

1 **Anthropogenic VOCs in the Long Island Sound, NY Airshed**
2 **and their Role in Ozone Production**

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20 **Abstract**

21 Cities like New York and areas downwind have poor air quality due to pollution
22 generated by human activity. Ozone (O_3), a harmful pollutant, is produced in the atmosphere
23 from photochemical reactions involving volatile organic compounds (VOCs) and nitrogen oxides
24 (NO_x). Aircraft measurements of VOCs obtained during an O_3 event in May 2017 over the New
25 York City metropolitan area, Long Island Sound, and Connecticut show concentrations of O_3
26 exceeding 100 ppb between ~200 – 500 meters above the surface. During this campaign, 35
27 whole air canisters were collected, and 50 VOC species were measured. Biogenic isoprene often
28 dominates VOC reactivity in summer months, but analysis of the chemistry in the middle of the
29 Planetary Boundary Layer (PBL) indicates that primary anthropogenic VOCs such as propylene
30 and isopentane played a major, possibly dominant role in this Spring 2017 O_3 event and are
31 suitable targets for control.

32 The most important measured anthropogenic VOCs ranked by OH reactivity were
33 propylene and isopentane, with n-butane and 1-pentene also playing a substantial role. The most
34 important anthropogenic VOCs ranked by Ozone Formation Potential (OFP) were again
35 propylene and isopentane, but aromatic species like toluene were also large contributors.
36 Isoprene ranked tenth in importance for OFP. For VOC species with short lifetimes, like isoprene
37 and propylene, mixing of air parcels throughout the PBL means air above the surface may
38 quickly become depleted of these VOCs, allowing longer-lived, anthropogenic VOCs to have a
39 proportionally greater impact on photochemical O_3 production aloft. Our observations indicate
40 that VOC reactivity was much larger in the morning than in the afternoon, suggesting controls of
41 propylene, isopentane, and aromatics overnight may be useful for reducing morning
42 concentrations of VOCs and therefore possibly reduce maximum concentrations of O_3 observed
43 in the afternoon. Finally, because formaldehyde (HCHO) was not measured during these flights,

44 and since oxygenates also contribute to O₃ formation, we estimate the relative contribution of
45 isoprene and propylene to HCHO formation using a box model. Propylene produces ~25% and
46 isoprene produces ~4% of HCHO within the box model.

47 **Keywords:**

48 Air Quality, Long Island Sound Tropospheric Ozone Study (LISTOS), Ozone production
49 regime, Anthropogenic VOCs, Maximum Incremental Reactivity, Ozone Formation Potential

50 **1: Introduction**

51 The National Ambient Air Quality Standards (NAAQS), established by the United States
52 Environmental Protection Agency (EPA) under the federal Clean Air Act, set national levels for
53 six criteria pollutants deemed harmful to human health and welfare [EPA, 1970; 1990]. EPA
54 reconsiders the NAAQS periodically and has increased the stringency of the surface ozone (O₃)
55 standard over time, with the most recent revision occurring in 2015. As of October 2022, the
56 New York City (NYC) metropolitan area is in non-attainment of both the older 2008 and more
57 recent 2015 O₃ NAAQS, which are 75 parts per billion (ppb) and 70 ppb for an 8 hour average,
58 respectively [EPA, 2016a; *Federal Register*, 2022a; b]. Areas with particularly poor air quality,
59 such as the NYC region, must use EPA-approved air quality models to develop strategies
60 demonstrating how they will reach attainment. Observational studies can support these modeling
61 efforts by identifying the meteorological and chemical conditions, emission sources, and the
62 specific atmospheric species most important for the photochemical production of tropospheric
63 O₃.

64 Located downwind from NYC, Connecticut (CT) frequently receives on-shore
65 atmospheric circulation of anthropogenic emissions exported from NYC over Long Island Sound
66 (LIS) [Caicedo *et al.*, 2021; Han *et al.*, 2022; Torres-Vazquez *et al.*, 2022; Wu *et al.*, 2021],

67 similar to bay-breezes advecting polluted air from the Chesapeake Bay over the Baltimore, MD
68 region [*Loughner et al.*, 2014; *Mazzuca et al.*, 2017; *Stauffer et al.*, 2015]. All of CT, including
69 its coastline, is presently designated as in non-attainment of the 2015 O₃ standard
70 [*Federal_Register*, 2022a]. CT, along with New Jersey and New York, will need to perform air
71 quality modeling that identifies the upwind sources, including those from states outside the NYC
72 metropolitan area, that contribute to poor air quality throughout the region.

