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#### **Key Points:**

- A nitrogen-to-phosphorus ratio budget method is used to quantify nutrient sources and sinks at two subtropical ocean study sites
- Vertical phosphate supply is the dominant source of phosphorus to the surface of the North Pacific and the North Atlantic study site
- Dissolved organic phosphorus transport and zooplankton excretion are more important than sinking particles as nutrient sinks

#### **[Supporting Information:](https://doi.org/10.1029/2023GL103213)**

[Supporting Information may be found in](https://doi.org/10.1029/2023GL103213)  [the online version of this article.](https://doi.org/10.1029/2023GL103213)

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# **Subtropical Gyre Nutrient Cycling in the Upper Ocean: Insights From a Nutrient-Ratio Budget Method**

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**Abstract** We use a nutrient-ratio budget method to investigate the relative importance of different nutrient source and sink terms at time-series Station ALOHA and Bermuda Atlantic Time-series Study (BATS) in the North Pacific and North Atlantic subtropical gyres, respectively. At mean state conditions over annual and multi-year time scales, vertical phosphate ( $PO_4^{3-}$ ) supply from the subsurface accounts for ~60% of the total phosphorus supply at both sites. Dissolved organic matter transport and zooplankton excretion are more important phosphorous export pathways than sinking particles at Station ALOHA and BATS. The nutrient-ratio budget approach provides quantitative, observation-based constraints on nutrient sources and sinks in the surface ocean, which helps improve our understanding of the biological carbon pump in oligotrophic oceans.

**Plain Language Summary** In this study, we explore the cycling of nutrients that support primary production in the surface ocean and its subsequent export to depth using observed elemental ratios of nitrogen to phosphorus for various nutrient sources and sinks. We use nutrient observations from long-term oceanographic time-series studies at Station ALOHA near Hawaii and the Bermuda Atlantic Time-series Study near Bermuda. We assume that both stations are under conditions of steady state in which nutrient concentrations are not changing over long time periods, and therefore, that the nitrogen-to-phosphorus ratio between inputs and outputs should be balanced. We apply a mathematical model to estimate the relative contribution of each input and output term. Our results suggest that nutrient input is driven primarily by the vertical transport of subsurface water at both study sites. Nutrient output (loss) is driven by the gravitational sinking of large particles, the downward mixing of dissolved constituents, and the active transport of migrant animals. The loss due to the latter two processes is more important in magnitude. Our simple methodology provides quantitative, observational constraints of nutrient sources and sinks to the upper ocean, contributing improved understanding of the biological carbon pump in the oligotrophic subtropical ocean.

# **1. Introduction**

The transfer of organic matter produced by phytoplankton in the surface ocean to the ocean interior exerts a strong control on the surface ocean carbon dioxide gas  $(CO<sub>2</sub>)$  concentration, which in turn has a significant impact on the atmospheric CO<sub>2</sub> level (Kwon et al., [2009;](#page-8-0) Volk & Hoffert, [1985](#page-9-0)). Despite the importance of this vertical transfer process, termed the ocean's biological carbon pump, regional variations in its strength and sensitivity to environmental changes are poorly characterized (Quay et al., [2020\)](#page-8-1). Our understanding of organic matter elemental composition has long been based on the Redfield ratio (Redfield, [1958](#page-8-2)), a constant carbon/nitrogen/phosphorus (C/N/P) ratio of 106/16/1, which is commonly used to link marine nutrient and carbon fluxes. Over the past 20 years, however, many laboratory (e.g., Geider & La Roche, [2002](#page-7-0)), field (Karl et al., [2001](#page-8-3); Lomas et al., [2022](#page-8-4); Martiny, Pham, et al., [2013;](#page-8-5) Martiny, Vrugt, et al., [2013](#page-8-6)), and modeling studies (Deutsch & Weber, [2012;](#page-7-1) Teng et al., [2014](#page-9-1); Wang et al., [2019;](#page-9-2) Weber & Deutsch, [2010](#page-9-3)) have indicated that the C/N/P stoichiometry of organic matter is significantly more variable both in time and space.

In the subtropical gyres, sinking particles collected by sediment traps near the surface exhibit N/P ratios generally >25 at time-series Station ALOHA and BATS (Karl et al., [2001;](#page-8-3) Lomas et al., [2022](#page-8-4); Singh et al., [2015](#page-9-4)), located in the North Pacific Subtropical Gyre (NPSTG) and the North Atlantic Subtropical Gyre (NASTG), respectively. Nitrogen fixation is required to support the export flux and high N/P of sinking particles at Station ALOHA (Karl et al., [1997;](#page-8-7) Quay, [2021](#page-8-8)). In addition to nitrogen fixation, lateral supply of both inorganic (nitrate + nitrite  $(NO_3^-)$ ) and phosphate  $(PO<sub>4</sub><sup>2</sup>)$ ) and organic (dissolved organic nitrogen (DON) and dissolved organic phosphorus (DOP))

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<span id="page-1-0"></span>**Figure 1.** Location of Station ALOHA and BATS (cyan pentagrams) and cruise tracks of dissolved organic nutrients measurements superimposed on surface nitrate concentrations from the WOA18 annual climatology (Garcia et al., [2018](#page-7-5)). Coordinates in the North Pacific are from Abell et al. ([2000\)](#page-7-2) (diamonds) and COOK-BOOK cruises (Church et al., [2008\)](#page-7-7) (squares). North Atlantic data are from Cavender-Bares et al. [\(2001](#page-7-6)) (CB; circles) and Torres-Valdés et al. [\(2009](#page-9-6)) (36N; triangles). Regions of WOA18 nutrient data used to calculate surface lateral supply ratios are highlighted with four green boxes that are  $30^{\circ} \times 5^{\circ}$  and twenty 1-degree latitude or longitude bands.

nutrients has been proposed as another important supply mechanism in the oligotrophic subtropical gyres (e.g., Abell et al., [2000](#page-7-2); Letscher et al., [2016;](#page-8-9) Lomas et al., [2010;](#page-8-10) Mahaffey et al., [2004](#page-8-11); Reynolds et al., [2014\)](#page-8-12). The relative importance of each nutrient supply mechanism and the associated impact on exported matter N/P has been difficult to quantify observationally, however.

A budget method, based on ratios of inorganic nutrient  $(PO_4^{3-}$ , NO<sub>3</sub> and dissolved inorganic carbon) supply, has been used to estimate the C/N/P ratio of organic matter exported from the surface ocean in the Pacific Ocean (Quay, [2021](#page-8-8)). We expand on this work by including the input and output of dissolved organic nutrients (e.g., Abell et al., [2000;](#page-7-2) Hopkinson & Vallino, [2005\)](#page-8-13) and the export of dissolved nutrients (inorganic and organic) through zooplankton excretion (Hannides et al., [2009](#page-7-3); Steinberg et al., [2002](#page-9-5)). The steady state framework of our budget method dictates that the N/P ratio of all dissolved nutrient sources to the mixed layer  $(\Delta NO_4/\Delta PO_4)$ , ΔDON/ΔDOP), supplied either vertically or laterally, should equal the N/P ratio of all nutrient sinks (sinking particles, dissolved organic matter (DOM) transport, zooplankton excretion) from the mixed layer. In the subtropical gyres, nitrogen fixation is an important source of nitrogen (Böttjer et al., [2017](#page-7-4); Karl et al., [1997](#page-8-7)) and needs to be incorporated into the budget method. The advantage of a "ratio" budget is that it does not require nutrient flux estimates, which typically carry large uncertainties. The nutrient-ratio budget method provides quantitative constraint on the relative importance of nutrient sources and sinks when the source or sink terms are characterized by a distinct N/P ratio.

