends to underpredict T, RH, and P with negative
 203 \overline{MB} (T is ~ -2.8 °C, RH is -6.8 % and P is -1.5 hPa). For WS, overall, the model tends to overpredict
 204 WS with \overline{MB} is ~ 1.5 ms⁻¹.

205 The underpredicted T and RH and overpredicted WS in the WRF model has been reported
 206 in other studies (Yerramilli et al., 2010). Yerramilli et al. (2010) also studied the effects of the
 207 planetary boundary layer (PBL) and land surface model (LSM) physics in WRF-Chem simulations
 208 of surface O₃ mixing ratios in the Central Gulf Coast region, southeast US. The results from their
 209 study indicated that a similar combination of the YSU-PBL and Noah-LSM schemes (Table 1)
 210 provided the best simulation of winds, temperature, humidity, mixed layer depth and O₃ mixing
 211 ratios diurnal variations, but the model normally underpredicted O₃ mixing ratios and temperature
 212 and overpredicted winds. WS are typically overpredicted in WRF due to unresolved topographical
 213 features and an underestimated surface drag parameterization in the model at these scales
 214 (Georgiou et al., 2018; Kumar et al., 2016; Yahya et al., 2014).

215 The evaluation of the daily Max 8-h O₃ mixing ratios show that across the different stations
 216 in and around the BMR, the total averages of observed and simulated daily Max 8-h O₃ are $25.8 \pm$
 217 9.8 ppb, and 20.8 ± 14.4 ppb, respectively. The r correlation ranges from -0.3 to 0.6 , MB from $-$
 218 24 to 15 ppb, RMSE from 12 to 57 ppb, and MNBE from -75 to 95 %. Large discrepancies and
 219 poor correlations that affect the statistical ranges are strongly impacted by the sites in coastal areas
 220 (at 19T, 26T, 27T and 34T monitoring station) and 60T monitoring station. Specifically for the
 221 other monitoring stations (19T, 26T, 27T, 34T and 60T are excluded), the model underpredicts
 222 daily Max 8-h O₃ with a \overline{MB} of ~ -2 ppb a \overline{RMSE} of ~ 17.6 ppb and \overline{MNBE} of $\sim -3.2\%$.

223 The baseline evaluation results show that the WRF-Chem model reasonably predicts the
 224 meteorological variables, except for WS, and overpredicts WS for all monitoring stations. The
 225 model performance for daily Max 8-hr O₃ predications near coastal areas are not good, but for
 226 other monitoring stations, the model overall provides good performance (i.e., MNBE $< \pm 10\%$).

227 3.2 Model evaluation during the O₃ event (5 to 6 March 2010) in the BMR

228 On 5 to 6 March 2010, O₃ exceedances were observed at nine monitoring stations located
 229 in the BMR (six monitoring stations are located in Bangkok (BKK) including 3T, 10T, 11T, 15T,
 230 52T and 61T sites, and three monitoring stations are located in the outskirts of BKK including
 231 14T, 22T and 27T site). The exceedances observed across the different monitoring sites ranged in
 232 mixing ratio from 101 to 169 ppb. This period contained the highest hourly O₃ mixing ratios
 233 observed (169 ppb) in 2010.

234 Figure 3 shows the comparison between the hourly observed and simulated O₃ mixing
 235 ratios at the nine BMR monitoring stations before (1 to 4 March 2010) and during the O₃ event (5

236 to 6 March 2010). The model shows an ability to capture the O₃ event and reasonably reproduce
 237 the diurnal variations of O₃ at all the monitoring stations during the event. The hourly simulated
 238 O₃ is frequently underpredicted in the evening and the early morning (19:00 to 7:00 LT), but
 239 overpredicted during afternoon (13:00 to 18:00 LT). Across the sites the correlation coefficients
 240 range from 0.58 to 0.89 (\bar{r} is ~ 0.8), MBs range from -24 to 10.2 ppb (\overline{MB} is ~ -11.2 ppb), RMSEs
 241 range from 28.7 to 41.4 ppb (\overline{RMSE} is ~ 35.5 ppb) and MNBEs range from -62.4 to -19.1 %
 242 (\overline{MNBE} is ~ -44.2 %).

243 **3.3 Spatial distribution of O₃ from baseline simulation**

244 Figure 4 shows spatial distributions of simulated daytime O₃ in d01 (“regional” scale; top
 245 panels) and in d03 (“local” scale; bottom panels) from the baseline simulation. The monthly-
 246 average spatial distributions of O₃ over East Asia and Southeast Asia shows an increase of O₃
 247 mixing ratios from January to March where the domain-wide monthly-average O₃ is 37.9 ± 12.1
 248 ppb in January, 40.3 ± 12.7 ppb in February, 43.3 ± 12.9 ppb in March and the episodic-average
 249 is 43.6 ± 17.8 ppb (Fig. 4 (a)-(d)). Strong gradients of O₃ occurred from western China (the Tibetan
 250 Plateau) to eastern China. Ni et al. (2018) reported a similar pattern of O₃ over China during the
 251 spring 2008, in which the O₃ mixing ratios over the southern Tibetan Plateau ranged from 75 to
 252 80 ppb and those over the North China Plain ranged from 25 to 40 ppb. Their study suggested that
 253 ~ 80 to 90% of O₃ mixing ratios over Tibet (low anthropogenic emissions) were contributed by
 254 biogenic O₃ precursors sources.

255 Our baseline simulation shows that local O₃ (d03) generally increases from January to
 256 March. The domain-wide monthly-average of O₃ are ~ 30 ppb in January, ~ 50 ppb in February, \sim
 257 61 ppb in March and the episodic-average is ~ 67 ppb (Fig 4 (e)-(h)). Very high O₃ mixing ratios
 258 frequently occur over the Tenasserim range (the west of the domain) and hilly regions (the
 259 southeast of the domain).

260 **3.4 Spatial distribution of O₃ precursors**

261 Table S1.1 and S1.2 summarize the domain-wide monthly- and episodic-average mixing
 262 ratios of NO_x, CO, VOC, PAN, nitric acid (HNO₃), and total odd oxygen (where O_x = O₃ + NO₂
 263 + 2 × nitrogen trioxide (NO₃) + HNO₃ + PAN) in the regional (d01), and local (d03) scales.

264 The regional (d01) spatial distributions of NO_x are similar to its emission sources (Fig.
 265 5a)). High mixing ratios of NO_x are present over east China and the central region of Thailand
 266 (including the BMR), while very low mixing ratios are present over remote areas of the region.
 267 Spatial distributions of pollutants are effectively controlled by their lifetimes. Therefore, for a short
 268 lifetime species, for example, NO_x, high mixing ratios are normally confined near the emission
 269 sources (Amnuaylojaroen et al., 2014). The domain-wide monthly-average NO_x shows a decrease
 270 in NO_x from January to March, from ~ 52 ppb in January to ~ 25 ppb in March, but a higher value
 271 during the episodic-average of ~ 44 ppb. There are also high mixing ratios of CO near their source
 272 regions. CO has a longer lifetime (~ 1 month) than NO_x, however, and thus CO can be transported
 273 continentally and contribute to both regional and global burdens (NASA, 2014). CO decreases
 274 from January to March, from ~ 199 ppb in January to ~ 173 ppb in March, but also has a higher
 275 value for the episodic-average of ~ 200 ppb. High mixing ratios of VOC occur near source regions
 276 of both anthropogenic (i.e. east China, the central region of Thailand and eastern India [Fig 5b]))

277 and biogenic sources (i.e. Malaysia and Indonesia). where the total VOCs range from ~8 to ~9 ppb
 278 during January to March. PAN is a nitrogen reservoir species, which may be transported over
 279 distances and release NO_x away from its source (Jacob, 1999). The mixing ratios of PAN in a clean
 280 environment are ~2 to ~100 ppt and in polluted air are ~10 to ~20 ppb (Jacobson, 2012). In our
 281 study, high mixing ratios of PAN occur over Burma, Thailand, Malaysia and Indonesia, and can
 282 range on average from ~0.5 to ~0.7 ppb during January to March (Fig S1.2). HNO₃ is formed from
 283 the reaction of OH and NO₂ during daytime (Jiménez et al., 2012), and our model simulations
 284 suggest high mixing ratios of HNO₃ over east China and Thailand, ranging from ~3 to ~3.5 ppb
 285 during January to March.

286 On the local scale (d03), there are high mixing ratios of CO and NO_x around the BMR. The
 287 domain-wide monthly-average CO, NO_x and O_x mixing ratios decrease from January to March,
 288 and range from ~300 ppb for CO, ~121 ppb for NO_x and ~162 ppb for O_x in January, ~259 ppb
 289 for CO, ~93 ppb for NO_x and ~161 for O_x in February, ~224 ppb for CO, ~74 ppb for NO_x and
 290 ~151 for O_x during March and the episodic-average are ~242 ppb for CO and ~76 ppb for NO_x.

291 There are locally widespread high values of VOC and HNO₃ for all months, while the
 292 highest mixing ratios of PAN are over the Tenasserim range to the west, and over the mountainous
 293 areas on the east and southeast of the BMR. The domain-wide monthly-average VOC, PAN and
 294 HNO₃ mixing ratios increase from January to March, on average ranging from ~28 to ~31 ppb for
 295 VOC, ~0.4 to ~1.2 ppb for PAN and ~10.6 to ~15.3 ppb for HNO₃.

296 Spatial distribution of domain-wide (d03) monthly-average delta NO_x, CO, VOC, PAN and
 297 HNO₃ are shown in supplement material (Fig. S1.1 to S1.4).

298 **4. Results from sensitivity analysis**

299 Figure 5 shows monthly-average NO_x and VOC emissions in January, February and
 300 March and the absolute differences (sensitivity – baseline) of the four NO_x and VOC emission
 301 reduction strategies (i.e 10 %, 20 %, and 40 % NO_x emission reductions, and 40 % NO_x emission
 302 reductions along with 40 % VOC emission reduction). Figure 6 shows monthly average wind
 303 fields in the outermost domain that are generated from the baseline simulation.

304 **4.1 Effects on regional-scale O₃ precursors and O₃**

305 The decreases in NO_x emissions over China (S1, S2, and S3) lead to NO_x mixing ratio
 306 reductions in eastern China during all the months, while there are some localized increases in NO_x
 307 levels over the central areas of Thailand in February and March. Table 4 summarizes delta (delta
 308 of species *i*, $(X_i) = \text{monthly-average } X_{i,\text{reduction_strategies}} - \text{monthly-average } X_{i,\text{baseline}}$) O₃ and its
 309 precursors, including NO_x, CO, VOC, PAN, HNO₃ and O_x in the regional scale (d01) due to
 310 China's emission reductions. There are widespread decreases in domain-wide monthly-average
 311 delta NO_x, on the order of about ~ 2.4 to 20.5 ppb (~ 10 to 40 %) compared with the baseline
 312 simulation . There are minor decreases in CO mixing ratios in eastern China, but slight increases
 313 in central Thailand . The domain-wide monthly-average delta CO are approximately ~ -0.5 to 0.03
 314 ppb (~ -0.1 to 0 %). Moderate VOC reductions mostly occur in eastern China due to S4, which
 315 are approximately ~ 0.2 to ~ 1.0 ppb (~ 7 to 14 %) during January to March (Fig S3.3). There is a
 316 moderate increase in PAN due to decreased NO_x emissions, with increases approximately ~ 0.01

317 to 0.03 ppb (~ 1 to 7 %). PAN increases are mitigated, however when incorporating the 40 % VOC
 318 emission reduction compared to the NO_x-only control strategies, with average changes of
 319 approximately ~ 0.01 ppb (2 %) . HNO₃ mixing ratios are also reduced by ~ 0.1 to ~ 0.5 ppb (~
 320 2% to 14 %) due to the NO_x emission reductions . O_x mixing ratios are also reduced by about ~
 321 0.9 to 7.8 ppb (~ 1.5 to 11.5 %) due to the NO_x emission reductions; however, the O_x decreases (~
 322 0.3 to 0.4 %) are significantly mitigated by including both 40 % NO_x and VOC emission reduction
 323 compared to the NO_x-only strategy . Spatial distribution of domain-wide (d01) monthly-average
 324 delta NO_x, CO, VOC, PAN and HNO₃ are shown in supplement material (Fig. S3.1 to S3.5).