73 Identifying volatile organic compounds (VOCs) most important for photochemical O₃
74 production is an area of active research. Previous studies have identified many biogenic and
75 anthropogenic VOCs that impact air quality, especially during summer months [*Baker et al.*,
76 2008; *Kleinman*, 2000]. *McDonald et al.* [2018] showed with a review of U.S. emissions
77 inventories and box modeling analyses that VOCs from pesticides, adhesives, personal care
78 products, cleaning agents, etc. (collectively called Volatile Chemical Products (VCPs)) are
79 important sources of urban VOCs. They showed that VCPs significantly impact ground-level
80 outdoor and indoor air quality in Pasadena, California and throughout the Los Angeles basin. In
81 the Long Island Sound Tropospheric Ozone Study (LISTOS) study region, *Coggon et al.* [2021]
82 showed that VCPs, especially anthropogenically produced monoterpenes, add significantly to the
83 total VOC burden in the NYC metropolitan region. Inclusion of chemistry associated with
84 oxygenated VCPs contributed to the production of peroxyacetyl nitrate (PAN), critical for the
85 model to achieve the high O₃ concentrations observed in the NYC metropolitan region during the
86 summer of 2018. Both *McDonald et al.* [2018] and *Coggon et al.* [2021] show the importance of
87 other anthropogenic VOCs in addition to those from biogenic and fossil fuel sources, especially
88 in densely populated metropolitan regions. Inclusion of VCPs along with mobile and
89 petrochemical VOC sources in model emissions inventories will be necessary to bring modeled

90 and observed OH reactivity and secondary organic aerosol formation in urban centers into
91 agreement [Seltzer *et al.*, 2021]. Monoterpenes along with sesquiterpenes and diterpenes also
92 have biogenic sources. These emissions occur throughout the day and evening, contributing to a
93 buildup of terpene concentrations at night [Li *et al.*, 2020]. High evening concentrations of
94 terpenes likely lead to O₃ scavenging and further contribute to the reduced concentrations of O₃
95 at night. Examination of terpenes in addition to isoprene is beyond the scope of this paper as
96 these VOCs were not measured during these aircraft flights, but more field missions and ground-
97 based monitoring of these VOCs are necessary because they are important components to
98 understanding atmospheric O₃ formation. Various research groups are working to improve
99 terpene chemistry in air quality models [Schwantes *et al.*, 2020], and observations of these
100 species will be useful for validation.

101 *Chen et al.* [2019] conducted a thorough analysis comparing VOCs measured in aircraft
102 campaigns over North America to model simulations. They found isoprene generally dominates
103 VOC reactivity across the US but recognized this result may not be true for polluted urban
104 regions. They explicitly excluded from their analysis observations of highly polluted air masses
105 (NO₂ > 4 ppbv) and recent biomass burning (CH₃CN > 0.2 ppbv) and called for more research to
106 understand the complex VOC chemistry in urban centers. Our study complements *McDonald et*
107 *al.* [2018], *Coggon et al.* [2021], and *Chen et al.* [2019] by examining airborne observations of a
108 suite of anthropogenic VOCs in the polluted NYC Metropolitan region that play an important
109 role in O₃ formation near this urban center and areas downwind. While it is clear from
110 *McDonald et al.* [2018] and *Coggon et al.* [2021] that VCPs and oxygenates are quite important
111 to air quality, measurements of these VOCs are not yet widely available due to methodology
112 limitations. This study focuses on primary VOCs with a long history of monitoring, well-

113 established methods of analysis, and VOCs that are inventoried, regulated and controllable. Our
114 flight campaign provides observations throughout the Planetary Boundary Layer (PBL) of O₃ as
115 well as these primary VOCs, typically observed only from the surface at monitoring locations
116 like the Photochemical Assessment Monitoring Stations (PAMS) sites, which *Hembeck et al.*
117 [2019] showed may be missing important VOCs.

