

# Article Detection of Strong NO<sub>X</sub> Emissions from Fine-scale Reconstruction of the OMI Tropospheric NO<sub>2</sub> Product

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Abstract: Satellite-retrieved atmospheric NO<sub>2</sub> column products have been widely used in assessing bottom-up NO<sub>X</sub> inventory emissions emitted from large cities, industrial facilities, and power plants. However, the satellite products fail to quantify strong NO<sub>X</sub> emissions emitted from the sources less than the satellite's pixel size, with significantly underestimating their emission intensities (smoothing effect). The poor monitoring of the emissions makes it difficult to enforce pollution restriction regulations. This study reconstructs the tropospheric NO<sub>2</sub> vertical column density (VCD) of the Ozone Monitoring Instrument (OMI)/Aura (13 × 24 km<sup>2</sup> pixel resolution at nadir) over South Korea to a fine-scale product (grid resolution of  $3 \times 3 \text{ km}^2$ ) using a conservative spatial downscaling method, and investigates the methodological fidelity in quantifying the major Korean area and point sources that are smaller than the satellite's pixel size. Multiple high-fidelity air quality models of the Weather Research and Forecast-Chemistry (WRF-Chem) and the Weather Research and Forecast/Community Multiscale Air Quality modeling system (WRF/CMAQ) were used to investigate the downscaling uncertainty in a spatial-weight kernel estimate. The analysis results showed that the fine-scale reconstructed OMI NO<sub>2</sub> VCD revealed the strong NO<sub>X</sub> emission sources with increasing the atmospheric NO<sub>2</sub> column concentration and enhanced their spatial concentration gradients near the sources, which was accomplished by applying high-resolution modeled spatial-weight kernels to the original OMI NO<sub>2</sub> product. The downscaling uncertainty of the reconstructed OMI NO<sub>2</sub> product was inherent and estimated by  $11.1\% \pm 10.6\%$  at the whole grid cells over South Korea. The smoothing effect of the original OMI NO<sub>2</sub> product was estimated by  $31.7\% \pm 13.1\%$  for the 6 urbanized area sources and  $32.2\% \pm 17.1\%$  for the 13 isolated point sources on an effective spatial resolution that is defined to reduce the downscaling uncertainty. Finally, it was found that the new reconstructed OMI NO<sub>2</sub> product had a potential capability in quantifying NO<sub>X</sub> emission intensities of the isolated strong point sources with a good correlation of R = 0.87, whereas the original OMI NO<sub>2</sub> product failed not only to identify the point sources, but also to quantify their emission intensities (R = 0.30). Our findings highlight a potential capability of the fine-scale reconstructed OMI NO<sub>2</sub> product in detecting directly strong NO<sub>X</sub> emissions, and emphasize the inherent methodological uncertainty in interpreting the reconstructed satellite product at a high-resolution grid scale.

**Keywords:** Air quality model; Anthropogenic emissions; CMAQ; NO<sub>2</sub> vertical column density; OMI; Spatial downscaling; WRF-Chem

## 1. Introduction

Nitrogen oxides ( $NO_X = NO + NO_2$ ) are major criteria pollutants forming photochemical ozone and particulate matters (PM) in the atmosphere, which are emitted from various anthropogenic



and biogenic sources such as fossil fuel combustion processes, road traffic exhaust, agricultural fertilizer use, biomass burning, and lightning (e.g., [1–3]). The nitrogen oxides and the secondary inorganic aerosols formed from NO<sub>X</sub> oxidation can impact on human health, ecosystems, and regional radiation budgets [4–6]. Park et al. [6] reported that the particulate matters formed by NO<sub>X</sub> contribute approximately 30.7% in aerosol optical depth (AOD) over East Asia. In Europe, it was shown that the effect of NO<sub>2</sub> from urban road traffic on premature mortality was 10 times higher than PM2.5 [7]. As the NO<sub>X</sub> emitted into the atmosphere rapidly reaches a photochemical equilibrium state given the meteorological and photochemical conditions, the atmospheric concentration measurements of NO<sub>2</sub> are used as a proxy of an urbanization level and NO<sub>X</sub> emission amount (e.g., [8–11]).

Satellite-based detection of the atmospheric NO<sub>2</sub> has been performed for past decades and has been used successfully to reveal anthropogenic and biogenic emission sources: Large cities, industrial complexes, and power plants emissions [12–14], agricultural emissions [15,16], lightning emissions [17,18], and wildfire emissions [19,20]. Many researchers have used the atmospheric NO<sub>2</sub> column concentration data in assessing the bottom-up anthropogenic emission inventories and in detecting temporal air quality changes in various geographical regions (e.g., [8–10,21–23]). Kim et al. [9] verified the bottom-up NO<sub>X</sub> emissions of major urban areas and power plants in the western United States using the atmospheric NO<sub>2</sub> column measurements from the Scanning Imaging Absorption Spectrometer for Atmospheric Chartography (SCIAMACHY) and the Ozone Monitoring Instrument (OMI)/Aura products and the regional air quality model of the Weather Research and Forecast-Chemistry (WRF-Chem). In the study, good agreement between the satellite-derived and modeled NO<sub>2</sub> columns of power plant emissions implied that the satellite retrievals are feasible to verify the bottom-up anthropogenic emissions monitored continuously in the U.S. Han et al. [21] compared the OMI tropospheric NO<sub>2</sub> column data from the Royal Netherlands Meteorological Institute (KNMI) to the Community Multiscale Air Quality modeling system (CMAQ) over the East Asia region in order to evaluate three different bottom-up emission inventories compiled in China, South Korea, and Japan. The study showed a capability of the satellite product in revealing high emission areas over the East Asia region and quantitatively verifying the bottom-up emissions in different seasons on a nationwide scale.

