

1 Changes in the macrobenthic infaunal community of the Southern California continental margin  
2 over five decades in relation to oceanographic factors

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## 15 Abstract

16 Climate change has altered the physiochemical conditions of the coastal ocean, but effects on  
17 infaunal communities have not been well-assessed. Here we used multivariate ordination to  
18 examine temporal patterns in benthic community composition from four southern California  
19 continental shelf monitoring programs that range in duration from 30 to 50 years. Temporal  
20 changes were compared to variations in temperature, oxygen and acidification using single-taxon  
21 random forest models. Species richness increased over time, coupled with a decline in overall  
22 abundance. Continental shelf macrobenthic communities from the 2010s were comprised of a  
23 broader array of feeding guilds and life history strategies than in the 1970s. Changing water  
24 temperature was associated with northward shifts in geographic distribution and increases in  
25 species abundance, while acidification was associated with southward shifts and declines in  
26 abundance of other species. Acidification was also associated with changes in depth distribution  
27 of benthic fauna, with shelled molluscs declining in abundance at depths most associated with  
28 increasing exposure to acidification. This broad-scale community-level analysis establishes  
29 causal hypotheses that set the stage for more targeted studies investigating shifts in abundance or  
30 distribution for taxa that appear to be responding to climate change-related disturbances.

31    **Keywords**

32    Ocean Acidification, ENSO, PDO, Range Shifts, Continental Shelf, Continental Slope

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## 1. INTRODUCTION

35      Climate change has led to marked effects on the coastal ocean, with increasing water temperature  
36      (Thomson & Krassovski 2010, Sutton & Bowen 2019, Johnson & Lyman 2020, Muff et al.  
37      2022) and enhanced intensity of El Niño and Pacific Decadal Oscillation events (Wang et al.  
38      2017, Freund et al. 2019, Cai et al. 2021). This has led to altered circulation patterns (Thomson  
39      & Krassovski 2010, Bograd et al. 2019) and shoaling of deep basin low dissolved oxygen and  
40      low pH waters (Bograd et al. 2008, Gilly et al. 2013). The carbonate chemistry of coastal waters  
41      has been further altered by acidification of surface waters due to increases in atmospheric CO<sub>2</sub>  
42      (Doney et al. 2009, Hauri et al. 2013, Feely et al. 2016).

43      Changes in oceanographic conditions and water quality/water chemistry have been linked to  
44      changes among planktonic, pelagic, and intertidal communities (Poloczanska et al. 2016, Stiasny  
45      et al. 2016, Algueró-Muñiz et al. 2017, Lemasson et al. 2017, Espinel-Velasco et al. 2018, Leis  
46      2018, Wang et al. 2018, Barclay et al. 2019, Pinsky et al. 2020). These changes have led to the  
47      alterations to food webs and to reduced fisheries productivity (Harley et al. 2006, Stiasny et al.  
48      2016, Jin et al. 2020, Wilson et al. 2020). Climate-related changes in ocean waters have also  
49      been linked to the habitat compression or geographic shifts of epibenthic organisms (Sato et al.  
50      2017).

51      However, the links between climate change or ocean acidification and infaunal benthic  
52      communities are less well understood. Warming waters may increase the ranges of tropical and  
53      sub-tropical fauna, while squeezing out colder water taxa (Wesławski et al. 2011, Pinsky et al.  
54      2013, Hiddink et al. 2015, McClatchie et al. 2016, Goransson 2017). However, the  
55      microenvironments infauna create in their sediment burrows and the naturally reduced conditions

56 of anerobic sediments may serve to de-couple infaunal adults from water column dynamics and  
57 insulate them from short-term water temperature shifts (Furukawa et al. 2001, Widdicombe &  
58 Spicer 2008, Silburn et al. 2017, Michaud et al. 2021). In contrast, there are clear effects of  
59 ocean acidification on the larvae of many benthic species, with decreases in aragonite saturation  
60 levels being particularly harmful to mollusc and echinoderm larvae (Waldbusser et al. 2015,  
61 Espinel-Velasco et al. 2018, Bednaršek et al. 2021). Furthermore there is growing evidence that  
62 overlying water with elevated temperatures and lower pH can negatively impact the functioning  
63 and productivity of adult and juvenile infauna established in the sediments – especially filter  
64 feeders (Green et al. 2009, Nixon et al. 2009, Widdicombe et al. 2009, Clements & Hunt 2017,  
65 Vlaminck et al. 2023).

66 Macrobenthic infauna are an ideal lens through which to view the influence of climate change on  
67 the resident biota of a location. Most species are relatively sessile and many live for multiple  
68 years, so patterns in abundance, biomass, and composition are an integrator of local  
69 environmental conditions (Gray & Elliott 2009). Furthermore, most marine benthic communities  
70 have relatively high taxonomic diversity, with a single sample containing dozens of species  
71 across multiple phyla (Ellingsen 2002, Villnäs & Norkko 2011, Gillett et al. 2021). This species  
72 diversity increases the likelihood of capturing differential responses to differing types of stress  
73 and therefore the ability to characterize stressor-specific assemblage responses (Lenihan et al.  
74 2003, Thrush et al. 2008, Rodil et al. 2013).

75 The coastal ocean of the Southern California Bight provides a unique setting to track temporal  
76 changes in biotic and abiotic conditions in the benthos. It is an ecologically and  
77 oceanographically complex region situated in a biogeographic transition zone, with colder water  
78 Oregonian fauna north of Point Conception, CA and warmer water Californian fauna to the

79 south, which leads to high biodiversity of benthic and pelagic fauna (e.g., Wares et al. 2001,  
80 Briggs & Bowen 2013, Claisse et al. 2018) (<https://scb.marinebon.org/>).

81 At broad scales, the region is oceanographically influenced at shallower depths by the cold-water  
82 California Current flowing to the south mixing with the warm-water Davidson Countercurrent  
83 flowing to the north (Bray et al. 1999), as well as seasonal upwelling of nutrient-rich water  
84 (Chhak & Di Lorenzo 2007). At depths below 300 to 400 m, the region is influenced by the  
85 northward flowing California Bottom Current, which transports relatively warmer, low oxygen  
86 Pacific Equatorial subsurface water along the continental slope (e.g., Thomson & Krassovski  
87 2010). The relative interplay of the bottom and surface water masses is influenced by El Niño /  
88 Southern Oscillation cycling from year-to-year, with La Niña events bringing greater amounts of  
89 subsurface water to the surface along the coastal zone and El Niño events having the opposite  
90 effect (e.g., Bograd et al. 2019).

91 The bottom topography of the continental slope and shelf, as well as the presence of the Channel  
92 Islands, create eddies in the northern most parts of the Southern California Bight (Oey 1996,  
93 Harms & Winant 1998, Kessouri et al. 2022). These meso-scale features contribute to the  
94 oceanographic heterogeneity of the region. There are distinct north – to – south gradients in  
95 water temperature of the shallow mixed layer that become more muted with increasing depth  
96 (Gelpi & Norris 2008). Similarly, northern waters of the region are exposed to acidic conditions  
97 more frequently and at shallower depths than the southern portions of the region (McLaughlin et  
98 al. 2018).

99 The Southern California Bight is also a region experiencing oceanographic change, with  
100 documented temporal alterations in oxygen (Booth et al. 2014) and acidification (McLaughlin et  
101 al 2018) dynamics. Coincidentally, the Southern California Bight is home to four well-curated

102 benthic invertebrate monitoring programs that span 30-50 years (City of Los Angeles  
103 Environmental Monitoring Division 2019, City of San Diego Ocean Monitoring Program 2020,  
104 Los Angeles County Sanitation Districts 2020, Orange County Sanitation District 2021). Here  
105 we used these four data sets to determine if there have been changes in the infaunal benthic  
106 community and to characterize the nature of those changes in the Southern California Bight. We  
107 also investigated relationships between changes in community composition and changes in  
108 oceanographic patterns and water characteristics like the El Niño-Southern Oscillation (ENSO),  
109 the Pacific Decadal Oscillation (PDO), dissolved oxygen, water temperature, and proxy-based  
110 estimates of ocean acidification.

111 **2. METHODS**

112 **2.1 Benthic Data**

113 Infaunal abundance and taxonomic data, as well as sediment grain size were obtained from the  
114 annual and semi-annual benthic monitoring programs of (northernmost to southernmost) the City  
115 of Los Angeles Environmental Monitoring Division (CLAEMD), the Los Angeles County  
116 Sanitation District (LACSD), the Orange County Sanitation District (OCSan) and the City of San  
117 Diego Ocean Monitoring Program (CSD). (Figure 1, Supplement 1 – Table S1). From these four  
118 programs, minimally disturbed sites used as reference condition benchmarks within these  
119 programs were selected at three different depths: mid-continental shelf (~60m), outer continental  
120 shelf (98-156m), and upper continental slope (~300m). Note that only the LACSD and OCSan  
121 locations had sites located at the continental slope depths. Summer (July – September) data were  
122 used for analysis.

123 The majority of the samples were collected with a 0.1m<sup>2</sup> modified Van Veen grab. Samples from  
124 the Los Angeles County Sanitation Districts collected before 1980 (i.e., eight sampling events)  
125 were collected as four replicate 0.04m<sup>2</sup> Shipek grabs. Given the smaller sample area of these  
126 older samples, benthic data from the first three replicate samples were summed together to  
127 approximate the samples collected with the Van Veen grab post-1980 – an approach developed  
128 by the Los Angeles County Sanitation Districts (S. Walther, pers. comm.).

129 Sediment from the grabs was sieved on a 1-mm screen, with the retained material fixed in  
130 buffered formalin before transfer to ethanol for preservation. All fauna were enumerated and  
131 identified to lowest possible taxonomic level, typically species. The names of all fauna from the  
132 different years were harmonized and updated to meet Southern California Association of Marine  
133 Invertebrate Taxonomists (SCAMIT) edition 12 (The Southern California Association of Marine  
134 Invertebrate Taxonomists 2018) conventions, with ambiguous taxa aggregated to higher  
135 taxonomic levels on a sample-wise basis to help ensure comparability of fauna across the  
136 temporal span of the dataset.

137 Sediment grain size data were available from LACSD, CLAEMD, and CSD data sets. Grain size  
138 composition by dry weight were measured as % sand (Phi -0.5 – 4.0), % silt (Phi 4.5 – 8.0), and  
139 % clay (Phi >8.5). The % sand was measured as material retained on a series of nested sieves  
140 between 2-mm and 63-µm. Percentages of silt and clay were measured by the pipette method  
141 (e.g., Plumb 1981) prior to 1990 and via a laser sediment analyzer thereafter (e.g., Beuselinck et  
142 al. 1998). Grain size data from CSD were only reported as % sand and % mud, where % mud is  
143 the sum of silts and clays.

144 2.2. Oceanographic Data

145 Patterns in ENSO and PDO were characterized using the National Oceanographic and  
146 Atmospheric Administration National Centers for Environmental Information (NOAA NCEI)  
147 SOI and PDO indices (Mantua & Hare 2002). Data were obtained from 1/1/1970 – 5/1/2020  
148 from <https://www.cpc.ncep.noaa.gov/data/indices/soi> (SOI) and  
149 <https://www.ncdc.noaa.gov/teleconnections/pdo/> (PDO). Monthly values for each index were  
150 averaged into quarterly seasonal values.

151 Sea surface and bottom water temperatures, dissolved oxygen, and salinity were obtained from  
152 the California Oceanic Fisheries Investigations (CalCOFI) hydrographic bottle dataset  
153 (<https://calcofi.org/ccdata.html>), which were collected quarterly using either Niskin/Wally  
154 bottles deployed on a vertical wire until 1993 or by a CTD+ bottle rosette frame after 1993.  
155 Detailed methods for collection and quality control are available at <https://calcofi.org/about-calcofi/methods.html>. Hydrographic data were matched to benthic sampling locations by first  
156 determining the closest CalCOFI monitoring station to each benthic station with the most  
157 available data (Figure 1). Horizontal proximity to the benthic stations was determined by  
158 calculating the geodesic distance between bottle sampling station and benthic fauna station using  
159 latitude and longitude via the distGeo function in the geosphere package (v1.5-10 [Hijmans  
160 2019]) in R (v3.6.1). CalCOFI stations were between 16.1 (LACSD Upper-Slope) and 24.6 km  
161 (OCSan Mid-Shelf) away from their respective benthic data stations, distances that – while not  
162 directly overlying each other – represented the most reasonable approximations of the ocean  
163 water masses the benthic sites were exposed to (e.g., Dong et al. 2009, Watson et al. 2011,  
164 Bograd et al. 2015, Kekuewa et al. 2022). Data from all bottles within +/- 20m of the benthic

166 station depth were selected and averaged to represent the bottom values for that benthic station.

167 The temperature from the 25m bottle was used as the surface water temperature.

168 Ocean acidification metrics were estimated from CalCOFI temperature, salinity, and oxygen data

169 using the proxy approach developed by Juranek et al. (2009) and Alin et al. (2012), which have

170 been previously applied to CalCOFI time-series data in Alin et al. (2012) and McClatchie et al.