325 The decreases in NO_x emissions over China (S1, S2, and S3) slightly increase the
 326 monthly-average O₃ mixing ratios on the regional scale ~ 0.3 to 2 ppb (~1 % to ~5 %) from the
 327 baseline simulation) perhaps owing to the region being VOC limited (Fig. 7). The spatial
 328 distributions of monthly-average delta O₃ shows moderate increases in O₃ mixing ratios from
 329 eastern China to southeast Asia in NE/SW directions. The domain-wide monthly-average delta
 330 O₃ during January to March are ~ 0.3 to 0.4 ppb (0.7 to 0.9%) in S1, ~ 0.6 to 0.7 ppb (1.3 to 1.9
 331 %) in S2, ~ 1.2 to 1.5 ppb (2.7 to 4.1 %) in S3, and ~ 0.2 to 0.6 ppb (0.4 to 0.5 %) in S4. The
 332 comparison between S3 and S4 shows that the combined 40 % NO_x and VOC emission reduction
 333 acts to mitigate the O₃ increases compared with S3 (NO_x-only reductions), and further
 334 demonstrates the implications for including VOC emission reduction strategies in China. The
 335 changes of O₃ mixing ratio in the regional scale correlate with the changes of PAN mixing ratio
 336 (increases in PAN increases in O₃).

337 The increase in the mixing ratios of O₃ and PAN (in a VOC-limited regime, a reduction
 338 of NO_x lead to an increasing of PAN mixing ratios [Jiménez et al., 2012]) due to the 10 to 40 %
 339 NO_x emission reduction strategies suggest that eastern China and northern Thailand are likely
 340 VOC-limited during January to March, while some areas of central Thailand are VOC-limited
 341 during January to February and become less VOC-limited in March. This is further supported by
 342 the mitigation in O₃ increases when including both 40 % decreases in NO_x and VOC emissions
 343 compared to the NO_x-only strategy. The results in this section also indicate the significance of
 344 China's emission reductions on regional-scale pollution, as far as southern Thailand and the
 345 BMR region.

346 **4.2 Effects on BMR O₃ pollution**

347 It was previously shown that emission reductions affect the regional-scale (d01) average
 348 O₃ and its precursor mixing ratios. In this case, we analyze these effects on the local-scale (d03)
 349 O₃ and precursor mixing ratios in the BMR.

350 Table 5 summarizes delta O₃ and its precursors in the local scale due to China's
 351 emission reductions. The NO_x-only control strategies in China generally lead to decreases in
 352 domain-wide monthly-average delta NO_x in January and February, on the order of ~ 3 to 19 ppb
 353 (~ 1 to 9 %) from the baseline simulation, but moderate increases in the central regions of the
 354 domain (including northwestern BMR) in March, on the order of ~ 22 to 28 ppb (~ 8 to 9 %) .
 355 There are minor increases in CO (< 11 ppb) and VOC (< 1.4 ppb) mixing ratios in January and
 356 February, on the order of < 1.5 %, and moderate increases in those species in the central regions
 357 of the domain in March (~ 36 to 40 ppb [~ 6.0 to 6.5 %] for domain-wide monthly-average delta

358 CO and ~ 4 to 6 ppb [~ 4 to 5 %] for domain-wide monthly-average delta VOC). There are
 359 increases in PAN due to the NO_x -only control strategies, with increases ~ 0.01 to 0.12 ppb (~ 1 to
 360 6 %). There are minor increases in domain-wide monthly-average delta HNO_3 in January, on the
 361 order of ~ 0.03 to 0.14 ppb (~ 0.6 to 3 %), but widespread decreases (~ 0.2 to 0.5 ppb [< 4 %]) in
 362 this species in February. In March, the reduction of HNO_3 mostly occurs in central domain (~ 0.4
 363 to 0.5 ppb [~ 3 to 4 %] for domain-wide monthly-average delta HNO_3). Widespread increase in
 364 O_x occur for all the months, on the order of ~ 1.1 to 6.1 ppb (~ 0.5 to 2 %). The modeling difference
 365 between S3 (40% China's emission reduction) and S4 (NO_x and VOC reduction each by 40%)
 366 shows that O_3 precursors (i.e. CO, HNO_3 , PAN and O_x) mixing ratios are mitigated during January
 367 and February; while in March, the incorporating of VOC reduction (S4) does not reduce the mixing
 368 ratio of the O_3 precursors. Spatial distribution of domain-wide (d03) monthly-average delta NO_x ,
 369 CO, VOC, PAN and HNO_3 are shown in supplement material (Fig. S3.7 to S3.12).

370 The NO_x -only control strategies in China mostly increase the monthly-average O_3 mixing
 371 ratios on the local scale (d03), on the order of 0.1 to 3 ppb (~ 1 to 6 %) from the baseline simulation
 372 (Fig. 8). The distributions of monthly-average delta O_3 shows widespread increase in O_3 mixing
 373 ratios in January and February, but some O_3 reductions occur in the central areas of the domain
 374 (including northern BMR). The domain-wide monthly-average delta O_3 during January to March
 375 are ~ -0.34 to 3.0 ppb (-0.4 to ~ 2 %) in S1, ~ 0.2 to 2.0 ppb (~ 0.2 to 3.5 %) in S2, ~ 1.0 to 3.0
 376 ppb (~ 1 to 6 %) in S3, and ~ 1.0 to 3.0 ppb (~ 1 to 7 %) in S4. The comparison between S3 and S4
 377 shows that the combined 40 % NO_x and VOC emission reduction acts to slightly mitigate the O_3
 378 increases compared with S3 in January and February, but the VOC emission reduction does not
 379 improve the O_3 levels in March. The increases in the mixing ratios of O_3 and PAN due to the NO_x -
 380 only control strategies suggest that the central, western and eastern regions of Thailand are likely
 381 VOC-limited during January to February. However, in March, the increases in O_3 are not mitigated
 382 by the reductions of VOC suggesting that these areas become less VOC-limited in this month.

383 Figure 9 illustrates the changes of delta O_3 due to the China's emission control strategies
 384 at 10 monitoring stations located in the BMR (3T, 10T, 11T, 14T, 15T, 19T, 22T, 27T, 52T and
 385 61T sites). In January and February (Fig. 9a) to 9b)), NO_x emission reduction mostly increases the
 386 monthly-average mixing ratio of O_3 . However, incorporating 40 % VOC emission reduction in
 387 China decreases O_3 at most monitoring stations. During March (Fig. 9c)), NO_x emission reduction
 388 in China enhance the mixing ratio of O_3 at 14T, 19T, 22T, and 27T monitoring stations. However,
 389 the increases in O_3 are mitigated by 40% VOC emission reduction in China. On the other hand,
 390 the NO_x -only control strategies reduce the mixing ratios of O_3 at 3T, 10T, 11T, 52T and 61T
 391 monitoring stations, while incorporating 40 % VOC reduction increases O_3 levels at most sites
 392 (Fig. 9c)). The results in this section further support that the BMR is VOC-limited in January and
 393 February, on the other hand, this area becomes less VOC-limited or even NO_x -limited in March in
 394 some parts of the BMR (NRC, 1991).

395 5. Analysis of NO_x -VOC indicators of O_3 formation

396 There are several photochemical indicators used to predict NO_x - and VOC-limited O_3
 397 formation regimes including $\text{H}_2\text{O}_2/\text{HNO}_3$, O_3/NO_x , O_3/NO_y (where $\text{NO}_y = \text{NO}_x + \text{HNO}_3 + \text{N}_2\text{O}_5 +$
 398 $\text{NO}_3 + \text{PAN} + \text{HONO} + \text{HNO}_4$), O_3/NO_z (where $\text{NO}_z = \text{HNO}_3 + \text{N}_2\text{O}_5 + \text{NO}_3 + \text{PAN} + \text{HONO} +$

399 HNO₄), HCHO/NO₂ and HCHO/NO_y (Sillman, 1995; Hammer et al., 2002; Lam et al., 2005;
400 Zhang et al., 2009; Campbell et al., 2015). The ratio of H₂O₂/HNO₃ is an important indicator of
401 NO_x- and VOC-limited regimes of O₃ formation (Lam et al., 2005 and Hammer et al., 2002). Under
402 low NO_x mixing ratios, H₂O₂ is the major radical sink, while HNO₃ is the major radical sink under
403 high NO_x mixing ratios. Therefore, a high and low ratio of H₂O₂/HNO₃ indicate a NO_x-limited
404 regime and a VOC-limited regime, respectively (Lam et al., 2005 and Jacob, 1999). The transition
405 value of H₂O₂/HNO₃ is 0.2 (Hammer et al., 2002; Lam et al., 2005; Liu et al., 2010; Sillman, 1995;
406 Zhang et al., 2009). When a value is lower than the transition value, a VOC-limited regime is more
407 indicated; otherwise, NO_x-limited regime is may be favored (Lam et al., 2005).

408 Figure 10a) to 10d) illustrate the spatial distributions of monthly-average H₂O₂/HNO₃ in
409 the regional scale during daytime (10:00 to 17:00 LT) with the transition value of 0.2. The spatial
410 distributions indicate that, in January, the majority of this region is NO_x-limited, except for
411 southern China, northwestern China, eastern India and some regions of Thailand (central and
412 northern Thailand), where those areas are VOC-limited. During February to March, most of
413 eastern China tends to become more VOC-limited; however, big cities (i.e. Beijing and Shanghai)
414 continue to remain NO_x-limited.

415 Figure 10e) to 10h) show the spatial distributions of monthly-average and O₃-event-period-
416 average H₂O₂/HNO₃ on the local-scale (d03) during daytime with the transitional value of 0.2. The
417 spatial distributions of the indicator suggest that during January to February, nearly the entire
418 domain, including the BMR, are VOC-limited. During March, the central regions of the domain
419 becomes NO_x-limited.

420 The VOC-NO_x indicator for the formation of ozone support the results in section 5 (the
421 change of O₃ mixing ratios due to China's NO_x and VOC emission reductions); which, in general,
422 cover eastern China and northern to central Thailand including the BMR, are likely VOC-limited.
423 However, in March, reversals from VOC-limited to NO_x-limited may be found in some parts of
424 the BMR.