118

119 **2: Methods and Data**

120 *Carter* [1994] detailed the development of a metric known as the Maximum Incremental
121 Reactivity (MIR) to represent the relative importance of individual VOCs to surface O₃
122 production. The MIR scale is derived from smog chamber analysis and photochemical box-
123 modeling of VOC-sensitive O₃ production scenarios, meaning VOCs have the maximum impact
124 on O₃ production. To calculate MIR, *Carter* [1994] performed two model simulations: a baseline
125 VOC-sensitive simulation, and one with VOCs added to the emissions that produced a linear
126 increase in O₃. The difference in the amount of O₃ produced between the two simulations,
127 divided by the amount of VOC added, yields the MIR (g O₃ produced / g of VOC added)
128 [*Carter*, 1994]. A significant refinement of previously calculated MIR values for VOCs was
129 performed by *Carter* [2010], and further improvements to MIR values using representative urban
130 atmospheres were conducted by *Venecek et al.* [2018]. All subsequent analysis in this study is
131 conducted using the updated MIR values from *Venecek et al.* [2018].

132 To identify the anthropogenic VOCs that may have been major contributors to the high
133 O₃ event in Spring 2017, we calculate the Ozone Formation Potential (OFP) of each VOC. Here,
134 OFP represents the amount of O₃ (ppbv) that could potentially be formed under ideal
135 atmospheric circumstances based on the amount of VOC present and its associated MIR value

136 (equation 1) [Tang *et al.*, 2007]. The MIR calculation is based on varying VOC emissions and
137 determining the resulting increase in O₃, so this quantity is best applied to VOC emissions. The
138 formula for OFP is:

$$139 \qquad \qquad \qquad OFP = MIR \times [VOC] \qquad (1)$$

140 In this work, MIR is used to calculate OFP based on observed VOC concentrations.
141 Therefore, OFP found in this manner is a lower limit to the potential O₃ formed in the actual
142 atmosphere, especially for the VOCs with short lifetimes such as alkenes since much of the
143 original emission has already reacted in the atmosphere before it can be measured. For those
144 VOCs with longer lifetimes, this method may be a better approximation of the actual OFP. This
145 metric we use places MIR into quantifiable terms, allowing air quality planners to establish
146 bounds for the potential production of O₃ based on each VOC and identify target species and
147 emission sources to control. Unlike OH reactivity, OFP accounts for compounds and radicals
148 produced in subsequent oxidation steps.

149 We analyze data collected from three research flights conducted over LIS and CT in May
150 2017 that provide *in situ* measurements of VOCs and other atmospheric species during an O₃
151 exceedance event. The reactivity with OH and the OFP for each VOC are calculated and ranked
152 in order of importance, providing air quality policy makers with a list of VOCs most critical to
153 the production of O₃ in the LIS region during this exceedance event.

154 **2.1: Aircraft Data**

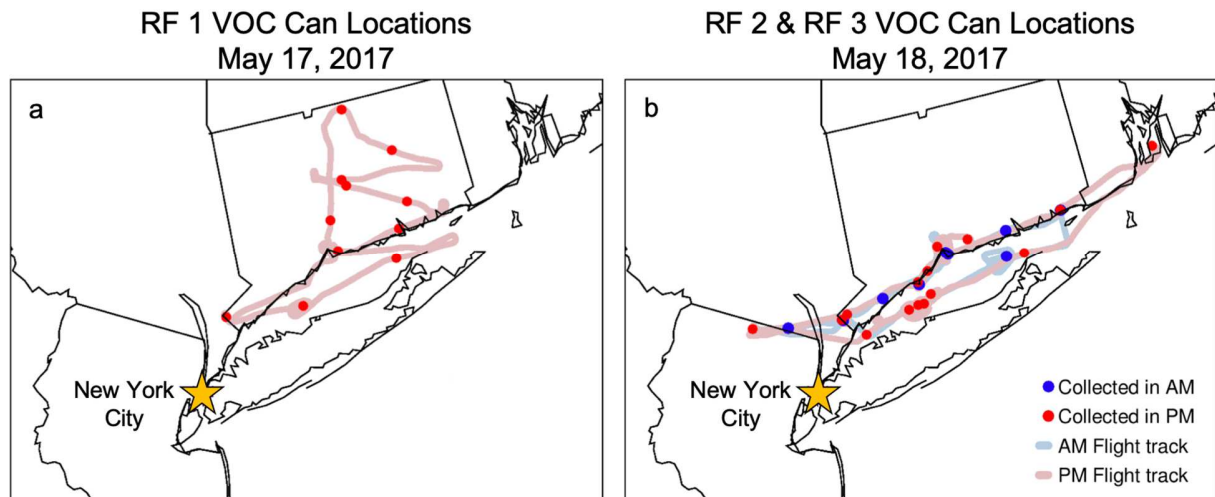
155 For the past two decades, the University of Maryland has led a multi-faceted effort to
156 understand air quality in the Eastern U.S. called the Regional Atmospheric Measurement
157 Modeling and Prediction Program (RAMMPP) (<http://www.atmos.umd.edu/~rammpp>). *In situ*
158 aircraft measurements are collected during the O₃ season (April – October) providing