171 (2016). Multiple linear regression equations were developed using the nortest (Gross & Ligges

172 2015) and robust (Wang et al. 2022) packages within R to estimate pH on the total scale ( $\text{pH}_T$ ),

173  $p\text{CO}_2$ , aragonite saturation state ( $\Omega_{\text{arag}}$ ), and calcite saturation state ( $\Omega_{\text{calc}}$ ) from proxy

174 temperature, salinity, and oxygen data using a calibration dataset consisting of inorganic carbon,

175 oxygen, and CTD observations collected on NOAA West Coast Ocean Acidification (WCOA)

176 cruises from 2007 to 2016 (Feely & Sabine 2013, Feely et al. 2015a,b, 2016, Alin et al. 2017).

177 Models were tuned to 25–300 m depth, with separate empirical relationships for the northern

178 (34°N–33.2°N) and southern (33.2°N–32.5°N) portions of the region. Final equations were

179 selected on the basis of having the lowest root mean squared errors (RMSE) and highest adjusted

180  $R^2$  values for the estimated parameters in the calibration data set (Supplement 1 – Table S2).

181 In comparison with the methods and results in Alin et al. (2012): 1) we did not standardize

182 variables within equations by subtracting mean parameter values, 2) higher standard error (SE)

183 values on coefficients likely reflect the very small numbers of calibration observations for this

184 study ( $n = 21$  and 42 samples, in the northern and southern regions, respectively), and 3) RMSE

185 and  $R^2$  values remain excellent, as the calibration data were tightly constrained to bottom depths

186 of 25–300 m within northern and southern regions so that the calibration data would be most

187 appropriate for this nearshore, benthos-focused study. It should be noted that the application of

188 these equations to data collected prior to the calibration dataset cruises (i.e., 2007) will likely

189 underestimate  $\Omega$  and pH and overestimate  $p\text{CO}_2$  values to some degree for the earliest parts of  
190 the time-series because we did not account for the increase in anthropogenic  $\text{CO}_2$  content across  
191 the decades of this study (cf. Feely et al. 2016). As a consequence, the magnitude of change in  
192 ocean acidification metrics ( $\Omega$ , pH, and  $p\text{CO}_2$ ) across the decades are likely underestimated by  
193 our results, as similarly noted in McClatchie et al. (2016). Furthermore, there was most likely  
194 some degree of inter-relatedness between ENSO and PDO cycles with water temperature,  
195 dissolved oxygen and low pH bottom water (via upwelling) at the sampling sites. However,  
196 separating the source of changes in carbonate chemistry and temperature (atmospheric vs.  
197 upwelled) were beyond the scope of this work, which was focused on their (aggregate) effects on  
198 the fauna.

199 For most of the years within the 1970-2020 period of interest, water quality data were collected  
200 during four quarterly cruises per year. However, due to changes within the CalCOFI program,  
201 there were no measures of temperature, dissolved oxygen, or acidification variables for 1970,  
202 '71, '73, '77-'80, and '82-'83. For bottom dissolved oxygen, surface temperature, and bottom  
203 water temperature, summer (July – September) values (i.e., concurrent with the benthic sampling  
204 period) were selected for comparison to benthic faunal patterns. For the acidification variables,  
205 the least acidified estimates (i.e., highest pH,  $\Omega_{\text{arag}}$ , and  $\Omega_{\text{calc}}$ , or lowest  $p\text{CO}_2$ ) across the four  
206 quarterly sampling events in a given year were selected to compare to benthic faunal patterns.

### 207 [2.3. Analytical Approach](#)

208 Changes in benthic faunal community composition were characterized using non-metric  
209 Multidimensional Scaling (nMDS) ordination. Bray-Curtis dissimilarity values were calculated  
210 based on presence-absence data (equivalent to Sørensen dissimilarity) through time at each  
211 sample site. Dissimilarity values were ordinated 2-d nMDS across a minimum of 250 iterations.

212 A 1-way Permutation Analysis of Variance (PERMANOVA) was then used to quantify the  
213 influence of time on community structure from each sampling site. PERMANOVAs were  
214 conducted (10,000 permutations) on Bray-Curtis dissimilarities as the response variable and year  
215 of collection as the predictor variable. The correlation between year of collection and the pattern  
216 in 2-d ordinations was also calculated across 1,000 permutations. All similarity calculations,  
217 ordinations, PERMANOVAs, and correlations were conducted in the vegan package (v2.6-2)  
218 (Oksanen et al 2022) in R (v3.6.1) using the metaMDS, adonis2, and envfit functions.

219 Taxa from each sample site were grouped in categories of shelled and non-shelled organisms.  
220 The relationships of these taxonomic groups with water quality/chemistry and oceanographic  
221 patterns at each site were quantified using random forest regression. Regression models were  
222 structured with faunal abundance as the response variable and the water quality and  
223 oceanographic variables detailed in Table 1 as the potential predictor variables across 10,000  
224 trees per relationship. The relative influence of each predictor variable on the abundance of the  
225 taxonomic groups through time was quantified from % Mean Square Error change values for  
226 each predictor in the models. Random forests regressions were calculated using the  
227 randomForest and importance functions within the randomForest package (v4.6-16) (Liaw &  
228 Wiener 2002) in R (v3.6.1).

229 Changes in the presence and absence of frequently observed individual taxa throughout time  
230 were characterized using logistic regression. The presence of each taxon observed in 10 or more  
231 years at each sampling site was treated as the response variable, with year of collection as the  
232 predictor variable. An alpha value of 0.05 was used to select taxa whose presence significantly  
233 changed over the span of the sampling period. Taxa were characterized as increasers – more  
234 likely to be observed in modern samples – or decreasers – less likely to be observed in modern

235 samples – based upon the sign of the beta term in the logistic regression equation. Polychaetes,  
236 crustaceans, echinoderms, and molluscs that were identified to the level of family or higher were  
237 omitted from consideration. The abundance of each increaser or decreaser taxon was then  
238 modeled with the suite of water quality/chemistry, sediment, and oceanographic variables  
239 detailed in Table 1 as predictor variables across 10,000 trees in a random forest regression. The  
240 relative influence of each predictor variable on the abundance of each taxon through time was  
241 quantified from % Mean Square Error change values for each predictor in the models. Random  
242 forests regressions were calculated using the randomForest and importance functions within the  
243 randomForest package (v4.6-16) (Liaw & Wiener 2002) in R (v3.6.1).

244 Changes in species richness and total abundance at each sample site and depth zone were both  
245 characterized using simple linear least squares regression. Species richness or total abundance  
246 were set as the response variable, with year of collection as the predictor variable. Regressions  
247 were done using the lm function in R (v3.6.1).

248 Geographic changes through time among frequently observed taxa were characterized using  
249 linear least squares regression. Abundance-weighted latitude for each taxon observed in 10 or  
250 more years was treated as the response variable and the year of collection was the predictor  
251 variable. Year-specific, abundance-weighted latitude for each taxon was calculated by summing  
252 across all four sampling sites the product of latitude at a sample site multiplied by relative  
253 abundance (sample abundance/total abundance) of a taxon within that year (EQ 1). An alpha  
254 value of 0.05 was used to select taxa whose location significantly changed over the span of the  
255 sampling period. Taxa were characterized as shifting northwards – taxa centered in higher  
256 latitudes in more modern samples – or shifting southwards – taxa centered in lower latitudes in  
257 more modern samples – based upon the sign of the beta term in the linear regression equation.

258

259 EQ1 - *Abundance Weighted Latitude (Sample Site i(CLAEMD → CSD)*

267 
$$\sum_i Latitude at Sample Site * \frac{Abundance at Sample Site}{Total Abundance}$$

268

260 Abundance of each northward or southward taxon was then modeled with the suite of water  
261 quality/chemistry and oceanographic variables detailed in Table 1 as predictor variables across  
262 10,000 trees in a random forest regression. The relative influence of each predictor variable on  
263 the abundance of each taxon through time was quantified from % Mean Square Error change  
264 values for each predictor in the models. Random forests regressions were calculated using the  
265 randomForest and importance functions within the randomForest package (v4.6-16) (Liaw &  
266 Wiener 2002) in R (v3.6.1).269 

### 3. RESULTS

270 

#### 3.1 Patterns in Oceanographic Data

271 From the 1970s through the 2010s there was greater spatial variability in the selected  
272 oceanographic measures across the different depth zones and sampling sites than temporal  
273 variability at any given location. (Figure 2). Spatial patterns largely followed expectations, with  
274 the water masses being colder, saltier, denser, more acidified, less oxygenated, and with more  
275 nitrate as depth increased. The mid-shelf depth waters (~60m) were the most variable from year  
276 to year, while the slope depth waters (~300m) were the most stable. Inspection of the time-series  
277 data in Figure 2 suggests an increasing trend in temperature of mid-shelf depth waters at the  
278 northern sampling sites and in the upper-slope depth waters at the southern sampling site.

279 Furthermore, there is a suggestion of declining dissolved oxygen across all depths and sampling  
280 sites from the early 1980's through 2010's. The pattern amongst the modelled carbonate  
281 chemistry variables suggests a more acidified/corrosive conditions in the outer-shelf depth  
282 waters (~150m) of the northern sample sites. The carbonate chemistry variables fluctuated year-  
283 to-year from the 1970s – 2010s, but were non-corrosive at mid-shelf depths across all three  
284 sampling sites.

285 **3.2 Patterns in Benthic Data**

286 The nMDS plots (Figure 3) illustrate a pattern of changes in benthic community composition  
287 through time in all four sampling regions and across all depth zones. While there is year-to-year  
288 variation in ordination space, the general trend, as illustrated by the year correlation vector, is a  
289 relatively unidirectional change from the 1970s through the 2010s. The most visually distinct  
290 patterns are observable in the northernmost sample locations (LACSD and CLAEMD) (Figure 3,  
291 panels A-E). Quantitatively, both the multivariate correlation (Table 2) and the PERMANOVA  
292 (Table 3) results support the visual interpretation that the year of collection was significant ( $\alpha =$   
293 0.05) from the perspective of both the 2-D ordination (correlation) and the underlying  
294 dissimilarity relationships among samples (PERMANOVA). Analysis of these patterns using  
295 either presence/absence data, as done here, or abundance data (Supplement 2) did not change the  
296 clear pattern of community composition change through time.

297 The top-ten most abundant taxa for each decade within each depth zone from the LACSD and  
298 OCSan samples sites are presented in Table 4. Tracking these taxa through time presents a  
299 similar pattern to the nMDS ordinations of Figure 3, where there was a shift in community  
300 dominants across the decades. Of the 10 taxa that were community dominants at a given depth  
301 and location in the 1970s, an average of only 3.6 of these across the two programs were still

302 dominant taxa in the 2010s. At mid-shelf depths, there was a shift from bivalves, polychaetes  
303 (cirratulids, lumbrinerids, and terebellids), and ostracods in the 1970s, giving way to ophiuroids,  
304 amphipods, and spionid polychaetes in the 1980s and 1990s. The mid-shelf communities of the  
305 2000s and 2010s were dominated by ophiuroids and a variety of polychaetes, with no shelled  
306 molluscs among the top-ten most abundant taxa. The outer-shelf communities show a similar  
307 break between 1970s and 1980s taxa in contrast to more modern samples. The outer-shelf was a  
308 bivalve and ophiuroid dominated community before 1990 that shifted to an ophiuroid and  
309 polychaete dominated community in the 2000s and 2010s. The dominant taxa of the upper-slope  
310 from the 1970s and 1980s were relatively persistent through 2000s compared to the other depth  
311 zones. The upper-slope communities had a mix of polychaetes and amphipods consistently in the  
312 top-10 taxa through the 2000s, with a shift to different species of polychaetes and molluses in the  
313 2010s.

314 Many taxa were inconsistently observed through time, with only 8 to 25% of the taxa at a given  
315 site occurring in 10 or more years (Table 5a). It is important to note that the consistency of  
316 detection of a given taxon through time within our data set may have potentially been impaired  
317 for some taxa due to the regional standard operating practice of using only a single benthic grab  
318 to characterize benthic communities combined with high biodiversity of the region. Despite this,  
319 the logistic regressions of taxa presence vs. absence indicated that 105 taxa had an increasing  
320 probability of being observed in more modern years – increaser taxa – and 44 had a decreasing  
321 probability of being observed in more modern years – decreaser taxa (details in Supplement 3 –  
322 Table S5). These patterns can be illustrated by the ampeliscid amphipod *Ampelisca hancocki* or  
323 the travisiid polychaete *Travisia brevis* – both increasers – versus the tellinid bivalve *Macoma*  
324 *carlottensis* or entropneusts – both decreasers (Figure 4). Nearly all of the taxa that displayed a

325 temporal trend had consistent increaser/decreaser patterns across all sampling locations and  
326 depth zones, with the exception of three taxa that increased in some sites and decreased in others:  
327 the bivalve *Axinopsida serricata* decreased in the mid-shelf and upper-slope of northern sites  
328 (CLAEMD and LACSD), while increasing in the outer-shelf of the southern location (CSD); the  
329 sigalionid polychaete *Sthenelanella uniformis* decreased in southern outer-shelf depths while  
330 increasing at the northern mid-shelf locations; lastly, the spionid polychaetes in the  
331 *Spiochaetopterus costarum* complex decreased in mid-shelf central locations (OCSan) but  
332 increased in outer-shelf depths.