425 **6. Effects of local flows and topography on O₃ levels in the BMR.**

426 China's emission control strategies affect regional and local chemical species, including
427 O₃ and its precursors. These effects are more obvious in March than in the other months. The
428 indicator analysis suggested that in the northeastern and eastern regions of China, and the central
429 areas of Thailand and the BMR, are normally in VOC-limited which a decrease in NO_x emissions
430 lead to an increase in O₃. In this section, we examine possible processes affecting O₃ formation,
431 accumulation and transport in this area, especially local flows and topography in the innermost
432 domain, since this domain has a very complex topography (flat terrain in the center, the Tenasserim
433 range on the west, mountainous areas on the east and southeast, and coastal areas and sea over the
434 south). The complex topography of the innermost domain produces a flow channeling in valleys
435 and very complicated flow structures due to the superposition of the different scale of flows (i.e.
436 mountain winds, breezes and flows from nonuniform land use) that influence mixing ratios, and
437 model simulation accuracy (Gustin et al., 2015; Jiménez et al., 2012).

438 We investigate the diurnal cycle of O₃ against wind vector in the lowest model layer during
439 the O₃ event (5 to 6 March 2010) to describe the pollution dynamics and examine the effects of

440 local flows on pollutants mixing ratios in d03. Figure 11 illustrates the dynamics of the daily cycle
441 of O₃ in d03. In the early morning (5:00 LT), mountain winds from the Tenasserim range and the
442 mountainous areas, and land- and sea-breeze circulations around the coastal areas help establishing
443 south-north flows over the flat plain. At this time, there are widespread low O₃ mixing ratios, while
444 higher O₃ mixing ratios are found around the mountainous areas, and around the coastal areas on
445 the south. Since there is no new O₃ is formed during this time, the spatial distribution of O₃ shows
446 the effect of topography (i.e. mountainous areas) on O₃ accumulation, and the effect of local flows
447 (i.e. land- and sea-breeze circulations) on O₃ transport, which the circulation recirculates O₃ to
448 inland. During 10:00 to 14:00 LT, the transition between sea- and land-breeze occurs over the gulf
449 of Thailand. Strong sea-breezes are the dominant flow over the coastal areas and penetrate the
450 inland. During this time, there are widespread increases in O₃ with high O₃ mixing ratios, which
451 are mostly found over the Tenasserim range and the mountainous areas, where biogenic VOC
452 (BVOC) emissions are high. The presence of BVOC emissions favor the formation of O₃ (Chatani
453 et al., 2015; Sillman, 1999) in this area, since the indicators analysis show that this area is VOC-
454 limited. O₃ and BVOC from the Tenasserim range are transported to the flat terrain by mountain
455 winds, which later converge with the sea-breezes and bring the pollutants to the northern regions
456 of the domain. In the late afternoon to the evening (18:00 to 23:00 LT), the mixing ratios of O₃
457 start decreasing due to a lack of sunlight to promote photochemical reactions, the titration of O₃
458 by NO at night (Jiménez et al., 2012; Tokarek et al., 2017) and the reaction of O₃ with unsaturated
459 biogenic hydrocarbons (Tokarek et al., 2017). These O₃ depletion mechanisms are further
460 supported by the diurnal cycles of VOC, NO and NO₂. At night, the increases in NO₂ are mostly
461 found over the central areas of the domain, while the mixing ratios of NO is near zero. A
462 combination of the presence of high VOC, especially over the flat terrain and strong sea-breezes
463 during nighttime, may be a cause of underpredicted O₃ at night over this area (d03), especially at
464 coastal monitoring stations (i.e. 19T and 27T station). Spatial distribution of O₃, BVOC, NO and
465 NO₂ during the O₃ event are shown in the supplement material (Fig. S4.1 to S4.4).

466 7. Summary and discussion

467 We examine O₃ behavior and the influence of China's NO_x and VOC emissions
468 and the subsequent transport of O₃ and O₃ precursors on O₃ levels in the BMR by the WRF-Chem
469 model, since elevated O₃ in the BMR was affected by both local and regional contribution of
470 precursor pollutants. Generally, the model reasonably predicts O₃ mixing ratios and performs well
471 in capturing the O₃ event. During the O₃ season, the decreases in China's NO_x emissions (10 %,
472 20 %, and 40 % NO_x emission reductions) lead to decreases in the regional O₃ precursors (i.e.
473 NO_x, CO, VOC, HNO₃ and O_x), but increases in the regional PAN and O₃. The increases in PAN
474 and O₃ are mitigated by incorporating 40 % China's VOC emission reduction.

475 The NO_x-only emission reductions in China mostly lead to widespread increases in local
476 scale (d03) O₃ and O₃ precursors (except NO_x that the minor decreases are found in January and
477 February), while incorporating 40 % China's VOC emission reduction mostly decrease O₃ mixing
478 ratios in this area. The increase in the mixing ratios of O₃ and PAN in the regional- and local-scale,
479 including the BMR, due to the NO_x-only emission reductions in China suggest that eastern China
480 and northern to central Thailand are likely VOC-limited; however, in March, these regions become
481 less VOC-limited. This is further supported by the mitigation in O₃ increases when including 40

482 % VOC emissions reduction and by the results from the indicator analysis. The discrepancy
483 between O₃/NO_z and the other indicator may come from very low simulated mixing ratios of NO_z
484 (approximately on the order of 10⁻² to 10⁻³ ppb), although we do not have an observation to
485 compare with the simulation for this species.

486 A change of VOC- and NO_x-limited regime may occur due to i) Emissions intensity, ii)
487 Cumulative solar radiation, iii) Intensity of solar radiation, and iv) VOC reactivity (Kannari and
488 Ohara, 2010). However, the use of modeled O₃ photochemical indicators averaged over regional
489 and monthly-to-seasonal time scales has been shown to be a well-established method used to
490 differentiate between NO_x- and VOC-limited O₃ chemical regimes, which are based on transition
491 values in North America (Zhang et al., 2009, Campbell et al., 2015), Europe (Walaszek et al.,
492 2017), and Asia (Liu et al., 2010). While the split between NO_x- and VOC-limited conditions
493 may indeed shift in space and time (as explained above) the use of particular O₃ indicator ratios
494 are valid tools for diagnosing if NO_x or VOC emission controls are beneficial or detrimental (Liang
495 et al., 2006; Jin et al., 2017; Walaszek et al., 2017). It is noteworthy that ambient O₃ mixing ratios
496 have a complex nonlinear relationship with VOC and NO_x mixtures; therefore, it is important that
497 the emission inventories be accurate (Fujita et al., 2013).

498 Our study depicts that China's emissions play an important role in controlling the pollutant
499 levels in this region. The changes in regional NO_x mixing ratios correspond directly to the changes
500 of China's NO_x emissions; however, controlling only NO_x emissions is not an effective strategy to
501 reduce O₃ mixing ratios in this region. The change in regional VOC mixing ratios varied from the
502 changes in China's VOC emission reduction. Therefore, the regional VOC mixing ratios is perhaps
503 influenced by other VOC emission sources (i.e. biogenic VOC emission sources) where the
504 decreases in VOC emissions provide additional benefits for controlling O₃ mixing ratios.

505 In the BMR, our analysis suggests that not only local anthropogenic emissions, but also
506 long-range transport of pollutants, as far as originating from China, influence O₃ formation in the
507 BMR. Results from this study with different scenarios show that the WRF-Chem model could be
508 a useful tool to evaluate O₃ formation and the relationship between O₃-NO_x-VOCs and O₃
509 formation in the BMR; however, the practical implementation or control strategies based on this
510 study are complicated. It is difficult to quantify the proportion of pollutants contributing to local
511 emissions and from those associated with long-range transport due to the lack of emission data in
512 the BMR. BVOC transported from the Tenasserim range by mountain winds also favor the O₃
513 formation in this area; since this area is indicated as a VOC-limited regime, the increases in VOC
514 lead to the increases in O₃ levels. Another factor influences O₃ and its precursors levels is land
515 and sea-breeze circulations over the Gulf of Thailand, which the circulations do not only dilute
516 the mixing ratios over the east side of coastal areas, but may also recirculate pollutants to inland,
517 resulting in the increases in pollutants level. Thus, issues associated with the complex terrain and
518 coastal vicinity of some of the sites may prevent the model from having accurate predictions.

519 It must be emphasized that uncertainties in this study may be from uncertainties in
520 biogenic and anthropogenic emissions. In general, uncertainties in spatial input (i.e. land use,
521 plant functional type and leaf area index) contribute to the accuracy of biogenic emission
522 estimation, tropospheric O₃ and its precursors variation (Porter and Heald, 2019). For

523 anthropogenic emission inventories, uncertainties in regional emissions may be expected as large
524 as a factor of 2 or even larger (Bond et al. (2004, 2007); Smith et al. (2010)). Uncertainties may
525 come from meteorological simulation and chemistry from the WRF-Chem model. In general,
526 windspeed is more likely to be the largest source of uncertainty; however, chemical reaction
527 rates and stoichiometry may also be a major cause of simulated O₃ uncertainty (~20%) (Sillman,
528 1999). In our study, the finest horizontal resolution is 4 km which may be too coarse to deal with
529 very complicated flow structures. Jiménez et al. (2012) suggested that a high-resolution model
530 (resolution of ≤ 1 km) was required in order to study air pollution in a very complex topography.
531 However, the high-resolution model had to be nested in a larger model domain in order to
532 investigate the effects of larger-scale transport.

533 Simulated NO_x and VOCs analysis offers important information to distinguish the
534 ozone—precursor relationship in terms of a split between a NO_x-limited and VOCs-limited
535 regimes (or NO_x-saturated). However, the simulations of these species are also subject to large
536 uncertainties. Sillman (1999) reported five factors that influenced estimated NO_x- and VOCs-
537 limited regimes were i) VOC/NO_x ratios, ii) VOC reactivity, iii) biogenic hydrocarbons, iv)
538 photochemical aging of the air mass, and v) rates of meteorological dispersion. Furthermore, this
539 study examines the formation of O₃ in the BMR primarily during January to March (dry season),
540 and the model evaluations are also done during the same study period. However, the study on O₃
541 formation and model evaluation need to be extended for at least a complete one-year study.

542 This article presents insights derived from WRF-Chem model and field measurements in
543 BMR; and seeks to demonstrate the relationships between ozone—NO_x—VOC chemistry and
544 transport. This analysis of ozone—NO_x—VOC sensitivity has close connection to air quality
545 management and regulatory policy for BMR.

546

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557 **Disclaimer**

558 The scientific results and conclusions, as well as any views or opinions expressed herein, are
559 those of the author(s) and do not necessarily reflect the views of NOAA or the Department of
560 Commerce.

561 Code and Data availability

562 The WRF-Chem model (all version) is available at WRF Users page
563 (http://www2.mmm.ucar.edu/wrf/users/download/get_sources.html#WRF-Chem).

564 Meteorology data is available at NCAR UCAR Research Data Archive computational &
565 Information System Lab website (<https://rda.ucar.edu/>).

566 Data to prepare initial and boundary conditions is available at NCAR UCAR
567 Atmospheric Chemistry Observations & Modeling website (<https://www.acom.ucar.edu/wrf-chem/mozart.shtml>).

569 EDGAR-HTAP data is available at European Commission, joint Research Center,
570 EDGAR-Emissions Database for Global Atmospheric Research (<https://edgar.jrc.ec.europa.eu/>).

571 Scripts in this study are written in the NCAR Command Language (NCL). The program
572 is available at GitHub NCAR/ncl (<https://github.com/NCAR/ncl>). More information about NCL
573 can be found at <http://www.ncl.ucar.edu/>.

574 Data is available with the primary author.

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Table 1. WRF-Chem physical and chemical options used in this study.