## **2. Materials and Methods**

#### **2.1. Nutrient Data**

Nutrient observations (dissolved, particulate, and sediment trap) used in this study come from time-series Station ALOHA (22.8°N, 158.0°W) and the Bermuda Atlantic Time-series Study (BATS; 31.6°N, 64.2°W) (Figure [1](#page-1-0)) (Karl et al., [2021;](#page-8-14) Lomas et al., [2013\)](#page-8-15). The annual mean N/P ratio of sinking particles (excluding N/P  $\geq$  100) from the shallowest sediment trap (150 m) at Station ALOHA and BATS is calculated as an integrated N/P ratio over the course of each year, weighted by the particulate P flux. The multi-year mean N/P ratio of sinking particles (Sinking N/P) is computed as the mean of all annual mean N/P ratios, weighted by the number of sediment trap flux observations in each year. Likewise, the multi-year mean surface N/P ratio of suspended particles (Suspended N/P) is calculated as the mean of all annual mean N/P ratios (0–150 m), weighted by the number of suspended particle observations in each year. Additional data sources include  $NO_3^-$  and  $PO_4^{3-}$  data from World Ocean Atlas (WOA18) monthly climatology (Garcia et al., [2018](#page-7-5)), and DON and DOP observations from cruises near Station ALOHA and BATS (Abell et al., [2000;](#page-7-2) Cavender-Bares et al., [2001;](#page-7-6) Church et al., [2008;](#page-7-7) Torres-Valdés et al., [2009](#page-9-6)) (Figure [1\)](#page-1-0).





<span id="page-2-0"></span>**Figure 2.** Surface ocean N and P budgets. Nutrient supply (sup) and external N source (src) terms: vertical and lateral supply of dissolved inorganic and organic N and P; nitrogen fixation and atmospheric (atm.) deposition. Nutrient export (exp) terms: sinking (Sinking) and suspended (Suspended) particles; exported dissolved organic matter (DOM); and zooplankton (Zoo.) excretion.

#### **2.2. Nutrient-Ratio Budget Method**

The relative importance of external N sources ( $N_{\rm src}$ ) to the N export flux ( $N_{\rm exp}$ ) can be estimated by combining the N and P budgets for the surface ocean (Quay, [2021](#page-8-8)) (Figure [2](#page-2-0)):

$$
\frac{N_{src}}{N_{exp}} = 1 - \left(\frac{N}{P}\right)_{sup} / \left(\frac{N}{P}\right)_{exp}
$$
\n(1)

<span id="page-2-1"></span>where  $N_{src}/N_{exp}$  is the fraction of exported organic nitrogen supported by external N sources to the surface ocean, (N/P)<sub>sup</sub> is the dissolved N and P nutrient supply ratio to the surface ocean, and (N/P)<sub>exp</sub> is the N/P ratio of exported organic matter from the surface ocean.

The dissolved nutrient supply term is further decomposed into vertical (v) and lateral (l) components. As in Quay ([2021\)](#page-8-8), the vertical nutrient supply (flux) ratio from the subsurface,  $(\Delta NO_{3}/\Delta PO_{4})_{\nu}$ , is estimated from the slope of  $NO_3^-$  to  $PO_4^{3-}$  data between the mean mixed layer depth (MLD) and 300 m for each year of observations at Station ALOHA and BATS. Robust regression is used to minimize the influence of outliers (*robustfit* in MATLAB). The vertical nutrient supply ratio incorporates the influence of preferential P remineralization in the subsurface. The multi-year mean  $(\Delta NO_3/\Delta PO_4)$ , between 1988 and 2020 is averaged from the annual vertical inorganic supply ratios (Figure S1 in Supporting Information S1). The sensitivity of  $(\Delta NO_{3}/\Delta PO_{4})_{v}$  to the depth interval used for the regression is discussed in Text S1 in Supporting Information S1. The steady-state assumption is supported by the lack of significant changes in the N/P of annual vertical supply and sinking particles with time using all data available in the past 30 years at Station ALOHA and BATS (*p* > 0.05; Figure S1 in Supporting Information S1). The mean lateral nutrient supply ratio,  $(\Delta NO_3/\Delta PO_4)_1$ , is calculated from the slope of  $NO_3^$ to  $PQ_4^{3-}$  within the annual mean MLD using WOA18 monthly climatology data (Garcia et al., [2018\)](#page-7-5) across specified meridional (30° latitude  $\times$  5° longitude) and zonal (5° latitude  $\times$  30° longitude) transects centered at Station ALOHA and BATS (Figure S2 in Supporting Information S1). Each transect is divided into five 1-degree longitude or latitude bands to estimate the variability of lateral supply ratios meridionally or zonally, respectively.

Concurrent DON and DOP measurements near the time-series stations are limited (Hansell et al., [2021](#page-7-8); Knapp et al.,  $2021$ ). In the North Atlantic, we use observations from individual cruises ( $N = 5$ ) to calculate the slope of the DON to DOP within the annual mean MLD across meridional (26–41°N and 62–71°W; Cavender-Bares et al., 2001) and zonal (35–38°N and 8–75°W; Torres-Valdés et al., [2009\)](#page-9-6) transects to estimate lateral supply ratios, (ΔDON/ΔDOP)<sub>1</sub>, at BATS (Figure [1](#page-1-0)). In the North Pacific, we use DON and DOP observations from three cruises between 10 and 45°N along 152–158°W (Abell et al., [2000](#page-7-2)), and between 24–55°N along 152–157°W (Church et al., [2008](#page-7-7)) to estimate the meridional lateral organic nutrient supply ratio at Station ALOHA (Figure [1](#page-1-0)). We were unable to find cruise data appropriate for determining the zonal  $(\Delta$ DON/ $\Delta$ DOP)<sub>1</sub> ratio for Station ALOHA. The vertical supply of semilabile DON and DOP from the subsurface is minimal (Lomas et al., [2010;](#page-8-10) Mahaffey et al., [2004\)](#page-8-11) and not considered in our budgets.



<span id="page-3-0"></span>**Figure 3.** All N/P supply (N/P)<sub>sup</sub> and export (N/P)<sub>exp</sub> terms in the surface ocean at Station ALOHA and BATS. Error bars represent ±1 standard deviations. "Supply" and "Export" boxes represent all nutrient supply and export terms, respectively.