Despite the successful evaluation studies of bottom-up anthropogenic emissions on a coarse resolution (e.g., [24–27]), the satellite-retrieved NO<sub>2</sub> column data have various uncertainty factors associated with the data retrieval processes [9,28–30]. In order to improve the OMI NO<sub>2</sub> column product, Russell et al. [28] applied a priori NO<sub>2</sub> profiles obtained from the WRF-Chem model and high-resolution terrain information. Goldberg et al. [30] used high-resolution a priori NO<sub>2</sub> profiles obtained from the CMAQ model along with a spatial downscaling technique, resulting in a better agreement against in-situ NO<sub>2</sub> column measurements of the Pandora and Airborne Compact Atmospheric Mapper (ACAM) spectrometers than the original product. The improvement led to an increase of NO<sub>2</sub> column densities by 160% in urban areas and a decrease by 20–50% in rural areas compared to the operational NASA OMI product.

Another intrinsic limitation in the satellite products is a spatial smoothing effect caused by its coarse pixel resolution. The ground nadir pixel sizes of operational satellites detecting atmospheric NO<sub>2</sub> column abundance are  $40 \times 320$  km<sup>2</sup> in the Global Ozone Monitoring Experiment (GOME) instrument on board the second European Remote Sensing satellite (ERS-2) [13],  $30 \times 60$  km<sup>2</sup> in the SCIAMACHY instrument on board ENVISAT [12,31],  $13 \times 24$  km<sup>2</sup> in the OMI instrument on board the Aura [14], and  $40 \times 80$  km<sup>2</sup> in the GOME-2 instrument on board Metop-A1 [32]. Meanwhile, anthropogenic emission sources (e.g., cities, industrial complexes) are generally smaller in size than the satellites' single pixel, consequently steep gradients in atmospheric NO<sub>2</sub> columns are poorly represented near the source regions (spatial smoothing effect) (e.g., [8,33,34]). Heue et al. [33] compared the OMI and SCIAMACHY NO<sub>2</sub> column data over an industrial area in South Africa against an airborne NO<sub>2</sub> column measurement. The satellites detected spatial enhancement in atmospheric NO<sub>2</sub> column concentration over the source region, but failed to represent strong spatial gradients found in

the aircraft measurement which showed 4–9 times higher  $NO_2$  column densities over the source region than the OMI product.

The spatial smoothing effect embedded in the low-resolution satellite products is a critical barrier in quantifying strong emission sources, especially narrow-isolated sources with a scale of less than a few kilometers. The oversampling approach has been developed and improved in order to enhance spatial and temporal variations of satellite-detected pixel data at finer grid scales [35–38]. Various gridding algorithms have been applied with different complexities to generate gridded products from satellite-detected orbital signals. The simplest method is based on common spatial interpolation methods such as linear interpolation, spline interpolation, and kriging methods [39]. Meanwhile, the physics-based oversampling approach proposed by Sun et al. [38] is one of the advanced methods, in which a generalized 2-dimensional Gaussian function is used to represent the spatial response function of each pixel detected by a satellite sensor. Many studies have been done based on the gridded products, such as an emission estimate (e.g., [40,41]), an environmental exposure assessment (e.g., [42,43]), and source identification (e.g., [44–46]). The oversampling approach is able to generate spatially enhanced satellite products finer than their original pixel resolution by virtue of the temporally averaging multiple pixel data over a target grid, which simultaneously results in decreased temporal resolution. This may be conceptually valid when the concentrations detected by a satellite are assumed to be stationary during the averaging period. Directional composite averaging of chemical plumes following local wind direction in a specific source is a method that can be applied to overcome the limitation [37]. The spatial response function in Sun et al. [38] represents a sensitivity distribution of a satellite's sub-pixel in terms of the sensor's detection ability, and the information about spatial concentration distribution of chemical plumes within a satellite pixel is not taken into account, which is an apparent limitation of the oversampling method. On the other hand, Kim et al. [47] proposed a conservative downscaling method to reconstruct the spatial distribution of concentration within a satellite pixel, for which the spatial-weight kernel is calculated from high-resolution spatial concentration data simulated by a high-fidelity air quality model. Unlike the oversampling methods, the method is able to reconstruct the fine-scale spatial distribution of atmospheric concentration on each satellite pixel without the temporal averaging process. Meanwhile, the downscaled products depend on the fidelity of the air quality model applied to calculate the spatial-weight kernels of the satellite pixels, which causes additional downscaling errors in generating the fine-scale satellite products. Recent studies found the conservative downscaling method useful to represent strong spatial gradients over large source areas on a city scale, comparing well with in-situ measurements [27,30,48]. However, the downscaling errors associated with an air quality model are not investigated in those studies which used a single air quality model. This study aims to investigate the methodological fidelity of the conservative spatial downscaling method in quantifying strong  $NO_X$  emission sources over South Korea and the downscaling uncertainties associated with inclusion of multiple air quality models. To do this, the OMI NO<sub>2</sub> column data were reconstructed over South Korea using independent high-fidelity air quality models of WRF-Chem and CMAQ.

The manuscript is structured as follows: Section 2 describes the air quality models, the OMI NO<sub>2</sub> columns data, and the conservative spatial downscaling method. Section 3 presents the original and newly reconstructed OMI NO<sub>2</sub> products over South Korea. Further discussed are the methodological uncertainty, the smoothing effect, and the potential capability of the reconstructed OMI NO<sub>2</sub> product in quantifying the major point sources over South Korea. The summary and conclusions follow in Section 4.

