

Nitrification in the water column of Lake Erie: seasonal patterns, community dynamics, and competition with cyanoHABs

Authors: Hoffman, D.K., McCarthy, M.J., Zastepa, A., Boedecker, A.R., Myers, J.A., Newell, S.E.

ABSTRACT

This study reports directly measured nitrification rates in the water column of western Lake Erie, which is affected by annual cyanobacterial harmful algal blooms (cyanoHABs), and across all three Lake Erie basins. Over three field seasons, $^{15}\text{NH}_4^+$ stable isotope tracers were employed to quantify nitrification rates, and relative abundances of ammonia-oxidizing bacteria (AOB) and archaea (AOA) were determined via qPCR. Nitrification rates ranged from undetectable to $1,270 \text{ nmol L}^{-1} \text{ d}^{-1}$ and were generally greatest in the western basin near the Maumee River mouth (a major nutrient source). Nitrification rates were highest in early summer, and often lowest during peak cyanoHAB months (August and September), before increasing again in October. In the western basin, nitrification was negatively correlated with cyanobacteria biomass. There were no consistent differences in nitrification rates between the three Lake Erie basins. Over the three years in western Lake Erie, AOB and AOA were often present in high and similar abundances, but overall, AOB exceeded AOA, particularly in 2017. No relationships were observed between nitrification rates and AOB and AOA abundances. Thus, despite abundant ammonia-oxidizer DNA, lower nitrification rates during cyanoHABs suggest that nitrifiers were poor competitors for regenerated and available NH_4^+ during cyanoHABs, as also observed in similar systems. Low nitrification rates during cyanoHABs could limit system N removal via denitrification, a natural pathway for excess nitrogen (N) removal and a valuable ecosystem service. Lower denitrification rates allow more bioavailable N to remain in the system and support biomass and microcystin production; therefore, these results help explain how non-N-fixing cyanoHABs persist, despite low bioavailable N concentrations during cyanoHABs, and support management efforts to reduce external N loading to eutrophic systems.

Keywords: nitrogen, cyanobacteria, eutrophication, freshwater

INTRODUCTION

Lake Erie is the smallest (by volume), shallowest, and warmest of the Laurentian Great Lakes and is subject to ecological and human health concerns due to annual and seasonal harmful cyanobacterial blooms (cyanoHABs; Watson et al., 2016). The proliferation of western basin cyanoHABs in summer since the mid-1990s has been linked to complex interactions of physicochemical factors, including the availabilities of phosphorus (P; e.g., Scavia et al., 2014) and chemically reduced nitrogen (N) forms, such as ammonium (NH_4^+ ; Newell et al., 2019; Hoffman et al. 2022). Since the late 1980s, all three basins have been invaded by *Dreissenid* mussels (zebra and quagga; Barbiero and Tuchman, 2004), which improve water clarity but also promote cyanoHABs via rapid nutrient cycling (Conroy et al., 2005) and selective filter-feeding (Vanderploeg et al., 2001). The central basin is subject to cyanoHABs dominated by *Dolichospermum* (capable of N fixation) in June and July and non-N-fixing *Microcystis* in August and September (Chaffin et al., 2019). *Microcystis* are particularly competitive for NH_4^+ (Blomqvist et al., 1994; Monchamp et al., 2014) and may spread to or be transported into the central basin (Chaffin et al., 2019).

Nitrification is a key link between organic matter remineralization (producing NH_4^+) and N removal via denitrification (producing N_2 gas, which leaves the system). Nitrification thus helps mitigate excess N loading by producing the substrate for, and often coupling with, denitrification (e.g., Boedecker et al. 2020). In canonical nitrification, ammonia oxidation converts NH_4^+ to nitrite (NO_2^-), which is then oxidized to nitrate (NO_3^-) via NO_2^- oxidation (Prosser, 1990; Hart et al., 1994; Kuypers et al. 2018). In this two-step process, ammonia oxidation is the rate-limiting step, in which ammonia monooxygenase (encoded by the *amoA* gene) transforms NH_4^+ to an intermediate compound, hydroxylamine, which is then converted to NO_2^- (Ward, 2008). Ammonia oxidation is performed by chemolithoautotrophic

52 bacteria (ammonia-oxidizing bacteria, AOB; Wagner et al., 1995) and archaea (ammonia-oxidizing
53 archaea, AOA; Francis et al., 2005). Together, they comprise the community of ammonia-oxidizing
54 organisms (AOO). The second step of nitrification, NO_2^- oxidation, is carried out by NO_2^- -oxidizing
55 bacteria (NOB). Historically, NO_2^- oxidation has been considered separate from (although often closely
56 linked to) ammonia oxidation, wherein the product of ammonia oxidation serves as the substrate for NO_2^-
57 oxidation (Ward, 2008). The canonical understanding of nitrification has been updated by the discovery
58 of certain NOB that can perform the entire nitrification pathway (complete ammonia oxidation, or
59 comammox), and their role in the environment is still being described (Daims et al., 2015; Kuypers et al.,
60 2018).

61 Nitrification links chemically reduced and oxidized N forms, and, thus, represents the link
62 between the most favorable N form for many primary producers, including cyanoHABs (Glibert et al.,
63 2016), and the substrate for natural N removal (via denitrification). *Microcystis* dominates cyanoHABs in
64 western Lake Erie, cannot fix atmospheric N, and are excellent competitors for NH_4^+ (Takamura et al.,
65 1987; Blomqvist et al., 1994; McCarthy et al. 2007); as such, they may quickly deplete bioavailable N
66 pools. CyanoHABs in eutrophic waterbodies, such as Lake Erie, may threaten the ability of the system to
67 compensate for excess N inputs by suppressing nitrification via substrate competition (Hampel et al.,
68 2018). *Microcystis* in culture exhibits a half-saturation constant (K_m) related to maximum specific growth
69 rate for NH_4^+ of 0.5–37 μM , higher than that of AOA but within the range for AOB (Nicklisch and Kohl,
70 1983). Therefore, AOO community structure during cyanoHABs may influence the ability of nitrifiers to
71 compete with bloom-forming taxa for NH_4^+ and alter the fate of NH_4^+ via removal or recycling (Hampel
72 et al., 2018, 2020).

73 As a result of high external nutrient loading from adjacent watersheds, Lake Erie experiences
74 annual, seasonal cyanoHABs dominated by non-N fixing *Microcystis*, which are largely restricted to the
75 western basin (Watson et al., 2016). Two studies have shown that AOB greatly outnumber AOA in the
76 sediments (Bollman et al., 2014) and water column (Mukherjee et al., 2016) of Lake Erie. A previous
77 study estimated that 96% of the N removed in Lake Erie sediments was due to coupled nitrification-
78 denitrification (Small et al., 2014), a finding supported by subsequent work showing that NO_3^- enrichment
79 did not stimulate direct denitrification in western basin sediments (Boedecker et al., 2020). Given high
80 nutrient loads and the physiological importance of NH_4^+ , the lack of measured nitrification rates
81 represents a knowledge gap, particularly in freshwater eutrophic lakes.

82 Direct measurements of water column nitrification in lakes are few compared to rates from
83 coastal systems and the open ocean (Damashek and Francis, 2018). Nitrification in the water column of
84 eutrophic lakes has been quantified in Lake Taihu (Hampel et al., 2018), Lake Mendota (Hall, 1986), and
85 Lake Okeechobee (Hampel et al., 2019, 2020). Nitrification often represented a small proportion of total
86 microbial NH_4^+ demand in these lakes, which are all affected by cyanoHABs, indicating that nitrifiers
87 often are not competitive for NH_4^+ during cyanoHABs. Along with nitrification rates, characterizations of
88 the AOO community in the water column of freshwater systems is lacking compared with those from
89 sediments in the same systems (Damashek and Francis, 2018; Hampel et al., 2020). Accordingly, there is
90 a need to quantify both nitrification rates and the abundance of AOB and AOA in freshwater systems.