# **3. Results and Discussion**

#### **3.1. N/P Ratios at Station ALOHA and BATS**

#### **3.1.1. N/P of Nutrient Supply**

The mean vertical inorganic nutrient supply ratios ( $\Delta NO_3/\Delta PO_4$ ), at Station ALOHA and BATS are 14.7  $\pm$  0.6 and 19.6  $\pm$  5.2, respectively (Figure [3](#page-3-0) and Figure S1 in Supporting Information S1). The lateral inorganic nutrient supply ratios ( $\Delta NO_3/\Delta PO_4$ )<sub>l</sub> are much smaller than the vertical ratios (Figure S2 in Supporting Information S1). Near Station ALOHA, the lateral surface transport comes mainly from the equatorial region to the north (Abell et al., [2000;](#page-7-2) Quay & Stutsman, [2003](#page-8-17)). A meridional  $(\Delta NO_3/\Delta PO_4)_1$  of  $1.6 \pm 0.6$  is estimated using WOA18 nutrient data between 7–22.5°N and 155.5–160.5°W (Figure S3 in Supporting Information S1). Near BATS, the largest lateral surface transport comes from the north (Palter et al., [2011](#page-8-18); Siegel & Deuser, [1997](#page-9-7); Williams & Follows, [1998](#page-9-8)). A meridional  $(\Delta NO_3/\Delta PO_4)$  of 5.3  $\pm$  0.5 is estimated between 15–45°N and 62–67°W (Figure S2 in Supporting Information S1). We hereafter refer to the lateral supply ratio as that estimated from meridional  $( \Delta NO_3/ \Delta PO_4)$ <sub>l</sub> terms at both locations. These estimates are likely an upper limit of  $( \Delta NO_3/ \Delta PO_4)$ <sub>l</sub> since lateral gradients into the subtropics are generally largest near the coastal margins (McGillicuddy Jr. et al., [2003](#page-8-19); Williams & Follows, [1998](#page-9-8); Williams et al., [2011](#page-9-9)) and  $PO_4^{3-}$  concentrations in the subtropical gyre compiled by the World Ocean Atlas are higher compared to observations derived from high-sensitivity methods (Martiny et al., [2019\)](#page-8-20).

Semilabile DON and DOP transported from coastal regions into the interior ocean has been identified as a potential nutrient source for the oligotrophic gyres (e.g., Abell et al., [2000](#page-7-2); Letscher et al., [2016](#page-8-9); Lomas et al., [2010;](#page-8-10) Mahaffey et al., [2004;](#page-8-11) Reynolds et al., [2014](#page-8-12)). Abell et al. [\(2000](#page-7-2)) estimated the net supply flux of DON and DOP into the NPSTG to be 20  $\pm$  2 mmol N and 4.4  $\pm$  1.0 mmol P m<sup>-2</sup> yr<sup>-1</sup>, respectively, which yields a meridional (ΔDON/ΔDOP)<sub>1</sub> supply ratio of 4.5  $\pm$  2.4 from the Equatorial region to Station ALOHA. The meridional ratio is 3.4 ± 1.6 estimated using DON and DOP data from the COOK-BOOK cruises (Church et al., [2008](#page-7-7)). We therefore use the average meridional ratio of  $4.0 \pm 2.9$  for Station ALOHA. Near BATS, the lateral ( $\Delta$ DON/ $\Delta$ DOP)<sub>1</sub> supply ratio is  $6.7 \pm 4.4$  meridionally (Cavender-Bares et al., [2001\)](#page-7-6) and  $6.9 \pm 1.9$  zonally (Torres-Valdés et al., [2009](#page-9-6)) (Figure S4 in Supporting Information S1).

#### **3.1.2. N/P of Nutrient Export**

Export of organic matter is categorized into three primary pathways, that is, the gravitational pump, mixing pump, and migrant pump (Siegel et al., [2023](#page-9-10)). The N/P representing the gravitational pump is determined from sinking particles collected in the 150-m sediment trap (Sinking N/P) which yield mean  $\pm$  s.d. of 28.5  $\pm$  7.8 and 40.5 ± 15.7 at Station ALOHA and BATS, respectively. Sinking N/P at both time-series stations are much higher than the Redfield ratio as well as the vertical and lateral supply ratios (Figure [3](#page-3-0) and Tables S1 and S2 in Supporting Information S1).

The "mixing pump" represents the downward transport of both suspended particulate and DOM by various physical mechanisms (Dall'Olmo et al., [2016;](#page-7-9) Hansell et al., [2009](#page-7-10); Levy et al., [2013;](#page-8-21) Omand et al., [2015](#page-8-22); Roshan & DeVries, [2017;](#page-8-23) Resplandy et al., [2019\)](#page-8-24). The multi-year weighted mean N/P ratios of suspended particles in the upper 150 m (Suspended N/P) are  $23.1 \pm 5.4$  at Station ALOHA and  $37.2 \pm 5.8$  at BATS, consistent with previous estimates (Karl et al., [2001;](#page-8-3) Singh et al., [2015](#page-9-4)) and higher than the Redfield ratio. The lifetime of refractory DON and DOP (>4,000 years; Letscher et al., [2015](#page-8-25)) far exceeds the ventilation timescale of waters at 300 m (≥4–10 years; Bullister et al., [2006](#page-7-11); Stanley et al., [2012\)](#page-9-11), so we only consider the export of semilabile DON and DOP. The N/P ratio used to represent transport of semilabile DON and DOP export at both Station ALOHA and BATS (20  $\pm$  3; ( $\Delta$ DON/ $\Delta$ DOP)<sub>exp</sub>) comes from the mean of incubation experiments at Station ALOHA, Georges Bank, and the Mid-Atlantic Bight (Hopkinson & Vallino, [2005\)](#page-8-13). This value aligns with the global mean N/P ratio of ∼20 for semilabile DON and DOP export in a modeling study (Letscher & Moore, [2015](#page-8-26)). Depth gradients of suspended particle concentrations are on average ∼10%–20% of the depth DOM concentration gradients at Station ALOHA and BATS (Figure S5 in Supporting Information S1) so we ignore the transport of suspended particles in this study.

The "migrant pump" represents dissolved (in)organic nutrients actively transported by migrating zooplankton. At Station ALOHA, Hannides et al. [\(2009](#page-7-3)) found that migrant excretions of dissolved nutrients (organic + inorganic) are characterized by a mean N/P ratio of 12. Steinberg et al. [\(2002](#page-9-5)) measured the dissolved inorganic N/P excretion ratios averaging  $12.0 \pm 4.9$  for a few zooplankton species at BATS. The N/P ratios associated with migrant excretion (Zoo. N/P) in the NPSTG and NASTG are very similar, and are P-rich compared to sinking N/P.

#### **3.2. Nutrient-Ratio Budgets at Station ALOHA and BATS**

#### **3.2.1. Role of Nitrogen Fixation in the Subtropical Nutrient Budget**

At both Station ALOHA and BATS, the significantly greater N/P of exported organic matter than N/P of dissolved nutrient supply is due, in part, to external N supplied from nitrogen fixation. Given small atmospheric nitrogen deposition rates at Station ALOHA (Duce et al., [2008;](#page-7-12) Kim et al., [2014](#page-8-27)), the external N source must primarily be nitrogen fixation. Nitrogen fixation was previously estimated to supply about 30%–50% of exported organic N based on δ15N measurements at Station ALOHA (Böttjer et al., [2017;](#page-7-4) Karl et al., [1997\)](#page-8-7). Using a simplified version of the nutrient-ratio budget method discussed here, Quay ([2021\)](#page-8-8) estimated that nitrogen fixation could support 33%–43% of the exported organic N (i.e.,  $N_{\rm src}/N_{\rm exp}$ ).