	Option	Selection
Physics	Microphysics	Thompson
	Cumulus parameterization	Grell-Freitas (only domain 1 and 2)
	Long wave radiation	A new version of Rapid Radiative Transfer Model (RRTMG)
	Short wave radiation	RRTMG
	Planetary boundary layer	Yonsei University
	Surface layer	MM5 similarity
	Land surface	Noah Land-Surface model
Chemistry	Chemical mechanism	RADM2-MADE/SORGAM
	Photolysis	Madronich F-TUV photolysis
	Dry deposition of gas/aerosol species	On
	Subgrid convective wet scavenging and aqueous chemistry	On (only domain 1 and 2)
	Vertical turbulent mixing	On (only domain 1 and 2)
	Aerosol effect in radiation	On

Table 2. Summarizes baseline simulation and sensitivity simulations.

Simulations	Adjusted China's emissions		Note
	NO _x	VOC	
Baseline	No	No	
Sensitivity			
- Strategy 1 (S1)	10 % reduction	No	The national reduction target during China's 12 th FYP (Wang et al., 2014a)
- Strategy 2 (S2)	20 % reduction	No	To examine the responses of O ₃ and its precursors in the BMR due to different China's NO _x emission reductions.
- Strategy 3 (S3)	40 % reduction	No	
- Strategy 4 (S4)	40 % reduction	40 % reduction	To investigate implications for incorporating VOC emission reduction. (Wang and Hao, 2012)

Table 3. Summary of the averaged statistics for January to March 2010.

Variable	Statistics	Station ID															Overall	
		3T	10T	11T	14T	15T	19T	21T	22T	24T	26T	27T	34T	41T	52T	60T		61T
T (°C)	\overline{Obs}	29.59	29.39	29.71	28.93	30.12	28.71	28.12	29.14	28.97	28.91	29.24	30.87	30.61	29.43	27.98	30.09	29.36
	\overline{Sim}	26.74	25.93	26.18	26.27	26.72	26.9	27.25	26.27	27.07	26.59	26.94	26	26.75	26.47	26.58	25.99	26.54
	r	0.78	0.86	0.85	0.77	0.78	0.8	0.9	0.89	0.89	0.89	0.80	0.56	0.83	0.84	0.87	0.85	0.82
	MB	-2.85	-3.47	-3.53	-2.67	-3.39	-1.81	-0.88	-2.87	-1.91	-2.32	-2.30	-4.87	-3.86	-2.97	-1.40	-4.10	-2.83
	RMSE	4.21	4.6	4.61	4.1	4.72	3.58	2.1	3.87	2.74	3.02	3.92	6.00	4.59	4.28	2.77	5.29	4.03
	NMB	-9.62	-11.79	-11.91	-9.21	-11.27	-6.3	-3.13	-9.85	-6.59	-8.01	-7.86	-15.77	-12.61	-10.08	-5.01	-13.63	-9.54
WS (ms ⁻¹)	\overline{Obs}	2.12	2.01	1.02	1.57	1.14	2.53	1.5	1.23	1.78	2	2.06	1.93	1.09	1.71	2.03	1.64	1.71
	\overline{Sim}	3.29	3.03	2.73	3.04	3.37	3.19	3.25	3.26	4.15	3.08	3.23	4.09	4.17	3.04	3.26	3.47	3.35
	r	-0.03	0.04	0.01	0.16	0.13	0.01	0.26	0.35	0.32	0.48	0.15	0.04	0.23	0.12	0.24	0.07	0.16
	MB	1.17	1.02	0.49	1.46	2.22	0.66	1.25	2.03	2.38	1.13	1.17	2.17	3.08	1.33	1.23	1.82	1.54
	RMSE	2.19	1.84	1.17	2.05	2.75	2.03	2.13	2.45	3.34	1.75	2.01	2.91	3.65	20.1	3	2.53	3.49
	NMB	55.39	50.76	197.41	92.92	194.66	26.25	121.9	165.03	133.57	57.74	56.91	112.11	282.4	77.93	60.35	110.85	112.26
RH (%)	\overline{Obs}	73.11	67.08	65.19	69.81	75.46	66.38	74.46	68.19	68.93	72.4	72.53	62.12	60.14	74.54			69.31
	\overline{Sim}	61.99	65.97	65.6	62.49	59.21	61.48	64.63	60.47	64.31	61.08	59.89	64	63.67	59.71			62.46
	r	0.68	0.77	0.74	0.68	0.64	0.71	0.78	0.74	0.79	0.6	0.59	0.65	0.67	0.71			0.70
	MB	-11.11	-1.11	0.41	N/A	-7.32	-16.25	-4.94	-9.83	-7.70	-4.61	-11.31	-12.63	1.88	3.53	-14.82	N/A	-6.84
	RMSE	17.32	12.61	12.7	15.29	21.26	14.61	16.27	14.28	11.68	18.06	18.96	15.25	14.35	20.08			15.91
	NMB	19.61	-1.65	0.63	-10.49	-21.53	-7.44	-13.2	-11.3	-6.69	-15.65	-17.41	3.03	5.87	-19.88			-6.86
P (hPa)	\overline{Obs}	758.29	757.02	758.94	757.29	758.73	757.48	756.13	758.64	753.61	759.66	759.60	762.43	747.32	759.73	764.33	760.20	758.09
	\overline{Sim}	759.09	759.23	759.39	758.92	758.97	754.70	759.20	759.34	756.54	757.82	758.16	733.64	757.23	759.20	755.15	759.23	756.61
	r	0.76	0.83	0.80	0.78	0.78	0.64	0.81	0.74	0.78	0.83	0.77	0.79	0.53	0.81	0.65	0.76	0.75
	MB	0.79	2.22	0.45	1.63	0.24	-2.78	3.07	0.70	2.93	-1.84	-1.44	-28.78	9.91	-0.53	-9.18	-0.97	-1.47
	RMSE	1.63	2.53	1.37	2.47	1.36	3.27	3.32	1.63	3.23	2.16	1.98	28.81	10.11	1.33	9.83	1.65	4.79
	NMB	0.11	0.29	0.06	0.22	0.03	-0.37	0.41	0.09	0.39	-0.24	-0.19	-3.77	1.33	-0.07	-1.20	-0.13	-0.19
Max	\overline{Obs}	30.1	25.20	16.27	20.45	20.51	26.53	10.88	22.68	21.53	29.68	33.85	28.54	33.86	30.43	33.56	28.62	23.68*
8-h O₃ (ppb)	\overline{Sim}	15.85	19.25	19.41	17.26	17.28	13.61	18.89	14.32	20.01	5.62	16.05	31.22	48.90	26.75	16.68	31.03	22.63*
	r	0.42	0.38	0.16	0.26	0.26	-0.03	0.42	0.56	0.58	-0.27	-0.02	-0.09	0.30	0.44	0.10	0.53	0.39*
	MB	-14.25	-5.95	3.14	-3.19	-3.23	-12.92	3.07	-8.37	-1.53	-24.06	-17.80	2.68	15.04	-3.68	-16.88	2.40	-1.96*
	RMSE	18.31	14.91	17.44	16.88	16.88	19.87	7.46	12.92	12.44	27.56	56.68	23.75	25.13	17.55	22.07	19.18	17.16*
	MNBE	-46.98	-18.87	38.33	-5.77	-5.98	-29.77	39.13	-34.45	-1.88	-74.48	-48.07	22.08	49.66	-13.04	-47.77	6.81	-3.22*

Note: *data from 19T, 26T, 27T, 34T and 60T are excluded.

Table 4. Summarized delta O₃, NO_x, CO, VOC, PAN, HNO₃ and O_x in the regional scale (d01) due to the China's emission reductions during January to March 2010.

Species (ppb)	January				February				March			
	S1	S2	S3	S4	S1	S2	S3	S4	S1	S2	S3	S4
O ₃	0.35±0.54 (0.9%)	0.72±.78 (1.9%)	1.53±1.49 (4.1%)	1.34±1.39 (3.5%)	0.35±0.69 (0.9%)	0.62±1.07 (1.5%)	1.27±2.07 (3.2%)	1.11±1.87 (2.8%)	0.30±0.80 (0.7%)	0.55±1.06 (1.3%)	1.17±2.0 (2.7%)	1.01±1.80 (2.3%)
NO _x	-5.41±10.76 (-10.4%)	-10.98±20.57 (-21.1%)	-20.53±38.82 (-39.4%)	-20.45±38.76 (-39.2%)	-4.15±9.04 (-10.9%)	-7.94±16.73 (-20.8%)	-14.91±31.33 (-39.1%)	-14.84±31.19 (-38.9%)	-2.41±8.41 (-9.7%)	-4.64±12.22 (-18.6%)	-9.34±20.79 (-37.5%)	-9.24±20.79 (-37.1%)
CO	-0.10±4.58 (-0.1%)	-0.31±3.65 (-0.2%)	-0.30±3.99 (-0.2%)	-0.46±3.96 (-0.2%)	-0.16±3.96 (-0.1%)	-0.13±3.90 (-0.1%)	-0.25±4.03 (-0.1%)	-0.39±3.86 (-0.2%)	0.03±8.69 (0.0%)	0.03±9.42 (0.0%)	-0.16±9.07 (-0.1%)	-0.17±9.42 (-0.1%)
VOC	0.01±0.40 (0.1%)	-0.03±0.29 (-0.3%)	-0.05±0.35 (-0.5%)	-1.06±1.81 (-11.6%)	0.01±0.50 (0.1%)	0.0±0.53 (0.0%)	-0.03±0.55 (-0.3%)	-0.80±1.64 (-9.2%)	0.07±1.26 (0.7%)	0.06±1.32 (0.8%)	0.02±1.27 (0.2%)	-0.55±1.82 (-6.5%)
PAN	0.01±0.02 (1.4%)	0.01±0.02 (2.7%)	0.03±0.04 (6.5%)	0.02±0.03 (4.0%)	0.01±0.03 (0.9%)	0.01±0.04 (1.9%)	0.02±0.05 (4.3%)	0.01±0.04 (2.3%)	0.007±0.05 (1.2%)	0.01±0.05 (2.1%)	0.03±0.07 (4.7%)	0.02±0.06 (2.9%)
HNO ₃	-0.07±0.18 (-2.2%)	-0.18±0.24 (-5.5%)	-0.42±0.47 (-13.1%)	-0.46±0.48 (-14.0%)	-0.11±0.21 (-3.0%)	-0.22±0.32 (-6.3%)	-0.49±0.64 (-13.9%)	-0.50±0.65 (-14.3%)	-0.09±0.18 (-3.0%)	-0.16±0.23 (-5.6%)	-0.35±0.43 (-12.3%)	-0.36±0.45 (-12.4%)
O _x	-1.90±3.98 (-2.8%)	-3.80±7.57 (-5.6%)	-7.49±15.15 (-11.1%)	-7.76±15.20 (-11.5%)	-1.33±2.93 (-2.1%)	-2.67±5.63 (-4.2%)	-5.38±11.43 (-8.4%)	-5.59±11.57 (-8.8%)	-0.93±2.21 (-1.5%)	-1.82±4.03 (-3.0%)	-3.80±8.16 (-6.3%)	-3.98±8.33 (-6.6%)

Table 5. Summarized delta O₃, NO_x, CO, VOC, PAN, HNO₃ and O_x on the local scale (d03) due to the China's emission reduction strategies during January to March 2010.