## 2. Data and Method

#### 2.1. Regional Air Quality Models: WRF-Chem and WRF/CMAQ

The regional air quality models of WRF-Chem [49] and CMAQ [50] were applied to produce high-resolution spatial distribution of NO<sub>2</sub> concentration over the South Korean region. The WRF

model is a three-dimensional Eulerian meteorological model that solves dynamic/thermodynamic conservation equations of momentum, heat, moisture, and mass, and atmospheric physical processes of radiative transfer, turbulent mixing, surface-atmosphere interaction, and precipitation, thus being capable of simulating multiscale atmospheric phenomena [51]. The WRF-Chem model is an on-line air quality model that includes chemical transformation, source emissions, dry and wet deposition of chemical gaseous and aerosol species within the WRF model. The model simultaneously integrates the meteorological and chemical processes in a model's integration time step, thus is able to efficiently represent complex interactions among those physical and chemical processes. The CMAQ model is another three-dimensional Eulerian air quality model that has been developed primarily for regional air quality forecasts by the U.S. Environmental Protection Agency (EPA). It calculates the evolution of the atmospheric gaseous and aerosol species through representing meteorological and chemical processes. Meteorological fields are independently prepared for the model, and they are subsequently fed to the CMAQ model through a meteorology-chemistry interface processor (MCIP) package. In this study, the WRF model is used to produce meteorological fields required for various meteorological and (WRF/CMAQ). Both the air quality models have been widely applied for various meteorological and

and scientific investigations (e.g., [9,52–57]). The WRF-Chem (ver. 3.9.1) and WRF/CMAQ (ver. 3.6.1/ver. 4.7.1) models were configured with three domains to model atmospheric NO<sub>2</sub> column concentrations over South Korea (Figure 1). The outermost domain covered a large East Asia region including China, Japan, and South Korea, with a horizontal grid resolution of 27 km (174 × 128 mesh). The second and third nested domains were configured with a horizontal grid resolution of 9 km (69 × 90 mesh) and 3 km (180 × 225 mesh), respectively, covering the whole South Korean region. The vertical grids were stretched ranging from ~16 m above ground level at the lowest grid to ~20 km (50 hPa) at the domain top. The WRF-Chem model simulates both meteorological and chemical species at 35 sigma levels, while the CMAQ model calculates chemical species at 16 reduced sigma levels for computational efficiency. The MCIP produces the reduced meteorological fields compatible to the CMAQ using the WRF-simulated meteorological fields at 35 sigma levels.

environmental problems such as air quality forecasts, emissions evaluations, regulatory applications,



**Figure 1.** Simulation domains of the Weather Research and Forecast-Chemistry (WRF-Chem) and WRF/Community Multiscale Air Quality (CMAQ) models. The outermost domain has a grid resolution of 27 km and the nested domains denoted by the rectangles have a grid resolution of 9 km and 3 km, respectively. Color-coded are monthly mean anthropogenic NO<sub>X</sub> emission fluxes for April 2015.

The physical options used in the WRF simulations were identically configured with the Goddard scheme [58] and the rapid radiative transfer model (RRTM) [59] for shortwave and longwave radiation, the YSU scheme [60] for atmospheric boundary layer turbulence, the NoahLSM [61] for surface-atmosphere interactions, and the WSM3 scheme [62] and the Kain-Fritsch scheme [63] for

grid-scale microphysics and sub-grid convective cloud parameterization. The exceptions were the Dudhia shortwave radiation scheme [64] and the Grell-3D cumulus parameterization in the WRF-Chem. The WRF-Chem uses the regional atmospheric chemistry mechanism (RACM) [65] and the modal aerosol dynamics model for Europe/Secondary organic aerosol model (MADE/SORGAM) [66,67] for gaseous and aerosol chemical mechanisms, respectively, while the WRF/CMAQ uses the Statewide Air Pollution Research Center, Version 99 (SAPRC-99) [68] and the aerosol module version 5 AERO5 [69].

The Model Inter-Comparison Study for Asia 2010 (MICS Asia 2010) developed for the East Asia air quality project [70,71] was used to produce anthropogenic emissions. The MICS-Asia 2010 emissions were gridded compatible to the model domains using the sparse matrix operator kernel for emissions (SMOKE) [72] emission processing module. The chemical speciation of volatile organic compounds (VOCs) follows the SAPRC-99 for WRF/CMAQ and the RACM for the WRF-Chem, and the chemical conversion between the two mechanisms follows Lee et al. [55]. Biogenic emissions were produced by the model of emissions of gases and aerosols from nature, version 2 (MEGAN 2) [73]. Table 1 summarizes the physical and chemical processes used in simulations of the WRF-Chem and the WRF/CMAQ models.

Process	WRF-Chem (V3.9.1)	WRF/CMAQ (V3.6.1/V4.7.1)
Shortwave radiation	Goddard [58]	Dudhia [64]
Longwave radiation	RRTM [59]	
Turbulence	YSU [60]	
Land surface processes	Noah LSM [61]	
Microphysics	WSM3 [62]	
Cumulus parameterization	New Grell-3D	Kain-Fritsch [63]
Gas-phase chemistry	RACM [65]	SAPRC-99 [68]
Aerosol mechanism	MADE [66]/SORGAM [67]	AERO5 [69]
Anthropogenic emission	MICS-Asia 2010 [71]	
Biogenic emission	MEGAN-2 [73]	

**Table 1.** The configuration of the physical and chemical processes in the WRF-Chem and WRF/CMAQ models.

The simulations were carried for first 15 days of each January, April, July, and October (total 60 days) in 2015. The meteorological initial and boundary conditions of the outermost domain were obtained from the National Centers for Environmental Prediction-Final Analysis (NCEP FNL) that is a global reanalysis data with a spatial resolution of  $1^{\circ} \times 1^{\circ}$  and a temporal resolution of 6 hrs. The simulated meteorological fields for the outermost domain were nudged to the large-scale reanalyzed meteorological fields of air temperature, humidity, and winds using the 4-dimensional data assimilation (4DDA) technique [74]. The simulations of the nested domains were conducted by a one-way nesting approach.