At BATS, the difference in N/P between the dissolved nutrient supply and exported particles is even greater than at Station ALOHA (Figure [3](#page-3-0) and Tables S1 and S2 in Supporting Information S1). Unlike at Station ALOHA, however, nitrogen fixation at BATS likely accounts for  $\langle 10\%$  of export flux based on  $\delta^{15}N$  measurements (Altabet, [1988;](#page-7-13) Knapp et al., [2005\)](#page-8-28), N budgets (Helmke et al., [2010](#page-8-29); Jenkins & Doney, [2003](#page-8-30); Stanley et al., [2015;](#page-9-12) Tang et al., [2019](#page-9-13)), and models (Wang et al., [2019\)](#page-9-2). Rate estimates of atmospheric nitrogen deposition are approx-imately half the nitrogen fixation at BATS (Knapp et al., [2010;](#page-8-31) Singh et al., [2013;](#page-9-14) Zamora et al., [2010](#page-9-15)). Thus  $N_{\rm cr}/N_{\rm crn}$  is likely <15% at BATS, which is significantly lower than the >50% fraction needed to balance the N/P budget. Therefore, additional nutrient sources with higher N/P, and/or nutrient sinks with lower N/P ratios are needed to balance the N and P budgets at BATS.

#### **3.2.2. Using N/P to Estimate the Relative Importance of Nutrient Sources and Sinks**

We use the observed N/P of the nutrient sources and sinks coupled with observed constraints on  $N_{\rm src}/N_{\rm exp}$  at Station ALOHA and BATS to quantitatively estimate the relative importance of specific nutrient sources and sinks for the surface ocean. Using a Monte Carlo method, we randomly assign an N/P value for each of the three dissolved nutrient source terms  $((\Delta NO_3/\Delta PO_4)_v$ , meridional  $(\Delta NO_3/\Delta PO_4)_v$ , and meridional  $(\Delta DON/\Delta DOP)_v$ 





<span id="page-5-0"></span>**Figure 4.** Histograms of the fraction of various nutrient sources (a-c) and sinks (d-f) at Station ALOHA and BATS, determined from the Monte Carlo simulations. The *y* axis is the probability of occurrence which sums to 1 for all fractions. The *x* axis represents the fraction of various sources and sinks. Panels illustrate the fractions of (a) vertical inorganic nutrient supply, (b) lateral inorganic nutrient supply, (c) lateral organic nutrient supply, and nutrient export through (d) sinking particles, (e) semilabile DOM export, and (f) zooplankton excretion. Vertical solid and dashed lines show the mean fractional contributions for Station ALOHA and BATS, respectively.

and the three nutrient sink terms (Sinking N/P, transported DOM N/P, and Zoo. N/P) based on their observed means and standard deviations assuming normal distributions (Tables S1 and S2 in Supporting Information S1). Input values for the fractional contributions of each nutrient source (supply) and sink (export) term are assumed to be independent and are randomly selected, but are forced to sum to 1 in each case. The N/P of total nutrient supply (N/P)<sub>sup</sub> and export (N/P)<sub>exp</sub> is determined and used to calculate  $N_{src}/N_{exp}$  based on Equation [1.](#page-2-1) The process is repeated 100,000 times. The subset of these scenarios that yield  $N_{\rm src}/N_{\rm exp}$  values within the observed ranges (0.3–0.5 at Station ALOHA and 0–0.15 at BATS, as discussed above) are used for interpretation. The mean, median, and interquartile range of each subset is used to determine the likely fraction of each nutrient source and sink term (Tables S1 and S2 in Supporting Information S1). Fractions of all nutrient terms are not normally distributed (Lilliefors test;  $p \ll 0.05$ ) and are better described by medians and interquartile ranges than means and standard deviations (Sainani, [2012](#page-8-32)). However, medians of the total nutrient source or sink do not sum to 100% whereas the means do. Therefore, we present both means and medians in the following discussion.

At Station ALOHA, histograms of the fractional contributions of lateral inorganic and organic nutrient supply are similar and both positively skewed (Figures [4b](#page-5-0) and [4c\)](#page-5-0), whereas the vertical inorganic nutrient supply histogram is negatively skewed (Figure [4a](#page-5-0)). The mean (median; interquartile range) of  $(\Delta NO_3/\Delta PO_4)$ <sub>v</sub>,  $(\Delta NO_3/\Delta PO_4)$ <sub>l</sub>, and  $(\Delta$ DON/ $\Delta$ DOP)<sub>1</sub> fractions at Station ALOHA are 0.58 (0.60; 0.45–0.72), 0.17 (0.14; 0.07–0.25), and 0.25 (0.20; 0.10–0.36), respectively (Table S1 in Supporting Information S1). The fractional contribution to  $(N/P)_{\text{sun}}$ from vertical inorganic supply is significantly greater than the total lateral nutrient supply (Mann-Whitey

U test;  $p \ll 0.05$ ). For (N/P)<sub>exp</sub>, all three fractions are positively skewed (Figures [4d–4f\)](#page-5-0). The mean (median; interquartile range) of sinking particles, semilabile DOM export, and zooplankton excretion fractions are 0.25 (0.21; 0.09–0.37), 0.34 (0.31; 0.14–0.51), and 0.40 (0.39; 0.20–0.59), respectively (Table S1 in Supporting Information S1).

At BATS, similar nutrient supply patterns emerge with the fractions of lateral inorganic and organic nutrient supply being positively skewed, and the vertical inorganic nutrient supply being slightly negatively skewed (Figures [4a–4c](#page-5-0)). The mean (median; interquartile range) of  $(\Delta NO_3/\Delta PO_4)_{\nu}$ ,  $(\Delta NO_3/\Delta PO_4)_{\nu}$ , and  $(\Delta DON/\Delta DOP)_{\nu}$ fractions are 0.55 (0.57; 0.40–0.72), 0.21 (0.17; 0.07–0.31), and 0.25 (0.19; 0.09–0.35), respectively (Table S2 in Supporting Information S1). The fractional contribution to  $(N/P)_{\text{sun}}$  from the vertical inorganic nutrient supply at BATS is also significantly higher than the total lateral nutrient fraction ( $p \ll 0.05$ ). The distribution of export fractions due to zooplankton excretion is almost symmetrical at BATS (skewness  $= 0.04$ ) (Figure [4f](#page-5-0)). The mean (median; interquartile range) of sinking particles, semilabile DOM export, and zooplankton excretion fractions are 0.21 (0.15; 0.07–0.29), 0.34 (0.30; 0.13–0.52), and 0.45 (0.46; 0.24–0.65), respectively (Table S2 in Supporting Information S1).

The vertical inorganic nutrient supply term that dominates the total supply in our Monte Carlo analysis represents all processes contributing to the vertical nutrient flux, such as turbulent diffusion below the base of the surface layer or episodic eddy events as observed by Johnson et al. ([2010\)](#page-8-33) at Station ALOHA. The skewness of histograms of fractional contributions is caused by the different N/P ratios of the individual nutrient sources and sinks, and the imposed range of  $N_{\rm src}/N_{\rm exp}$  (Figures S6 and S7 in Supporting Information S1). The N/P of nutrient sinks (Station ALOHA: 12.0–28.5; BATS: 12.0–40.5) differs greatly from the N/P of lateral ( $\Delta NO_3/\Delta PO_4$ )<sub>l</sub> and  $(\Delta$ DON/ $\Delta$ DOP)<sub>1</sub> supply (1.6–6.7), while the N/P of vertical  $(\Delta NO_3/\Delta PO_4)$ <sub>v</sub> supply (14.7–19.6) falls within the range of sink term N/P ratios (Figure [3](#page-3-0) and Tables S1 and S2 in Supporting Information S1). The strong positive skewness for fractions of lateral inorganic and organic supply occurs because increasing the fraction of lateral supply decreases overall  $(N/P)_{\text{sup}}$  ratio which severely diminishes the number of scenarios that can yield a balance N/P budget constrained by the upper bound of  $N_{\rm src}/N_{\rm exp}$  constraints (Station ALOHA: 0.5; BATS: 0.15). Likewise, most of the scenarios with a higher fraction of sinking particles (with highest N/P of ∼30–40) are filtered out because they require an external N source that exceed the imposed upper bound of  $N_{\rm src}/N_{\rm ext}$ .