333 The random forest models of the individual increaser and decreaser taxa through time indicates  
334 that there were relatively distinct influences on the two types of taxa. The most influential  
335 predictors (based upon their effect on model mean square error) of the abundance of decreaser  
336 taxa through time were associated with ocean acidification or dissolved oxygen at all three depth  
337 zones, as well as ENSO/PDO oceanographic variables for decreaser taxa from the outer-shelf  
338 and sediment composition for decreaser taxa from the upper-slope (Figure 5). Conversely, the  
339 top three most influential variables for increaser taxa were related to changes in temperature in  
340 all depth zones and ENSO/PDO in the mid and outer-shelf zones. Relatively few increaser taxa  
341 were influenced by acidification or dissolved oxygen patterns. Sediment grain size was an  
342 important predictor for approximately 20% of the taxa, most commonly among polychaetes at  
343 upper-slope depths. Supplement 3 (Tables S6 and S7) contains the detailed variable importance  
344 data for each increaser or decreaser taxon.

345 The greater number of increaser taxa versus decreaser taxa across the sampling sites was echoed  
346 in the overall trend of species richness through time. All sites had significantly ( $\alpha = 0.1$ )  
347 increasing taxa richness through time except mid-shelf depths at CSD and upper-slope depths at

348 LACSD (Figure 6). In contrast, the total abundance within a given sample significantly ( $\alpha = 0.1$ )  
349 declined through time in the mid-shelf depths of the CLAEMD, OCSan, and CSD sites, as well  
350 as at the outer-shelf and upper-slope depths of the LACSD site (Figure 7). Total abundance  
351 increased at the OCSan outer-shelf sites and there was no discernable trend at the other sample  
352 sites.

353 Eighty-six taxa shifted their geographic center of abundance northward or southward across the  
354 study region (Supplement 3 – Table S8). Most shifts occurred among mid-shelf taxa, with 37  
355 taxa shifting northward and only 7 shifting southward (Table 5b). Conversely, the distributions  
356 of more taxa from the outer-shelf (29) and upper shelf (3) shifted southward than shifted  
357 northward (18 and 0, respectively). These patterns are illustrated in Figure 8 with the northward  
358 shifts in distribution of the spionid polychaete *Prionospio dubia* and the ampharetid polychaete  
359 *Asabellides lineata* (both mid-shelf). Conversely, southward shifts in distribution are illustrated  
360 by the spionid polychaete *Prionospio jubata*, the ampeliscid amphipod *Ampelisca pacifica*, and  
361 the thyasirid bivalve *Adontorhina cyclia* along the outer-shelf, or the scaphopod mollusc  
362 *Rhabdus rectius* on the upper-slope. The center of distribution for 51 of the 86 taxa shifted by  
363 more than 1 degree of latitude (a maximum of 1.24 degrees) north or south within our sampling  
364 area.

365 The random forest models of individual taxa whose distributions shifted northward or southward  
366 indicated a mix of influences on the two types of taxa (Figure 9). Of the taxa whose distribution  
367 shifted northward, temperature and acidification variables were the most influential predictors of  
368 abundance for more than 50% of the taxa on the mid-shelf and the outer-shelf, while  
369 oceanographic and dissolved oxygen predictors were less (<30%) frequently important to mid-  
370 shelf and outer-shelf taxa. Among those taxa whose distribution shifted southwards, acidification

371 was the only clearly important predictor for outer-shelf taxa and acidification and dissolved  
372 oxygen were similarly important for mid-shelf taxa. As noted above, only 3 upper-slope taxa had  
373 significantly southward shifted distributions and only 1 taxon had a northward shifted  
374 distribution. For these taxa acidification and dissolved oxygen predictors were important for all  
375 of the northward and southward taxa, with temperature also important for 2 of the 3 southward  
376 taxa (full details in Supplement 3 – Tables S9 and S10). Note that ENSO/PDO oceanographic  
377 variables were not influential predictors for any southward shifting taxa at upper-slope depths.

378 A visual inspection of the relative abundance of shelled and non-shelled taxa through time  
379 showed distinct changes in the fauna through time. Figure 10 highlights the pattern in all three  
380 depth zones from the LACSD sampling sites. The upper-slope location, where there is the  
381 greatest, consistent exposure to acidified waters, and the outer-shelf location, where exposure has  
382 increased in more recent decades, showed relatively high abundance of shelled organisms in the  
383 1970s and early 1980s. This was then followed by an initial sharp decline in the mid-1980s,  
384 short-term recovery, and a longer-term decline in the mid 1990's (Figure 10). In contrast, the  
385 abundance of non-shelled organisms showed a less distinct pattern through time, with fluctuating  
386 abundance that would be expected with most fauna. At mid shelf depths, where there is little  
387 expected exposure to corrosive waters, the relative abundance of shelled organisms declined  
388 briefly in the early 1990s, but recovered to normal levels soon after, while the abundance of non-  
389 shelled organisms was relatively consistent through time. Plots of shelled and non-shelled  
390 organisms from all depths and sites are presented in Supplement 4. The decline in shelled  
391 organisms at outer-shelf depth was not as drastic at the centrally located OCSan site, with a  
392 muted decline in the mid-1990s and recovery through the 2000s. It was not apparent at all at the  
393 southern-most, CSD sampling site.

394 The random forest regression models of total shelled fauna abundance indicated in the northern  
395 and centrally located mid-shelf stations (CLAEMD, LACSD, and OCSan) that acidification  
396 predictors were most frequently selected as influential variables. Temperature and oceanographic  
397 variables were most influential at the southern mid-shelf station. At outer-shelf depths,  
398 acidification variables were never identified as influential predictors of shelled fauna abundance,  
399 whereas DO, temperature and oceanographic variables were. In the upper-slope locations,  
400 acidification variables were most frequently identified as influential predictors, though DO,  
401 temperature, and oceanographic variables were all selected once too. Of the individual shelled  
402 taxa that could be classified as increaser or decrease taxa, the likelihood of observing 8 taxa  
403 decreased through time and 13 increased (Supplement 3 – Table S5) across all three depth zones.  
404 Of the individual shelled taxa that could be classified as northward or southward shifting taxa,  
405 the distribution of 4 shifted northwards and 8 shifted southwards (Supplement 3 – Table S8).

406 **4. DISCUSSION**

407 This study presents empirical evidence for ongoing temporal changes in the composition of the  
408 macrobenthic communities of the coastal ocean of Southern California. The changes appear to be  
409 gradual and relatively unidirectional at a decadal scale, with oscillation of community dominants  
410 and secondary taxa from year-to-year at all of the sampling locations and depth zones. Beyond  
411 demonstrating the change, we were able to characterize the nature of the change. Some taxa  
412 became more frequently observed in modern samples than in the past, while others commonly  
413 observed in the past were rarely observed in modern times. Similarly, there were detectable shifts  
414 in the geographic distribution of a number of taxa, shifting either northward or southward across  
415 the breadth of the Southern California Bight. Most significantly, we were able to provide insight

416 into the relative influence of changing ocean conditions on these appearances, disappearances, or  
417 geographic shifts. The bulk of the decreasing taxa were driven by changes in carbonate  
418 chemistry and dissolved oxygen regimes of their habitats, while the increasing taxa were more  
419 frequently linked to changes in water temperature and ENSO/PDO cycles.

420 Our findings for benthic fauna are similar to that of Hale et al. (2018) on the Atlantic Coast of  
421 the US. When considering the autecology and natural history of the taxa across the decades,  
422 there were shifts away from a bivalve (thyasirids, lucinids, tellinids) and polychaete (spionids  
423 and cirratulids) dominated community towards a community dominated by amphipodid  
424 ophiuroids, amphipods, and a more functionally diverse array of polychaetes. These data indicate  
425 that the communities of the continental shelf and slope of the Southern California Bight are,  
426 within recent record, comprised of a majority of deposit and interface feeding taxa, which makes  
427 sense given the depth of the water and separation from the photic zone. However, the shifts from  
428 a community dominated by lucinid bivalves (i.e., *Axinopsida serricata* and *Parvilucina*  
429 *tenuisculpta*) and deposit/interface-feeding polychaetes to one with an array of additional feeding  
430 modes from predatory polychaetes (e.g., *Lumbrineris* spp., *Scoletoma tetraura*) and crustaceans  
431 (e.g., *Metaphoxus frequens*, *Rhepoxynius* spp.) to true filter feeders (e.g., *Phoronis* sp.)  
432 represents a broadening of the realized ecological niche space in the habitat. The broadening of  
433 niche space occupied by the fauna, may in part account for the increasing species richness  
434 combined with declining total abundance observed across the decades in the dataset (Cardinale et  
435 al. 2009, Niklaus et al. 2017).

436 We would suggest that the scope of the temporal community composition changes we observed  
437 was ecologically relevant. To give that change context, the difference in composition between  
438 samples from the 1970s and the 2010s was equivalent to differences in composition between a

439 reference condition site and one disturbed by anthropogenic activities. As an illustration, the  
440 Bray-Curtis dissimilarity of presence/absence data between the 1974 and 2018 sample from the  
441 LACSD mid-shelf site was 0.75. The same dissimilarity measure between a mid-shelf reference  
442 condition site and a disturbed site from a 2018 Southern California Bight regional survey – using  
443 a similar type of grab, as well as the same sieve size and taxonomic standard – was 0.73 (Gillett  
444 et al. 2022). The nature of the changes was different – one is a shift from an older community to  
445 a more modern community, the other a shift from an intact modern assemblage to a pollution  
446 tolerant assemblage – but both pairs represent an approximately 75% difference in sample  
447 composition.

448 There are likely multiple mechanisms for the change in community composition and increase in  
449 species richness observed over the decades. All of the sample sites were selected to minimize  
450 influence of local human disturbance (i.e., wastewater outfalls, dredging, trawling). However, the  
451 dominance of the lucinid bivalves and deposit feeders in older samples (i.e., typical indicators of  
452 organic matter enrichment) versus a broader array of feeding types in newer samples could be  
453 reflective of regional recovery from anthropogenic pollution in the earlier part of the 20<sup>th</sup> century  
454 (Leonard-Pingel et al. 2019, Los Angeles County Sanitation Districts 2020, Orange County  
455 Sanitation District 2021). This recovery could partially explain the species richness and  
456 abundance patterns that were observed (e.g., Diaz et al. 2008). However, the degree of  
457 disturbance observed at the sampling sites across the length of the dataset was relatively  
458 minimal, with nearly all of the samples within a reference or low disturbance category (following  
459 Smith et al. 2001, Gillett et al. 2022) (Figure 11).

460 An alternative, non-environmental explanation one could posit for both the increased species  
461 richness and the differences in taxa observed across time might be changes in the science of

462 taxonomy (Isaac et al. 2004, Agapow & Sluys 2005, Morrison et al. 2009). Taxonomy changes  
463 as new species are erected from within old polyphyletic “species”, local taxonomic precision  
464 changes, or the names change due to refined precedence of descriptions. This can be of particular  
465 concern with data records as long as those used in this study, during which the taxonomist  
466 personnel in each of the monitoring program has changed numerous times. However, that does  
467 not appear to be a major contributing factor to the patterns observed in the present study. The  
468 nature and magnitude of community change we observed was similar across four independent  
469 programs whose taxonomists did not change at the same times. Furthermore, the declines in  
470 abundance that were observed concurrently with the increases in species richness are unlikely to  
471 occur if species names were just being split or refined from family to species. In fact, the relative  
472 taxonomic stability across our data sets is a tribute to the Southern California Association of  
473 Marine Invertebrate Taxonomists (SCAMIT, [www.scamit.org](http://www.scamit.org)) who work diligently to ensure  
474 consistency in nomenclature over time and who assisted in assembling these data sets prior to  
475 analysis to limit any taxonomic confounding.

476 Beyond any changes in regional pollution levels or shifts in taxonomic nomenclature, our  
477 analyses indicate that temperature and carbonate chemistry appear to have been important factors  
478 affecting the benthic composition over time. Temperature was an important predictor for taxa  
479 whose geographic distribution shifted, as well as those that increased in frequency of occurrence  
480 in the more modern samples. We would suggest that a combination of range expansions within  
481 the study area and from outside of the region to inside of it created the increase in overall species  
482 richness observed within the dataset. This would follow patterns observed in benthic  
483 communities along the Atlantic coast of US (Hale et al. 2017) and the Kattegat (Goransson  
484 2017), as well pelagic communities in the Atlantic and the Pacific (e.g., ter Hofstede et al. 2010,

485 McClatchie et al. 2016). Of the taxa whose range did shift, temperature was more frequently  
486 important for those moving northwards than southwards (predominantly crustaceans, as well as  
487 spionid and maldanid polychaetes), which follows with the encroachment of warmer waters into  
488 the northern parts of the Southern California Bight (e.g., Fumo et al. 2020) (Figure 2). The scale  
489 of range shifts for benthic species has been documented up to 70km per decade (Birchenough et  
490 al. 2015). These rates are in line with observations of the taxa in the present study, where more  
491 than two-thirds of the taxa that had northward or southward shifts in their distribution changed  
492 by 1-1.2 degrees latitude (approximately 111 – 133km) across the five decades of our study.