## 2.2. Satellite Measurement: OMI/Aura tropospheric NO<sub>2</sub> Columns

The OMI tropospheric NO<sub>2</sub> vertical column densities (VCD) retrieved by the Royal Netherlands Meteorological Institute Dutch-OMI-NO2 version 2.0 (KNMI DOMINO v2.0) algorithm [75] were used. The OMI instrument, onboard the NASA Aura satellite, passed the equator at ~13:45 local standard time (LST) with a cross-track field of view angle of 114°, a swath width of 2600 km, and a nadir pixel size of 13 × 24 km<sup>2</sup>. The OMI tropospheric NO<sub>2</sub> VCD product was made through a series of data processing procedures. The KNMI OMI NO<sub>2</sub> slant column densities were calculated by a differential optical absorption spectroscopy (DOAS) technique, from which the tropospheric NO<sub>2</sub> column densities were calculated by separating the stratospheric and tropospheric contribution and subsequently by applying the tropospheric air mass factor that was obtained from the global chemical transport model TM4 and the radiative transfer model DAK [75]. The errors estimated during the retrieval procedures were  $\sim 0.25 \times 10^{15}$  molecules cm<sup>-2</sup> in the separation of stratospheric and tropospheric contribution,  $\sim 1.0 \times 10^{15}$  molecules cm<sup>-2</sup> in the application of the air mass factor, and  $\sim 0.7 \times 10^{15}$  molecules cm<sup>-2</sup> in the spectral fitting process.

Figure 2 shows the frequency distribution of overpassing days in terms of the mean pixel size of tropospheric NO<sub>2</sub> VCD passing over the South Korean region ( $126^{\circ}E-130^{\circ}E$ ;  $34^{\circ}N-8^{\circ}N$ ). The tropospheric NO<sub>2</sub> VCD data were available for 33 days (~55%) before the data quality control among 60 days of the simulation period and the average pixel sizes ranged from 312–1653 km<sup>2</sup> in the South Korean region. Approximately 33% of the days had a pixel size close to the satellite's nadir resolution, while approximately 30% had a mean pixel size of greater than 1000 km<sup>2</sup>. Meanwhile, the Korean cities have small urbanized areas of ranging from 10.5–227.5 km<sup>2</sup> when compared to the satellite pixel sizes, and the spatial sizes of strong point sources such as electric power plants, industrial facilities are much smaller than those of the cities. This indicates that fine-scale spatial distribution of the NO<sub>X</sub> emissions at a sub-pixel scale should be hardly detected in the OMI NO<sub>2</sub> VCD product (shaded in Figure 2).



**Figure 2.** The size distribution of the Aura/OMI pixels passed over the South Korean region (126°E–130°E, 34°N–38°N). The solid line denotes cumulative frequency in percentage. A median size of the overpassing pixels on a day was selected as a representative pixel size of the day. Total 33 days are analyzed. The shaded area indicates a range of the cities' sizes in South Korea.

The quality control procedure was applied to the OMI tropospheric NO<sub>2</sub> VCD data because the satellite products were significantly contaminated in the conditions of high cloud fractions and long slant paths. Previous studies generally adopted the filtering conditions of 15–40% for cloud fractions and 1000 km<sup>2</sup> for slant path lengths according to the data availability in the area of interest (e.g. [10,21,76–78]). This study filtered the NO<sub>2</sub> VCD with a cloud fraction of less than 40% and a pixel size of less than 1000 km<sup>2</sup> considering frequent cloudy days. Finally, approximately 20% of the OMI NO<sub>2</sub> VCD were used in reconstructing the satellite product.

### 2.3. Conservative Spatial Downscaling Method

The conservative spatial downscaling method proposed by Kim et al. [47] was used, which consists of two-step procedures: Conservative spatial re-gridding and conservative downscaling. The conservative spatial re-gridding procedure is to interpolate the OMI NO<sub>2</sub> VCD pixel data to high-resolution air quality model grids by weighting the overlapped areas between the satellite pixel and the model grids. The resulting gridded OMI NO<sub>2</sub> VCD data ( $OMI_i$ ) is calculated as follows:

$$OMI_j = \frac{\sum_{i}^{N} OMI_i A_{i,j}}{\sum_{i}^{N} A_{i,j}},$$
(1)

where  $OMI_i$  is the *i*th raw OMI NO-<sub>2</sub> VCD pixel value,  $A_{i,j}$  is the area fraction of  $OMI_i$  overlapped on the *j*th model grid, and *N* is the number of the OMI pixels overlapped on the *j*th model grid. Each model grid meets  $\sum_{i}^{N} A_{i,j} = 1$ . For example, when two OMI NO<sub>2</sub> VCD pixels are overlaid on a model grid *j*, the gridded OMI NO<sub>2</sub> VCD is calculated as  $(OMI_1A_{1,j} + OMI_2A_{2,j})/(A_{1,j} + A_{2,j})$ . The OMI NO<sub>2</sub> VCD is simply re-gridded to the model grids without any changes in spatial distribution.

The conservative downscaling procedure is to calculate the gridded OMI NO<sub>2</sub> VCD data ( $OMI_DS_j$ ) by changing their spatial distribution according to the modeled spatial distribution as follows:

$$OMI\_DS_j = \frac{\sum_{i}^{N} OMI_i K_{i,j} A_{i,j}}{\sum_{i}^{N} A_{i,j}},$$
(2)

where  $K_{i,j}$  is a spatial-weight kernel defined by a ratio of modeled NO<sub>2</sub> VCD to the OMI NO<sub>2</sub> VCD. Here, the spatial-weight kernel was constructed on each *i*th OMI pixel using high-resolution modeled NO<sub>2</sub> VCD in order to conserve the OMI NO<sub>2</sub> VCD after applying the spatial-weight kernel. The averaged  $OMI_DS_j$  on the OMI pixel was identical to the OMI NO<sub>2</sub> VCD pixel value. In this study, the three-dimensional distribution and evolution of the atmospheric NO<sub>2</sub> VCD over South Korea were made by the two independent regional air quality models of the WRF-Chem and WRF/CMAQ so as to produce the spatial-weight kernel  $K_{i,j}$ . In the modeled NO<sub>2</sub> VCD calculation, the KNMI OMI averaging kernel (AK) was applied to the model's vertical layers of the atmospheric NO<sub>2</sub> VCD. It has been known that the averaging kernel is sensitive to the retrieved NO<sub>2</sub> VCD [21,78,79].