Species (ppb)	January				February				March			
	S1	S2	S3	S4	S1	S2	S3	S4	S1	S2	S3	S4
O ₃	0.53±1.66 (1.5%)	1.20±1.57 (3.5%)	2.15±2.03 (6.2%)	1.95±1.96 (5.6%)	1.32±3.10 (2.2%)	1.64±3.09 (2.7%)	2.98±3.42 (5.0%)	2.55±3.18 (4.2%)	-0.34±5.78 (-0.4%)	0.16±6.39 (0.2%)	0.90±6.46 (1.1%)	0.70±44.47 (1.2%)
NO _x	2.60±33.73 (0.6%)	-10.62±26.52 (-2.5%)	-18.40±30.20 (-4.4%)	-18.76±30.66 (-4.4%)	-2.93±43.36 (-0.8%)	-3.26±41.42 (-0.9%)	-5.33±42.45 (-1.5%)	-6.72±40.68 (-1.9%)	22.98±76.75 (7.6%)	28.15±82.10 (9.0%)	24.44±78.90 (7.8%)	22.02±406.89 (3.7%)
CO	10.59±43.58 (1.5%)	-0.01±33.85 (0.0%)	3.38±38.38 (0.5%)	1.80±39.58 (0.3%)	-1.03±56.70 (-0.2%)	1.46±54.35 (0.2%)	3.94±55.85 (0.6%)	1.12±53.43 (0.2%)	36.92±100.96 (6.0%)	39.81±109.17 (6.5%)	36.41±105.73 (5.9%)	39.47±108.83 (6.4%)
VOC	1.32±5.03 (1.2%)	0.32±4.20 (0.3%)	1.28±4.59 (1.2%)	-0.55±4.56 (-0.5%)	0.13±8.09 (0.1%)	0.34±8.21 (0.3%)	0.81±8.37 (0.7%)	0.12±7.97 (0.1%)	4.44±15.90 (3.4%)	6.22±15.96 (4.7%)	5.25±15.66 (4.0%)	-4.56±113.39 (-2.3%)
PAN	0.01±0.17 (1.1%)	0.02±0.16 (2.6%)	0.05±0.20 (5.2%)	0.04±0.18 (4.5%)	0.04±0.38 (2.2%)	0.06±0.37 (2.8%)	0.12±0.39 (5.9%)	0.09±0.36 (4.7%)	0.06±0.70 (1.7%)	0.09±0.72 (2.6%)	0.09±0.71 (2.5%)	0.09±0.70 (2.6%)
HNO ₃	0.03±0.26 (0.62%)	0.06±0.24 (1.4%)	0.14±0.40 (3.0%)	-0.04±0.33 (-0.9%)	0.06±0.70 (0.5%)	-0.16±0.69 (-1.4%)	-0.38±0.77 (-3.3%)	-0.45±0.78 (-3.9%)	-0.39±0.95 (-2.9%)	-0.49±0.97 (-3.7%)	-0.49±0.97 (-3.7%)	-0.51±0.95 (-3.8%)

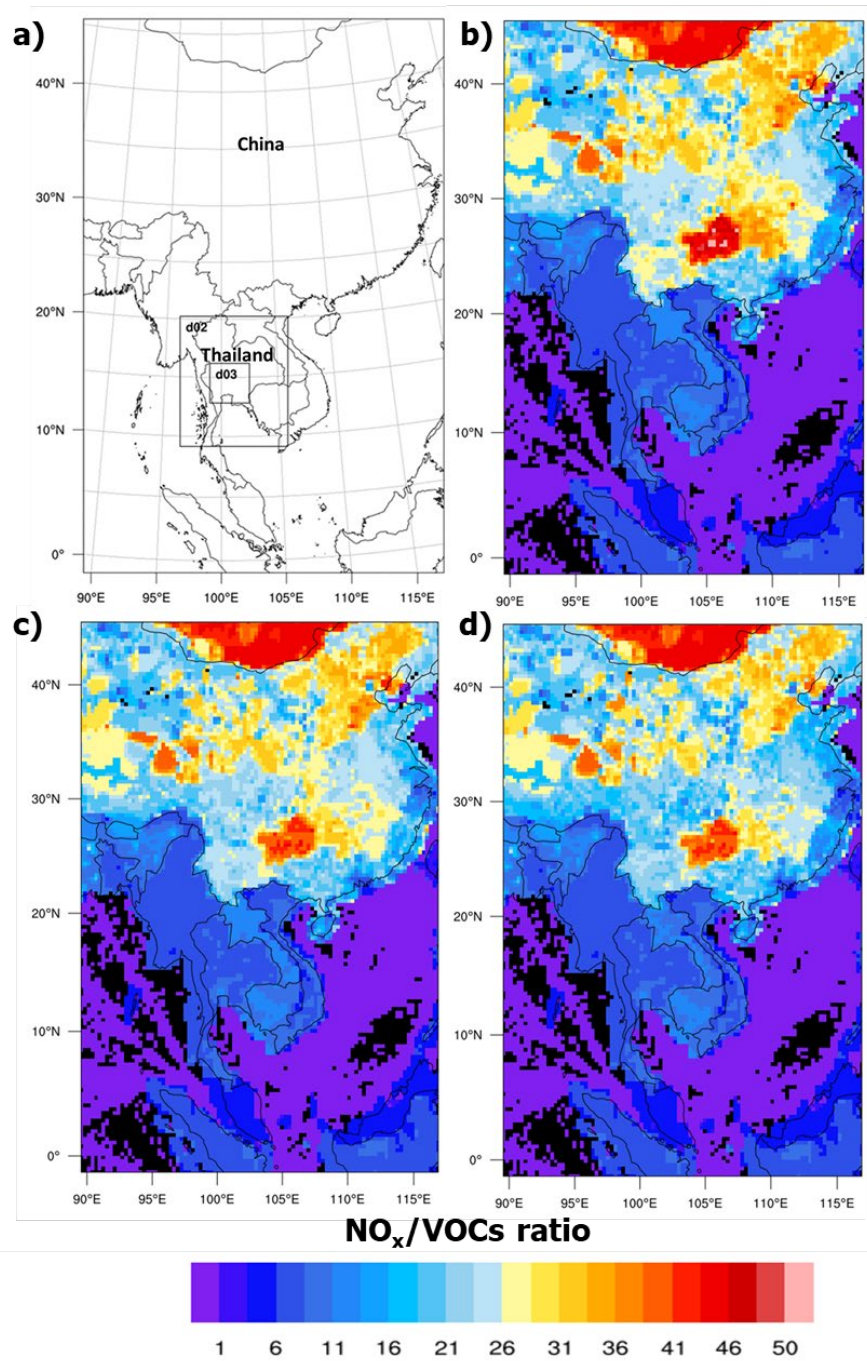


Figure 1. a) A triple-nested domain, including outermost domain (d01), the second domain (d02) and the innermost domain (d03) with 36-, 12- and 4-km horizontal resolutions, respectively; b) to d) NO_x/VOCs emission during January, February and March, respectively.

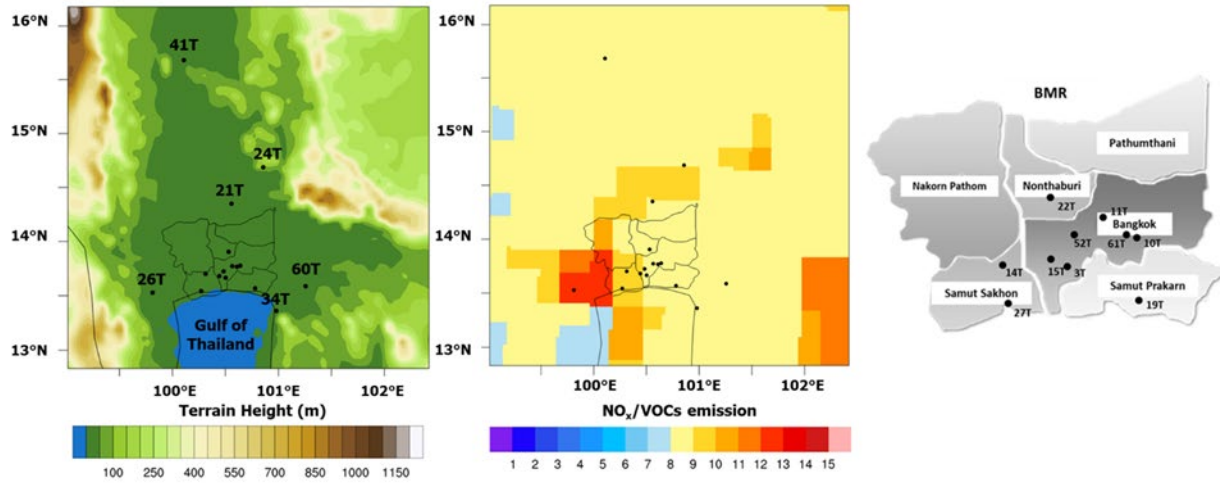


Figure 2. Terrain height in d03, NO_x/VOCs emission in d03 and the location of 16 monitoring stations which six stations located outside of the BMR and 10 stations located inside of the BMR for the model evaluation.

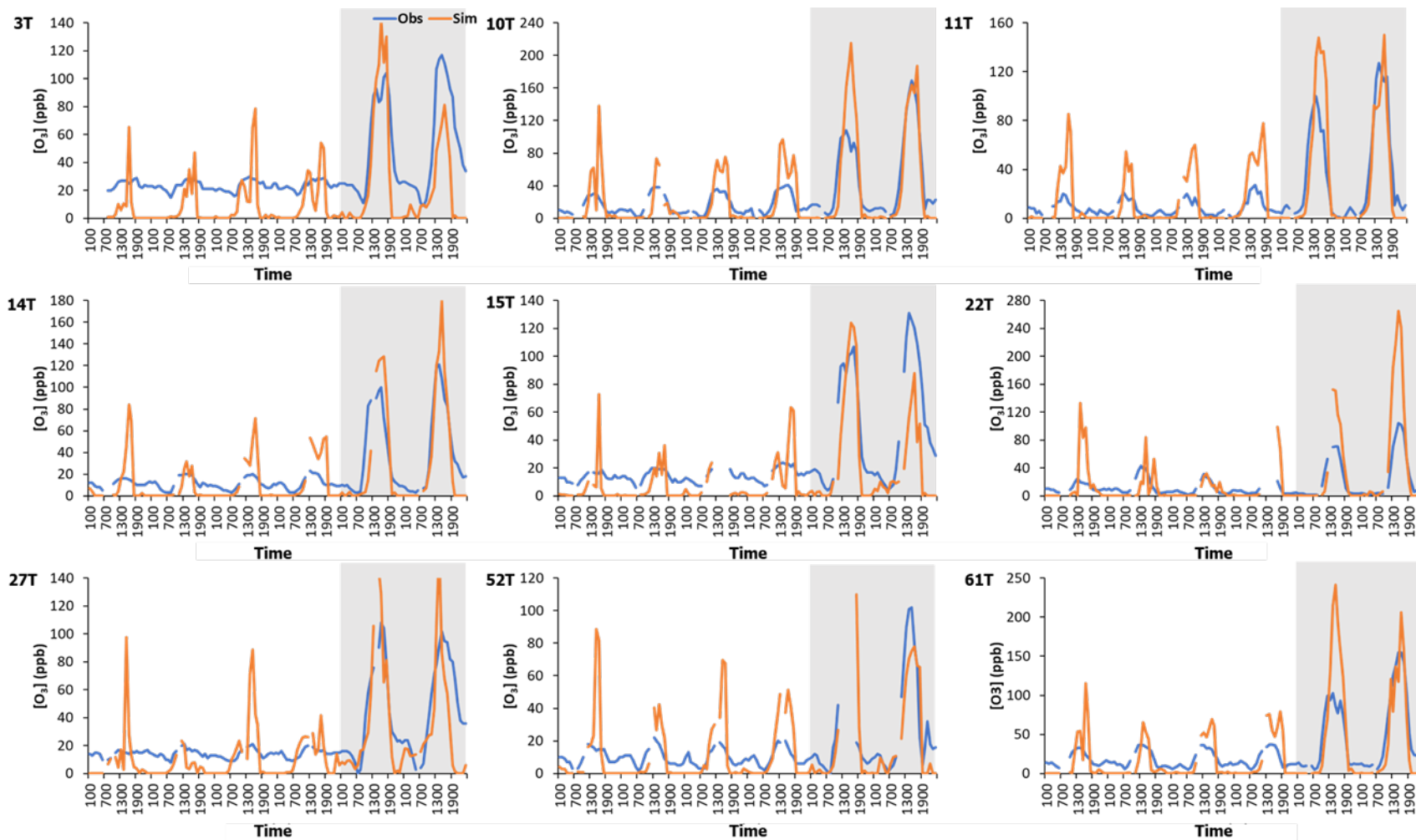


Figure 3. Temporal variations of hourly observed (blue) and simulated (orange) O_3 concentrations during 1 to 6 March 2010 at nine monitoring stations. The shaded areas refer to the O_3 episode. Simulated hourly O_3 concentrations are not presented when observations are missing.