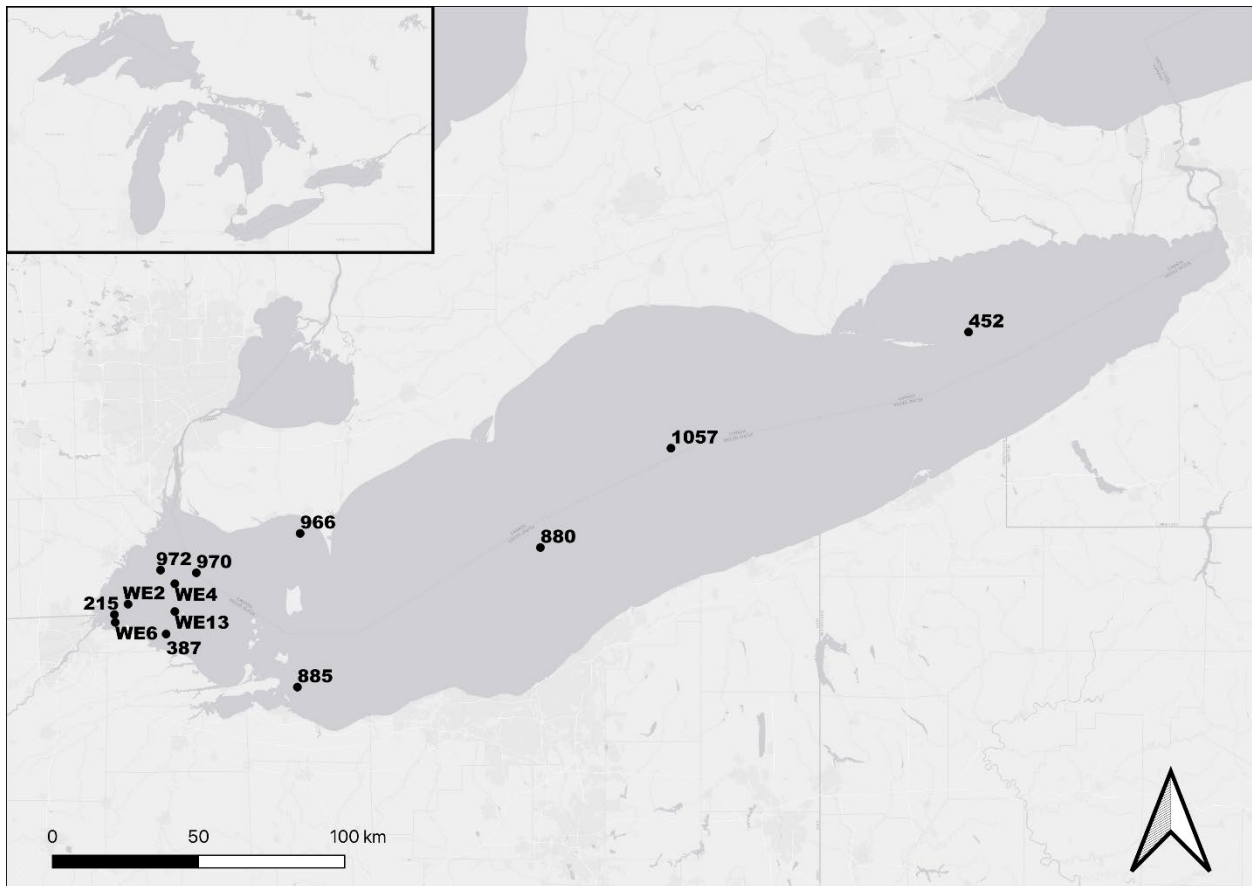
91 Quantifying the fate of NH_4^+ is crucial for modeling the ability of denitrification to remove excess
92 bioavailable N from aquatic systems. Seasonally, models and incubations both show that NO_3^- limits
93 direct denitrification in many lakes (Rissanen et al., 2013; Powers et al., 2017; Cavaliere et al., 2018),
94 including Lake Erie (Small et al., 2016; Boedecker et al., 2020). However, many ecosystem models for
95 Lake Erie do not include N or within-system N dynamics (e.g., Bertani et al., 2016; Scavia et al., 2023).
96 The objective of this study was to quantify nitrification rates (as accumulation of $^{15}\text{NO}_2^-$ and $^{15}\text{NO}_3^-$ from
97 $^{15}\text{NH}_4^+$ used as a stable isotope tracer) along seasonal and spatial gradients in Lake Erie. The abundance
98 of *amoA* genes for AOB and AOA were determined to assess AOO community structure in the western
99 basin. This project combined biogeochemical and molecular techniques to determine the relative
100 contribution of nitrifiers to total community NH_4^+ demand (as reported in Hoffman et al., 2022) and the
101 generation of substrate for denitrification. Considering the findings of previous work describing the
102 impact of *Microcystis*-dominated cyanoHABs on nitrification in other freshwater systems, we

103 hypothesized that nitrification rates in Lake Erie would be highest during pre- and post-bloom sampling
104 periods, while rates during cyanoHABs would be lower. It was also expected that AOO community
105 abundances would be lower during peak cyanoHABs and within blooms in the western basin.
106

107
108 **METHODS**

109 **Site description**

110 From west to east, Lake Erie spans a length of approximately 350 km, with a coastline covering
111 nearly four times that length (~1,400 km) and including four U.S. states (Ohio, Michigan, New York, and
112 Pennsylvania) and one Canadian province (Ontario). Its waters represent a surface area of more than
113 25,000 km², with an estimated volume of 483 km³ and a watershed drainage area of 78,000 km² (United
114 States Environmental Protection Agency (US EPA)). This water volume is spread across three
115 morphologically and trophically distinct basins. The western basin, eutrophic and prone to disruptive
116 cyanoHABs, is the shallowest and most well-mixed, with an average depth of 7.4 m. The central basin is
117 mesotrophic, experiences seasonal bottom-water hypoxia and cyanoHABs (Chaffin et al., 2019), and has
118 an average depth of 18.3 m (PA DCNR 2010). The eastern basin is the deepest and least productive, with
119 an average depth of 27 m (Great Lakes Fisheries Commission, 2017), and has the most oligotrophic
120 profile in offshore regions (Edwards et al., 1990; Barbiero and Tuchman, 2004; Depew et al. 2006; Wang
121 et al., 2008).
122



123
124 Figure 1. Sampling stations across Lake Erie. Stations WE2, WE4, WE6, and WE13 were sampled with
125 NOAA GLERL/CIGLR. Numbered stations were sampled with Environment and Climate Change
126 Canada (ECCC) aboard the *CCGS Limnos*.
127

128 **Field sampling**

129 Field sampling procedures, in situ water quality measurements, sampling stations, and sampling
130 dates for nitrification rates were previously described in Hoffman et al. (2022). Briefly, water sampling in
131 western Lake Erie was conducted monthly, during the growing season, alongside the NOAA Great Lakes
132 Environmental Research Laboratory's (GLERL) and Cooperative Institute for Great Lake Research's
133 (CIGLR) weekly HABs monitoring program at three or four stations at both ~1 m below the water surface
134 and ~1 m above sediments (Fig. 1; station and month combinations varied depending on year). Ambient
135 physicochemical parameters, including water temperature, dissolved oxygen concentration, Secchi depth,
136 conductivity, photosynthetically active radiation (PAR), turbidity, and total suspended solids (TSS) were
137 collected by NOAA GLERL and CIGLR. These data, plus dissolved and particulate microcystin
138 concentrations, are published by NOAA GLERL and publicly available at
139 [https://www.ncei.noaa.gov/access/metadata/landing-page/bin/iso?id=gov.noaa.nodc:GLERL-CIGLR-](https://www.ncei.noaa.gov/access/metadata/landing-page/bin/iso?id=gov.noaa.nodc:GLERL-CIGLR-HAB-LakeErie-water-qual)
140 [HAB-LakeErie-water-qual](https://www.ncei.noaa.gov/access/metadata/landing-page/bin/iso?id=gov.noaa.nodc:GLERL-CIGLR-HAB-LakeErie-water-qual). 12 ml of water was filtered (0.22 μm Nylon syringe filters; Thermo Fisher
141 Target 2) immediately upon collection (Reed et al., 2023) for analysis of dissolved nutrient concentrations
142 and stored frozen until analysis. A Lachat Quikchem 8500 was used to measure concentrations of NH_4^+
143 (alkaline phenol and DCIC method 31-107-06-1-G), NO_2^- , NO_3^- (sulfanilamide/NED Cd reduction
144 method 31-107-04-1-E), ortho-phosphate (ortho-P; molybdate method 31-115-01-1-I), and urea (diacetyl
145 monoxime/thiosemicarbazide method 31-206-00-1-A). Water for nitrification rate incubations was
146 collected with a 10-L Niskin bottle, decanted into 20-L polyethylene carboys, and transported to Wright
147 State University for isotope amendment, incubation, and analysis.

148 To determine AOO community structure and *amoA* abundance in the western basin, lake water
149 was filtered (60–360 ml) through Sterivex filters (0.2 μm , EMD Millipore). Air was flushed through
150 filters to clear residual water, then filled with RNALater (Thermo Fisher) and capped to preserve genetic
151 material. Filters were placed on dry ice in the field, then frozen at -80 °C until DNA extraction.

152 For the larger lake surveys, sampling was conducted aboard the *CCGS Limnos* in conjunction
153 with the Environment and Climate Change Canada (ECCC) HABs program. Cruises occurred in October
154 2015, August/September 2017, and October 2017. Six stations (215, 885, 387, 880, 1057, and 452) were
155 sampled in 2015, seven (215, 966, 970, 972, 880, 1057, and 452) in Aug/Sept 2017, and two (880 and
156 452) in Oct 2017 (Fig. 1). During each cruise, stations were selected to establish a gradient across the
157 east-west span of Lake Erie. During Oct 2017, inclement weather limited sampling to two stations, one in
158 the central basin (880; Fig. 1), and one in the eastern basin (452; Fig. 1). At each station, water was
159 collected from two depths (~1 m below the water surface and ~2 m above the sediment surface) using a
160 10-L Niskin bottle. 10 L of water was collected from each depth and transferred into 3-L polyethylene
161 containers. Ambient nutrient and DNA samples were immediately filtered and stored frozen until
162 analysis, as described above. Other physicochemical parameters (water temperature, conductivity, oxygen
163 concentration, oxygen saturation, turbidity, transmittance, chlorophyll-*a*, phycocyanin, and pH) were
164 collected by ECCC, along with chlorophyll-*a* specific estimates of phycocyanin-Cyanobacteria,
165 Heterokontophyta and Pyrrophyta, which includes diatoms and chrysophytes), and total algae biomass via
166 fluoroprobe (Zastepa et al., 2023).

167 168 **Incubations**

169 Nitrification rates were quantified via the addition of $^{15}\text{NH}_4^+$. A control bottle containing site
170 water, but no isotope amendment, was used to ensure no ^{15}N contamination of experiments. For
171 nitrification rates, site water was decanted into triplicate, colorless, translucent, 125 ml Nalgene bottles
172 and amended with trace amounts of $^{15}\text{NH}_4\text{Cl}$ (98 atom%), which equated to approximately 20% of the
173 ambient NH_4^+ pool, except in August and September, when NH_4^+ pools were often depleted (SI Table 1).
174 Samples from each bottle were filtered immediately (0.22 μm) and frozen. Bottles were incubated for 16–
175 25 hours (SI Table 1) in simulated lake conditions (outdoor water bath) before being sampled again.
176 Samples were filtered (0.22 μm) into polystyrene tubes for total NH_4^+ concentration (i.e., $^{14}\text{N} + ^{15}\text{N}$) and
177 immediately frozen. Additional 30 mL samples for $^{15}\text{NO}_x$ accumulation from tracer additions were
178 filtered into 50 ml centrifuge tubes and frozen until pre-analysis preparation (see below).