## 2.4. Area and Point Sources of NO<sub>X</sub> Emissions in South Korea

Several major cities and isolated point sources with high NO<sub>X</sub> emissions over South Korea were selected in order to investigate methodological fidelity of the conservative spatial downscaling method in quantifying the emissions from the sources. The sources were selected based on the annual NO<sub>X</sub> emissions reported in the South Korean emission inventory data of the Clean Air Policy Support System (CAPPS). Figure 3 shows the spatial distribution of the monthly mean NO<sub>X</sub> emissions over South Korea along with geographic locations of the major sources. The major Korean cities are denoted by area sources (A), while electric power plants and industrial facilities are denoted by point sources (P) (Figure 3a). The NO<sub>X</sub> emissions of the area sources were well correlated with the sizes of the urbanized area, while the point sources' emissions range according to their activities (Figure 3b).



**Figure 3.** (a) The spatial distribution of anthropogenic  $NO_X$  emissions used for the simulations of the WRF-Chem and WRF/CMAQ models. A and P denote the highly urbanized area sources and the isolated point sources in South Korea, respectively. (b) A comparison of the annual  $NO_X$  emissions of the selected area and point sources with their urbanized area. The source area of the point emissions is fixed by  $3\times3$  km<sup>2</sup> for clarity.

#### 3. Results

## 3.1. OMI-detected and Modeled NO<sub>2</sub> VCD

Figure 4 compares the spatial distributions of the re-gridded OMI NO<sub>2</sub> VCD and the modeled NO<sub>2</sub> VCD by the WRF-Chem and WRF/CMAQ models. The modeled NO<sub>2</sub> VCD were averaged for the satellite's valid pixels applying with the KNMI OMI's averaging kernels. Large source areas are clearly seen with enhanced concentrations both in the OMI-detected and the modeled NO<sub>2</sub> VCD, but distinguishable differences are also apparent. The OMI NO<sub>2</sub> VCD is able to discriminate strong NOx emissions from large urbanized areas, but fails to identify strong point emissions (Figure 4a). Previous studies that showed a similar spatial pattern of the KNMI OMI NO<sub>2</sub> VCD over the South Korea region at different time periods (e.g., [21,80]) considered the nation-wide spatial average of the OMI NO<sub>2</sub> VCD in their analysis. Meanwhile, both the modeled NO<sub>2</sub> VCD show more clearly enhanced signals of the strong point sources as well as the urbanized area sources than the OMI NO<sub>2</sub> product (Figure 4b,c), which is also spatially well correlated with the distribution of NO<sub>X</sub> emissions in Figure 3. The two models simulate similarly the spatial distributions over the South Korea, but the NO<sub>2</sub> VCD modeled by the WRF-Chem tends to be higher than the WRF/CMAQ.



**Figure 4.** The spatial distributions of tropospheric NO<sub>2</sub> VCD over the South Korean region averaged for a period of January, April, July, and October 2015. (a) KNMI OMI NO<sub>2</sub> VCD, (b) WRF-Chem modeled NO<sub>2</sub> VCD, and (c) WRF/CMAQ modeled NO<sub>2</sub> VCD.

Figure 5 compares the OMI-detected and modeled NO<sub>2</sub> VCD to evaluate the model performance of the WRF-Chem and WRF/CMAQ in simulating the atmospheric NO<sub>2</sub> column concentrations over the South Korean region. The monthly mean OMI NO<sub>2</sub> VCD was  $3.6 \times 10^{15}$  molec. cm<sup>-2</sup> in July and  $9.0 \times 10^{15}$  molec. cm<sup>-2</sup> in January, with an annual mean value of  $6.7 \times 10^{15}$  molec. cm<sup>-2</sup>. Han et al. [21] reported a similar seasonal change of the KNMI OMI NO<sub>2</sub> VCD ranging from  $3.4-6.7 \times 10^{15}$  molec. cm<sup>-2</sup> over the South Korean region in 2006. Kim et al. [80] reported approximately  $4.0 \times 10^{15}$  molec. cm<sup>-2</sup> in summer and  $7.0 \times 10^{15}$  molec. cm<sup>-2</sup> in winter 2010 in the KNMI OMI NO<sub>2</sub> VCD Level 3 product. Meanwhile, the WRF-Chem model simulated the higher NO<sub>2</sub> VCD levels by 50–145% than the WRF/CMAQ model during the periods. The NO<sub>2</sub> VCD modeled by the WRF/CMAQ compared better to the satellite-retrieved atmospheric NO<sub>2</sub> column concentrations than the WRF-Chem model. This discrepancy may partly be attributed to the deactivation of aqueous chemistry processes and the configuration of higher vertical grid resolution in the WRF-Chem simulation, but both the models' NO<sub>2</sub> VCD values were within the 1-sigma error range of the satellite product. The atmospheric NO<sub>2</sub> VCD simulated by the WRF-Chem and WRF/CMAQ models were used to calculate the spatial-weight kernels in the conservative spatial downscaling method.



**Figure 5.** A comparison of the monthly mean KNMI OMI-detected and modeled NO<sub>2</sub> VCD over South Korea. The atmospheric NO<sub>2</sub> VCD were averaged over the South Korean region (126°E–130°E; 34°N–38°N). The vertical bar in the KNMI OMI data denotes 1-sigma error.