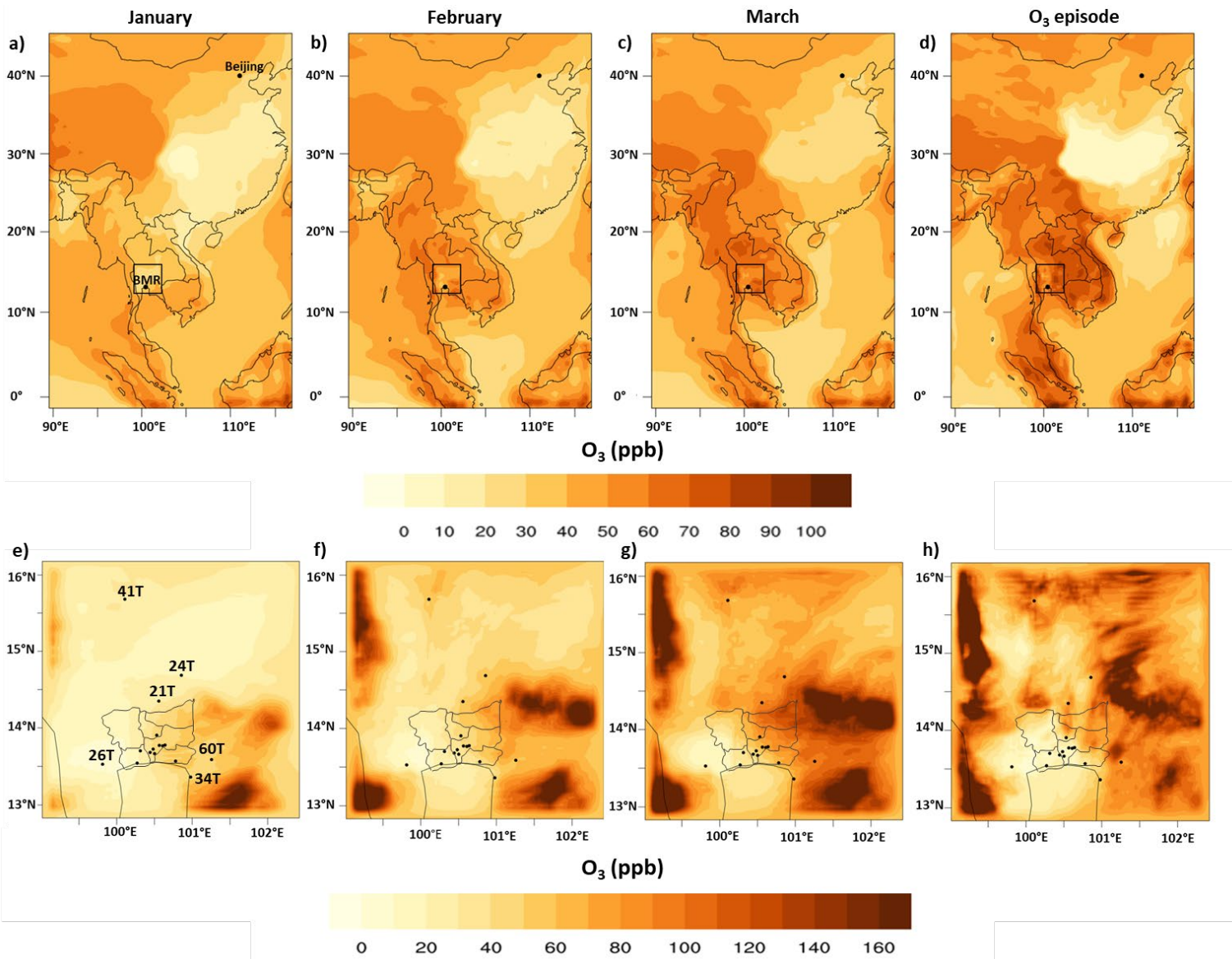
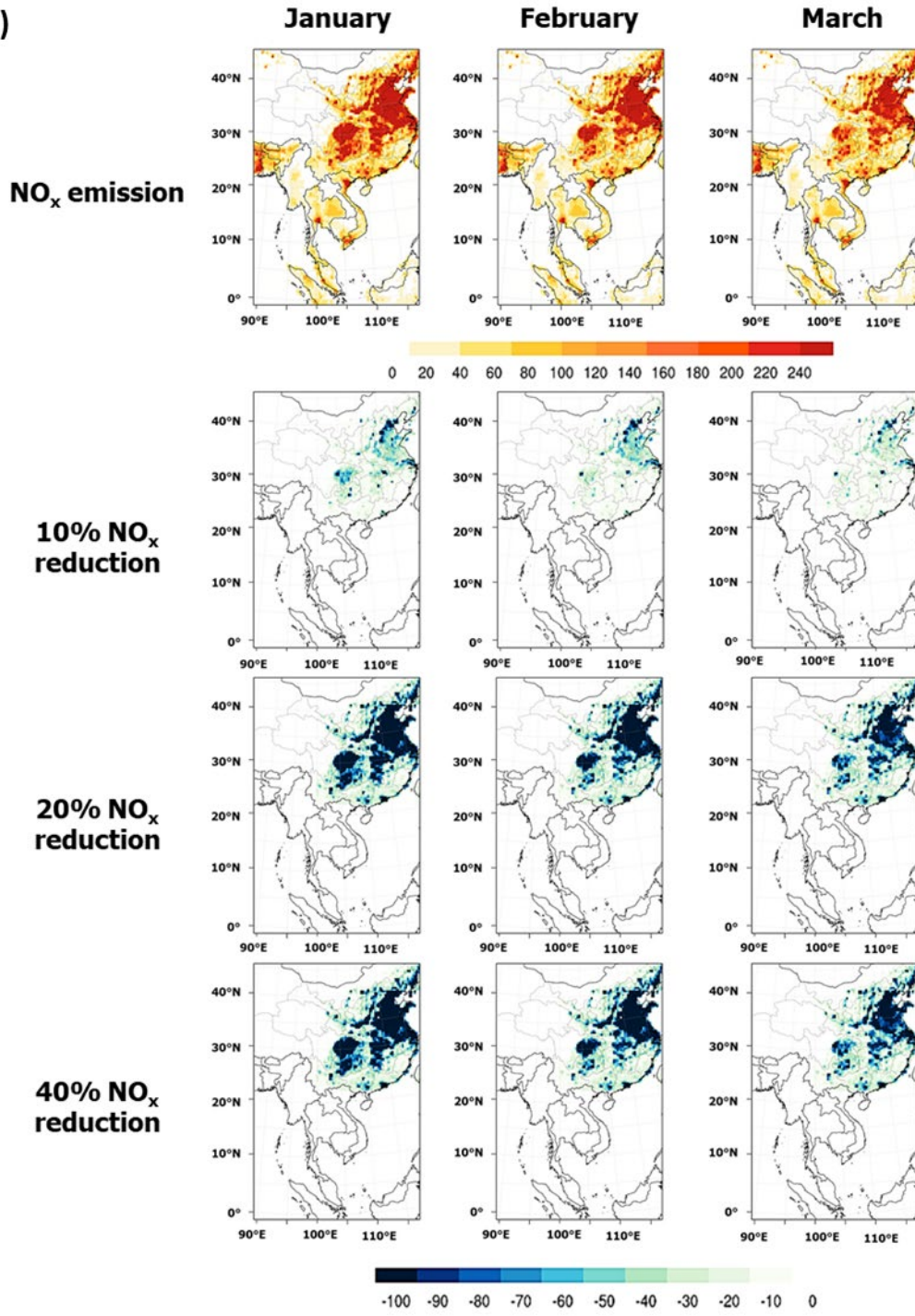


Figure 4. Spatial distributions of average O_3 from the baseline simulation in d01 and in d03 for the months of a) and e) January, b) and f) February, c) and g) March, and d) and h) the episodic O_3 event period. Small boxes indicate the boundary of d03 compared with d01.

a)



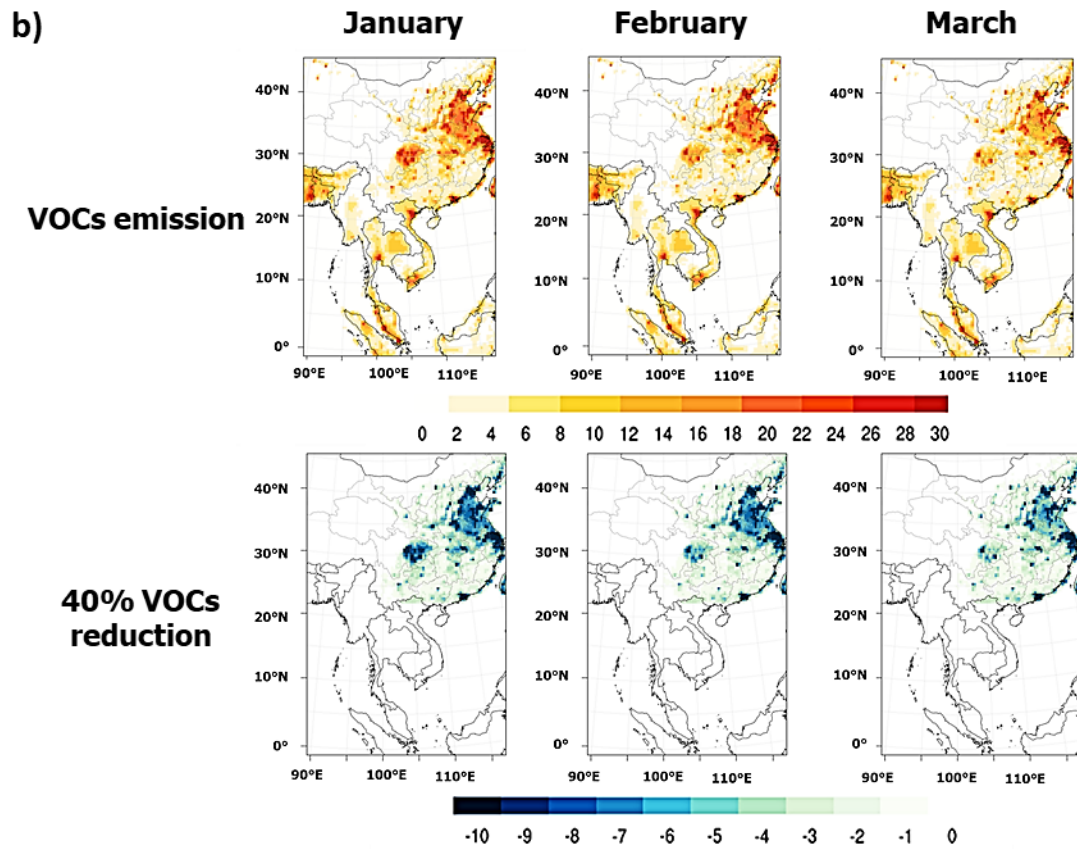


Figure 5. Monthly-average a) NO_x and b) VOC emissions and absolute differences of emissions (sensitivity – baseline) based on S1 to S4 in January to March 2010.