179 Initially, the amount of added $^{15}\text{NH}_4^+$ that accumulated as $^{15}\text{NO}_2^-$ and $^{15}\text{NO}_3^-$ was determined
180 separately (as described below). Based on results from 2015 and 2016, samples from 2017 were analyzed
181 for total $^{15}\text{N-NO}_x$ ($\text{NO}_2^- + \text{NO}_3^-$) accumulation, rather than for $^{15}\text{NO}_2^-$ and $^{15}\text{NO}_3^-$ separately.
182

183 **Sample preparation and nitrification rate calculations**

184 Nitrification rates were measured as the accumulation of $^{15}\text{NO}_2^-$ and $^{15}\text{NO}_3^-$ from added $^{15}\text{NH}_4^+$
185 tracer. To quantify $^{15}\text{NO}_2^-$ accumulation, a sodium azide (NaN_3) solution was used to convert all NO_2^- in
186 each sample to N_2O (McIlvin and Altabet, 2005). For $^{15}\text{NO}_3^-$ accumulation, a cadmium (Cd) reduction
187 step, coupled to the azide method, was used to first reduce all NO_3^- to NO_2^- before converting NO_2^- to
188 N_2O (Granger and Sigman, 2009; Heiss and Fulweiler, 2016). Post-transformation, samples were sent to
189 the University of California - Davis Stable Isotope Facility for analysis of ^{15}N -labeled (masses 45, 46) and
190 unlabeled (mass 44) N_2O . Control (no amendment) incubations were included at all steps to ensure lack
191 of ^{15}N contamination. Nitrification rates were calculated according to Heiss et al. 2022.

192 Rates were corrected for reduction efficiency of the NaN_3 reaction and are reported in units of
193 nM day^{-1} .

194 In 2015 and 2016, $^{15}\text{NH}_4^+$ converted to and accumulated as NO_2^- and NO_3^- was quantified
195 separately. Accumulation rates of $^{15}\text{NO}_2^-$ were frequently, but not always, detectable and were generally
196 $\sim 5\%$ of $^{15}\text{NO}_3^-$ accumulation rates. Therefore, in 2017, only total $^{15}\text{NO}_x$ accumulation was measured, and
197 for 2015 and 2016, nitrification rates are presented as the sum of ^{15}N accumulation in both NO_2^- and NO_3^-
198 pools.
199

200 **Molecular analysis**

201 DNA was extracted from Sterivex filters with the Qiagen Puregene© Core A kit and a modified
202 version of the manufacturer's protocol (based on Ward et al. 2000). Residual RNA Later was removed
203 from the filters with a phosphate buffer solution rinse, followed by the addition of cell lysis buffer.
204 Proteinase K was then added, and filters were incubated for consecutive 1 h cycles at 55 and 65 °C.

205 Following extraction, DNA yields were quantified via spectrophotometry (Thermo Fisher
206 Nanodrop 2000) and used to determine the abundance of functional genes for ammonia oxidation (*amoA*)
207 in both bacterial and archaeal lineages via qPCR. Cleaned and cloned PCR amplicons were used to create
208 plasmid standards for qPCR. For archaeal *amoA*, a 635 base-pair (bp) region was amplified with Arch-
209 amoAF and Arch-amoAR primers (Francis et al., 2005, SI Table 2). A 491 bp region was amplified with
210 amo-AF and amo-A2R primers (Rotthauwe et al., 1997, SI Table 2) for determination of bacterial *amoA*.

211 qPCR analyses were performed on 96-well plates. Serial dilutions of standards were plated
212 alongside samples and negative (non-template containing) controls. Negative controls, five standards, and
213 samples were plated in triplicate. The reaction mixture was formulated with Luna Universal qPCR Master
214 Mix (New England Biolabs) following manufacturer instructions. 5–25 ng of template was added, and
215 each reaction was performed on an Eppendorf realplex² thermocycler. Archaeal and bacterial *amoA* qPCR
216 programs were carried out as follows:

217 AOA *amoA*: 40 cycles of 95 °C for 2 min, 95 °C for 30 s, 53 °C for 45 s, 72 °C for 1 min; 72 °C
218 for 5 min; melting curve (Francis et al., 2005).

219 AOB *amoA*: 40 cycles of 95 °C for 2 min, 94 °C for 45 s, 56 °C for 30 s, 72 °C for 1 min; 72 °C
220 for 5 min; melting curve (Beman et al., 2008).

221 Gene copy number was determined as:

222 $\text{Copy number} = (\text{ng} * \text{number mol}^{-1}) / (\text{bp} * \text{ng g}^{-1} * \text{g mol}^{-1} \text{ of bp})$

223 and is given in units of gene copies per ml of sample water.
224

225 **Statistical analysis**

226 All statistical analyses were performed in R version 4.2.2 (R Core Team, 2022). Duplicate and
227 triplicate nitrification samples were heterogeneous within each station and sampling event and were thus
228 treated as individual data points. Western basin data collected during NOAA GLERL and CIGLR weekly

229 cruises were analyzed separately from data collected during ECCC cruises due to different temporal
230 sampling patterns and accompanying metadata for comparative analysis.

231 For western basin samples, nitrification rates failed assumptions of normality, even after log-
232 transformation and normalization. Accordingly, negative binomial general linear models (“mass”
233 package; Ripley et al., 2013) were used to determine variability in spatial and temporal gradients for rates.
234 The effect of month as a random effect on nitrification rates from each station was explored via mixed
235 models; month was only responsible for 3% of variance in rates by station, so temporal conditions (month
236 and year) were analyzed as fixed effects on measured nitrification rates. For ECCC cruise samples, due to
237 small sample size ($n = 2$ or 3 for each station/depth at each time point), and based on the hypothesis of
238 basin-driven patterns, a nonparametric Kruskal-Wallis test was used to determine spatial differences in
239 rates, which were grouped by depth or basin instead of station.

240 For *amoA* gene abundance, standard linear models (lm function in base R “stats” package) were
241 employed following log transformation. Relationships between nitrification rates, *amoA* gene copies, and
242 environmental variables were explored with Spearman’s rank correlations (“Hmisc” package; Harrel Jr.,
243 2019). Tukey’s HSD post-hoc tests for each model were performed using the “emmeans” package (Lenth
244 et al, 2019).

245

246 RESULTS

247 Ambient conditions during western Lake Erie sampling

248 Environmental parameters for sampling stations and dates for nitrification experiments occurred
249 simultaneously with NOAA GLERL monitoring and are also reported in Hoffman et al. (2022). Data
250 tables for ambient water quality parameters used in statistical analyses are included in SI Tables 3 and 4.

251 Water column dissolved inorganic N (DIN), urea, and ortho-P concentrations in the western basin
252 generally followed expected patterns, with concentrations decreasing with distance from the Maumee
253 River. Phycocyanin (a diagnostic pigment for cyanobacteria), particulate and dissolved microcystins, and
254 chlorophyll *a* followed similar patterns. Greatest phytoplankton biomass concentrations were measured at
255 the westernmost station (WE6) at the height of the bloom, corresponding with depleted or undetectable
256 ambient NH_4^+ concentrations.

257

258 Ambient conditions during ECCC cruise sampling

259 In October 2015, six stations ranging in depth from 5.5 m (387) to 52 m (452) were sampled, and
260 water quality data are reported in supplemental tables. Nutrient concentrations were greatest at station 215
261 nearest the Maumee River inflow to western Lake Erie and decreased with distance, except for NH_4^+ and
262 NO_3^- , which had concentration spikes in bottom waters at stations 1057 and 452 (SI Table 5).

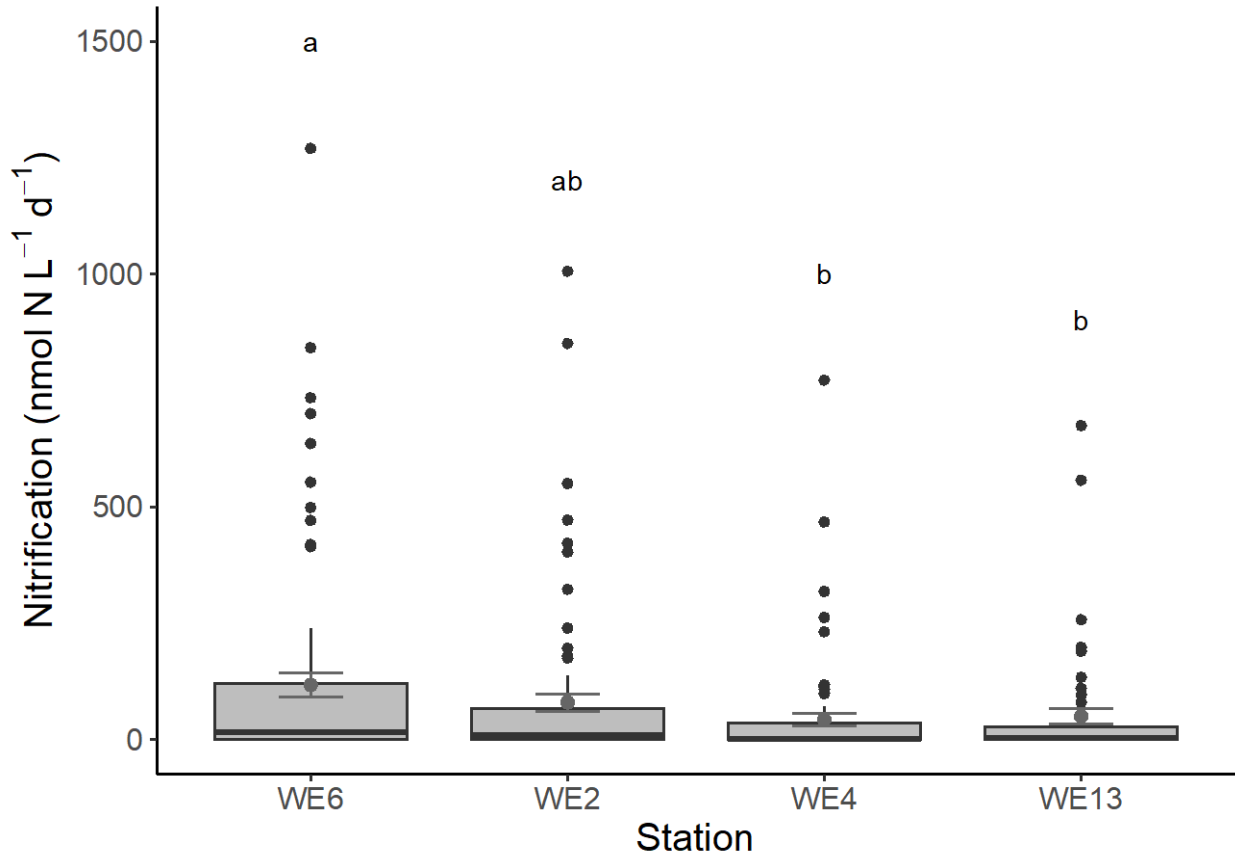
263 Phytoplankton biomass was concentrated in the western basin and decreased moving eastward, with
264 cyanobacteria and diatoms comprising most of the algal biomass at most stations and depths (SI Table 6).