159 information about long-term trends in the Mid-Atlantic and Eastern U.S. [*He et al.*, 2013;
160 *Taubman et al.*, 2004]. For this study, a small portion of the RAMMPP data are used to analyze
161 VOCs in the LIS region of New York on May 17 and 18, 2017. These data were collected during
162 a high O₃ event, over the course of three research flights with the intent of sampling air within
163 the boundary layer, from near the surface to ~1500 m altitude.

164 Trace gas and aerosol measurements were collected by University of Maryland
165 researchers aboard a Cessna 402B aircraft. A complete description of the aircraft's
166 instrumentation setup can be found in *Ren et al.* [2018]. A total of 35 whole air canisters
167 (hereafter simply called cans) sampling ambient air mostly between 250 – 500 m above the
168 surface were collected, and 50 VOC species were reported for each can during this field
169 campaign. The list of species can be found in the supplemental information (Table S1). All can
170 samples were processed by the Maine Department of Environmental Protection (DEP) using
171 Standard Operating Procedures (SOP) [*MaineDEP*, 2017] in accordance with the National Air
172 Toxics Trends Station technical assistance document [*EPA*, 2016b]. Samples were analyzed
173 using a Gas Chromatograph (Agilent 7890A GC System) coupled to a Mass Spectrometer
174 (Agilent 5975C Inert MSD) and referenced to a NIST traceable certified TO-15 VOC standard.
175 Preceding any analysis, the GC/MS is tested to ensure attainment of the instrument performance
176 standard and is then calibrated within 24 hours prior to processing the can samples. The reported
177 VOC concentrations (ppbv) have an uncertainty of $\pm 20\%$ [*EPA*, 2016b; *MaineDEP*, 2017].

178 Two reference or blank cans were collected during the sample period. The average
179 reference can concentration for each VOC species was calculated and subtracted from the
180 corresponding measured VOC concentration in each sample can, correcting for any

181 contamination from the can interior or sample inlets. Maps of the three research flights and VOC
182 can collection locations are shown in Figure 1.