To identify which N/P terms contribute the most uncertainty to the mean fractions of different nutrient sources and sinks, we repeated our Monte Carlo simulation while excluding the standard deviation on each N/P source or sink term, one at a time. Based on this sensitivity analysis (Table S3 in Supporting Information S1), the N/P terms that yield the most uncertainty in the estimated mean fractional contributions are lateral organic nutrient supply, sinking particles, and zooplankton excretion at both Station ALOHA and BATS, which suggests that better constraining N/P ratios of these terms should be a focus of future research efforts.

Note that the fractions of different N/P terms represent the relative contributions of different P sources and sinks to the surface P budget. The vertical supply of inorganic  $PO_4^{3-}$  is the dominant dissolved P source and almost identical at Station ALOHA (mean: 58%) and BATS (55%). The combined export of semilabile DOP and zooplankton excretion far exceed (∼3 folds) the export by sinking particles at Station ALOHA (25 vs. 75%) and BATS (21 vs. 79%). The significant P fluxes from DOM export and zooplankton excretion may explain why oxygen and dissolved inorganic carbon budget based estimates of total organic matter export (Emerson, [2014](#page-7-14); Quay, [2021;](#page-8-8) Stanley et al., [2015\)](#page-9-12) consistently exceed sediment trap fluxes at 150 m at Station ALOHA and BATS (Church et al., [2013;](#page-7-15) Karl et al., [2021;](#page-8-14) Lomas et al., [2013](#page-8-15), [2022\)](#page-8-4).

# **4. Conclusions**

A nutrient budget method leveraging distinct elemental N/P ratios at Station ALOHA and BATS in the oligotrophic subtropical gyres was used to estimate the importance of different nutrient supply and export pathways in the surface layer at these two time-series stations. We found that vertically supplied  $PO_4^{3-}$  from below is the dominant nutrient source at both stations, accounting for ∼60% of the total phosphorus supply. At both stations, the combined export of organic matter via DOM transport and zooplankton migration and excretion is more important than export via sinking particles, which supports previously observed discrepancies between chemical tracer-budget estimates of export and sediment trap fluxes at these sites.



Elemental stoichiometry (e.g., N/P) is a powerful tool to study nutrient cycling in the ocean, because it avoids the need to estimate rates of nutrient supply and organic matter export which are difficult to make. To the best of our knowledge, these budgets provide the first observation-based constraints on the fractional contributions of multiple nutrient source and sink processes associated with organic matter export from the surface ocean. Contemporaneous measurements of the C/N/P elemental stoichiometry of particles, DOM, and zooplankton excretion across additional ocean biomes would improve characterization of nutrient cycling and the biological carbon pump at a global scale.

## **Data Availability Statement**

The nutrient data used for estimating N/P ratios of vertical nutrient supply, sinking and suspended particles in the study are available at the home pages of the Hawaii Ocean Time-series (HOT) Project and Bermuda Atlantic Time-series Study. Station ALOHA data were obtained via the Hawaii Ocean Time-series HOT-DOGS application (<https://hahana.soest.hawaii.edu/hot/hot-dogs/interface.html>) from the University of Hawai'i at Mānoa under National Science Foundation Award # 1756517. The BATS data were obtained from the Bermuda Atlantic Time-series Study (<http://bats.bios.edu/data/>) supported by the NSF Chemical and Biological Oceanography Programs. The WOA18 nutrient data used to estimate horizontal nutrient supply ratios were obtained from the National Centers for Environmental Information website [\(https://www.ncei.noaa.gov/access/world-ocean-at](https://www.ncei.noaa.gov/access/world-ocean-atlas-2018/)[las-2018/](https://www.ncei.noaa.gov/access/world-ocean-atlas-2018/)). This study also uses data from the cruise RRS Charles Darwin CD171, provided by the British Oceanographic Data Centre and funded by the Natural Environment Research Council. DOP and DON data from the COOK-BOOK cruises can be found at <https://hahana.soest.hawaii.edu/cookbook/cookbook.html>. The code for the Monte Carlo approach used in this study can be found at<https://doi.org/10.5281/zenodo.7568622>.