265 During the August 2017 cruise, seven stations with a depth range of 8 m at the shallowest station
266 in the western basin (215) to 52 m in the eastern basin (452) were sampled. O_2 concentrations in surface
267 waters were near or above 100% saturation, and deep waters from stations ≤ 11 m deep remained
268 oxygenated. However, at the two central basin stations (880 and 1057), O_2 concentrations decreased to
269 hypoxic (and near anoxic) levels in bottom water (SI Table 7). As in the previous cruise, cyanobacteria
270 and diatoms dominated the phytoplankton community at most sampling sites (SI Table 8). Nutrient
271 concentrations were lower in the westernmost part of the western basin. East of station 215, NO_3^-
272 concentrations in the western basin were 13–15 μM compared to 6–8 μM in the central basin, regardless
273 of depth. In bottom waters at station 452 (52 m), there was a deep pool of NO_3^- , with higher
274 concentrations than any others measured that week (SI Table 7).

275 In October 2017, due to inclement weather, only one central basin (880) and one eastern basin
276 (452) station were sampled. NH_4^+ concentrations were 3x greater at station 880 than at 452, but NO_3^-
277 concentrations at station 452 were 2–7x greater than at 880, with the largest pool of NO_3^- observed at
278 depth, as also observed in August (SI Table 7).

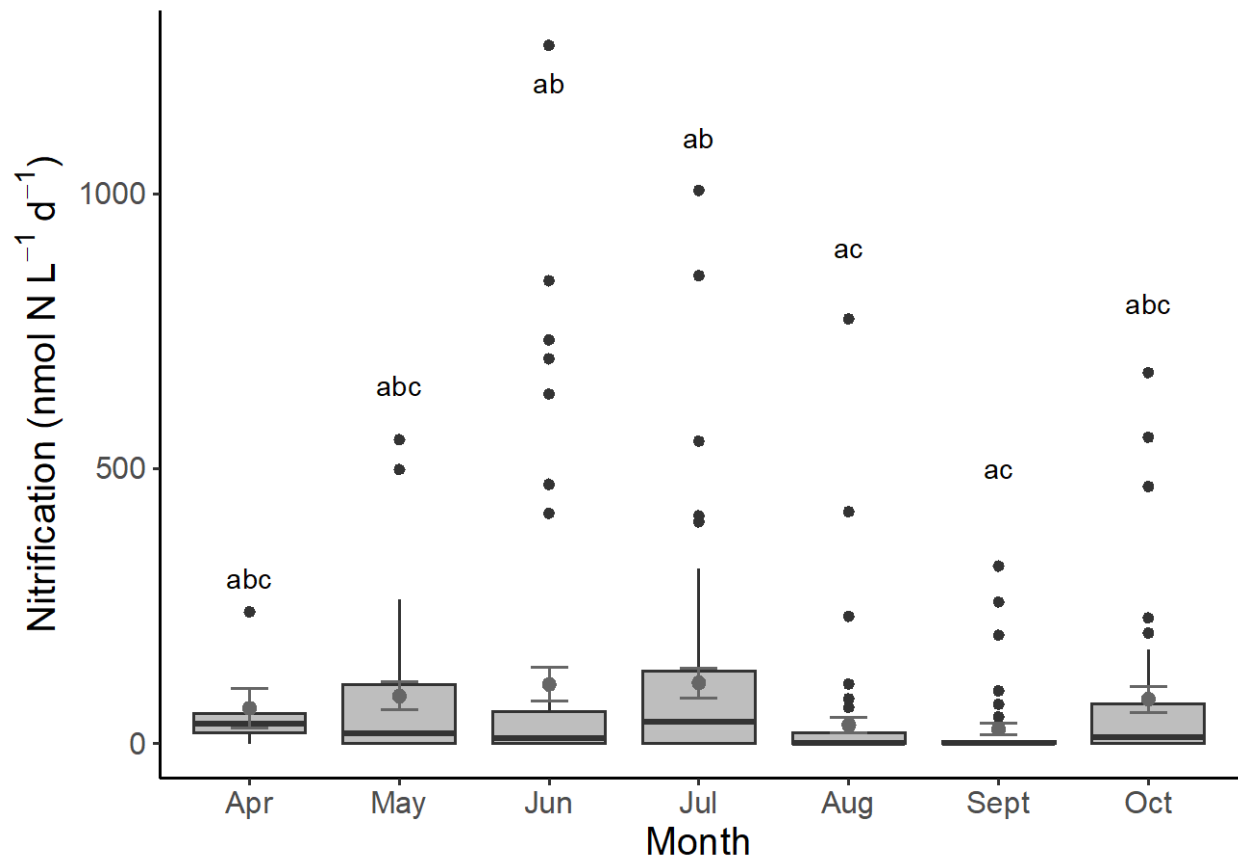
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280 Nitrification rates in western Lake Erie
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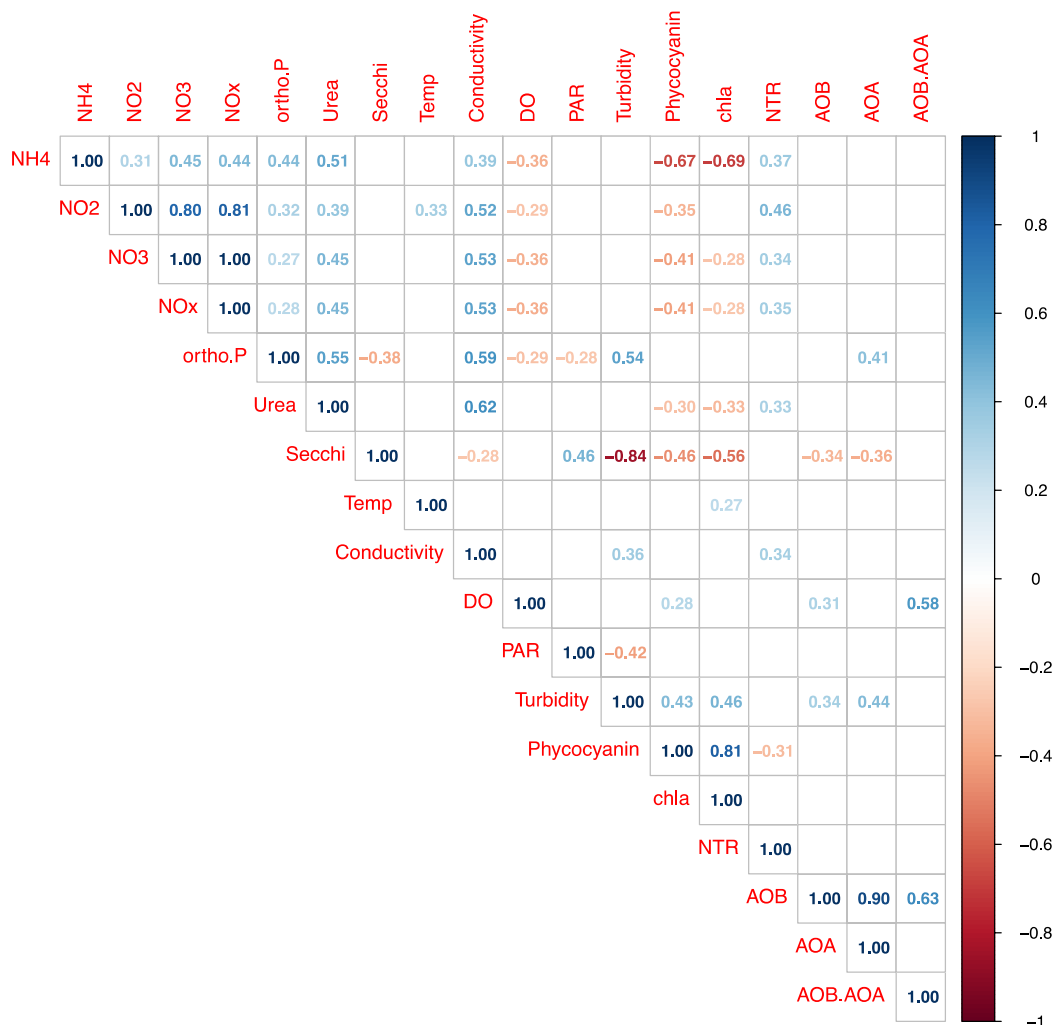
283 Figure 2. Nitrification rates ($\text{nmol L}^{-1} \text{d}^{-1}$) at four western Lake Erie sampling stations. Means (\pm SE) are
284 indicated on each boxplot. Letters reflect differences in nitrification rates between stations (Tukey's HSD
285 post-hoc tests). n for each station is as follows: WE6 = 90, WE2 = 96, WE4 = 90, WE13 = 60.
286
287

288 Nitrification rates were not influenced by depth across all stations and months, but rates from
289 samples collected ~ 1 m above the sediment were greater ($p = 0.006$) than surface measurements in 2015.
290 Nitrification rates at WE6 were greater than those at WE4 and WE13 ($p = 0.0025$ and 0.030 ,
291 respectively), but stations WE2, WE4, and WE13 were not different from each other ($p > 0.20$; Fig. 2).
292 Nitrification peaked in late spring/early summer, decreased to low or undetectable levels during peak
293 cyanobacteria blooms (August and September), and then increased in October once the bloom dissipated (Fig. 3).
294 When all three sampling years were considered together, nitrification rates in June and July were greater
295 than those in August and September ($p \leq 0.03$ for all pairings). Within each year, nitrification was greater
296 in July than in September in 2015 ($p < 0.001$, Fig. 3), while June and July had greater rates than August in
297 2016 ($p < 0.001$ and $p = 0.048$, respectively; Fig. 3). In 2017, there were no differences in nitrification
298 rates between sampling months (SI Fig. 1). Differences in nitrification rates at individual stations between
299 years were largely driven by rates at WE4 but were not statistically robust; however, these relationships
300 may be ecologically relevant ($p = 0.053 - p = 0.058$) for the following: at WE4, nitrification rates in June
301 2016 were greater than those in June 2015 and June 2017; July 2017 nitrification rates were greater than
302 those in July 2015 or July 2016; and nitrification rates in August 2017 were greater than those in August
303 2015 (Fig. 3 and SI Fig. 1).
304



305
 306 Figure 3. Nitrification rates in each month at each station across all three years. Means (\pm SE) are
 307 indicated on each boxplot. Letters reflect differences in nitrification rates between months (Tukey's HSD
 308 post-hoc tests). n for each month is as follows: April = 6, May = 36, June = 66, July = 60, August = 66,
 309 September = 54, October = 48.