493 There is less evidence in the literature for geographic shifts in benthic infauna related to changes  
494 in carbonate chemistry, though the patterns from our study suggest that carbonate chemistry and  
495 dissolved oxygen can shape the distribution of benthic taxa as strongly as the more well  
496 documented changes related to water temperature noted above. Sato et al (2017) demonstrated  
497 habitat compression for motile benthic epifauna on the continental shelf of the Southern  
498 California Bight as a product of changes in carbonate chemistry and dissolved oxygen. The  
499 gradient in exposure to low pH waters across relatively short spatial scales in the region (Hauri et  
500 al. 2013, McLaughlin et al. 2018, Kessouri et al. 2022) may contribute to the range shifts we  
501 observed. Conversely, the influence of carbonate chemistry on changes in community  
502 composition related to local extirpations and appearances we observed in the macrobenthos has  
503 been predicted or observed in other systems (e.g., Kroeker et al. 2011, Busch et al. 2013,  
504 Nagelkerken & Connell 2022). A large number of the decrease taxa influenced by carbonate  
505 chemistry were bivalve and gastropod molluscs, which makes sense given their calcium  
506 carbonate shells (Green et al. 2009, Clements & Hunt 2017). In contrast, the increase taxa  
507 influenced by carbonate chemistry were mostly spionid or cirratulid polychaetes and crustaceans,

508 most of which were also deposit/interface feeders and could be filling the niche and physical  
509 space vacated by the disappearing bivalves and gastropods.

510 The specific impacts of the observed changes in macrobenthic community composition on the  
511 functioning of the soft-sediment continental shelf and slope ecosystem are hard to quantify  
512 within the scope of the present study. However, experiments exposing different types of infauna  
513 to altered temperature and pH conditions have demonstrated changes in behavior and allocation  
514 of energetic resources within individual fauna (Wood et al. 2008, Widdicombe et al. 2009,  
515 Christensen et al. 2017). When these climate change- and ocean acidification-driven impacts are  
516 extrapolated to the whole of the benthic ecosystem (e.g., Busch et al. 2013, Morley et al. 2022,  
517 Weinert et al. 2022), decreases in rates of secondary production, nutrient cycling, and carbon  
518 sequestration are predicted. Our data indicate that the macrobenthic community of the Southern  
519 California continental margin has changed compositionally and that total abundance has  
520 declined. When viewed through the lens of the aforementioned studies, the patterns we observed  
521 could be suggestive of the potential for a reduction in ecosystem functioning (e.g.,  
522 bioturbation/nutrient cycling and secondary productivity). However, the accompanying increase  
523 in taxonomic and feeding guild diversity suggests that there may have been some degree of  
524 community compensation against the influence of climate change and acidification (e.g.,  
525 Hendriks et al. 2010, Lavergne et al. 2010, Kroeker et al. 2011).

526 One of the more striking examples of community change we observed was the decline of shelled  
527 organisms over time in the northern portions of the region (see also Tomašových & Kidwell  
528 2017, Leonard-Pingel et al. 2019). The northern parts of our sample area have greater exposure  
529 to acidic waters (Hauri et al. 2013, McLaughlin et al. 2018) due to the oceanographic currents of  
530 the region (Harms & Winant 1998, Bray et al. 1999). However, the abundance patterns of shelled

531 fauna, especially in the 1970s and early 1980s, at the northern sample sites did not perfectly track  
532 the patterns in  $\Omega_{\text{arag}}$ . This is, in part, due to the unfortunate lack of consistent data at our water  
533 quality/chemistry sites during the early years of the benthic monitoring data record that prevents  
534 us from drawing quantitative conclusions. McClatchie et al. (2010) highlight the 1970s and  
535 1980s as a period of increased oxygen concentrations in the Southern California Bight as a  
536 whole, from which one could infer generally lower pCO<sub>2</sub> and less acidic conditions. The four  
537 data points prior to 1984 where we estimate low  $\Omega_{\text{arag}}$  values may represent local anomalies or  
538 influence of legacy organic matter pollution in the continental shelf of the region.

539 An additional level of complexity to consider when interpreting these patterns is that water  
540 column acidification most likely affects the veliger larvae and freshly settled infaunal molluscs  
541 more acutely than the adults buried deeper in the sediment (Green et al. 2009, Widdicombe et al.  
542 2009, Waldbusser et al. 2015). This differential impact would create a lag where the adults –  
543 which are detected in the benthic monitoring data – persist through acidified conditions but are  
544 not being replaced at the population level by juveniles and larvae – which are not detected in the  
545 benthic monitoring data. Illustrating this potential lag in population-level response, the  
546 abundance of a many of the molluscs classified as decreasers in our data set (e.g., *Acila*  
547 *castrensis*, *Chaetoderma* sp, *M. carlottensis*, *P. tenuisculpta*) were influenced in our random  
548 forest analyses by ocean acidification variables from one and three years prior to their collection  
549 rather than measurements from the same year of their collection (Supplement 3 – Table S7).

550 There has been increasing recognition among managers of the need to monitor and track both the  
551 exposure to and potential effects of climate change and acidification in coastal waters (Boehm et  
552 al 2015, Cross et al. 2019, Tilbrook et al 2019). The present study could represent the first steps  
553 towards developing specific benthic indicators of dissolved oxygen, temperature, or

554 acidification. We have identified a number of species that appear to be either sensitive to or  
555 indicative of exposure to these different water chemistry/quality stressors. These species could  
556 be used as the subjects of focused exposure or physiological studies to support the statistical  
557 relationships we observed in this retrospective study, much as Bednarsek et al. (2017) suggested  
558 pteropods as sentinels for midwater taxa. Alternatively, the patterns in abundance of the benthic  
559 species could be combined into multi-species metrics or used to create a stressor-specific  
560 assemblage models in a benthic index of acidification exposure. These types of experiments and  
561 assessment tools would combine nicely with existing chemical and biological monitoring efforts  
562 in the water column to create a holistic perspective on the exposure and effects of climate change  
563 on the coastal ocean.

564

## 565 5. ACKNOWLEDGEMENTS

566 The authors thank SCCWP's Commission Technical Advisory Group for reviewing the  
567 manuscript and providing feedback on its content and readability. They also thank Valerie  
568 Goodwin for assistance in data preparation. The authors acknowledge the efforts of the present  
569 and former staff working at the LACSD, CLAEMD, OCSan, and CSD ocean monitoring  
570 programs over the decades who have spent, and continue to spend, their careers collecting the  
571 samples and producing the benthic data used in this study. Finally, the authors thank Richard  
572 Feely and two anonymous reviewers who helped to improve the science and readability of the  
573 manuscript. This is PMEL contribution number 5467. This work was funded by SCCWRP  
574 internal funds, while the contributions of SRA were funded by PMEL.

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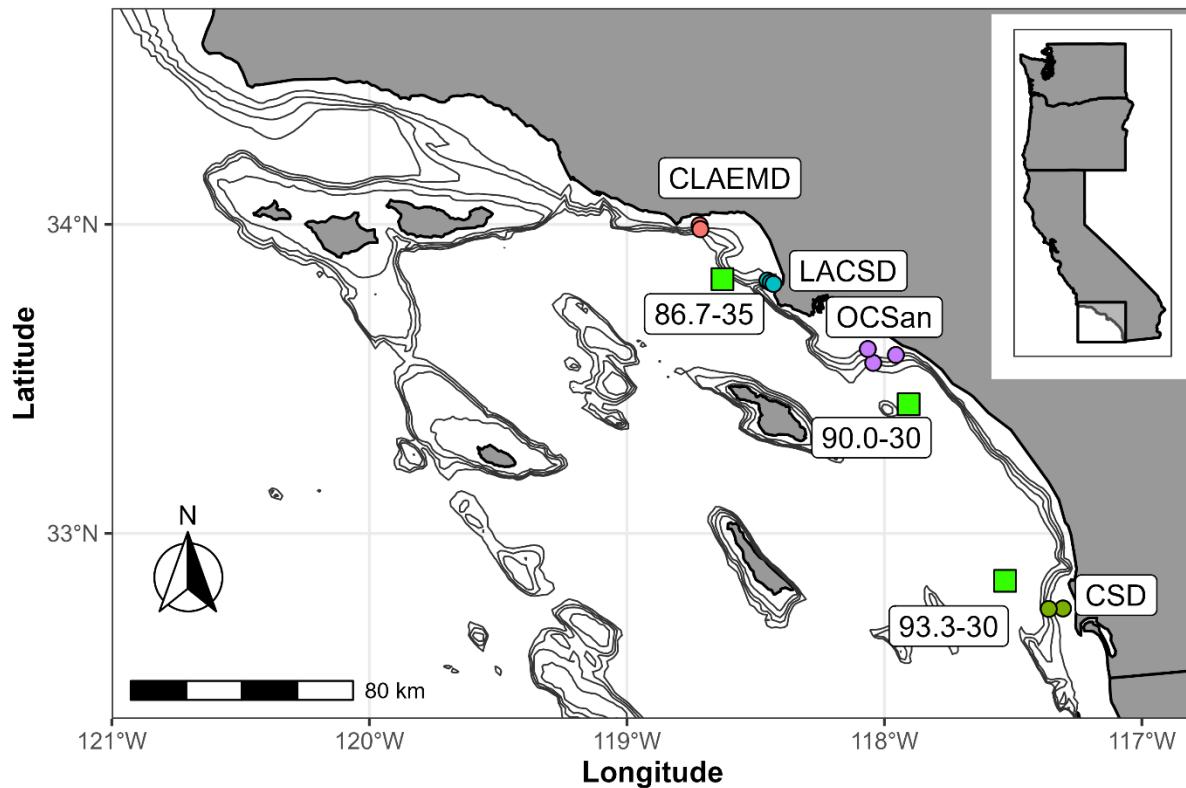
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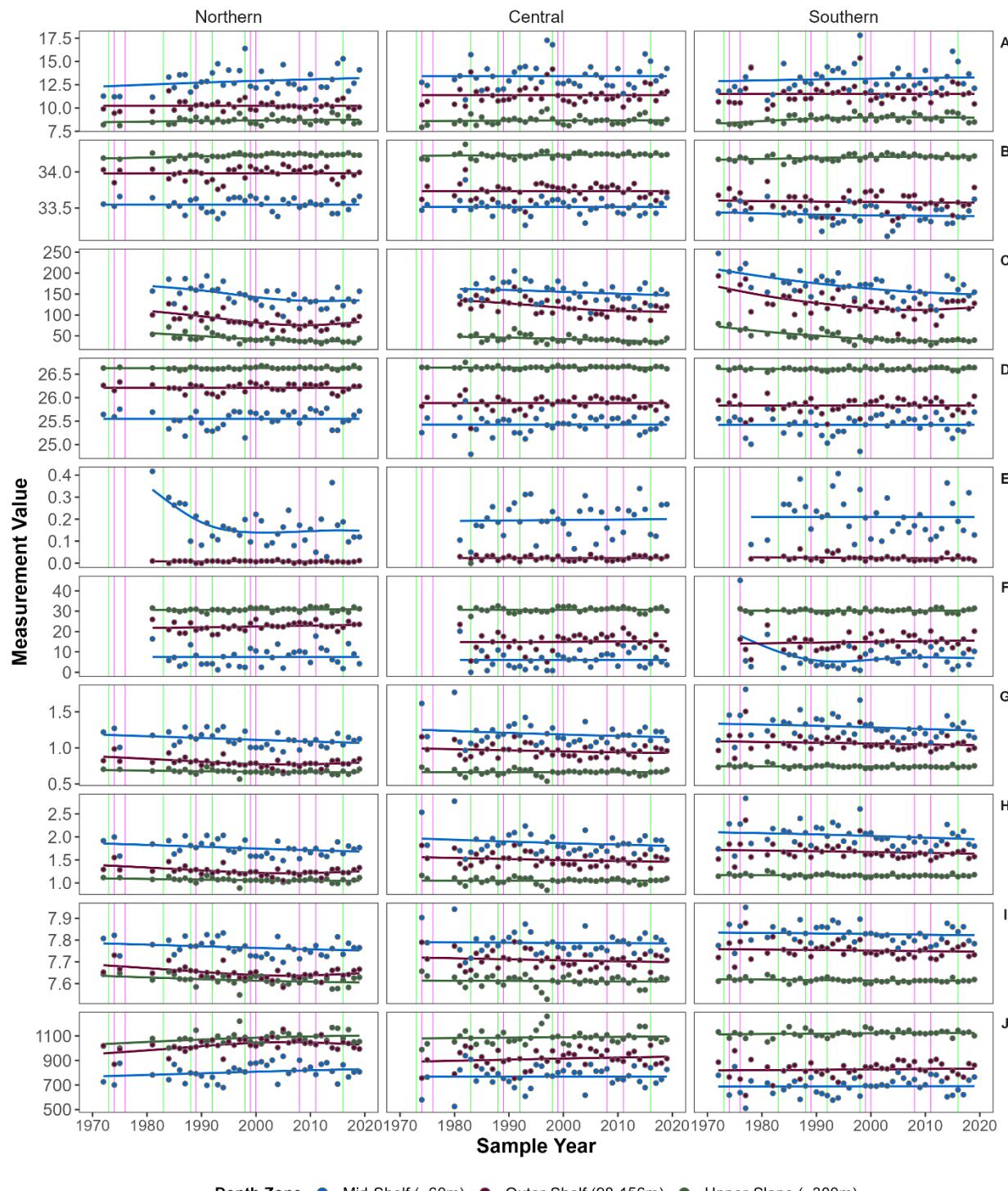
## 889 Figures



890

891 Figure 1 A map of the Southern California Bight illustrating the locations of the ten benthic  
892 sampling sites, color-coded by their sampling locations/data sources. The green squares represent  
893 the location of the CalCOFI water quality monitoring stations labelled with their transect  
894 number-station ID. The black lines represent the 100-m isobaths between 100 and 400 m deep.  
895 The inset shows the position of the Southern California Bight relative to the California and  
896 eastern Pacific coasts.