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#### **References**

- <span id="page-7-2"></span>Abell, J., Emerson, S., & Renaud, P. (2000). Distributions of TOP, TON and TOC in the North Pacific subtropical gyre: Implications for nutrient supply in the surface ocean and remineralization in the upper thermocline. *Journal of Marine Research*, *58*(2), 203–222. [https://doi.](https://doi.org/10.1357/002224000321511142) [org/10.1357/002224000321511142](https://doi.org/10.1357/002224000321511142)
- <span id="page-7-13"></span>Altabet, M. A. (1988). Variations in nitrogen isotopic composition between sinking and suspended particles: Implications for nitrogen cycling and particle transformation in the open ocean. *Deep-Sea Research, Part A: Oceanographic Research Papers*, *35*(4), 535–554. [https://doi.](https://doi.org/10.1016/0198-0149(88)90130-6) [org/10.1016/0198-0149\(88\)90130-6](https://doi.org/10.1016/0198-0149(88)90130-6)
- <span id="page-7-4"></span>Böttjer, D., Dore, J. E., Karl, D. M., Letelier, R. M., Mahaffey, C., Wilson, S. T., et al. (2017). Temporal variability of nitrogen fixation and particulate nitrogen export at Station ALOHA. *Limnology & Oceanography*, *62*(1), 200–216. <https://doi.org/10.1002/lno.10386>
- <span id="page-7-11"></span>Bullister, J. L., Wisegarver, D. P., & Sonnerup, R. E. (2006). Sulfur hexafluoride as a transient tracer in the North Pacific Ocean. *Geophysical Research Letters*, *33*(18), L18603.<https://doi.org/10.1029/2006GL026514>
- <span id="page-7-6"></span>Cavender-Bares, K. K., Karl, D. M., & Chisholm, S. W. (2001). Nutrient gradients in the western North Atlantic Ocean: Relationship to microbial community structure and comparison to patterns in the Pacific Ocean. *Deep Sea Research Part I: Oceanographic Research Papers*, *48*(11), 2373–2395. [https://doi.org/10.1016/S0967-0637\(01\)00027-9](https://doi.org/10.1016/S0967-0637(01)00027-9)
- <span id="page-7-7"></span>Church, M. J., Björkman, K. M., Karl, D. M., Saito, M. A., & Zehr, J. P. (2008). Regional distributions of nitrogen-fixing bacteria in the Pacific Ocean. *Limnology & Oceanography*, *53*(1), 63–77. <https://doi.org/10.4319/lo.2008.53.1.0063>
- <span id="page-7-15"></span>Church, M. J., Lomas, M. W., & Muller-Karger, F. (2013). Sea change: Charting the course for biogeochemical ocean time-series research in a new millennium. *Deep Sea Research Part II: Topical Studies in Oceanography*, *93*, 2–15.<https://doi.org/10.1016/j.dsr2.2013.01.035>
- <span id="page-7-9"></span>Dall'Olmo, G., Dingle, J., Polimene, L., Brewin, R. J. W., & Claustre, H. (2016). Substantial energy input to the mesopelagic ecosystem from the seasonal mixed-layer pump. *Nature Geoscience*, *9*(11), 820–823.<https://doi.org/10.1038/ngeo2818>
- <span id="page-7-1"></span>Deutsch, C., & Weber, T. (2012). Nutrient ratios as a tracer and driver of ocean biogeochemistry. *Annual Review of Marine Science*, *4*(1), 113–141. <https://doi.org/10.1146/annurev-marine-120709-142821>
- <span id="page-7-12"></span>Duce, R. A., LaRoche, J., Altieri, K., Arrigo, K. R., Baker, A. R., Capone, D. G., et al. (2008). Impacts of atmospheric anthropogenic nitrogen on the open ocean. *Science*, *320*(5878), 893–897.<https://doi.org/10.1126/science.1150369>
- <span id="page-7-14"></span>Emerson, S. (2014). Annual net community production and the biological carbon flux in the ocean. *Global Biogeochemical Cycles*, *28*(1), 14–28. <https://doi.org/10.1002/2013GB004680>
- <span id="page-7-5"></span>Garcia, H. E., Weathers, K., Paver, C. R., Smolyar, I., Boyer, T. P., Locarnini, R. A., et al. (2018). World Ocean Atlas 2018, Volume 4: Dissolved inorganic nutrients (phosphate, nitrate and nitrate+nitrite, silicate). NOAA Atlas NESDIS, *84*, 35. Retrieved from [https://www.ncei.noaa.gov/](https://www.ncei.noaa.gov/archive/accession/NCEI-WOA18) [archive/accession/NCEI-WOA18](https://www.ncei.noaa.gov/archive/accession/NCEI-WOA18)
- <span id="page-7-0"></span>Geider, R., & La Roche, J. (2002). Redfield revisited: Variability of C:N:P in marine microalgae and its biochemical basis. *European Journal of Phycology*, *37*(1), 1–17.<https://doi.org/10.1017/S0967026201003456>
- <span id="page-7-3"></span>Hannides, C. C. S., Landry, M. R., Benitez-Nelson, C. R., Styles, R. M., Montoya, J. P., & Karl, D. M. (2009). Export stoichiometry and migrant-mediated flux of phosphorus in the North Pacific Subtropical Gyre. *Deep Sea Research Part I: Oceanographic Research Papers*, *56*(1), 73–88. <https://doi.org/10.1016/j.dsr.2008.08.003>
- <span id="page-7-8"></span>Hansell, D. A., Carlson, C. A., Amon, R. M. W., Álvarez-Salgado, X. A., Yamashita, Y., Romera-Castillo, C., & Bif, M. B. (2021). Compilation of dissolved organic matter (DOM) data obtained from global ocean observations from 1994 to 2021 version 3 [Dataset]. National Centers for Environmental Information. <https://doi.org/10.25921/s4f4-ye35>
- <span id="page-7-10"></span>Hansell, D. A., Carlson, C. A., Repeta, D. J., & Schlitzer, R. (2009). Dissolved organic matter in the ocean: A controversy stimulates new insights. *Oceanography*, *22*(4), 202–211.<https://doi.org/10.5670/oceanog.2009.109>