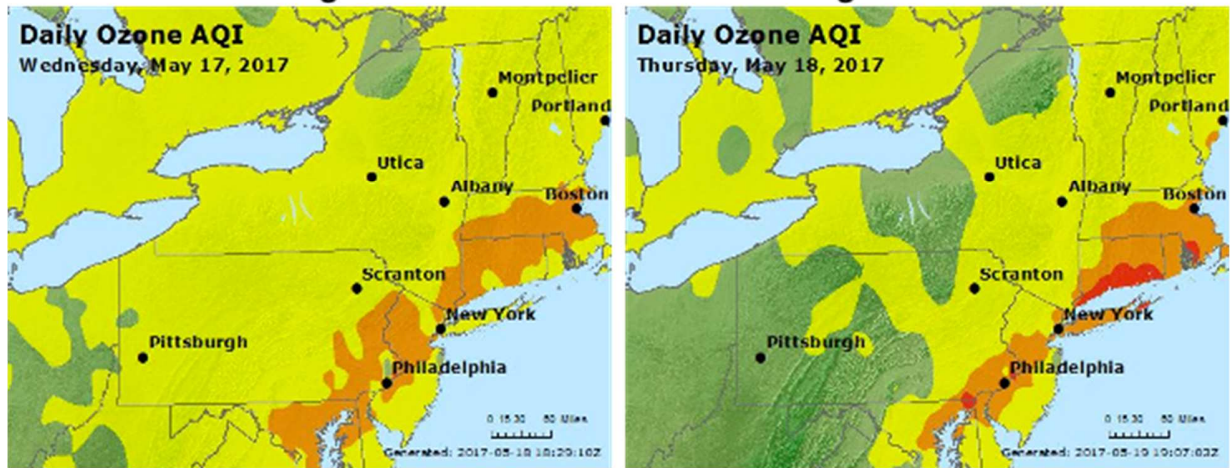


183
184 **Figure 1:** VOC can locations and research flight (RF) tracks over Connecticut and Long Island Sound (a) from
185 14:50 to 18:10 LST on May 17, 2017 and (b) from 7:40 to 10:25 LST (AM) and from 13:10 to 16:40 LST (PM) on
186 May 18, 2017. VOC cans were collected in flight and at various altitudes, with the vast majority within the boundary
187 layer. Color distinctions represent cans collected in the morning (blue) and in the afternoon (red). A total of 9 cans
188 were collected in the morning and a total of 26 cans were collected in the afternoon.

189
190 Figure 2 shows the archived Air Quality Index (AQI) maps for May 17 and 18, 2017
191 from the AirNow website. This forecasted O₃ exceedance event, early in the O₃ season,
192 motivated these research flights. Surface O₃ concentrations were forecast to reach unhealthy
193 levels for sensitive groups on May 17, escalating to unhealthy levels for everyone on May 18
194 along the CT coast north of the LIS [AirNow, Accessed 2018]. The measured surface 8-hour
195 average maximum O₃ and 1-hour maximum O₃ on May 17 and 18, 2017 for selected Air Quality
196 System (AQS) sites in NY and CT around LIS shown in Figure S1 are provided in Table S2. Of
197 the 18 total AQS sites examined, 9 sites were in exceedance of the NAAQS 2015 surface O₃
198 standard of 70 ppb on May 17, 2017, and 14 sites were in exceedance on May 18, 2017. The
199 maximum O₃ value measured along RF1 was 116 ppb, along RF2 was 81 ppb, and along RF3

200 was 130 ppb. Ambient temperatures peaked for the morning flight (RF2) at 82°F (27.8°C) and at
201 91°F (32.8°C) for both afternoon flights (RF1, RF3). Surface winds were westerly, generally
202 moving the NYC pollution plume over LIS and CT.

Archived Ozone AQI Images for Long Island Sound Flight Locations and Surrounding Area



204 **Figure 2:** AirNow O₃ AQI forecast images for May 17 and 18, 2017. Yellow coloring indicates air quality is
205 acceptable however there may be a moderate health concern for people who are especially sensitive to pollution
206 (AQI: 51-100). Orange coloring indicates air quality is unsafe for sensitive groups like the elderly, children, and
207 people with respiratory problems (AQI: 101-150). Red coloring indicates air quality is unsafe and all people may
208 begin to experience health effects (AQI: 151-200). More information about AQI and the designations described
209 above can be found on the AirNow website <https://www.airnow.gov>. The images above can be found in the archives
210 located at <http://files.airnowtech.org/?prefix=airnow/2017/>.