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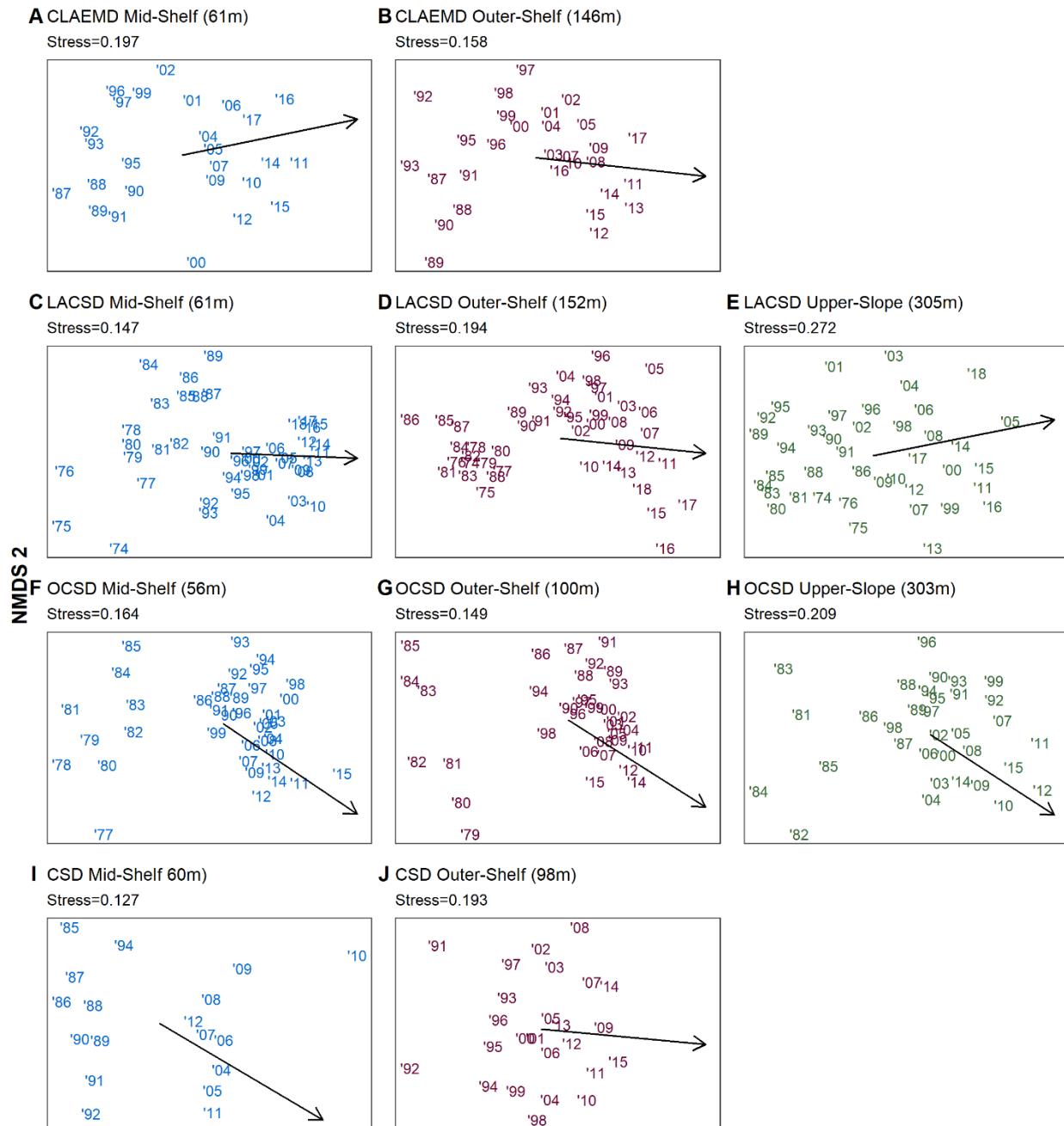


898

899 Figure 2 Water quality and water chemistry (summarized from quarterly bottle samples) at each  
 900 of the three CalCOFI water monitoring stations (see Figure 1) used in subsequent analyses of  
 901 benthic fauna. A trend line (general additive model for smoothing) for each set of data is

902 presented only to help the reader follow the patterns amongst the individual points. Bottles were  
903 selected within 20m of the depth of the associated benthic sample locations. A = maximum  
904 bottom water temperature (C), B= median salinity (PSU), C= minimum dissolved oxygen ( $\mu\text{mol}$   
905  $\text{kg}^{-1}$ ), D=median water density ( $\sigma_T$ ), E=median chlorophyll a ( $\mu\text{g L}^{-1}$ ), F= median nitrate ( $\mu\text{g L}^{-1}$ ),  
906 G=minimum estimated aragonite saturation State, H=minimum estimated calcite saturation State,  
907 I=minimum estimated pH, and J=maximum estimated  $\text{pCO}_2$ . The green vertical lines indicate  
908 years with strong to very strong El Niño events. The magenta lines indicate years with strong La  
909 Niña events. ENSO designations based upon NOAA Oceanic Niño Index  
910 (<https://psl.noaa.gov/data/correlation/oni.data>).

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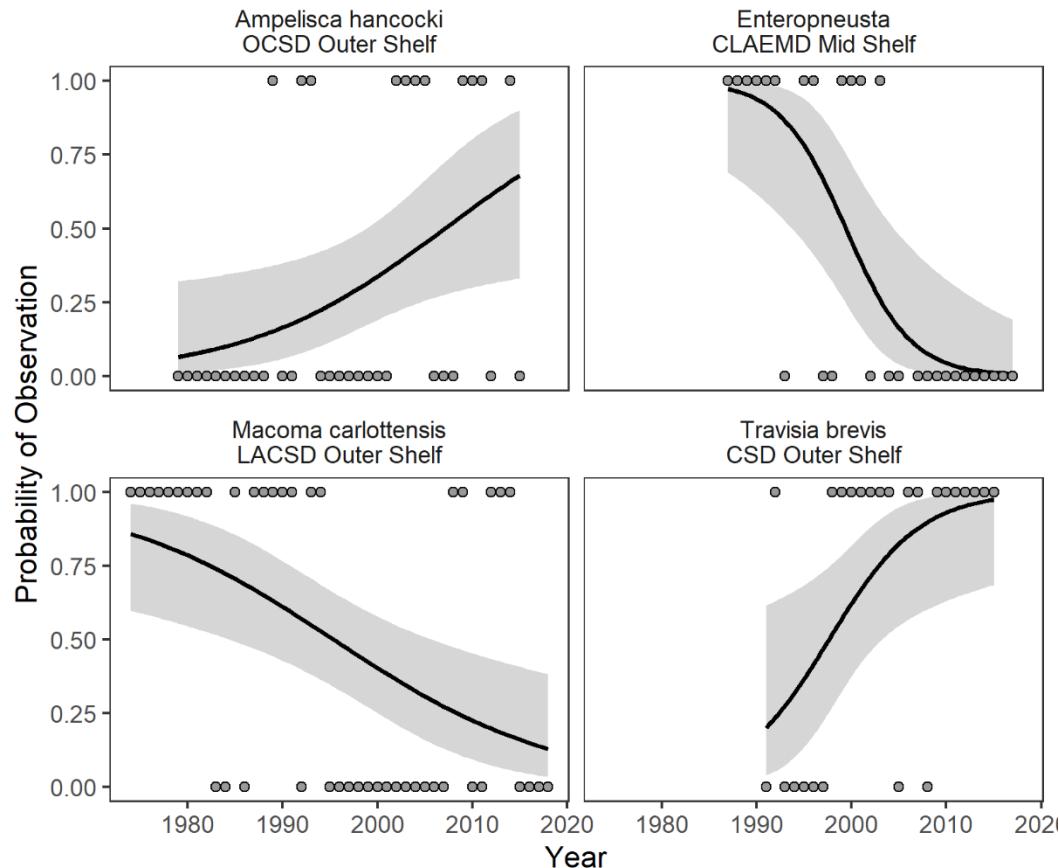


912

913 Figure 3 A series of 2-D nMDS ordination plots of Bray-Curtis dissimilarity values of benthic  
 914 infauna communities at each of the sampling sites from each year across the breadth of the data  
 915 set. The two-digit number represents the year of collection (i.e., 1998=98, 2001=01). The black

916 arrows indicate the trend of time across the different ordinations based upon multivariate  
917 correlations (Table 2). Ordinations based upon dissimilarities of presence-absence transformed  
918 community data with a minimum of 250 iterations.

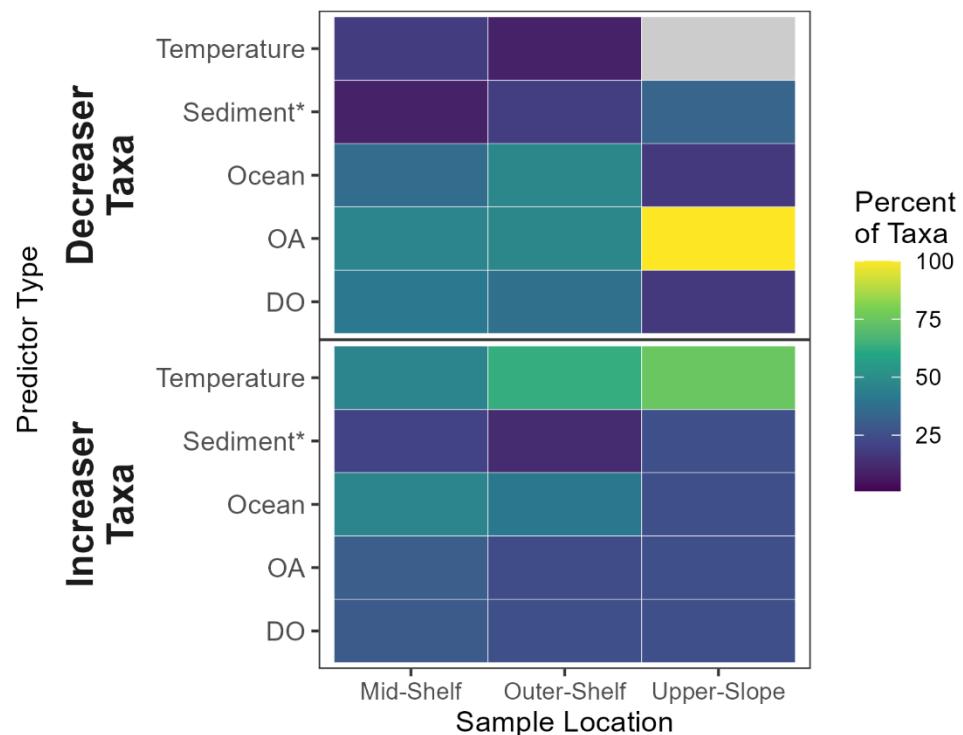
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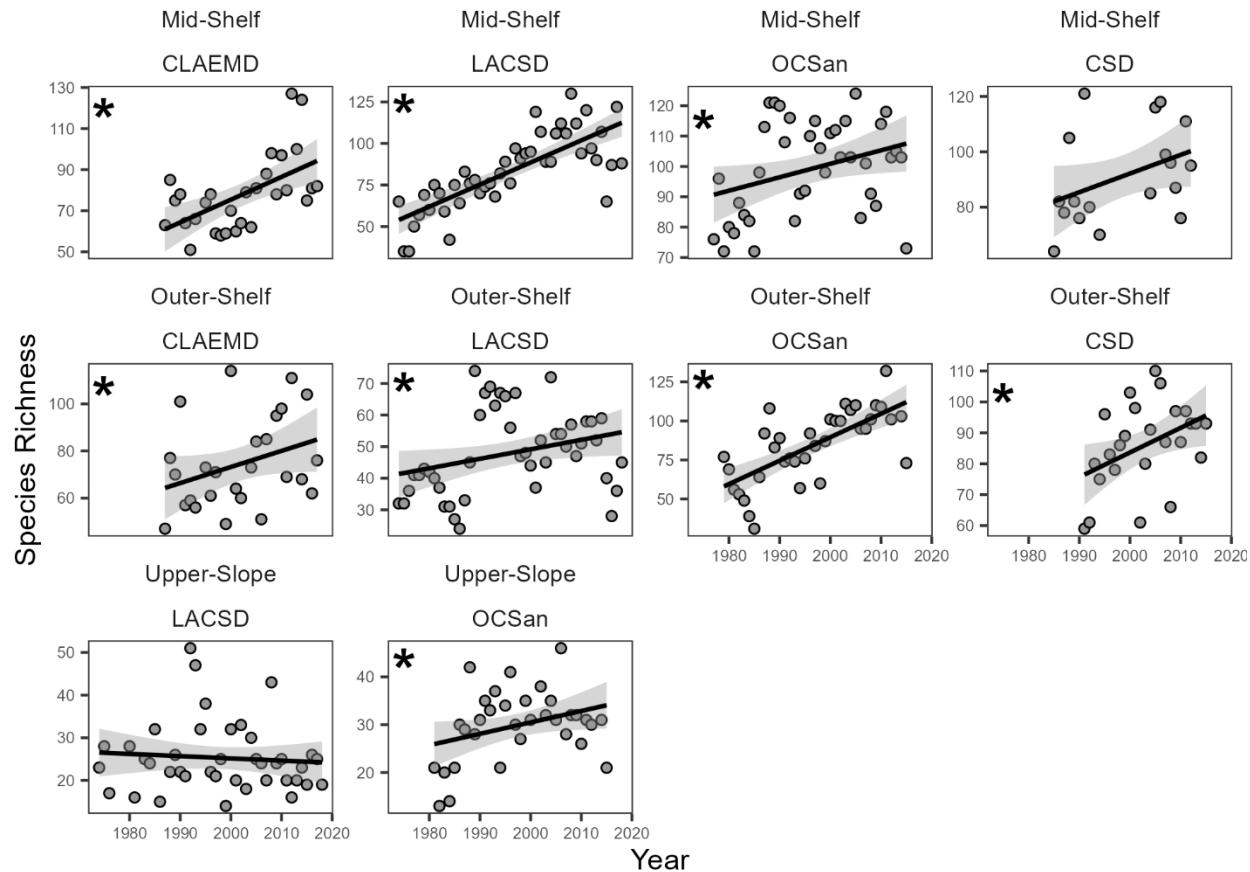
921 Figure 4 Logistic regression curves of four example taxa illustrating patterns of increasing  
 922 probability of observation through time (*Ampelisca hancocki* and *Travisia brevis*) and decreasing  
 923 probability of observation through time (*Enteropneusta* and *Macoma carlottensis*). Each of  
 924 these taxa, as well as all taxa identified in Supplement 3 – Table S5, had a regression with a beta  
 925 term significantly different than 0 ( $\alpha=0.05$ ). The grey ribbon represents the standard error of the  
 926 probability estimate. The grey dots in the rug represent the presence (1) or absence (0) of that  
 927 taxon in a given year at the site.