310
 311 In 2015, nitrification rates ranged from undetectable to 851 nmol L⁻¹ d⁻¹ (48.5 ± 15.0 , mean \pm SE;
 312 $n = 72$), with the greatest rates measured at WE2 in July (Figs. 2 and 3 and SI Fig. 1). Nitrification rates
 313 in August and September were much lower and often undetectable, although rates as high as 230 nmol L⁻¹
 314 d⁻¹ were observed at WE4. In 2016, rate maxima (1270 nmol L⁻¹ d⁻¹ at WE6 in June) were nearly double
 315 those in 2015 (mean 84.8 ± 18.0 ; $n = 119$). During July and August, nitrification rates remained below 35
 316 nmol L⁻¹ d⁻¹ but increased to 200–300 nmol L⁻¹ d⁻¹ in September and 675 nmol L⁻¹ d⁻¹ at WE13 in October.
 317 Nitrification rates up to 1,000 nmol L⁻¹ d⁻¹ were measured in 2017 (79.7 ± 16.1 ; $n = 106$); as in 2015, this
 318 rate maximum was observed at WE2 in July. Compared to previous years, some of the greatest rates (up
 319 to 772 nmol L⁻¹ d⁻¹; SI Fig. 1) were observed at stations furthest from the Maumee River (WE4 and
 320 WE13) in August and September 2017, while rates at the westernmost stations increased again in
 321 October. Nitrification rates were positively correlated with ambient NH₄⁺ ($p = 0.007$), NO₂⁻ ($p = 0.0002$),
 322 NO₃⁻ ($p = 0.0051$), and urea ($p = 0.023$) concentrations, as well as conductivity ($p = 0.022$), and were
 323 negatively correlated with phycocyanin ($p = 0.025$, Fig 4).
 324

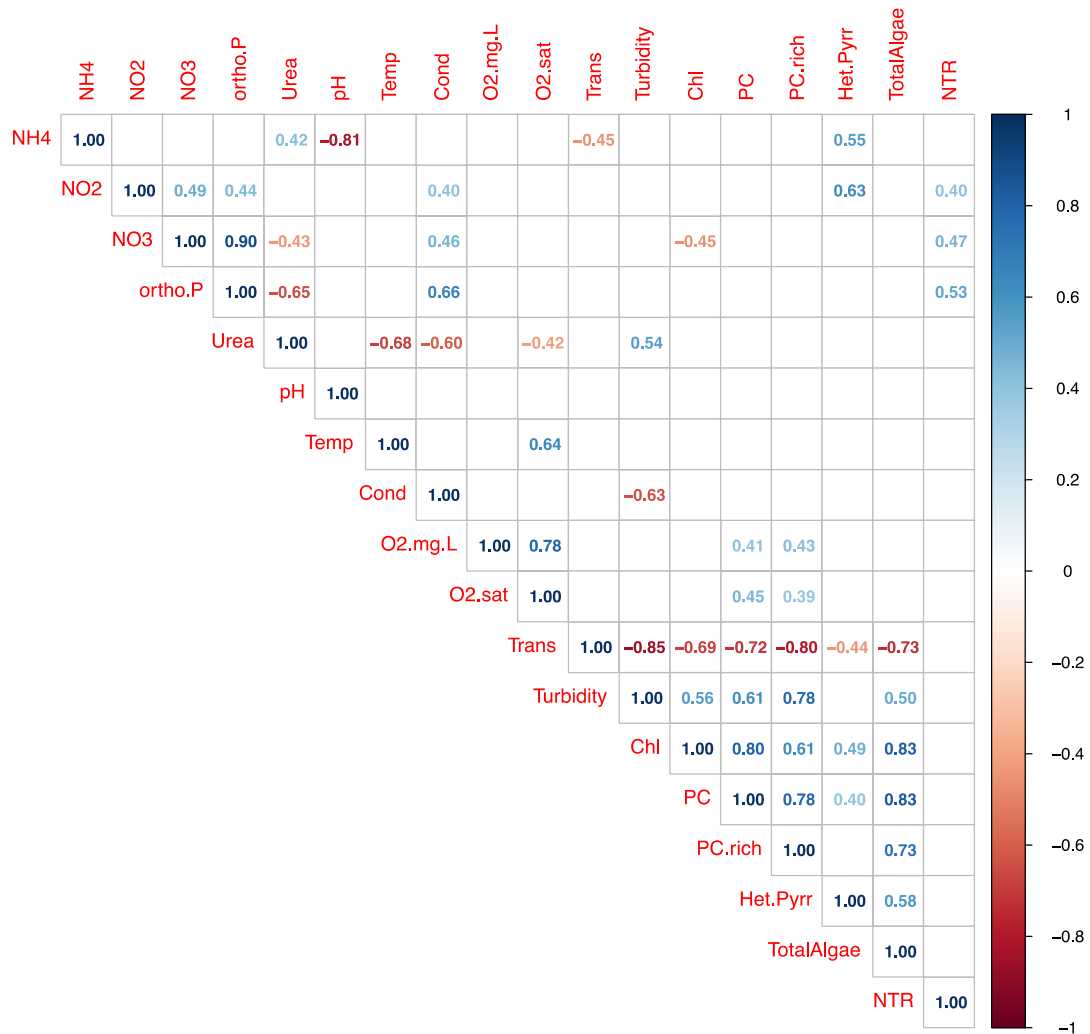


325
 326 Figure 4. Spearman's rank correlations for environmental variables, nitrification rates, and gene counts in
 327 WLE. Abbreviations are as follows: ammonium (NH₄), nitrite (NO₂), nitrate (NO₃), nitrate + nitrite
 328 (NO_x), water temperature (Temp), dissolved oxygen (DO), photosynthetically active radiation (PAR),
 329 chlorophyll-*a* (chl_a), nitrification rate (NTR), copies of *amoA* in ammonia-oxidizing bacteria and archaea
 330 (AOB and AOA, respectively), and the ratio of AOB to AOA (AOB.AOA). Values shown for each
 331 pairing are the correlation coefficient (rho) at $p < 0.05$.

332
 333 **Nitrification rates across Lake Erie on ECCC Cruises**

334 During cruise events, nitrification rates ranged from undetectable to 878 nmol L⁻¹ d⁻¹, with the
 335 greatest rates measured in August 2017 in the western basin. In October 2015, there was no effect of
 336 depth on nitrification rates, but rates tended to be greater in deep water samples at stations 387, 885, and
 337 880 (SI Fig. 2). Nitrification rates at depth were higher overall in August 2017 ($p = 0.027$), and at the
 338 deepest station (452; ~52 m). Nitrification rates in bottom water samples were 100 times greater than in

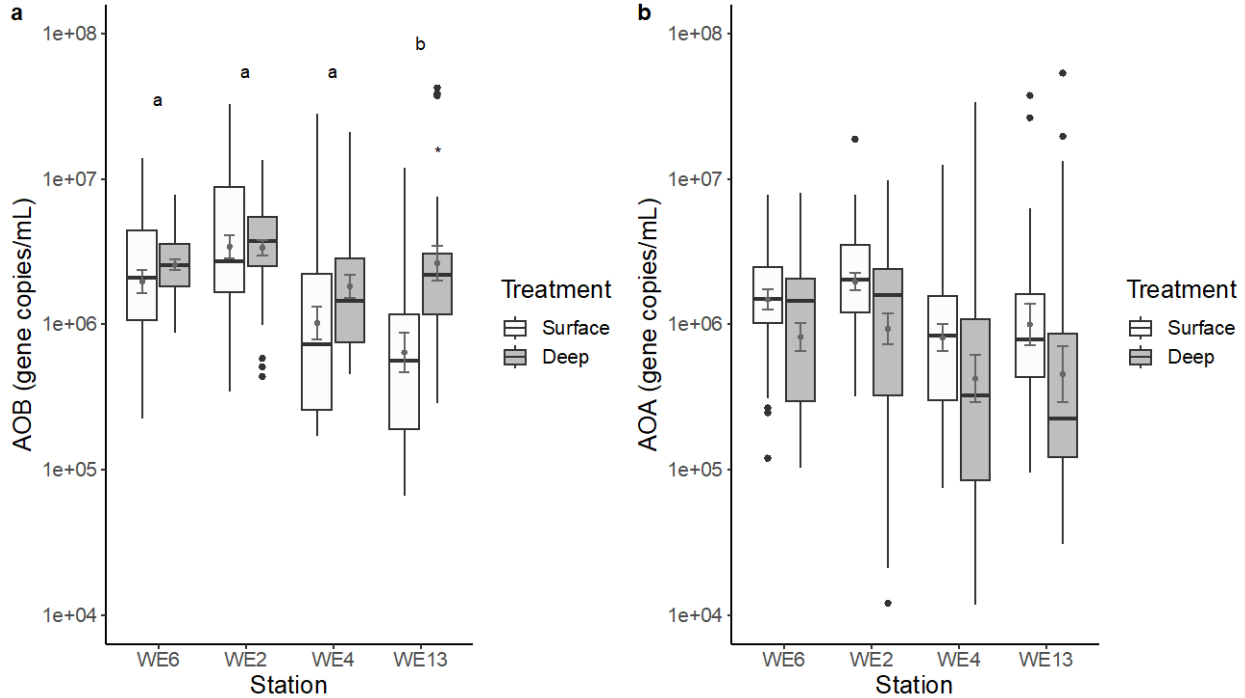
339 surface water. In October 2017, nitrification rates were greater ($p = 0.003$) at the eastern basin station
 340 (452) than the central basin station (880). When all cruise data is considered, nitrification rates were
 341 positively correlated with reaction products (NO_2^- and NO_3^-), as well as ortho-P, but not with any other
 342 physicochemical or biological parameter (Fig. 5).
 343