#### 3.2. Fine-scale Reconstructed OMI NO<sub>2</sub> VCD

Figure 6 presents the reconstructed OMI NO<sub>2</sub> VCD over the South Korean region by applying the spatial-weight kernels obtained from the independent air quality models. The newly reconstructed OMI NO<sub>2</sub> VCD have a spatial resolution of 3 km in accordance with the models' grid resolution. Figure 7 compares the re-gridded and reconstructed OMI NO<sub>2</sub> VCD over a large urbanized area (Seoul metropolitan area) and an area with a few isolated strong point sources. The reconstructed OMI NO<sub>2</sub> VCD can reveal the highly urbanized area sources as well as the strong isolated point sources with enhanced spatial gradients and the higher level of column concentrations (Figure 6), which is well compared to the  $NO_X$  emissions (Figure 3a). In the Seoul metropolitan area, the OMI NO<sub>2</sub> VCD downscaled by the WRF-Chem and WRF/CMAQ had values of  $3.27 \pm 0.75 \times 10^{16}$  molec.  $cm^{-2}$  (maximum 4.53 × 10<sup>16</sup> molec.  $cm^{-2}$ ) and 3.43 ± 0.68×10<sup>16</sup> molec.  $cm^{-2}$  (maximum 4.89 × 10<sup>16</sup> molec. cm<sup>-2</sup>), respectively, whereas the re-gridded OMI NO<sub>2</sub> VCD had relatively low values of  $2.63 \pm 0.40 \times 10^{16}$  molec. cm<sup>-2</sup> (maximum of  $3.33 \times 10^{16}$  molec. cm<sup>-2</sup>) (Figure 7a–c). During the Megacity Air Pollution Studies-Seoul (MAPS-Seoul) field campaign conducted in May-June 2015, the atmospheric NO<sub>2</sub> VCD detected by the Pandora spectrometer ranged from 1.7-1.9 Dobson units  $(DU, 1 DU \approx 2.7 \times 10^{16} \text{ molec. cm}^{-2})$  (4.59-5.13 × 10<sup>16</sup> molec. cm<sup>-2</sup>) during 12-14 LST at a site within the Seoul metropolitan area [81]. The fine-scale reconstructed OMI NO<sub>2</sub> VCD products compares better to the surface in-situ measurement than the original OMI NO<sub>2</sub> VCD. Meanwhile, the capability of the reconstructed OMI NO<sub>2</sub> products are apparent in quantifying the emissions from the strong isolated point sources (Figure 7d-f). The strong point sources with a similar amount of NO<sub>X</sub> emissions were cement industrial facilities separated with a distance of ~20 km. The reconstructed OMI NO<sub>2</sub> VCD by the WRF-Chem and WRF/CMAQ had values of  $2.23 \pm 0.86 \times 10^{16}$  molec. cm<sup>-2</sup> (maximum  $4.55 \times 10^{16}$  molec. cm<sup>-2</sup>) and  $2.13 \pm 0.89 \times 10^{16}$  molec. cm<sup>-2</sup> (maximum  $4.35 \times 10^{16}$  molec. cm<sup>-2</sup>), respectively. The signals of the strong point sources were clearly separated in the reconstruct OMI NO<sub>2</sub> products, whereas the re-gridded original OMI NO<sub>2</sub> product failed to identify the point sources showing relatively lower concentration levels of  $1.50 \pm 0.12 \times 10^{16}$  molec. cm<sup>-2</sup> (maximum  $1.70 \times 10^{16}$ molec. cm<sup>-2</sup>).

Figure 8 compares the simulated tropospheric NO<sub>2</sub> VCD and the reconstructed OMI NO<sub>2</sub> VCD by the WRF-Chem and WRF/CMAQ models. The tropospheric NO<sub>2</sub> VCD modeled by the WRF-Chem were spatially well correlated with those by the WRF/CMAQ (R = 0.84), but the WRF-Chem model simulated higher NO<sub>2</sub> VCD than the WRF/CMAQ (Figure 8a). On the other hand, a better spatial correlation (R = 0.97) was found in the reconstructed OMI NO<sub>2</sub> VCD products (Figure 8b). The similarity in the reconstructed OMI products shows that the conservative spatial downscaling method is less sensitive to the concentration level of the modeled NO<sub>2</sub> VCD.



**Figure 6.** The spatial distributions of the fine-scale reconstructed OMI NO<sub>2</sub> VCD over the South Korean region obtained by applying the spatial-weight kernels from (**a**) WRF-Chem and (**b**) WRF/CMAQ.



**Figure 7.** The spatial distributions of the re-gridded and the reconstructed OMI NO<sub>2</sub> VCD over (**a**–**c**) Seoul metropolitan area and (**d**–**f**) an area with industrial complexes. (**a** and **d**) re-gridded KNMI OMI NO<sub>2</sub> VCD, (**b** and **e**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstructed OMI NO<sub>2</sub> VCD using the WRF-Chem, (**c** and **f**) reconstr



**Figure 8.** A comparison of (**a**) the modeled NO<sub>2</sub> VCD and (**b**) the reconstructed OMI NO<sub>2</sub> VCD by the WRF-Chem and WRF/CMAQ models over the South Korean region. The modeled NO<sub>2</sub> VCD in (**a**) were calculated by applying the KNMI OMI averaging kernels to the modeled vertical profiles of NO<sub>2</sub> concentrations.