928



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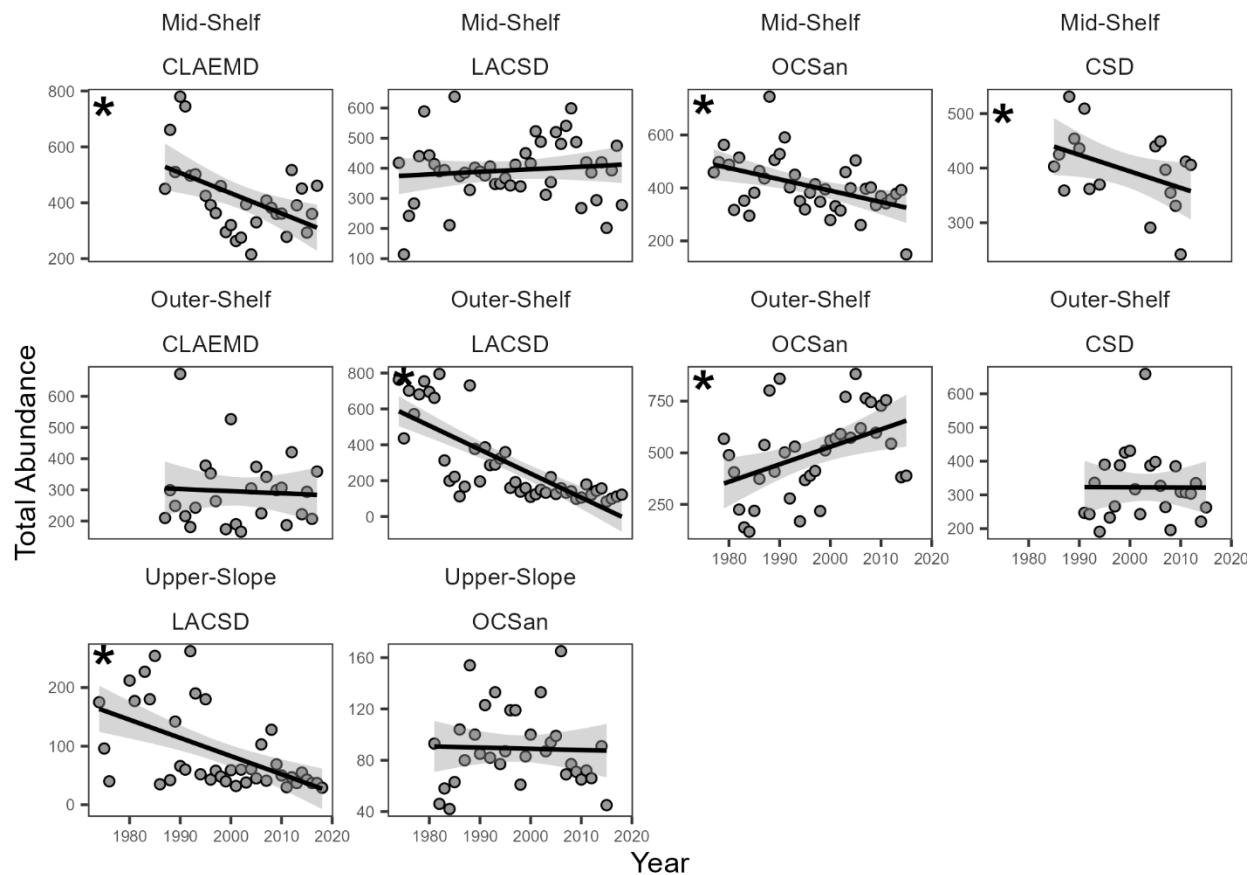
930 Figure 5 A heat map summarizing the most important variables in predicting the abundance  
 931 patterns of benthic infauna classified as decreasing or increasing in abundance through time (see  
 932 Supplement 3 – Tables S6 and S7) across all four sample locations and divided by depth zone.  
 933 Predictors included measures of temperature (surface water temperature, bottom water  
 934 temperature), oceanographic patterns (ENSO, PDO), bottom water ocean acidification (aragonite  
 935 saturation, calcite saturation, pH, pCO<sub>2</sub>), bottom water dissolved oxygen, or sediment grain size  
 936 (%Sand, %Silt, %Clay). Their predictive importance was derived from random forest regression  
 937 variable importance outputs. Warmer colors indicate a predictor that was important for more taxa  
 938 within a given depth zone. Cooler colors indicate a predictor that was important for fewer taxa.  
 939 No temperature variables were important predictors for decrease taxa from the upper-slope. See  
 940 Table 1 for a full list of all potential predictor variables. Note that sediment\* grain size data were  
 941 only available for the CLAEMD, LACSD, and CSD sample sites.



942

943 Figure 6 Least squares linear regression plots of taxa richness through time at each of the  
 944 sampling sites at the three depth zones. An asterisk indicates a slope significantly ( $\alpha = 0.1$ )  
 945 different than zero. The grey ribbon indicates the standard error of the predicted values.

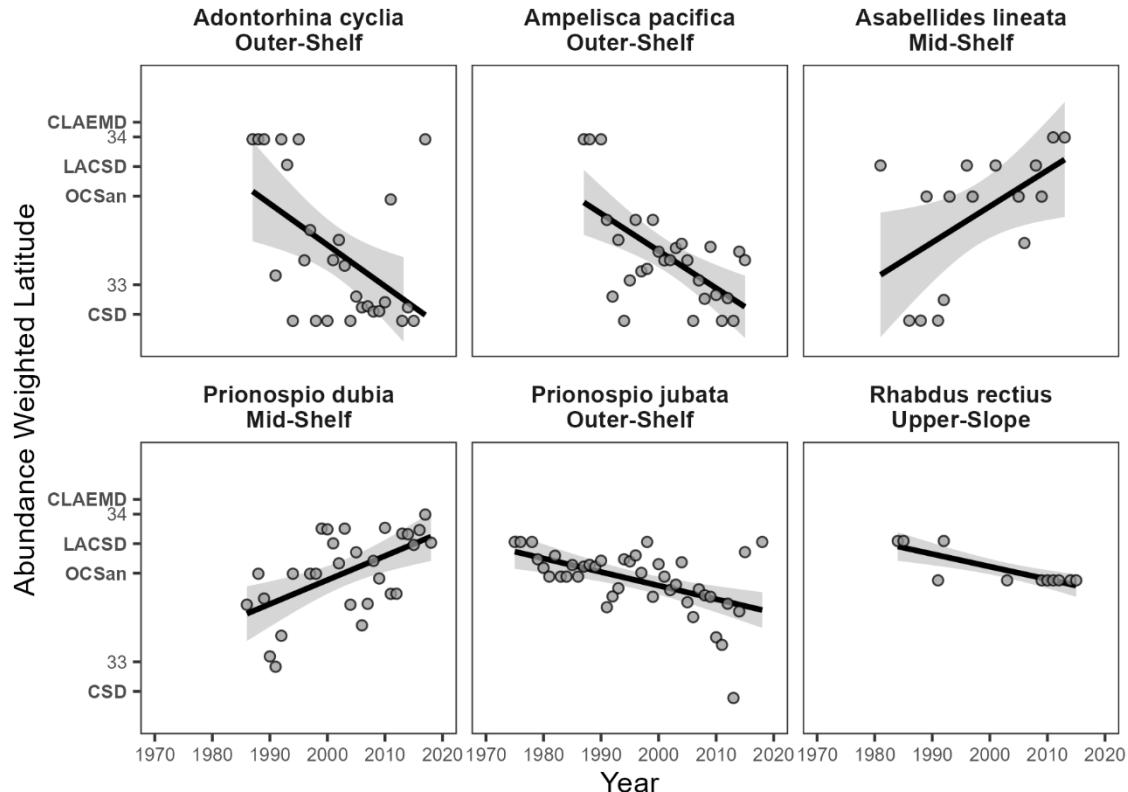
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948 Figure 7 Least squares linear regression plots of sample abundance through time at each of the  
 949 sampling sites at the three depth zones. An asterisk indicates a slope significantly ( $\alpha = 0.1$ )  
 950 different than zero. The grey ribbon indicates the standard error of the predicted values.

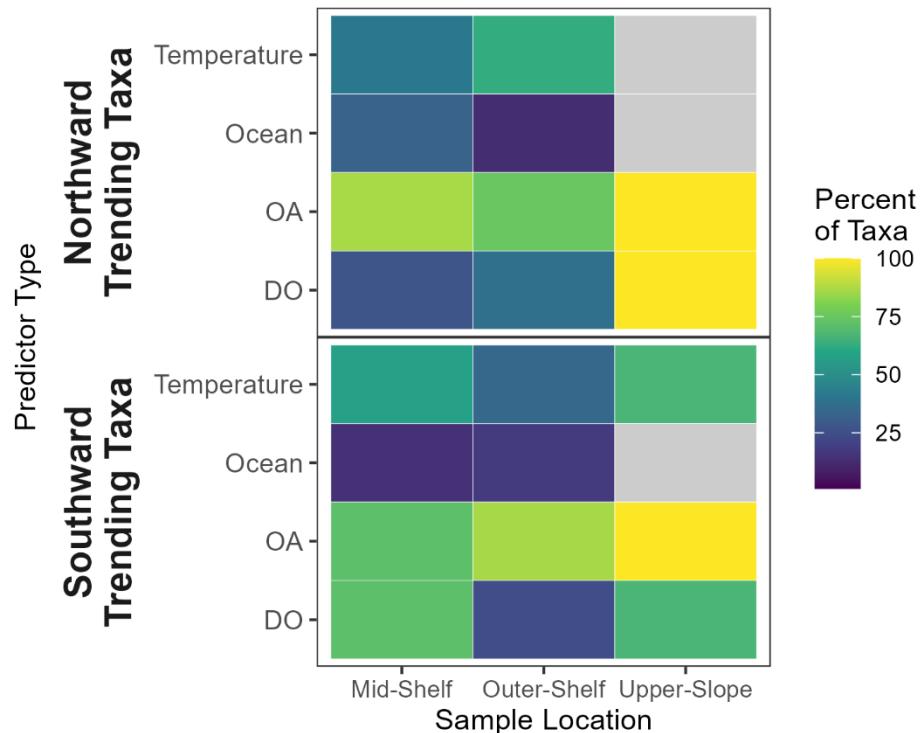
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953 Figure 8 Least squares linear regression of abundance weighted latitude through time for six taxa  
 954 illustrating a trend of a taxon whose distribution has shifted northwards (*Asabellides lineata* and  
 955 *Prionospio dubia*) or a taxon whose distribution has shifted southwards through time (*Prionospio*  
 956 *jubata*, *Adontorhina cyclia*, *Ampelisca pacifica*, and *Rhabdus rectius*). Each of these taxa, as well  
 957 as all taxa identified in Supplement 3 – Table S8, had a regression with a slope significantly  
 958 different than 0 ( $\alpha=0.05$ ). The grey ribbon indicates the standard error of the predicted values.  
 959 The average latitudinal position of the four sampling sites is indicated on the y-axis for reference.