344
 345 Figure 5. Correlation plot of nitrification rates and environmental variables in October 2015, August
 346 2017, and October 2017 from ECCC cruises. Abbreviations/subgroups are as follows: water temperature
 347 (Temp), conductivity (Cond), oxygen concentration (O2.mg.L), oxygen saturation (O2.sat), transmittance
 348 (Trans), chlorophyll-*a* (chl_a), phycocyanin (PC), phycocyanin-Cyanobacteria (PC.rich), Het.Pyrr
 349 (Heterokontophyta and Pyrrophyta which includes diatoms, chrysophytes, and dinoflagellates;
 350 nomenclature from Zastepa et al. 2023b), ammonium (NH₄), nitrite (NO₂), nitrate (NO₃), and
 351 nitrification rate (NTR). Values shown are coefficients (rho) from Spearman's correlations at $p < 0.05$.

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Ammonia oxidizer abundances in western Lake Erie



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Figure 6. (a) AOB and (b) AOA *amoA* copies at each station across all three sampling years. Means (\pm SE) are indicated on each boxplot. Letters reflect differences in *amoA* copies between stations (Tukey's HSD post-hoc tests). n for each *amoA* group at each station is as follows: WE6 = 72, WE2 = 84, WE4 = 72, WE13 = 47.

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For bacterial *amoA*, spatial and depth gradients influenced gene counts in the western basin. Across all three years, and with increasing distance from the Maumee River, bacterial *amoA* gene counts at WE6 and WE2 were not different from each other, nor were WE4 and WE13, but each of the former grouping had greater AOB *amoA* copies than the latter ($p < 0.03$, Fig. 6a). The two deepest stations (WE4 and WE13) had greater AOB *amoA* gene copy numbers in near-bottom versus surface samples ($p < 0.05$, Fig. 6a). Archaeal *amoA* gene copies were not different between stations, but within each station, there were more AOA *amoA* gene copies in surface versus near-bottom samples (Fig. 6b).

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AOB gene copies were greater in 2016 than the other two sampling years, and there was no difference between 2015 and 2017 (SI Fig. 3). AOA *amoA* abundances were also greatest in 2016, but all three sampling years were different from each other, with 2017 having lower abundances than the other two (SI Fig. 3). In 2016 and 2017, AOB (Fig. 6a) and AOA (Fig. 6b) *amoA* gene copies in bloom months (July–September) were not different from each other ($p > 0.10$), nor were they different from those in mid-spring (May). However, *amoA* abundances in June were lower than in other months ($p < 0.003$ for all pairings). AOB *amoA* abundances were not correlated with nitrification rates but were positively related to AOA gene copies ($p < 0.0001$), dissolved oxygen concentrations ($p = 0.038$), and turbidity ($p = 0.023$), and negatively correlated with Secchi depth ($p = 0.024$; Fig. 4). AOA gene counts were similarly not related to nitrification rates, were positively correlated with turbidity ($p = 0.0026$) and ortho-P concentrations ($p = 0.004$), and negatively related to Secchi depth ($p = 0.014$; Fig. 4).

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AOB vs. AOA in western Lake Erie

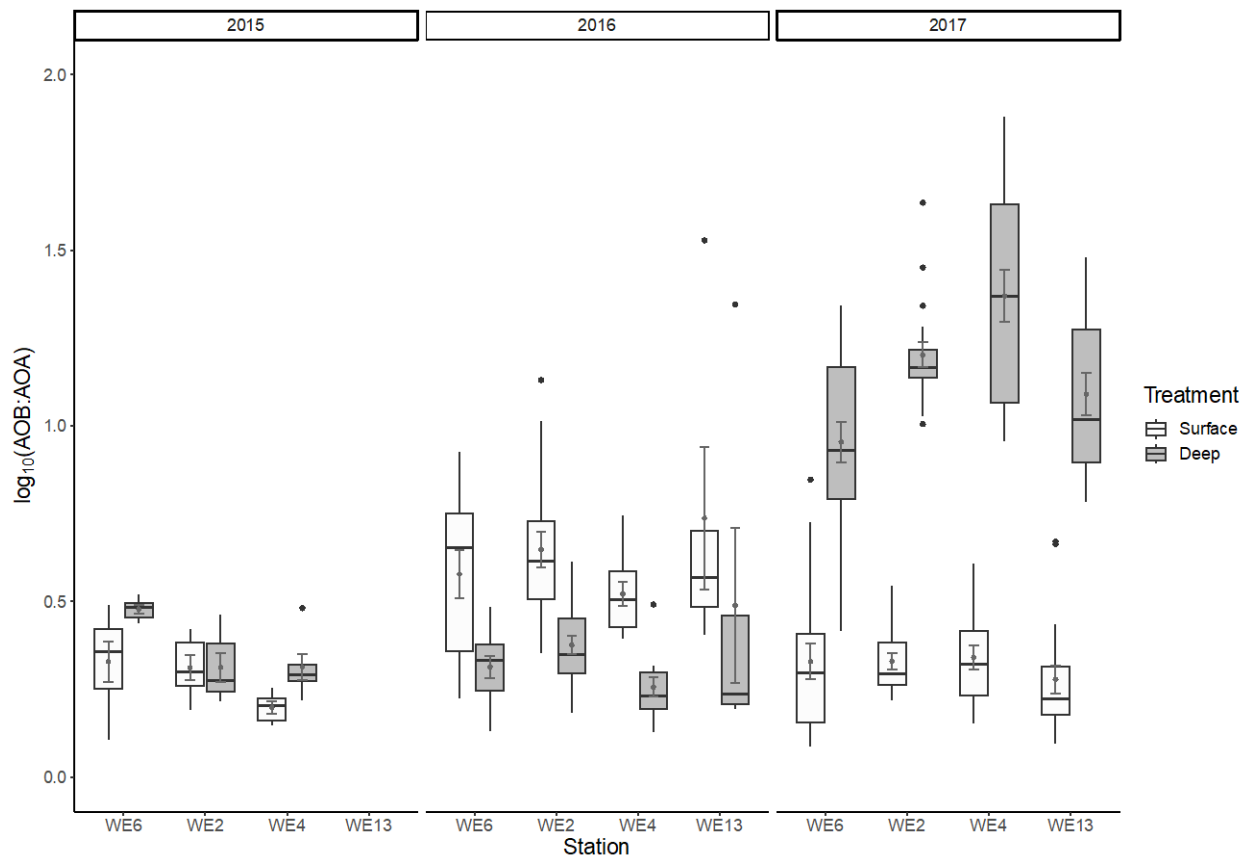
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Overall, AOB *amoA* gene copies ml^{-1} were greater ($p < 0.001$) than AOA in the western basin. To explore potential differences in community composition, the ratio of AOB:AOA gene copies were calculated for each sampling event. The median ratio across the entire dataset was 1.74. In 2015 (when we

383 only collected DNA in August and September), AOO communities were near equal in abundance, with a
 384 median ratio of 0.97. In 2016 and 2017, the median ratio increased to 1.55 and 3.78, respectively. Across
 385 all three years, there was no difference in ratios by station, but within each station, higher ratios were
 386 observed in near-bottom versus surface waters (Fig. 7), an effect driven mostly by high bottom water
 387 ratios in 2017 (Fig. 7). Seasonally, there was no difference across months, but the greatest ratios were
 388 observed in September at WE4 in 2017 (Fig. 7). This peak ratio did not coincide with greatest nitrification
 389 rates measured that year.

390 Ratios in 2015 were not different from those in 2016, but ratios in 2017 were greater at depth ($p <$
 391 0.001) than those in other years. The AOB:AOA ratio ranged from 0.28 – 2.30 in 2015, 0.34 – 33 in 2016,
 392 and 0.22 – 74 in 2017. The effect of depth differed by year, with surface samples having generally greater
 393 ratios than deep samples in 2016, while samples taken at depth in 2017 had markedly greater AOB:AOA
 394 (Fig. 7, SI Fig. 4). *amoA* ratios were correlated with AOB gene copies ($p < 0.0001$) but not with AOA
 395 gene copies or any other environmental parameter (Fig. 4).

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 399 Figure 7. Ratio of AOB:AOA gene copies at each station and depth in each year. In September 2017, the
 400 ratio in near-bottom water at WE4 reached 74, but the y-axis was limited at 40 to preserve scale. n for
 401 each year is as follows: 2015 = 36, 2016 = 96, 2017 = 143.

402
 403 **DISCUSSION**

404 During three study seasons (2015–2017), nitrification rates across Lake Erie and abundances of
 405 genes associated with nitrification (western basin only) were measured to identify spatial and temporal
 406 patterns and possible drivers of nitrification. Results from this study indicate that, despite the ubiquitous
 407 presence of AOO, nitrification in western Lake Erie is greatest when cyanoHABs are not present,
 408 suggesting that nitrification as a link to denitrification is unlikely to compensate for external and internal

409 (e.g., NH_4^+ regeneration in the water column; Hoffman et al., 2022) N loading with regard to mitigating
410 cyanoHABs. These results provide new insights into the spatial dynamics surrounding NH_4^+ cycling and
411 trophic status across all three basins of Lake Erie.
412

413 **Nitrification dynamics in western Lake Erie**

414 Nitrification rates were expected to follow seasonal trends, peaking in pre- and post-bloom
415 periods, and decreasing when cyanoHABs were present (as in Hampel et al., 2018). During summer,
416 when cyanoHAB biomass was greatest (August and September), nitrification rates were the lowest.
417 Ambient water column NH_4^+ concentrations were depleted during cyanoHABs, suggesting substrate
418 limitation of nitrification. Although nitrification was not correlated with chlorophyll *a*, nitrification rates
419 were negatively correlated with phycocyanin (cyanobacteria pigment), which peaked during August and
420 September. During the pre-bloom season, greatest nitrification rates were measured in the westernmost
421 part of the basin, nearest to the primary external nutrient source (Maumee River). However, when blooms
422 were present, greatest nitrification rates were observed at stations furthest from the river inflow.