The relative differences between the reconstructed OMI NO<sub>2</sub> VCD products in Figures 6-8 are interpreted as a methodological uncertainty associated with the conservative spatial downscaling

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method. The downscaling uncertainty is inherent, but has not been quantitatively analyzed in previous studies (e.g., [30,47]). Figure 9 presents the spatial distribution of the reconstructed OMI NO<sub>2</sub> VCD averaged by the WRF-Chem and WRF/CMAQ models and the frequency distribution of grid-scale normalized absolute difference (NAD). The NAD is calculated as a downscaling uncertainty metric  $|OMI_DS_{WRF-Chem}-OMI_DS_{WRF/CMAQ}|$ by  $\frac{100M-2WRF-Chem}{0.5(OMI_DS_{WRF-Chem}+OMI_DS_{WRF/CMAQ})}$ . It is clear that the averaged OMI NO<sub>2</sub> product reveals highly urbanized areas and strong point sources over the South Korea region (Figure 9a). The downscaling uncertainty in the reconstructed OMI NO<sub>2</sub> VCD is relatively large near the strong area and point sources such as cities, industrial complexes, and power plants, resulting from a dissimilar shape of the simulated plumes (Figures 6 and 7). The NAD ranged from  $11.1\% \pm 10.6\%$  (maximum 131.2%) for the whole grid cells and  $12.0\% \pm 11.5\%$  (maximum 108.9%) at the high-concentration grid cells of >8 × 10<sup>15</sup> molec. cm<sup>-2</sup>. The grid-scale NAD values are less than 10% (20%) at approximately 58% (85%) of the whole grid cells, which is similar in the high-concentration grid cells (Figure 9b). These results indicate that the fine-resolution reconstructed OMI NO<sub>2</sub> product has a potential capability in detecting steep concentration gradients found near the strong area and point sources with a reasonable error range compared to the original OMI NO<sub>2</sub> VCD (Figure 4a).



**Figure 9.** (a) The spatial distribution of the reconstructed OMI NO<sub>2</sub> VCD over the South Korean region and (b) the frequency distribution of the grid-scale normalized absolute difference defined as  $\frac{|OMI\_DS_{WRF-Chem}-OMI\_DS_{WRF/CMAQ}|}{0.5(OMI\_DS_{WRF-Chem}+OMI\_DS_{WRF/CMAQ})}.$ 

## 3.3. Effective Spatial Resolution and Detection Capability of Strong NOx Emissions

Despite the strong emissions with enhanced spatial gradients and the concentrations being successfully identified in the reconstructed satellite product, the reconstructed OMI NO<sub>2</sub> VCD had somewhat large downscaling uncertainty at the grid resolution of 3 km. Therefore, the effective spatial resolution was further investigated by changing an averaging box size near the sources to see how much the downscaling uncertainty reduces. The averaging box sizes of the emission sources increased gradually from  $3 \times 3$  grid cells ( $9 \times 9$  km<sup>2</sup>) to  $23 \times 23$  grid cells ( $69 \times 69$  km<sup>2</sup>). Figure 10a,b compare the re-gridded and the reconstructed OMI NO<sub>2</sub> VCD and the associated downscaling uncertainties over the Seoul metropolitan area and an industrial facility site as a function of the averaging box size. For the large urbanized area (Figure 10a), the reconstructed OMI NO<sub>2</sub> VCD had higher concentration levels than the re-gridded satellite product by a maximum of 30.0%. The difference between the re-gridded and reconstructed OMI NO<sub>2</sub> VCD decreases gradually with an averaging box size until the averaging area is large enough to cover the source region. The downscaling uncertainty in NAD also gradually decreases from  $\sim 7\%$  at 225 km<sup>2</sup> (15 × 15 km<sup>2</sup>). For the isolated industrial facility site (Figure 10b), the tropospheric NO<sub>2</sub> VCD enhancement in the reconstructed satellite product is apparent by 72.8% at the maximum and decreases rapidly with the averaging area. The estimated NAD is approximately 19% at 9 km<sup>2</sup> ( $3 \times 3$  km<sup>2</sup>) as a maximum value, and it decreases rapidly with an averaging box size. The downscaling uncertainty of each area and point sources estimated in NAD are presented for the 6 urbanized area sources and the 13 isolated point sources in Figure 10c,d. The NAD decreases gradually from 17% to 2%, on average, for the urbanized area sources as the averaging area increases, while it decreases more rapidly from 26% to 2% for the isolated point sources. It is also

found that the downscaling uncertainty depends on the sources' characteristics such as the spatial size and distribution of emissions, and the complexities in local meteorology.



**Figure 10.** A comparison of the re-gridded and reconstructed OMI NO<sub>2</sub> VCD in terms of an averaging box size for (**a**) Seoul metropolitan area and (**b**) a site with cement industrial facilities in South Korea. The normalized absolute difference of the reconstructed OMI NO<sub>2</sub> VCD is given for (**c**) 6 urbanized area sources and (**d**) 13 isolated point sources. Solid lines in (c) and (d) represent the mean normalized absolute difference for the urbanized area sources and the isolated point sources.

The smoothing effect of the original OMI  $NO_2$  product in quantifying the atmospheric  $NO_2$ column enhancements for the 6 urbanized area and the 13 isolated point sources over South Korea was investigated. A large variability in the downscaling uncertainties of the area and point sources leads to consideration of the effective spatial resolution that is determined by an appropriate averaging box size to reduce the downscaling uncertainties in a reasonable level. The NAD values reduced to less than  $\sim 10\%$  for the urbanized area sources when the averaging area was 4–5 times of the source's aerial size. Meanwhile, the averaging area of  $5 \times 5$  grid cells ( $15 \times 15$  km<sup>2</sup>) for the most isolated point sources led to a significant reduction in NAD ranging within 3-25% except for a few point sources where the downscaling uncertainty still remained as large as 37-40%. Figure 11 compares the re-gridded and the reconstructed OMI  $NO_2$  VCD for the 6 urbanized area sources and the 13 isolated point sources over South Korea on the effective spatial resolution. The reconstructed OMI  $NO_2$ VCD were well correlated with the original product for the urbanized area sources (R = 0.99) and the isolated point sources (R = 0.87). However, the reconstructed OMI NO<sub>2</sub> VCD were higher by 18-54% for the urbanized area sources and by 7-60% for the isolated point sources than the original OMI NO2 product, which is caused by the inability of the satellite's pixels to resolve the spatial heterogeneity of the sources. In this study, the smoothing effect of the original OMI NO<sub>2</sub> product was evaluated by  $31.7\% \pm 13.1\%$  for the urbanized area sources and  $32.2\% \pm 17.1\%$  for the isolated point sources on the effective spatial resolution.