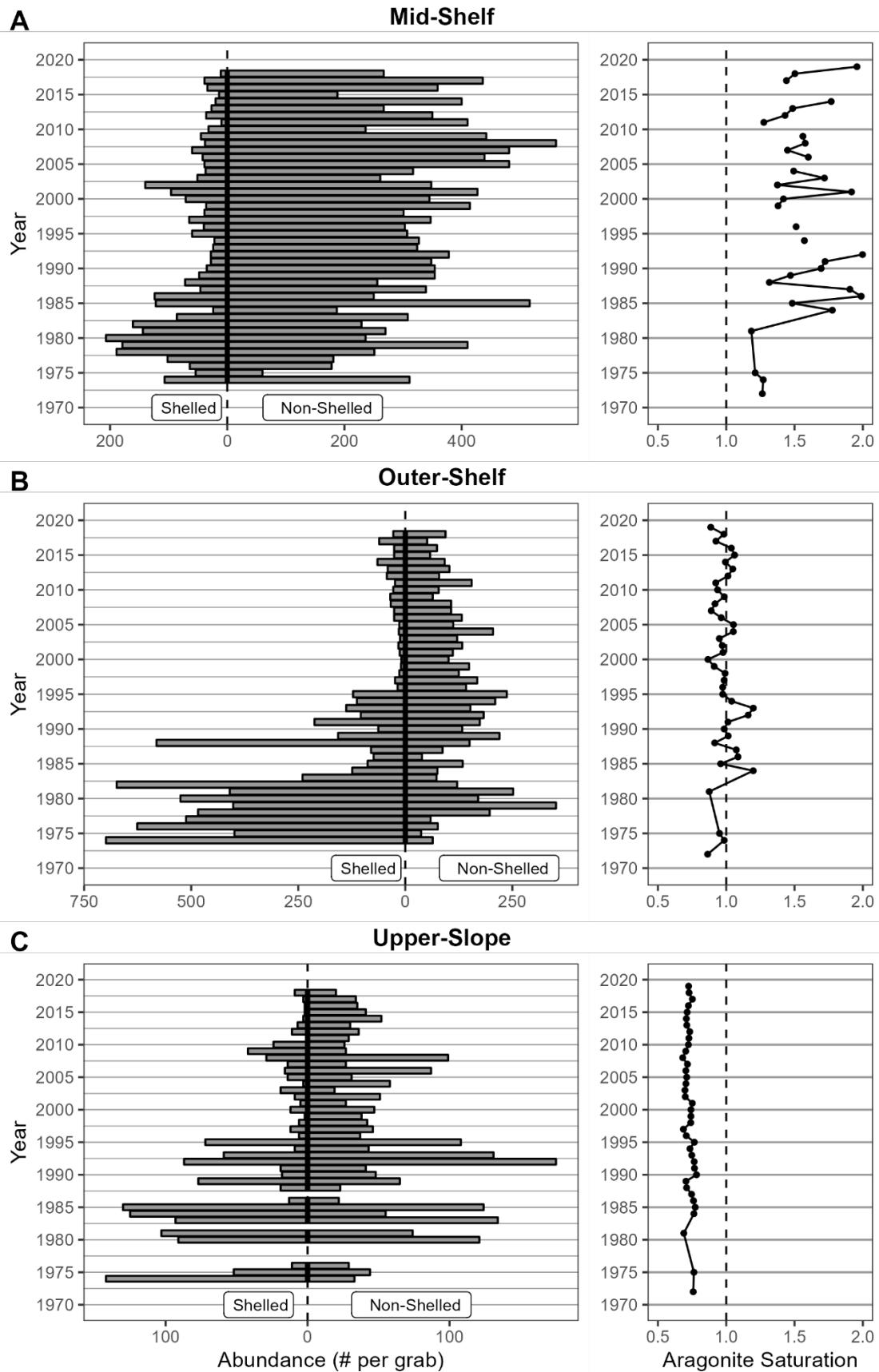
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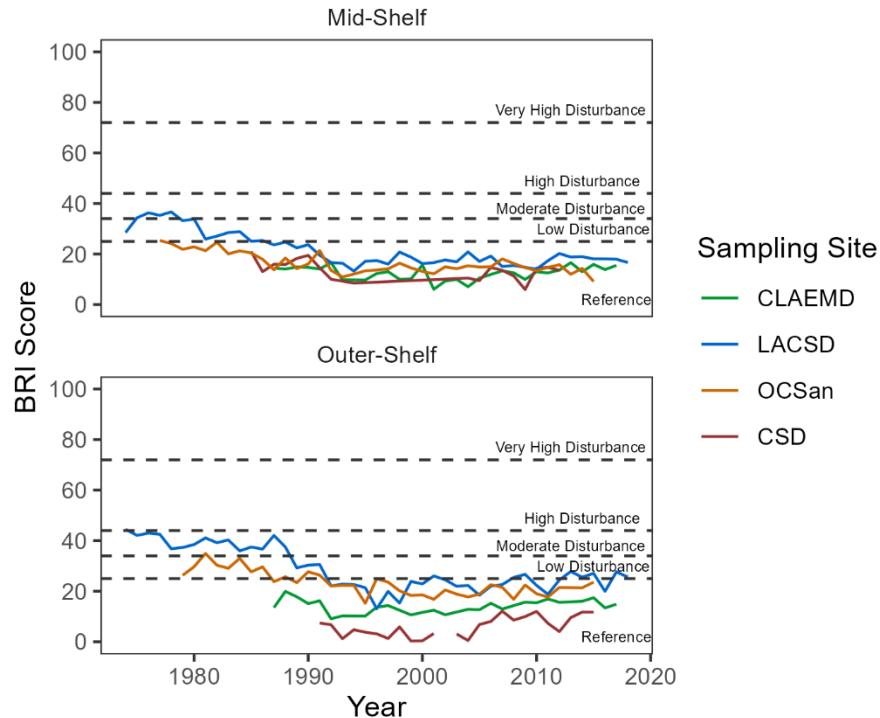
962 Figure 9 A heat map summarizing the most important variables in predicting the abundance  
 963 patterns of benthic infauna classified as northward or southward shifting taxa through time (see  
 964 Supplement 3 – Tables S9 and S10) within each of the three depth zones. Predictors included  
 965 measures of temperature (surface water temperature, bottom water temperature), oceanographic  
 966 patterns (ENSO, PDO), bottom water ocean acidification (aragonite saturation, calcite saturation,  
 967 pH,  $\text{pCO}_2$ ), or bottom water dissolved oxygen. Their predictive importance was derived from  
 968 random forest regression variable importance outputs. Warmer colors indicate a predictor that  
 969 was important for more taxa within a given depth zone. Cooler colors indicate a predictor that  
 970 was important for fewer taxa. Only one northward trending taxon was observed from the upper-  
 971 slope and oceanographic variables were not important predictors for southward shifting taxa  
 972 from the upper-slope. See Table 1 for a full list of all potential predictor variables.

973



975 Figure 10 An illustration of the change in abundance of shelled and non-shelled organisms at  
976 Mid-Shelf (A), Outer-Shelf (B), and Upper-Slope (C) sample sites relative to modeled aragonite  
977 saturation values at the LACSD location through time. The dashed vertical line on the aragonite  
978 saturation plots highlight a value of 1.0, below which aragonite will dissolve into seawater from  
979 a solid form.

980



981

982 Figure 11 Benthic condition scores through time at each of the sampling sites in the mid-shelf  
 983 and outer-shelf depth zones using the Benthic Response Index (BRI) of Smith et al. (2000). A  
 984 higher score indicates a more disturbed sample. The dashed lines indicate the different thresholds  
 985 of community impact that were linked to compositional changes in addition to loss of taxonomic  
 986 and functional diversity with high levels of disturbance. The BRI is not calibrated to Upper-  
 987 Slope habitats.

988 [Tables](#)

989 **Table 1.** Environmental variables used as potential predictors of change in macrobenthic community composition through time, the  
 990 sources of those data, and the coverage of those variables at each of the four sampling locations. PDO - Pacific Decadal Oscillation  
 991 index value. SOI - Southern Oscillation Index value. \*CSD sediment data were reported in % sand and % mud, where % mud = % silt  
 992 +% clay. 1- and 3-year time lags were selected for consideration, as most of the benthic infauna are thought to live from 1 to 3 years at  
 993 a maximum, though some of larger bivalves may persist longer.

<b>Data Type</b>	<b>Potential Predictor Variables</b>	<b>Data Source</b>	<b>CLAEMD</b>	<b>LACSD</b>	<b>OCSan</b>	<b>CSD</b>
Local Sediment	% Sand	Discharger	X	X		X
	% Silt	Discharger	X	X		X*
	% Clay	Discharger	X	X		X*
Regional Oceanography	Mean Springtime PDO	NOAA	X	X	X	X
	Mean Spring PDO One Year Prior	NOAA	X	X	X	X
	Mean Spring PDO Three Years Prior	NOAA	X	X	X	X
	Mean Springtime SOI	NOAA	X	X	X	X
	Mean Springtime SOI One Year Prior	NOAA	X	X	X	X
	Mean Springtime SOI Three Years Prior	NOAA	X	X	X	X
Local Temperature	Mean Summer Surface Water Temperature	CalCOFI	X	X	X	X
	Mean Summer Surface Water Temperature One Year Prior	CalCOFI	X	X	X	X
	Mean Summer Surface Water Temperature Three Years Prior	CalCOFI	X	X	X	X
	Mean Summer Bottom Water Temperature	CalCOFI	X	X	X	X
	Mean Summer Bottom Water Temperature One Year Prior	CalCOFI	X	X	X	X
	Mean Summer Bottom Water Temperature Three Years Prior	CalCOFI	X	X	X	X

Local Dissolved Oxygen	Mean Summer Bottom Water Dissolved Oxygen	CalCOFI	X	X	X	X
	Mean Summer Bottom Water Dissolved Oxygen One Year Prior	CalCOFI	X	X	X	X
	Mean Summer Bottom Water Dissolved Oxygen Three Years Prior	CalCOFI	X	X	X	X
	pCO2	Modelled	X	X	X	X
Local Acidification	pCO2 One Year Prior	Modelled	X	X	X	X
	pCO2 Three Years Prior	Modelled	X	X	X	X
	pH	Modelled	X	X	X	X
	pH One Year Prior	Modelled	X	X	X	X
	pH Three Years Prior	Modelled	X	X	X	X
	Aragonite Saturation	Modelled	X	X	X	X
	Aragonite Saturation One Year Prior	Modelled	X	X	X	X
	Aragonite Saturation Three Years Prior	Modelled	X	X	X	X
	Calcite Saturation	Modelled	X	X	X	X
	Calcite Saturation One Year Prior	Modelled	X	X	X	X
	Calcite Saturation Three Years Prior	Modelled	X	X	X	X

995

996 **Table 2** Multivariate correlation of year of sampling with the nMDS ordinations presented in  
 997 Figure 3. Correlations were calculated across 1,000 permutations.

Depth Zone	Sample Location	r	p-value
Mid-Shelf	CLAEMD	0.93	0.002
	LACSD	0.96	0.002
	OCSan	0.92	0.002
	CSD	0.91	0.002
Outer-Shelf	CLAEMD	0.95	0.002
	LACSD	0.91	0.002
	OCSan	0.84	0.002
	CSD	0.88	0.002
Upper-Slope	LACSD	0.80	0.002
	OCSan	0.86	0.002

998

999

1000      **Table 3.** Outputs of 1-way PermANOVAs testing the differences in macrobenthic community  
 1001      structure through time at the different depth zones of each of the four sampling locations based  
 1002      upon Bray-Curtis dissimilarities of taxon presence-absence over 10,000 permutations.

Depth Zone	Sample Location	Term	Df	SS	R <sup>2</sup>	pseudo-F	p-value
Mid-Shelf	CLAEMD	Year	1	0.92	0.20	6.79	<0.0001
		Residual	27	3.64	0.80		
	LACSD	Year	1	1.99	0.23	12.87	<0.0001
		Residual	43	6.65	0.77		
	OCSan	Year	1	1.12	0.16	6.86	<0.0001
		Residual	37	6.04	0.84		
	CSD	Year	1	0.76	0.26	5.56	<0.0001
		Residual	16	2.20	0.74		
Outer-Shelf	CLAEMD	Year	1	0.84	0.17	4.97	<0.0001
		Residual	24	4.05	0.83		
	LACSD	Year	1	1.94	0.19	10.31	<0.0001
		Residual	43	8.09	0.81		
	OCSan	Year	1	0.98	0.17	6.77	<0.0001
		Residual	34	4.91	0.83		
	CSD	Year	1	0.50	0.15	3.92	<0.0001
		Residual	23	2.92	0.85		
Upper-Slope	LACSD	Year	1	0.94	0.09	3.79	<0.0001
		Residual	38	9.41	0.91		
	OCSan	Year	1	0.96	0.13	4.44	<0.0001
		Residual	31	6.72	0.87		

1003

1004

1005 **Table 4.** Top-10 most abundant taxa in descending rank order (with ties) for each decade 1970s  
1006 – 2010s at Mid-Shelf, Outer-Shelf, and Upper-Slope depth zones from the LACSD and OCSan  
1007 sampling sites. Taxa are color-coded based upon the decade during which they first appeared in  
1008 the top-10 of the sample site. ^ denotes taxa indicative of disturbed conditions based upon Smith  
1009 et al. (2001). # denotes taxa indicative of non-disturbed conditions based upon Smith et al.  
1010 (2001). (c) = crustacean, (e) = echinoderm, (h) = hemichordate, (l) = phoronid, (m) = mollusc,  
1011 and (p) = polychaete. Note that that no samples were collected from the OCSan Upper-Slope site  
1012 in the 1970s.