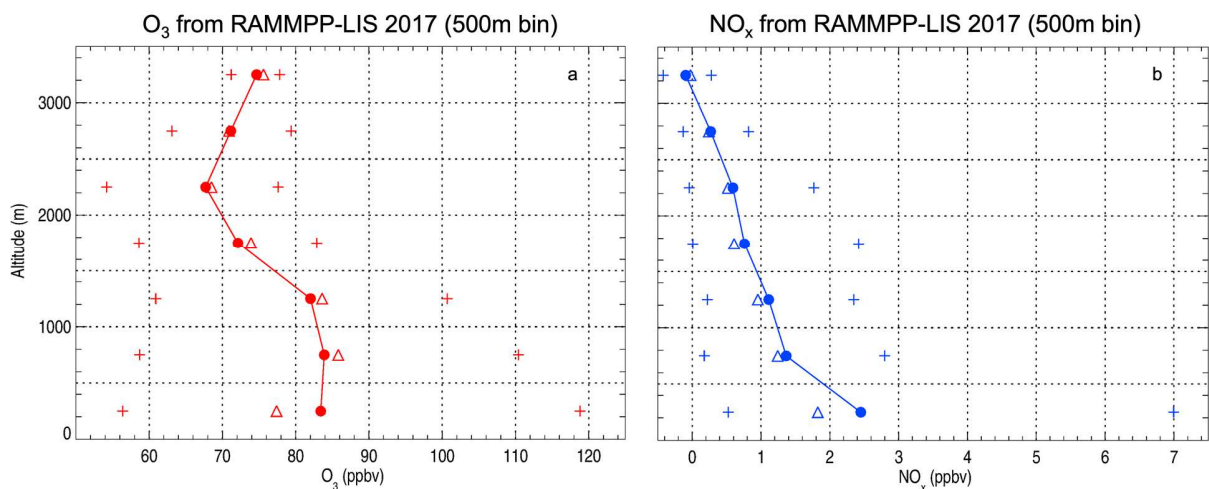
211 Since 1998, approximately 14% of all O₃ exceedances in NJ, NY, and CT occur during
212 the meteorological spring (before June 1), and approximately 28% of exceedances occur during
213 the astronomical spring (before June 21). A table showing the spring exceedances, total number
214 of exceedances and percentage of spring exceedances for all three states for every year from
215 1998-2021 can be found in supplemental information (Table S3). A map of the monitoring sites
216 used to generate this table is shown in Figure S2. While there are fewer exceedances during the
217 spring than the summer months, exceedance events like the one in this study are not uncommon.

218 **3: Results and Discussion**

219 On May 17 and 18, 2017, the NYC metropolitan area suffered an O₃ event, with
220 Maximum Daily Average 8-hr O₃ (MDA8) reaching 84 ppb at the Stafford Hollow, CT site on
221 May 17 and 91 ppb at the Bridgeport, CT site on May 18, consistent with the forecasted Daily
222 Ozone AQI shown in Figure 2. Table S2 provides MDA8 and 1-hr Max values of O₃ at all AQS
223 sites in the LIS area. Figure 3 shows the observed O₃ and NO_x for all three RFs, binned every
224 500 meters. Both O₃ and NO_x were measured continuously along the flight tracks, providing
225 observations with high temporal resolution. As shown in Figure 3a high concentrations of O₃
226 were observed aloft, indicating photochemical O₃ production occurred throughout the PBL.
227 Further analysis shows elevated levels of SO₂ in conjunction with the high O₃ above 3000 meters
228 (Figure 3a), indicating a possible influence from powerplants via long-range transport. Figure S3
229 shows the vertical 500 meter binned profile of SO₂ and the temperature profile of the aircraft
230 spiral with the highest SO₂ measurement. A temperature inversion is evident at ~2500 meters
231 indicating the top of the PBL, providing further support for high SO₂ and O₃ transport in the
232 lower free troposphere. HYSPLIT ensemble back trajectories [Stein *et al.*, 2015] in Figure S4
233 provide further evidence for long-range transport of pollution by showing the air above the

234 temperature inversion with high SO₂ and O₃ originated over a region with numerous active
235 powerplants that emitted SO₂ during these flights. HYSPLIT trajectories are initialized at the
236 height of the highest SO₂ measurement and are conducted for two days using archived 3 km
237 High Resolution Rapid Refresh meteorological input. The black dots in Figure S4 represent the
238 power plants in operation on May 16 and 17 according to Continuous Emissions Measurement
239 System (CEMS) data. To confirm the presence of synoptic meteorological conditions favorable
240 for atmospheric transport, we conducted two additional ensemble back trajectories initialized at
241 2500 meters and 3500 meters (Figure S4). A consistent pattern is discernable among all three
242 ensemble trajectories.

243 The mixing time within the PBL is typically only a few hours on a warm sunny day,
244 meaning O₃ produced aloft or transported can mix down to the surface. Mixing of O₃, its
245 precursors and their intermediates throughout the PBL introduces uncertainty into determining
246 the O₃ production regimes in air parcels at the surface and higher in the PBL. Measurements of
247 species at the surface and aloft provide a more complete understanding of the local O₃
248 production regime, and therefore must be considered when trying to develop mitigation
249 strategies for O₃ exceedance events.