Finally, a potential capability of the reconstructed OMI NO<sub>2</sub> product in quantifying the NO<sub>X</sub> emission intensity for isolated strong point sources was investigated. Figure 12 compares the re-gridded and the reconstructed OMI NO<sub>2</sub> VCD of the 13 isolated point sources against the 2015 annual NO<sub>X</sub> emission of each point source. The NO<sub>X</sub> emissions are profiled in the Korean national emission inventory database from the direct measurements of stack emission intensity of the point sources. The reconstructed OMI NO<sub>2</sub> VCD were well correlated with the emission intensity of the point sources (R = 0.87), whereas the original OMI product showed only a weakly positive correlation (R = 0.30). The original OMI product shows significant limitations not only in identifying the strong emission sources

with a satellite's sub-pixel scale but also in quantifying the atmospheric  $NO_2$  column enhancements corresponding to the emission intensity. In contrast, the newly reconstructed OMI  $NO_2$  product shows a great potential capability in quantifying the emissions of the isolated point sources within a reasonable range of downscaling uncertainties.



**Figure 11.** A comparison of re-gridded and reconstructed OMI NO<sub>2</sub> VCD on an effective spatial resolution for (**a**) 6 urbanized area sources and (**b**) 13 isolated point sources in South Korea.



**Figure 12.** A comparison of the re-gridded and reconstructed OMI NO<sub>2</sub> VCD of the 13 major isolated point sources in South Korea against their annual NO<sub>X</sub> emissions. The vertical bar denotes the downscaling uncertainty estimated for the fine-scale reconstructed OMI NO<sub>2</sub> product.

## 4. Summary and Conclusions

Large spatial detection coverage of the satellite-retrieved NO<sub>2</sub> VCD is beneficial in evaluating air quality models and bottom-up inventory NO<sub>X</sub> emissions. However, the satellite's coarse pixel resolution remains a critical limitation in quantifying a strong area and/or point sources and steep gradients on a satellite's sub-pixel scale. This study reconstructed the KNMI OMI NO<sub>2</sub> VCD ( $13 \times 24$  km<sup>2</sup> at nadir) over South Korea on a spatial resolution of  $3 \times 3$  km<sup>2</sup> using a conservative spatial downscaling method to investigate the methodological fidelity in quantifying the atmospheric NO<sub>2</sub> column enhancements of the major area and point sources over the South Korean region. Multiple air quality simulations by the WRF-Chem and WRF/CMAQ models were performed over the South Korea region for four different months (60 days) to investigate the downscaling uncertainty in calculating spatial-weight kernels in the conservative spatial downscaling method.

The results showed that the newly reconstructed OMI NO<sub>2</sub> VCD based on high-fidelity fine-scale air quality modeling revealed clearly the strong area and point sources with enhanced spatial heterogeneity and steep gradients in accordance with their emission intensities. This is a critical weakness in the original OMI NO<sub>2</sub> product that significantly underestimates the atmospheric NO<sub>2</sub> column

enhancements over the area and point sources. The downscaling uncertainty of the reconstructed OMI NO<sub>2</sub> product was estimated by 11.1%  $\pm$  10.6% at the whole grid cells over South Korea due to the use of multiple air quality models in the spatial-weight kernel estimate, which is not negligible. The smoothing effect in the original OMI NO<sub>2</sub> product was evaluated, on average, by 31.7%  $\pm$  13.1% for the 6 urbanized area sources and 32.2%  $\pm$  17.1% for the 13 isolated point sources on the effective spatial resolution that is defined by an averaging box size of about 4 times of the urban size and 5  $\times$  5 grid cells (15  $\times$  15 km<sup>2</sup>) for the area and point sources, respectively. Without considering the geometric smoothing effect, the surface emissions are easily misleading in terms of their intensity and steep spatial gradient. Finally, the potential capability of the reconstructed OMI product was tested in quantifying the isolated strong point sources that are listed in the Korean emission inventory with direct monitored annual NO<sub>X</sub> emissions. This showed that the isolated strong point sources with different emission intensities are clearly identified in the reconstructed OMI NO<sub>2</sub> VCD with a good correlation of 0.87. In contrast, the original OMI NO<sub>2</sub> product totally fails to identify the strong point sources and quantify their emission intensities, showing a weakly positive correlation with the emission intensity (R = 0.30).

This study highlights a potential capability of the fine-scale reconstructed OMI NO<sub>2</sub> product in quantifying directly strong NO<sub>X</sub> emissions and their steep spatial gradients and thus in assessing bottom-up inventory emissions for isolated areas and/or point sources that have a spatial scale of less than the satellite's pixel resolution. It also emphasizes that the methodological uncertainty associated with a spatial-weight kernel estimate should be considered cautiously in interpreting the reconstructed satellite product at a high-resolution grid scale.

The traditional oversampling methods are more beneficial in identifying hidden  $NO_X$  emissions, but the spatial concentration distribution may be poorly represented due to a relatively coarse spatial and temporal resolution of the product. In contrast, the conservative downscaling method is better in representing fine-scale spatial distribution of identified  $NO_X$  emissions on a finer spatial and temporal resolution of product. However, the application may be limited to detect unknown sources. Therefore, the oversampling and conservative downscaling methods can be used complementarily.

More research is needed to understand the downscaling uncertainty associated with a priori inventory emissions used in the air quality models, which is another uncertainty factor in the conservative downscaling method. In addition, further validation of the fine-scale reconstructed OMI NO<sub>2</sub> product can be useful if high-resolution in-situ measurements and/or remote sensing data are available.

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