District	Depth Zone	1970s	1980s	1990s	2000s	2010s
LACSD	Mid-Shelf	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Spiophanes duplex</i> (p)	<i>Amphiodia</i> sp (e)	<i>Amphiodia</i> sp (e)	<i>Amphiodia</i> sp (e)
		<i>Aphelochaeta-Kirkegaardia</i> Cmplx <sup>^</sup> (p)	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Spiophanes duplex</i> (p)	<i>Spiophanes duplex</i> (p)	<i>Lumbrineris cruzensis</i> (p)
		<i>Axinopsida serricata</i> (m)	<i>Axinopsida serricata</i> (m)	<i>Amage scutata</i> (p)	<i>Axinopsida serricata</i> (m)	<i>Spiophanes duplex</i> (p)
		<i>Prionospio jubata</i> (p)	<i>Amphiodia</i> sp (e)	<i>Axinopsida serricata</i> (m)	<i>Mediomastus</i> sp (p)	<i>Aphelochaeta-Kirkegaardia</i> Cmplx <sup>^</sup> (p)
		<i>Mediomastus</i> sp (p)	<i>Myriochele striolata</i> (p)	<i>Phoronis</i> sp	<i>Thysira flexuosa</i> (m)	<i>Mediomastus</i> sp (p)
		<i>Euphilomedes producta</i> (c)	<i>Lumbrineris</i> sp (p)	<i>Pectinaria californiensis</i> (p)	<i>Paraprionospio alata</i> (p)	<i>Cossura candida</i> (p)
		<i>Lumbrineris</i> sp (p)	<i>Marphysa disjuncta</i> (p)	<i>Marphysa disjuncta</i> (p)	<i>Prionospio jubata</i> (p)	<i>Prionospio dubia</i> (p)
		<i>Hamatocalpium californicum</i> (c)	<i>Prionospio jubata</i> (p)	<i>Prionospio jubata</i> (p)	<i>Phoronis</i> sp (l)	<i>Prionospio jubata</i> (p)
		<i>Euphilomedes carcharodonta</i> (c)	<i>Heterophoxus</i> sp (c)	<i>Paraprionospio alata</i> (p)	<i>Lumbrineris cruzensis</i> (p)	<i>Gymnonereis crosslandi</i> (p)
		<i>Pectinaria californiensis</i> (p)	<i>Cossura candida</i> (p)	<i>Glycera nana</i> (p)	<i>Maldane sarsi</i> (p)	<i>Eclyssipe trilobata</i> <sup>#</sup> (p)
						<i>Lumbrineris</i> sp (p)
						<i>Phoronis</i> sp (l)
OCSan	Mid-Shelf	<i>Amphiodia</i> sp (e)				
		<i>Euphilomedes carcharodonta</i> (c)	<i>Pectinaria californiensis</i> (p)	<i>Prionospio jubata</i> (p)	<i>Prionospio jubata</i> (p)	<i>Prionospio jubata</i> (p)
		<i>Pectinaria californiensis</i> (p)	<i>Prionospio jubata</i> (p)	<i>Chloea pinnata</i> (p)	<i>Mediomastus</i> sp (p)	<i>Euphilomedes carcharodonta</i> (c)
		<i>Kirkegaardia serratiseta</i> <sup>^</sup> (p)	<i>Spiophanes duplex</i> (p)	<i>Pectinaria californiensis</i> (p)	<i>Aricidea (Acmira) catherinae</i> (p)	<i>Lumbrineris cruzensis</i> (p)
		<i>Heterophoxus oculatus</i> (c)	<i>Euphilomedes carcharodonta</i> (c)	<i>Rhepoxynius bicuspidatus</i> (c)	<i>Euphilomedes carcharodonta</i> (c)	<i>Chloea pinnata</i> (p)
		<i>Chloea pinnata</i> (p)	<i>Heterophoxus oculatus</i> (c)	<i>Amphideutopus oculatus</i> (c)	<i>Leptochelia dubia</i> Cmplx(c)	<i>Scoloplos armiger</i> Cmplx(p)
		<i>Lumbrineris</i> sp (p)	<i>Ampelisca brevisimulata</i> (c)	<i>Phoronida</i> (l)	<i>Rhepoxynius bicuspidatus</i> (c)	<i>Aricidea (Acmira) catherinae</i> (p)
		<i>Prionospio jubata</i> (p)	<i>Tellina modesta</i> (m)	<i>Ampelisca brevisimulata</i> (c)	<i>Chloea pinnata</i> (p)	<i>Sthenelanella uniformis</i> (p)
		<i>Sabellidae</i> (p)	<i>Goniada maculata</i> (p)	<i>Aricidea (Acmira) catherinae</i> (p)	<i>Lumbrineris ligulata</i> (p)	<i>Stereobalanus</i> sp (h)
		<i>Axinopsida serricata</i> (m)	<i>Rhepoxynius bicuspidatus</i> (c)	<i>Euclymeninae</i> sp A (p)	<i>Paraprionospio alata</i> (p)	<i>Photis californica</i> (c)
LACSD	Outer-Shelf	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Axinopsida serricata</i> (m)	<i>Paradiopatra parva</i> (p)	<i>Aphelochaeta-Kirkegaardia</i> Cmplx <sup>^</sup> (p)
		<i>Axinopsida serricata</i> (m)	<i>Axinopsida serricata</i> (m)	<i>Spiophanes fimbriata</i> (p)	<i>Paraprionospio alata</i> (p)	<i>Paraprionospio alata</i> (p)
		<i>Acila castrensis</i> (m)	<i>Acila castrensis</i> (m)	<i>Spiophanes kimbballi</i> (p)	<i>Spiophanes berkeleyorum</i> (p)	<i>Spiophanes kimbballi</i> (p)
		<i>Macoma carlottensis</i> <sup>^</sup> (m)	<i>Tellina</i> sp B (m)	<i>Paraprionospio alata</i> (p)	<i>Mediomastus</i> sp (p)	<i>Axinopsida serricata</i> (m)
		<i>Pectinaria californiensis</i> (p)	<i>Spiophanes berkeleyorum</i> (p)	<i>Aphelochaeta-Kirkegaardia</i> Cmplx <sup>^</sup> (p)	<i>Aphelochaeta-Kirkegaardia</i> Cmplx <sup>^</sup> (p)	<i>Nuculana</i> sp A (m)
		<i>Aphelochaeta-Kirkegaardia</i> Cmplx <sup>^</sup> (p)	<i>Pectinaria californiensis</i> (p)	<i>Pectinaria californiensis</i> (p)	<i>Tellina</i> sp B (m)	<i>Tellina</i> sp B (m)
		<i>Spiophanes berkeleyorum</i> (p)	<i>Aphelochaeta-Kirkegaardia</i> Cmplx <sup>^</sup> (p)	<i>Lumbrineridae</i> (p)	<i>Spiophanes fimbriata</i> (p)	<i>Chloea pinnata</i> (p)
		<i>Spiophanes duplex</i> (p)	<i>Lumbrineris</i> sp (p)	<i>Tellina</i> sp B (m)	<i>Axinopsida serricata</i> (m)	<i>Brisaster</i> sp (e)
		<i>Onuphis</i> sp (p)	<i>Onuphis</i> sp (p)	<i>Heterophoxus affinis</i> (c)	<i>Spiophanes duplex</i> (p)	<i>Mediomastus</i> sp (p)
		<i>Tellina</i> sp B (m)	<i>Macoma carlottensis</i> <sup>^</sup> (m)	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Marphysa disjuncta</i> (p)	<i>Polyschides quadrifissatus</i> (m)
1013			<i>Maldane sarsi</i> (p)		<i>Spiophanes kimbballi</i> (p)	
1014						

OCSan	Outer-Shelf	<i>Pectinaria californiensis</i> (p)	<i>Amphiodia</i> sp (e)	<i>Pectinaria californiensis</i> (p)	<i>Aphelochaeta glandaria</i> Cmplx <sup>^</sup> (p)
		<i>Amphiodia</i> sp (e)	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Spiophanes berkeleyorum</i> (p)	<i>Axinopsida serricata</i> (m)
		<i>Chloea pinnata</i> (p)	<i>Aphelochaeta glandaria</i> Cmplx <sup>^</sup> (p)	<i>Euryalida</i> (e)	<i>Spiophanes berkeleyorum</i> (p)
		<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Kirkegaardia serratiseta</i> <sup>^</sup> (p)	<i>Aphelochaeta glandaria</i> Cmplx <sup>^</sup> (p)	
		<i>Axinopsida serricata</i> (m)	<i>Spiophanes duplex</i> (p)	<i>Spiophanes duplex</i> (p)	<i>Petaloclymene pacifica</i> (p)
		<i>Lumbrineris</i> sp (p)	<i>Spiophanes berkeleyorum</i> (p)	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Prionospio jubata</i> (p)
		<i>Maldanidae</i> (p)	<i>Lumbrineris</i> sp (p)	<i>Prionospio jubata</i> (p)	<i>Pectinaria californiensis</i> (p)
		<i>Glycera nana</i> (p)	<i>Prionospio jubata</i> (p)	<i>Scoletoama tetraura</i> Cmplx (p)	<i>Lumbrineris cruzensis</i> (p)
		<i>Mediomastus</i> sp (p)	<i>Prionospio lighti</i> (p)	<i>Glycera nana</i> (p)	<i>Rhepoxynius bicuspidatus</i> (c)
		<i>Praxillella pacifica</i> (p)			
LACSD	Upper-Slope	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Lirobittium rugatum</i> (m)	<i>Lirobittium rugatum</i> (m)	<i>Cyclopecten catalinensis</i> (m)
		<i>Macoma carlottensis</i> <sup>^</sup> (m)	<i>Pectinaria californiensis</i> (p)	<i>Parapriionospio alata</i> (p)	
		<i>Onuphis</i> sp (p)	<i>Parvilucina tenuisculpta</i> <sup>^</sup> (m)	<i>Ampelisca unsocalae</i> (c)	<i>Yoldia seminuda</i> (m)
		<i>Bipalponephrys cornuta</i> (p)	<i>Axinopsida serricata</i> (m)	<i>Onuphis</i> sp (p)	<i>Onuphis</i> sp (p)
		<i>Parapriionospio alata</i> (p)	<i>Parapriionospio alata</i> (p)	<i>Delectopecten vancouverensis</i> (m)	
		<i>Harpiniopsis galera</i> (c)	<i>Bipalponephrys cornuta</i> (p)	<i>Axinopsida serricata</i> (m)	<i>Mediomastus</i> sp (p)
		<i>Amphipoda</i> (c)	<i>Macoma carlottensis</i> <sup>^</sup> (m)	<i>Bipalponephrys cornuta</i> (p)	
		<i>Malmgreniella</i> sp (p)	<i>Aoroides columbiae</i> <sup>#</sup> (c)	<i>Aoroides columbiae</i> <sup>#</sup> (c)	<i>Bipalponephrys cornuta</i> (p)
		<i>Goniada brunnea</i> (p)	<i>Aphelochaeta-Kirkegaardia</i> Cmplx (p)	<i>Myriochele gracilis</i> (p)	<i>Delectopecten vancouverensis</i> (m)
		<i>Pectinaria californiensis</i> (p)			
OCSan	Upper-Slope		<i>Pectinaria californiensis</i> (p)	<i>Aphelochaeta monilaris</i> (p)	<i>Aphelochaeta monilaris</i> (p)
			<i>Parapriionospio alata</i> (p)	<i>Ampelisca unsocalae</i> (c)	<i>Ampelisca unsocalae</i> (c)
			<i>Ampelisca unsocalae</i> (c)	<i>Parapriionospio alata</i> (p)	<i>Myriochele gracilis</i> (p)
			<i>Aphelochaeta glandaria</i> Cmplx (p)	<i>Melinna heterodonta</i> (m)	<i>Parapriionospio alata</i> (p)
			<i>Melinna heterodonta</i> (m)	<i>Spiophanes kimballi</i> (p)	<i>Spiophanes berkeleyorum</i> (p)
			<i>Kirkegaardia serratiseta</i> (p)	<i>Limifosser fratula</i> (m)	<i>Spiophanes kimballi</i> (p)
			<i>Prionospio ehlersi</i> (p)	<i>Eudorella pacifica</i> (c)	<i>Melinna heterodonta</i> (m)
			<i>Aphelochaeta monilaris</i> (p)	<i>Caecognathia crenulatifrons</i> (c)	<i>Eudorella pacifica</i> (c)
			<i>Lumbrineris</i> sp (p)	<i>Onuphis</i> sp A (p)	<i>Caecognathia crenulatifrons</i> (c)

1016 **Table 5** Inventories of the number of taxa from the data set whose likelihood of occurrence  
 1017 increased or decreased through time (A) or whose geographic distribution shifted northward or  
 1018 southward through time (B). Total Richness is the count of all distinct taxa observed at a site or  
 1019 depth zone. Frequent Taxa are those observed ten or more times at a site or depth zone and were  
 1020 considered for classification as increaser/decreaser or northward/southward.

A	Depth Zone	Sample Site	Total Richness	Frequent Taxa	Increaser	Decreaser	Other
Mid-Shelf	LACSD		547	129	51	9	69
	CLAEMD		415	73	18	4	51
	OCSan		602	135	33	13	89
	CSD		496	90	7	1	82
Outer-Shelf	LACSD		389	66	14	16	36
	CLAEMD		441	66	10	7	49
	OCSan		468	107	37	3	67
	CSD		460	115	15	9	91
Upper-Slope	LACSD		259	29	1	6	22
	OCSan		266	22	3	2	17

B	Depth Zone	Total Richness	Frequent Taxa	Northward	Southward	Neither
Mid-Shelf		861	246	37	7	202
Outer-Shelf		773	207	15	29	163
Upper-Slope		365	50	1	3	46

1021