423 As expected, positive relationships were observed between nitrification rates and substrate (NH_4^+)
424 and end product (NO_2^- and NO_3^-) concentrations. Relationships between nitrification rates and NO_x
425 concentrations are commonly reported (e.g., Pauer and Auer, 2000; Newell et al., 2011; Hampel et al.,
426 2018; Cavaliere and Baulch, 2019) and illustrates the importance of nitrification as a source of substrate
427 for denitrification. Molecular oxygen is obligate for ammonia oxidation, but no relationship was observed
428 between nitrification rates and dissolved oxygen concentrations. The western basin is shallow and
429 generally well-mixed to the sediment surface, and dissolved oxygen is uniform throughout the water
430 column. There was also no relationship between nitrification rates and temperature, which may be due to
431 large cyanoHABs in the warmest months and is consistent with observations from another eutrophic lake
432 afflicted with *Microcystis*-dominated blooms (Taihu; Hampel et al., 2018).

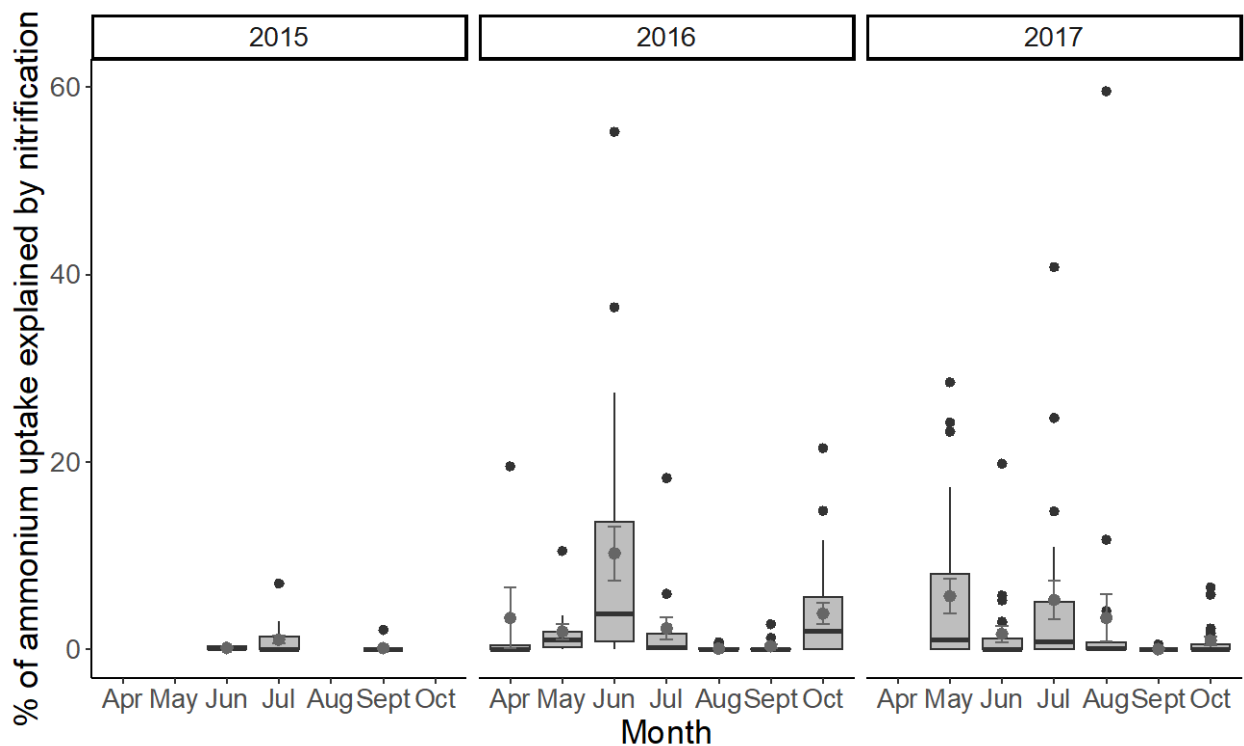
433 Relationships between nitrification rates and pH were also explored, since the optimal range for
434 nitrification generally falls between pH 7.5 and 9, depending on AOO community structure (Shammas et
435 al., 1986; Antoniou et al. 1990; Park et al. 2008). pH data for the water column was not collected as part
436 of the GLERL sampling, so real-time buoy data from WE2, WE4, and WE13 were accessed. No
437 discernable patterns, or much variation in pH levels, were observed, with values ranging from 7.7 to 9.1
438 but mostly remaining between 8 and 9 across sampling seasons and years (CIGLR and NOAA GLERL,
439 2019b). The stability of pH within the optimum range for nitrification indicates that pH is likely not a
440 main driver of observed seasonal differences in nitrification rates.

441 Compared to other lakes in which $^{15}\text{NH}_4^+$ tracer addition methods were employed, late
442 spring/early summer nitrification rates in western Lake Erie greatly exceeded those measured in
443 oligotrophic Lake Superior (USA, 0 – 51 $\text{nmol L}^{-1} \text{d}^{-1}$; Small et al., 2013) but were similar to those in
444 saline Mono Lake (USA, 0 – 480 $\text{nmol L}^{-1} \text{d}^{-1}$; Carini and Joye, 2008) and Lake Croche (Canada, 0 – 333
445 $\text{nmol L}^{-1} \text{d}^{-1}$; Massé et al., 2019). Nitrification rates in western Lake Erie were similar to the range of rates
446 from mesotrophic Lake Lacawac (1 – 568 $\text{nmol L}^{-1} \text{d}^{-1}$; Heiss et al., 2022) and other eutrophic lakes,
447 including lakes Okeechobee (30 – 120 $\text{nmol L}^{-1} \text{d}^{-1}$; James et al. 2011) and Taihu (China, 0 – 3,750 nmol
448 $\text{L}^{-1} \text{d}^{-1}$; Hampel et al. 2018). Rate maxima from western Lake Erie were about a third of the highest values
449 reported in Taihu (Hampel et al., 2018). Higher rates at depth were also observed in mesotrophic Lake
450 Lacawac (Heiss et al., 2022) and are commonly observed in the ocean (Ward et al., 2008) due to
451 inhibition of nitrification at high light intensities found in surface waters (Guerrero and Jones, 1996;
452 Merbt et al., 2012). One exception to these patterns is Lake Onondaga, where sediment nitrification was
453 observed, but apparently with no corresponding activity in the water column. The authors attribute this
454 lack of water column nitrification to low nitrifier densities, which contrasts with the present and many
455 other studies (Pauer and Auer, 2000).
456

457 **Western basin nitrification vs. community NH_4^+ demand and water column NH_4^+ regeneration**

458

459 In Hoffman et al. 2022, community potential NH_4^+ uptake rates were reported for the water
 460 column at the same western basin stations and time points as nitrification rates measured here.
 461 Nitrification rates were compared to potential NH_4^+ uptake to determine the proportion of this community
 462 NH_4^+ uptake attributable to nitrification. As expected, the percentage of NH_4^+ uptake explained by
 463 nitrification was greatest in late spring and early summer, but declined to, on average, less than 5% of
 464 total NH_4^+ demand during peak bloom months (Fig. 8). In 2015, the most intense cyanoHAB year, the
 465 percentage of NH_4^+ demand accounted for by nitrification ranged from 0 to 7.04% ($0.413 \pm 0.133\%$,
 466 mean \pm SE; $n = 71$). In 2016 and 2017, the maximum contribution of nitrification increased to 55.2%
 467 (3.61 ± 0.745 ; $n = 119$) and 59.6% (3.81 ± 1.10 ; $n = 143$), respectively (Fig. 8), early in the season (WE2
 468 in June 2016) and at stations furthest from the Maumee River mouth (WE13 in August 2017). The
 469 contribution of nitrification to community potential NH_4^+ uptake rates in western Lake Erie spans a much
 470 greater range than that reported in Lake Taihu, where values from 0.2 – 15% were observed (Hampel et
 471 al., 2018).
 472



473 Figure 8. The percentage of community potential NH_4^+ uptake explained by nitrification in each month
 474 within each sample year. n for each month in each year is as follows: 2015 (June = 18, July = 18,
 475 August = 18, September = 18); 2016 (April = 6, May = 12, June = 24, July = 18, August = 24,
 476 September = 12, October = 24); 2017 (May = 24, June = 24, July = 24, August = 24, September
 477 = 24, October = 24).
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479
 480 **Nitrification rates during ECCC cruises**

481 Nitrification rates did not vary by basin ($p > 0.15$), despite highest nutrient inputs in the western
 482 basin, perhaps because western basin nitrification rates were inhibited by summer cyanoHABs.
 483 Nitrification rates measured from these experiments ($118 \pm 31.0 \text{ nmol L}^{-1} \text{ d}^{-1}$) are within the ranges
 484 reported in the water columns of other eutrophic lakes, although rate maxima (up to $878 \text{ nmol L}^{-1} \text{ d}^{-1}$)
 485 were much lower than those from Lake Taihu ($3,750 \text{ nmol L}^{-1} \text{ d}^{-1}$; Hampel et al., 2018), Lake Okeechobee
 486 ($1,280 \text{ nmol L}^{-1} \text{ d}^{-1}$; Hampel et al., 2020), and Lake Mendota ($1,700 - 5,000 \text{ nmol L}^{-1} \text{ d}^{-1}$; Hall, 1986).
 487 Nitrification rates in Lake Erie also were similar to those in mesotrophic Lake Lacawac ($1 - 568 \text{ nmol L}^{-1}$

488 d⁻¹; Heiss et al., 2022) and saline Lake Mono (60 – 335 nmol L⁻¹ d⁻¹; Carini and Joye, 2008). However,
489 nitrification rates in Lake Erie, particularly those from the eastern basin, were much greater than those in
490 oligotrophic Lake Superior (0 – 51 nmol L⁻¹ d⁻¹; Small et al., 2013).