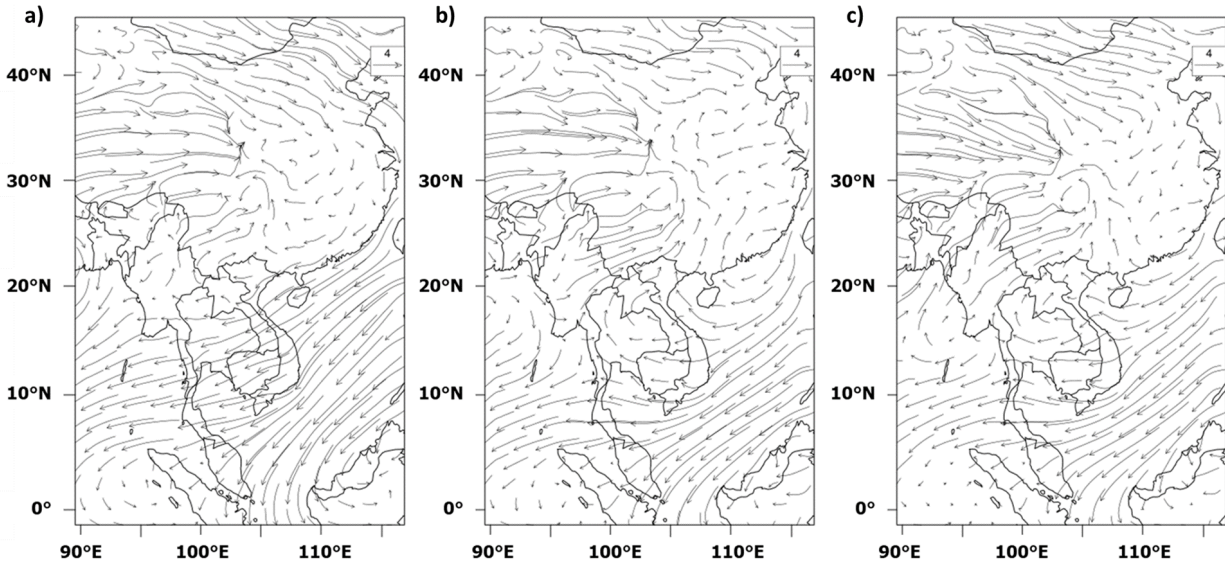


Figure 6. Monthly-average wind field during a) January b) February and c) March 2010 in d01.

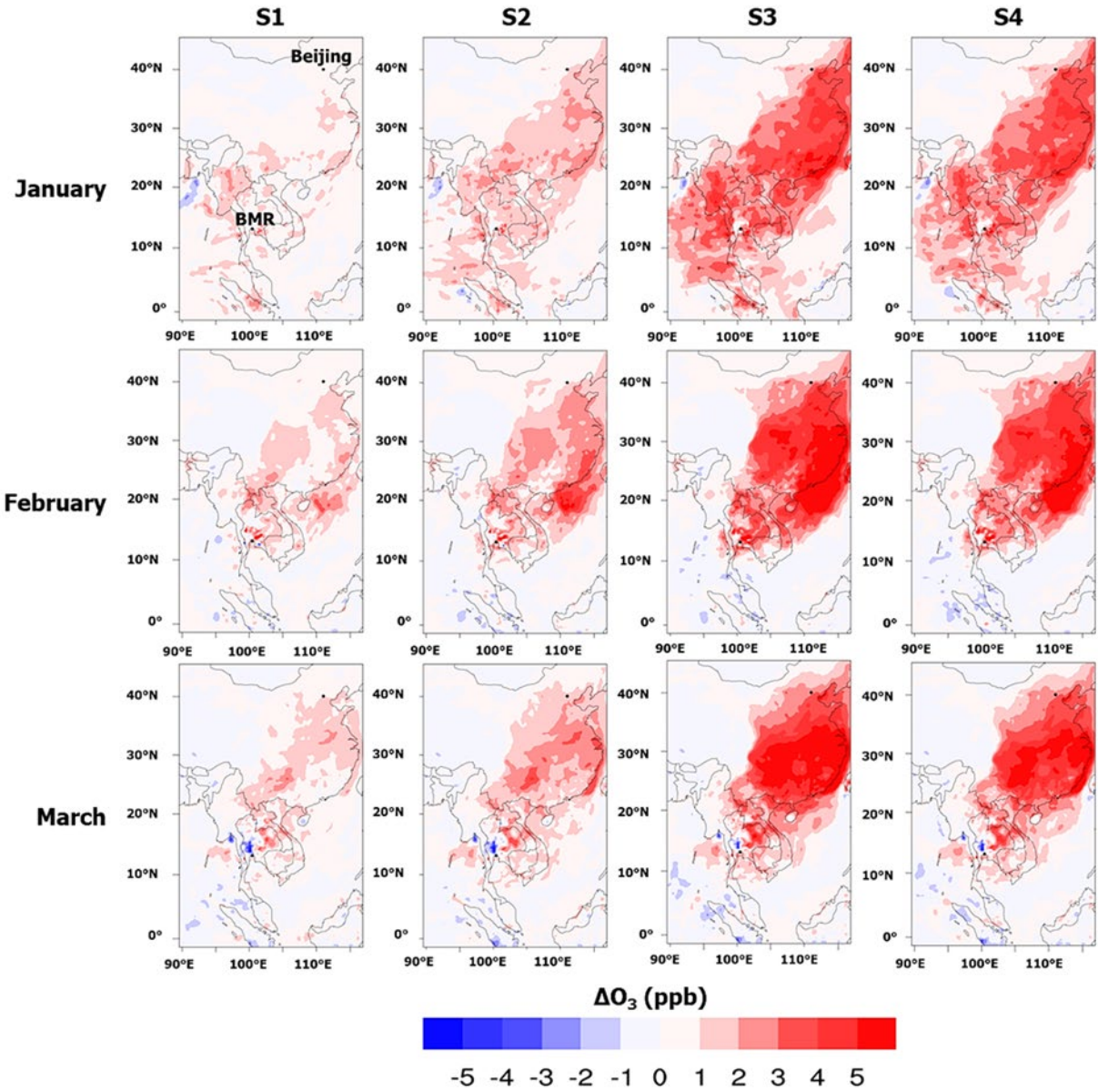


Figure 7. Spatial distributions of monthly-average delta O₃ on the regional scale (d01) due to China's emission reductions during January to March 2010.

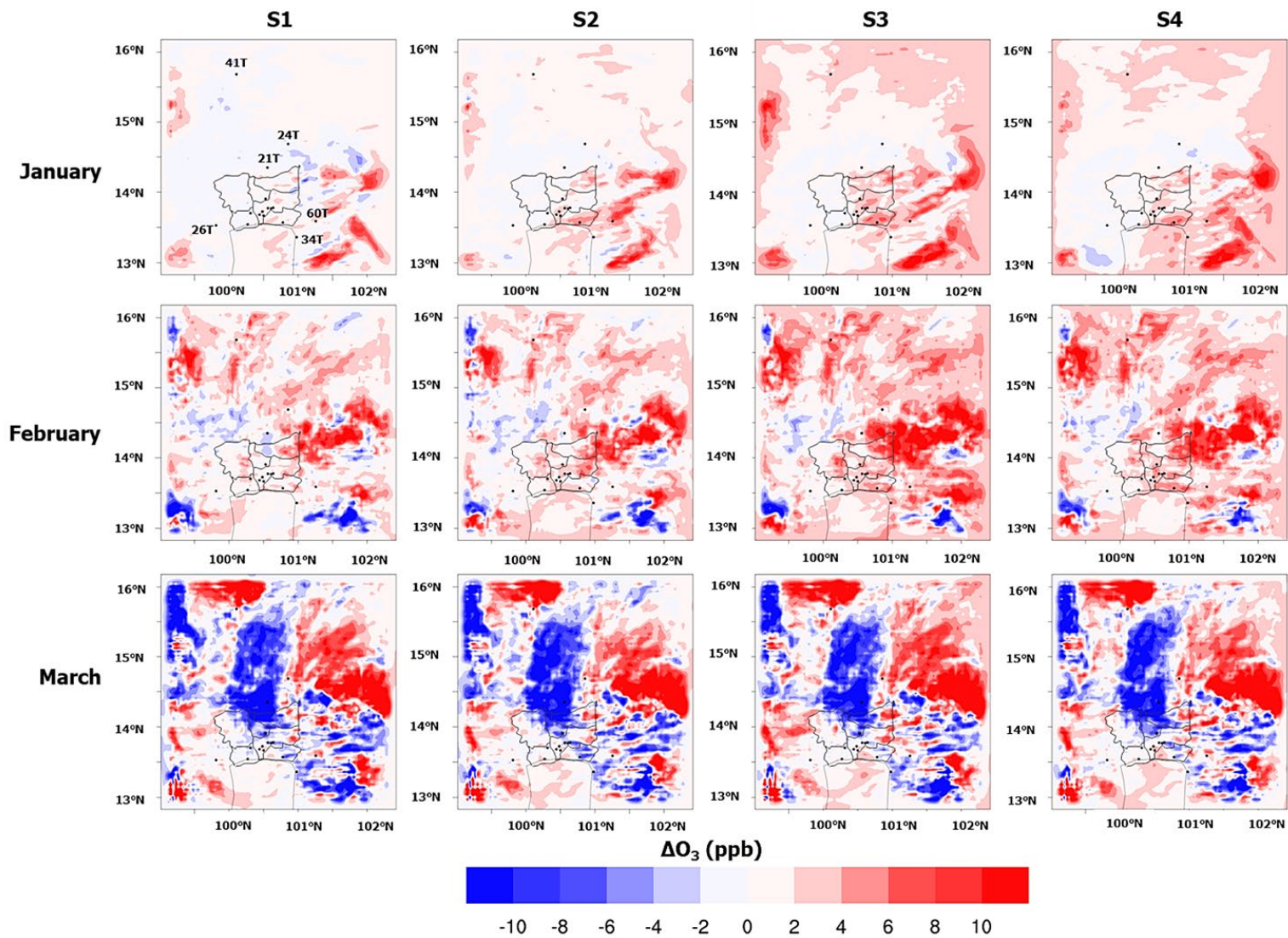


Figure 8. Spatial distribution of monthly-average delta O_3 on the local scale (d03) due to the China's emission reduction strategies during January to March 2010.