- <span id="page-8-29"></span>Helmke, P., Neuer, S., Lomas, M. W., Conte, M. H., & Freudenthal, T. (2010). Cross-basin differences in particulate organic carbon export and flux attenuation in the subtropical North Atlantic gyre. *Deep Sea Research Part I: Oceanographic Research Papers*, *57*(2), 213–227. [https://](https://doi.org/10.1016/j.dsr.2009.11.001) [doi.org/10.1016/j.dsr.2009.11.001](https://doi.org/10.1016/j.dsr.2009.11.001)
- <span id="page-8-13"></span>Hopkinson, C. S., & Vallino, J. J. (2005). Efficient export of carbon to the deep ocean through dissolved organic matter. *Nature*, *433*(7022), 142–145. <https://doi.org/10.1038/nature03191>
- <span id="page-8-30"></span>Jenkins, W. J., & Doney, S. C. (2003). The subtropical nutrient spiral. *Global Biogeochemical Cycles*, *17*(4), 1110. [https://doi.](https://doi.org/10.1029/2003GB002085) [org/10.1029/2003GB002085](https://doi.org/10.1029/2003GB002085)
- <span id="page-8-33"></span>Johnson, K. S., Riser, S. C., & Karl, D. M. (2010). Nitrate supply from deep to near-surface waters of the North Pacific subtropical gyre. *Nature*, *465*(7301), 1062–1065. <https://doi.org/10.1038/nature09170>
- <span id="page-8-7"></span>Karl, D., Letelier, R., Tupas, L., Dore, J., Christian, J., & Hebel, D. (1997). The role of nitrogen fixation in biogeochemical cycling in the subtropical North Pacific Ocean. *Nature*, *388*(6642), 533–538. <https://doi.org/10.1038/41474>
- <span id="page-8-3"></span>Karl, D. M., Björkman, K. M., Dore, J. E., Fujieki, L., Hebel, D. V., Houlihan, T., et al. (2001). Ecological nitrogen-to-phosphorus stoichiometry at station ALOHA. *Deep Sea Research Part II: Topical Studies in Oceanography*, *48*(8), 1529–1566. [https://doi.org/10.1016/](https://doi.org/10.1016/S0967-0645(00)00152-1) [S0967-0645\(00\)00152-1](https://doi.org/10.1016/S0967-0645(00)00152-1)
- <span id="page-8-14"></span>Karl, D. M., Letelier, R. M., Bidigare, R. R., Björkman, K. M., Church, M. J., Dore, J. E., & White, A. E. (2021). Seasonal-to-decadal scale variability in primary production and particulate matter export at Station ALOHA. *Progress in Oceanography*, *195*, 102563. [https://doi.](https://doi.org/10.1016/j.pocean.2021.102563) [org/10.1016/j.pocean.2021.102563](https://doi.org/10.1016/j.pocean.2021.102563)
- <span id="page-8-27"></span>Kim, I. N., Lee, K., Gruber, N., Karl, D. M., Bullister, J. L., Yang, S., & Kim, T. W. (2014). Increasing anthropogenic nitrogen in the North Pacific Ocean. *Science*, *346*(6213), 1102–1106. <https://doi.org/10.1126/science.1258396>
- <span id="page-8-31"></span>Knapp, A. N., Hastings, M. G., Sigman, D. M., Lipschultz, F., & Galloway, J. N. (2010). The flux and isotopic composition of reduced and total nitrogen in Bermuda rain. *Marine Chemistry*, *120*(1), 83–89.<https://doi.org/10.1016/j.marchem.2008.08.007>
- <span id="page-8-16"></span>Knapp, A. N., Letscher, R. T., & Liang, Z. (2021). DOP concentration observations from the global ocean between 1990 and 2021 (DOP N2 fixation and export production project) version 3 [Dataset]. Biological and Chemical Oceanography Data Management Office (BCO-DMO). <https://doi.org/10.26008/1912/bco-dmo.855139.3>
- <span id="page-8-28"></span>Knapp, A. N., Sigman, D. M., & Lipschultz, F. (2005). N isotopic composition of dissolved organic nitrogen and nitrate at the Bermuda Atlantic Time-series Study site. *Global Biogeochemical Cycles*, *19*(1), GB1018. <https://doi.org/10.1029/2004gb002320>
- <span id="page-8-0"></span>Kwon, E. Y., Primeau, F., & Sarmiento, J. L. (2009). The impact of remineralization depth on the air–sea carbon balance. *Nature Geoscience*, *2*(9), 630–635.<https://doi.org/10.1038/ngeo612>
- <span id="page-8-26"></span>Letscher, R. T., & Moore, J. K. (2015). Preferential remineralization of dissolved organic phosphorus and non-Redfield DOM dynamics in the global ocean: Impacts on marine productivity, nitrogen fixation, and carbon export. *Global Biogeochemical Cycles*, *29*(3), 325–340. [https://](https://doi.org/10.1002/2014GB004904) [doi.org/10.1002/2014GB004904](https://doi.org/10.1002/2014GB004904)
- <span id="page-8-25"></span>Letscher, R. T., Moore, J. K., Teng, Y. C., & Primeau, F. (2015). Variable C : N : P stoichiometry of dissolved organic matter cycling in the community Earth system model. *Biogeosciences*, *12*(1), 209–221. <https://doi.org/10.5194/bg-12-209-2015>
- <span id="page-8-9"></span>Letscher, R. T., Primeau, F., & Moore, J. K. (2016). Nutrient budgets in the subtropical ocean gyres dominated by lateral transport. *Nature Geoscience*, *9*(11), 815–819. <https://doi.org/10.1038/ngeo2812>
- <span id="page-8-21"></span>Levy, M., Bopp, L., Karleskind, P., Resplandy, L., Ethe, C., & Pinsard, F. (2013). Physical pathways for carbon transfers between the surface mixed layer and the ocean interior. *Global Biogeochemical Cycles*, *27*(4), 1001–1012.<https://doi.org/10.1002/gbc.20092>
- <span id="page-8-15"></span>Lomas, M. W., Bates, N. R., Johnson, R. J., Knap, A. H., Steinberg, D. K., & Carlson, C. A. (2013). Two decades and counting: 24-years of sustained open ocean biogeochemical measurements in the Sargasso Sea. *Deep Sea Research Part II: Topical Studies in Oceanography*, *93*, 16–32. <https://doi.org/10.1016/j.dsr2.2013.01.008>
- <span id="page-8-4"></span>Lomas, M. W., Bates, N. R., Johnson, R. J., Steinberg, D. K., & Tanioka, T. (2022). Adaptive carbon export response to warming in the Sargasso Sea. *Nature Communications*, *13*(1), 1211. <https://doi.org/10.1038/s41467-022-28842-3>
- <span id="page-8-10"></span>Lomas, M. W., Burke, A. L., Lomas, D. A., Bell, D. W., Shen, C., Dyhrman, S. T., & Ammerman, J. W. (2010). Sargasso Sea phosphorus biogeochemistry: An important role for dissolved organic phosphorus (DOP). *Biogeosciences*, *7*(2), 695–710.<https://doi.org/10.5194/bg-7-695-2010>
- <span id="page-8-11"></span>Mahaffey, C., Williams, R. G., Wolff, G. A., & Anderson, W. T. (2004). Physical supply of nitrogen to phytoplankton in the Atlantic Ocean. *Global Biogeochemical Cycles*, *18*(1), GB1034. <https://doi.org/10.1029/2003GB002129>
- <span id="page-8-20"></span>Martiny, A. C., Lomas, M. W., Fu, W., Boyd, P. W., Chen, Y. L., Cutter, G. A., et al. (2019). Biogeochemical controls of surface ocean phosphate. *Science Advances*, *5*(8), eaax0341. <https://doi.org/10.1126/sciadv.aax0341>
- <span id="page-8-5"></span>Martiny, A. C., Pham, C. T. A., Primeau, F. W., Vrugt, J. A., Moore, J. K., Levin, S. A., & Lomas, M. W. (2013). Strong latitudinal patterns in the elemental ratios of marine plankton and organic matter. *Nature Geoscience*, *6*(4), 279–283. <https://doi.org/10.1038/ngeo1757>
- <span id="page-8-6"></span>Martiny, A. C., Vrugt, J. A., Primeau, F. W., & Lomas, M. W. (2013). Regional variation in the particulate organic carbon to nitrogen ratio in the surface ocean. *Global Biogeochemical Cycles*, *27*(3), 723–731.<https://doi.org/10.1002/gbc.20061>
- <span id="page-8-19"></span>McGillicuddy, D. J., Jr., Anderson, L. A., Doney, S. C., & Maltrud, M. E. (2003). Eddy-driven sources and sinks of nutrients in the upper ocean: Results from a 0.1° resolution model of the North Atlantic. *Global Biogeochemical Cycles*, *17*(2), 1035.<https://doi.org/10.1029/2002GB001987>
- <span id="page-8-22"></span>Omand, M. M., D'Asaro, E. A., Lee, C. M., Perry, M. J., Briggs, N., Cetinić, I., & Mahadevan, A. (2015). Eddy-driven subduction exports particulate organic carbon from the spring bloom. *Science*, *348*(6231), 222–225. <https://doi.org/10.1126/science.1260062>
- <span id="page-8-18"></span>Palter, J. B., Lozier, M. S., Sarmiento, J. L., & Williams, R. G. (2011). The supply of excess phosphate across the Gulf Stream and the maintenance of subtropical nitrogen fixation. *Global Biogeochemical Cycles*, *25*(4), GB4007.<https://doi.org/10.1029/2010GB003955>
- <span id="page-8-8"></span>Quay, P. D. (2021). Impact of the elemental composition of exported organic matter on the observed dissolved nutrient and trace element distributions in the upper layer of the ocean. *Global Biogeochemical Cycles*, *35*(10), e2020GB006902.<https://doi.org/10.1029/2020GB006902>
- <span id="page-8-1"></span>Quay, P. D., Emerson, S., & Palevsky, H. (2020). Regional pattern of the ocean's biological pump based on geochemical observations. *Geophysical Research Letters*, *47*(14), e2020GL088098.<https://doi.org/10.1029/2020GL088098>
- <span id="page-8-17"></span>Quay, P. D., & Stutsman, J. (2003). Surface layer carbon budget for the subtropical N. Pacific: δ13C constraints at station ALOHA. *Deep Sea Research Part I: Oceanographic Research Papers*, *50*(9), 1045–1061. [https://doi.org/10.1016/S0967-0637\(03\)00116-X](https://doi.org/10.1016/S0967-0637(03)00116-X)
- <span id="page-8-2"></span>Redfield, A. C. (1958). The biological control of chemical factors in the environment. *American Scientist*, *46*(3), 205–221.
- <span id="page-8-24"></span>Resplandy, L., Lévy, M., & McGillicuddy, D. J., Jr. (2019). Effects of eddy-driven subduction on ocean biological carbon pump. *Global Biogeochemical Cycles*, *33*(8), 1071–1084. <https://doi.org/10.1029/2018GB006125>
- <span id="page-8-12"></span>Reynolds, S., Mahaffey, C., Roussenov, V., & Williams, R. G. (2014). Evidence for production and lateral transport of dissolved organic phosphorus in the eastern subtropical North Atlantic. *Global Biogeochemical Cycles*, *28*(8), 805–824. <https://doi.org/10.1002/2013GB004801>
- <span id="page-8-23"></span>Roshan, S., & DeVries, T. (2017). Efficient dissolved organic carbon production and export in the oligotrophic ocean. *Nature Communications*, *8*(1), 2036. <https://doi.org/10.1038/s41467-017-02227-3>