491

492 **Ammonia-oxidizer community abundance and competition dynamics in western Lake Erie**

493 We hypothesized that AOB would outnumber AOA due to differences in relative affinity for
494 NH₄⁺ and the AOB community's competitive advantage in high-NH₄⁺ environments. Ambient NH₄⁺
495 concentrations in western Lake Erie ranged from undetectable to 7.30 μM, within the established K_m for
496 AOB but thought to be greater than that of AOA (AOA isolates: 0.05–0.14 μM; Martens-Habbenha et al.,
497 2009). However, some AOA isolates have higher K_m values, which may explain some of the variability in
498 AOB:AOA (Jung et al. 2022). Higher abundances of AOO were expected during non-bloom periods, but
499 AOO gene copies were lowest in June, when some of the highest nitrification rates were measured. Both
500 AOB and AOA were positively correlated with turbidity, which may be related to incoming nutrient and
501 sediment loads from the Maumee River and/or sediment resuspension.

502 Comammox *amoA* gene abundances were not quantified in this study, but nitrification rates were
503 high at times in this study. All nitrifiers may play an important role in aquatic systems, and comammox
504 should be investigated further in future lake studies. Comammox is important in wastewater treatment
505 plants (Yang et al. 2020, Su et al. 2021) and lake sediments (Lu et al. 2020, Xu et al. 2020), but the few
506 comammox studies in lake water columns suggest that abundances are low (e.g., Harringer and Alfreider
507 2021).

508 Despite the presence of AOO at all sampling stations and times, nitrification rates were low
509 during summer months and constituted a small fraction of community potential NH₄⁺ uptake. There was
510 no correlation between measured nitrification rates and either bacterial or archaeal *amoA* gene copy
511 numbers. With a subset of the data (undetectable rates excluded), nitrification rates and AOA gene copies
512 were correlated ($\rho = 0.19$). A similar study in Lake Taihu found that nitrification rates were correlated
513 with AOB but not AOA (Hampel et al., 2018), while in Lake Okeechobee, nitrification rates were
514 strongly related to AOA abundance (Hampel et al., 2020). *amoA* abundance was quantified using DNA,
515 as opposed to gene transcription via RNA, which may explain discrepancies between biogeochemical and
516 molecular results. These discrepancies also support the hypothesis that, although present and presumably
517 active, the AOO community were not effective competitors for available NH₄⁺, especially during
518 cyanoHABs.

519 High abundances of *amoA* in western Lake Erie are not without precedent, as a previous study
520 found large numbers of AOB and AOA in Lake Erie sediments (Bollmann et al., 2014). Since western
521 Lake Erie is well-mixed and often turbid, particle-attached AOO communities may be introduced into the
522 water column via sediment loading from the river and/or sediment resuspension. Within the water
523 column, *Proteobacteria* (which includes AOB) can account for nearly 20% of the metagenome during
524 cyanoHABs in Lake Erie (and ~70% in hypereutrophic Lake Taihu; Steffen et al., 2012), and these
525 estimates do not account for co-occurring archaea. However, high gene abundances observed in our Lake
526 Erie study must also be considered relative to literature on copies of *amoA* per cell. *Betaproteobacterial*
527 AOB, particularly within the genus *Nitrosospira*, can have up to three copies of the *amoA* gene compared
528 to one in AOA (Norton et al., 2008; Blainey et al., 2011; Lagostina et al., 2015). Accordingly, *amoA*
529 abundances in Lake Erie do not equate to cell number. Specific community members were not identified
530 in this study, and attempts to correct AOB:AOA ratios with variations in copy number within each
531 genome do not always explain high AOB:AOA ratios (Lagostina et al., 2015). Thus, gene copy numbers
532 were not corrected based on potential genomic differences among AOB and AOA.

533 The only other known literature on water column AOO community abundances in Lake Erie
534 comes from a central basin location, where AOB outnumbered AOA (Mukherjee et al., 2015). However,
535 the authors were unable to detect any AOO in surface waters, which contrasts with the high surface water
536 abundances reported here. These discrepancies between studies may be partly attributable to different
537 methodologies for enumeration (CARD-FISH versus qPCR, where qPCR is more likely to detect inactive
538 cells and represents more than one gene copy per cell) or spatio-temporal variation in communities

539 (Murkherjee et al. sampled only in the central basin in July 2011 vs. full-season sampling in this study in
540 the western basin only).

541 Approximate parity between AOB and AOA abundances in the western Lake Erie water column
542 aligns with findings from work conducted in Lake Erie sediments but contrasts with results from other
543 eutrophic lakes. At a station near WE6, AOB and AOA numbered from 10^6 – 10^7 copies per g sediment,
544 and the two community abundances were not different from each other (Bollmann et al., 2014). In Lake
545 Taihu, AOA outnumbered AOB in the water column, with AOA and AOB copies ml^{-1} reaching as high as
546 10^7 and 10^5 , respectively (Hampel et al., 2018). In a Taihu mesocosm experiment, however, AOB
547 outnumbered AOA by up to two orders of magnitude (Chen et al., 2016). AOB gene copies in western
548 Lake Erie were often much greater than those in Taihu by an order of magnitude, but AOA gene copies
549 were similar between the two lakes. Similarly, AOA abundances in Lake Okeechobee exceeded those of
550 AOB (AOA up to 10^6 , AOB up to 10^5 ; Hampel et al., 2020), with maximum abundances in western Lake
551 Erie greater for both AOA and AOB by an order of magnitude or more. In general, maximum AOB and
552 AOA abundances were within the ranges of those from other lakes, including eutrophic Lake Dianchi
553 (China, maximum 10^5 for both AOA and AOB *amoA* abundances; Yang et al. 2016), mesotrophic Lake
554 Lacawac (USA, up to 10^6 copies ml^{-1} for both AOA and AOB; Heiss et al. 2022), and oligotrophic Lake
555 Constance (Europe, 10^5 AOA copies ml^{-1} ; Klotz et al. 2022).

556 While early studies of AOA indicated that they were important mostly in oligotrophic
557 environments (Beman et al., 2008; Pester et al., 2012), more recent work has demonstrated that some
558 AOA ecotypes thrive in eutrophic systems. AOA are abundant alongside AOB in wastewater treatment
559 plant bioreactors (e.g., Park et al., 2006; Limpiyakorn et al., 2013; Zhang et al., 2015), and new ecotypes
560 have been reported in eutrophic lakes, such as Lake Okeechobee (Hampel et al., 2020). Given evolving
561 knowledge and high abundances of AOA in Lake Erie and other eutrophic lakes, AOA may play a larger
562 role in these systems than previously thought (Zeng et al., 2012; Damashek et al., 2015; Hampel et al.,
563 2018, 2020).

564 Despite some discrepancies in AOO community dynamics between this study and others
565 conducted in eutrophic lakes, seasonal patterns of nitrification rates suggest that the presence of
566 cyanoHABs suppresses nitrification in WLE. This observation matches with findings reported in lakes
567 Taihu (Hampel et al., 2018) and Okeechobee (Hampel et al., 2019), both of which are also afflicted with
568 *Microcystis*-dominated cyanoHABs and may be explained by the strong affinity of *Microcystis* for NH_4^+ .
569 Studies in culture and in lakes have reported a broad range of K_m values for *Microcystis*, with maximum
570 values reaching 37 μM in culture (Nicklisch and Kohl, 1983) and 113 μM in lakes (Yang et al., 2017). In
571 Maumee Bay, community K_m values changed along a seasonal gradient, beginning at very low K_m (0.32
572 μM) in early summer, increasing to 3.53 μM in August and 8.52 μM in October (Hampel et al., 2019).
573 The authors suggested that this pattern illustrates community dominance of *Microcystis* during
574 cyanoHAB development and peak bloom biomass, which reflects the strong competitive abilities of
575 *Microcystis* for available NH_4^+ . Based on previously discussed literature values, coupled with seasonally
576 decreasing nitrification rates and NH_4^+ concentrations during the cyanoHAB season, it appears that
577 nitrifiers in western Lake Erie were unable to compete effectively with cyanoHAB communities for NH_4^+ .

578 579 **SYNTHESIS**

580 Despite high water column NH_4^+ demand in peak bloom months (Hoffman et al. 2022),
581 nitrification comprised a small proportion of this demand. Even though bioavailable N concentrations
582 were depleted during cyanoHABs, regeneration of the NH_4^+ pool is still occurring at rates exceeding
583 nitrification. However, where NH_4^+ regeneration is greatest (at WE6, Hoffman et al. 2022), nitrification
584 was also near-undetectable, further supporting the conclusion that nitrifiers cannot scavenge NH_4^+ as
585 efficiently as cyanoHAB microorganisms. The lack of relationship between *amoA* gene copies and
586 nitrification rates indicates that gene abundance is not a reliable indicator of activity in western Lake Erie,
587 and future studies should consider other methodology, such as RNA or metatranscriptomic analysis when
588 comparing rate measurements to genetic data. The ability of water column nitrification to act as a buffer
589 against external N loading by supplying the substrate for denitrification is impeded by cyanoHABs in

590 western Lake Erie and similar systems; therefore, management of external N loads is crucial for
591 mitigating the harmful effects of eutrophication. These findings support previous work demonstrating
592 suppression of nitrification in oxic waters via competition for NH_4^+ during cyanoHABs and suggest that
593 non-cyanoHAB biomass (e.g., other phytoplankton, such as diatoms and cryptophytes) in the central and
594 eastern basins may also outcompete nitrifiers for available NH_4^+ . Ecosystem models often neglect within-
595 system transformation pathways, such as nitrification, and as such may lack important information needed
596 to accurately make management decisions. Accordingly, results from this and other studies should be
597 considered when creating and validating models to help inform management decisions.

598
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