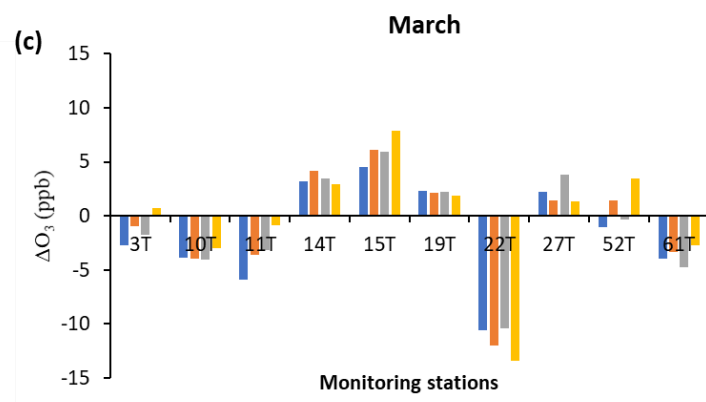
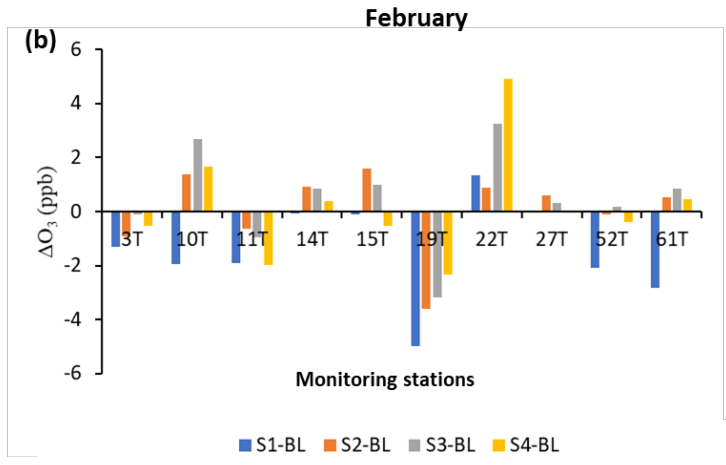
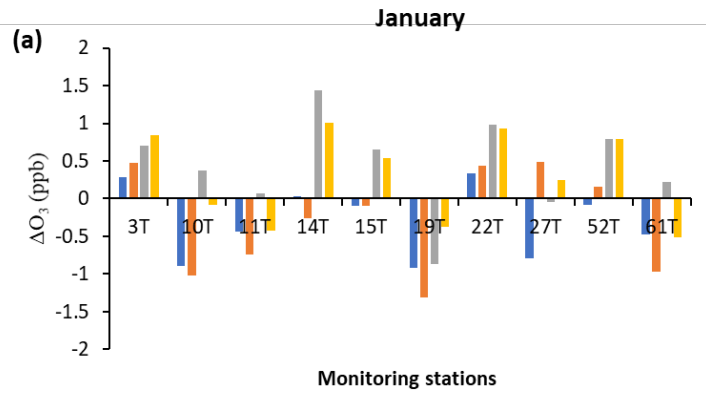


Figure 9. Change of O₃ levels due to China's emission control strategies at 10 monitoring stations in (a) January, (b) February and (c) March 2010 in the BMR. S1-BL, S2-BL, S3-BL and S4-BL refer to the differences between O₃ concentrations simulated from strategy 1, 2, 3 and 4 with baseline simulation, respectively.

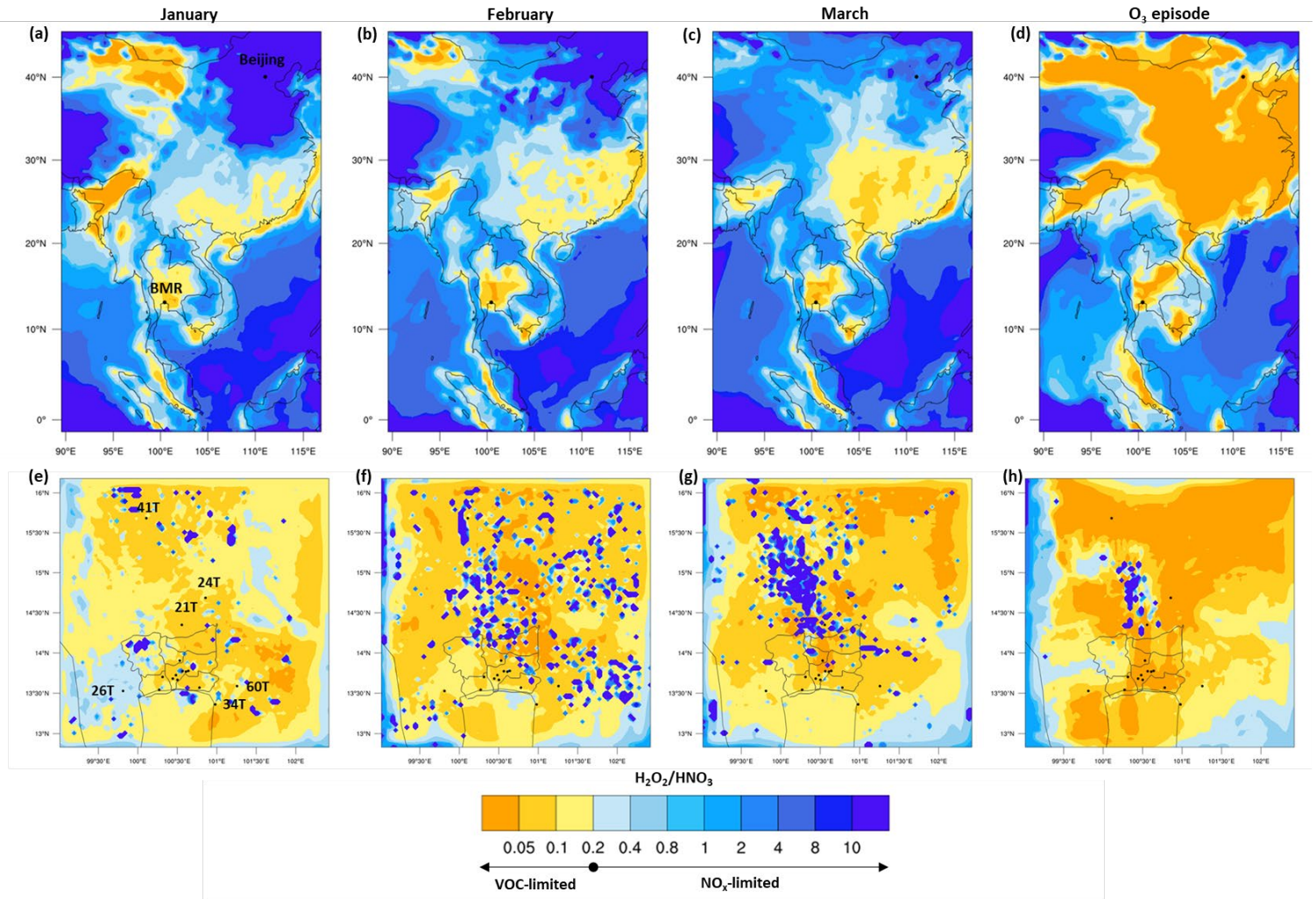


Figure 10. Spatial distributions of monthly-average and episodic-average $\text{H}_2\text{O}_2/\text{HNO}_3$ from the baseline simulation on the regional (d01) and local scales (d03) during a) and e) January, b) and f) February, c) and g) March and d) and h) the O_3 event 2010, respectively.

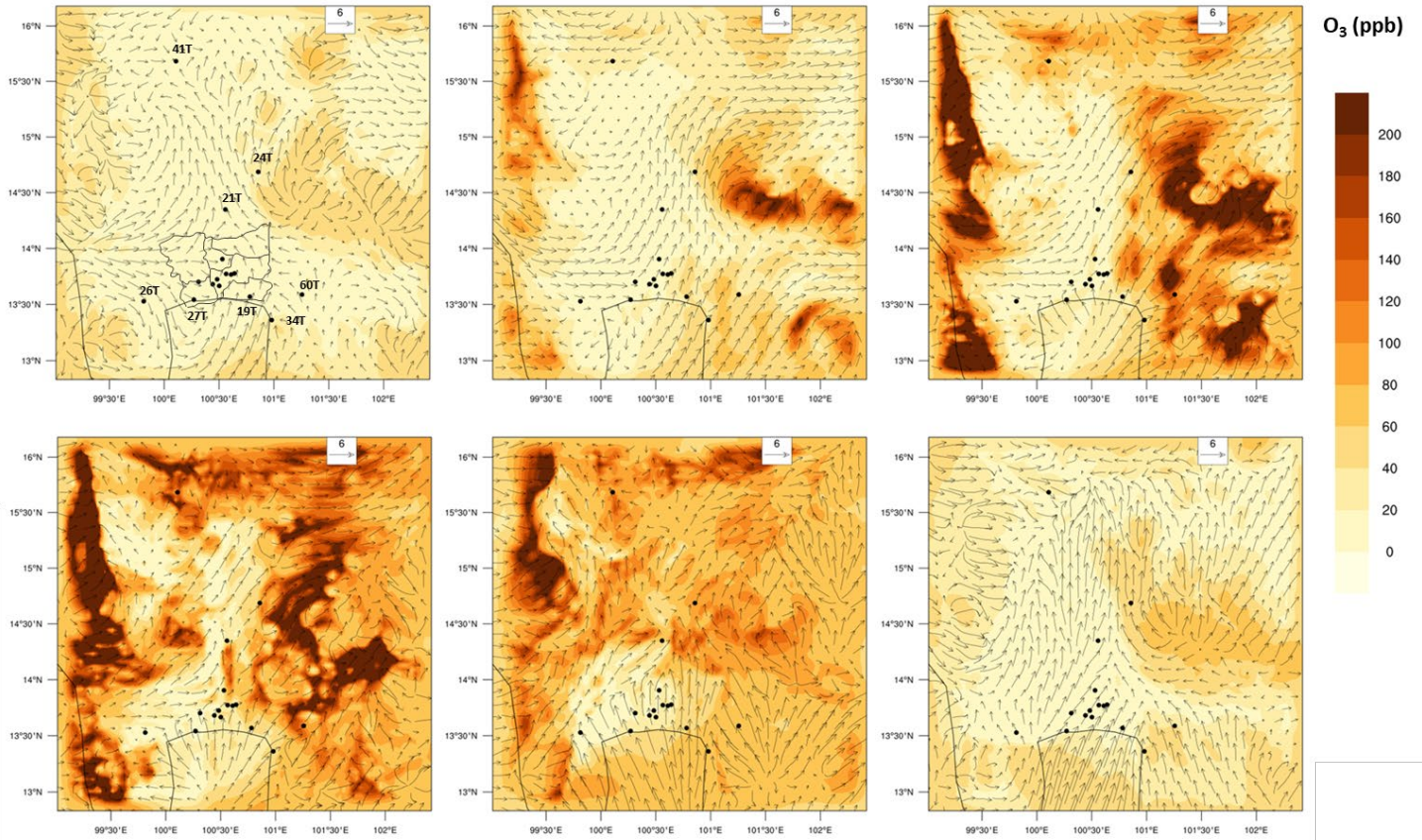


Figure 11. Diurnal cycle of O₃ in the local scale (d03) at (from left-right and top-down) 05:00 LT, 10:00 LT, 12:00 LT, 14:00 LT, 18:00 LT and 23:00 LT of 5 March 2010.