<span id="page-8-32"></span>Sainani, K. L. (2012). Dealing with non-normal data. *PM&R*, *4*(12), 1001–1005. <https://doi.org/10.1016/j.pmrj.2012.10.013>

- <span id="page-9-7"></span>Siegel, D. A., & Deuser, W. G. (1997). Trajectories of sinking particles in the Sargasso Sea: Modeling of statistical funnels above deep-ocean sediment traps. *Deep Sea Research Part I: Oceanographic Research Papers*, *44*(9), 1519–1541. [https://doi.org/10.1016/S0967-0637\(97\)00028-9](https://doi.org/10.1016/S0967-0637(97)00028-9) Siegel, D. A., DeVries, T., Cetinić, I., & Bisson, K. M. (2023). Quantifying the ocean's biological pump and its carbon cycle impacts on global
- <span id="page-9-10"></span><span id="page-9-4"></span>scales. *Annual Review of Marine Science*, *15*(1), 329–356.<https://doi.org/10.1146/annurev-marine-040722-115226> Singh, A., Baer, S. E., Riebesell, U., Martiny, A. C., & Lomas, M. W. (2015). C : N : P stoichiometry at the Bermuda Atlantic time-series Study
- <span id="page-9-14"></span>station in the North Atlantic Ocean. *Biogeosciences*, *12*(21), 6389–6403.<https://doi.org/10.5194/bg-12-6389-2015> Singh, A., Lomas, M. W., & Bates, N. R. (2013). Revisiting N2 fixation in the North Atlantic Ocean: Significance of deviations from the Redfield ratio, atmospheric deposition and climate variability. *Deep Sea Research Part II: Topical Studies in Oceanography*, *93*, 148–158. [https://doi.](https://doi.org/10.1016/j.dsr2.2013.04.008) [org/10.1016/j.dsr2.2013.04.008](https://doi.org/10.1016/j.dsr2.2013.04.008)
- <span id="page-9-11"></span>Stanley, R. H. R., Doney, S. C., Jenkins, W. J., & Lott, I. D. E. (2012). Apparent oxygen utilization rates calculated from tritium and helium-3 profiles at the Bermuda Atlantic Time-series Study site. *Biogeosciences*, *9*(6), 1969–1983. <https://doi.org/10.5194/bg-9-1969-2012>
- <span id="page-9-12"></span>Stanley, R. H. R., Jenkins, W. J., Doney, S. C., & Lott, D. E., III. (2015). The 3He flux gauge in the Sargasso Sea: A determination of physical nutrient fluxes to the euphotic zone at the Bermuda Atlantic time-series Site. *Biogeosciences*, *12*(17), 5199–5210. [https://doi.org/10.5194/](https://doi.org/10.5194/bg-12-5199-2015) [bg-12-5199-2015](https://doi.org/10.5194/bg-12-5199-2015)
- <span id="page-9-5"></span>Steinberg, D. K., Goldthwait, S. A., & Hansell, D. A. (2002). Zooplankton vertical migration and the active transport of dissolved organic and inorganic nitrogen in the Sargasso Sea. *Deep Sea Research Part I: Oceanographic Research Papers*, *49*(8), 1445–1461. [https://doi.](https://doi.org/10.1016/S0967-0637(02)00037-7) [org/10.1016/S0967-0637\(02\)00037-7](https://doi.org/10.1016/S0967-0637(02)00037-7)
- <span id="page-9-13"></span>Tang, W., Wang, S., Fonseca-Batista, D., Dehairs, F., Gifford, S., Gonzalez, A. G., et al. (2019). Revisiting the distribution of oceanic N2 fixation and estimating diazotrophic contribution to marine production. *Nature Communications*, *10*(1), 831. [https://doi.org/10.1038/](https://doi.org/10.1038/s41467-019-08640-0) [s41467-019-08640-0](https://doi.org/10.1038/s41467-019-08640-0)
- <span id="page-9-1"></span>Teng, Y. C., Primeau, F. W., Moore, J. K., Lomas, M. W., & Martiny, A. C. (2014). Global-scale variations of the ratios of carbon to phosphorus in exported marine organic matter. *Nature Geoscience*, *7*(12), 895–898. <https://doi.org/10.1038/ngeo2303>
- <span id="page-9-6"></span>Torres-Valdés, S., Roussenov, V. M., Sanders, R., Reynolds, S., Pan, X., Mather, R., et al. (2009). Distribution of dissolved organic nutrients and their effect on export production over the Atlantic Ocean. *Global Biogeochemical Cycles*, *23*(4), GB4019.<https://doi.org/10.1029/2008GB003389>
- <span id="page-9-0"></span>Volk, T., & Hoffert, M. I. (1985). Ocean carbon pumps: Analysis of relative strengths and efficiencies in ocean-driven atmospheric CO2 changes. In E. T. Sundquist & W. S. Broecker (Eds.), *The carbon cycle and atmospheric CO2: Natural variations archean to present* (Vol. 32, pp. 99–110). American Geophysical Union. <https://doi.org/10.1029/GM032p0099>
- <span id="page-9-2"></span>Wang, W. L., Moore, J. K., Martiny, A. C., & Primeau, F. W. (2019). Convergent estimates of marine nitrogen fixation. *Nature*, *566*(7743), 205–211. <https://doi.org/10.1038/s41586-019-0911-2>
- <span id="page-9-3"></span>Weber, T. S., & Deutsch, C. (2010). Ocean nutrient ratios governed by plankton biogeography. *Nature*, *467*(7315), 550–554. [https://doi.](https://doi.org/10.1038/nature09403) [org/10.1038/nature09403](https://doi.org/10.1038/nature09403)
- <span id="page-9-8"></span>Williams, R. G., & Follows, M. J. (1998). The Ekman transfer of nutrients and maintenance of new production over the North Atlantic. *Deep Sea Research Part I: Oceanographic Research Papers*, *45*(2), 461–489. [https://doi.org/10.1016/S0967-0637\(97\)00094-0](https://doi.org/10.1016/S0967-0637(97)00094-0)
- <span id="page-9-9"></span>Williams, R. G., McDonagh, E., Roussenov, V. M., Torres-Valdes, S., King, B., Sanders, R., & Hansell, D. A. (2011). Nutrient streams in the North Atlantic: Advective pathways of inorganic and dissolved organic nutrients. *Global Biogeochemical Cycles*, *25*(4), GB4008. [https://doi.](https://doi.org/10.1029/2010GB003853) [org/10.1029/2010GB003853](https://doi.org/10.1029/2010GB003853)
- <span id="page-9-15"></span>Zamora, L. M., Landolfi, A., Oschlies, A., Hansell, D. A., Dietze, H., & Dentener, F. (2010). Atmospheric deposition of nutrients and excess N formation in the North Atlantic. *Biogeosciences*, *7*(2), 777–793. <https://doi.org/10.5194/bg-7-777-2010>

# **References From the Supporting Information**

- Quay, P. D., & Wu, J. (2015). Impact of end-member mixing on depth distributions of δ13C, cadmium and nutrients in the N. Atlantic Ocean. *Deep Sea Research Part II: Topical Studies in Oceanography*, *116*, 107–116. <https://doi.org/10.1016/j.dsr2.2014.11.009>
- Quay, P. D., Cullen, J. T., Landing, W. M., & Morton, P. L. (2015). Processes controlling the distributions of Cd and PO4 in the ocean. *Global Biogeochemical Cycles*, *29*(6), 830–841.<https://doi.org/10.1002/2014GB004998>