250

251 **Figure 3:** Observations of O₃ and NO_x from all flights over Long Island Sound conducted on May 17 (RF1) and
252 May 18 (RF2, RF3) in 2017. (a) O₃ observations and (b) NO_x observations are binned every 500 meters. The solid
253 circles represent the average concentration, the open triangles are the 50th percentile, and the + are the 5th and 95th
254 percentiles for the observations within each altitude bin.

255

256 **3.1: Importance of HCHO**

257 Formaldehyde (HCHO) can be used as a proxy to estimate VOCs in the atmosphere from
258 satellite observations because HCHO can be measured from space and is a significant
259 intermediate in the atmospheric oxidation of VOCs. Space based observations of HCHO are not
260 used for this analysis because they have a footprint much larger than the aircraft sampling area,
261 typically occur during a single overpass once a day, and the retrieval on individual days can be
262 noisy, requiring averaging over longer timeframes to produce a cleaner signal [Harkey *et al.*,
263 2021; Wolfe *et al.*, 2016a; Zhu *et al.*, 2016]. Due to the propensity of HCHO to stick to the walls
264 of the sample cans and react quickly with other species, HCHO cannot be reliably stored in a
265 whole air sampler canisters, and therefore is not measured from the can samples. Therefore, the
266 direct measurement of HCHO, like with the NASA In-situ Airborne Formaldehyde Instrument
267 [Wolfe *et al.*, 2016a], is the best way to quantify the ambient concentration of HCHO. For these
268 flights, direct measurement of HCHO was not conducted, so we are not able to incorporate
269 observations of HCHO into the analysis.

270 During summer months, isoprene accounts for the majority of HCHO present in the
271 atmosphere over the NYC region [Lin *et al.*, 2012; Luecken *et al.*, 2012; Luecken *et al.*, 2018; Lv
272 *et al.*, 2019; Wolfe *et al.*, 2016a]. This flight campaign was conducted during the meteorological
273 spring, so rather than isoprene, anthropogenic VOCs and primary emissions of HCHO are likely
274 the dominant sources of HCHO in the atmosphere. Other VOCs not measured in this campaign
275 are also likely significant sources of atmospheric HCHO, like terpenes and ethylene [Franco *et*

276 *al.*, 2022; *Luecken et al.*, 2018]. Without direct measurement of these VOCs, we cannot
277 comment on their decomposition leading to HCHO production in the atmosphere.

278 A box model was used to assess the potential production of HCHO from propylene and
279 isoprene. These VOCs have similar atmospheric lifetimes (hours) and potential O₃ production,
280 but propylene has an anthropogenic and isoprene has a biogenic source. The Framework for 0-
281 Dimensional Atmospheric Modeling (FOAM) box model [*Wolfe et al.*, 2016b] uses the Master
282 Chemical Mechanism (MCM) version 3.3.1 chemical mechanism and accounts for the diurnal
283 cycle. The model is constrained to all measured atmospheric species at the time of day the can
284 was collected. Each simulation is initialized with 0 ppb of HCHO and calculates HCHO based on
285 the VOCs present and their reactions within the MCM. Three different simulations are conducted
286 for each can: 1. All VOCs are initialized with observed values (All VOC), 2. All VOCs except
287 isoprene are initialized with observed values (No Isoprene) and 3. All VOCs except propylene
288 are initialized with observed values (No Propylene). Concentrations of HCHO are reported for
289 each simulation 1 hour after model initialization. This timeframe allows important chemistry to
290 occur and produce HCHO quickly but limits concentrations from building up in the model due to
291 reactions that happen at long time scales. FOAM cannot reliably be used to estimate the
292 concentrations of HCHO that likely would have been sampled in each can because it cannot
293 account for primary emissions or the amount of HCHO already produced from VOC oxidation in
294 the atmosphere prior to sampling. As stated above, the model calculated HCHO is likely an
295 underestimation therefore, its strength resides in determining the relative contribution of
296 individual VOCs to potential HCHO production. By comparing the “No Isoprene” simulations to
297 the “All VOC” simulations, we determine the relative importance of isoprene to potential HCHO
298 production for each can. The same is done using the No Propylene simulations to estimate the

299 importance of propylene to potential HCHO production for each can. The percentage of HCHO
300 produced in the box model for each can based on isoprene and propylene concentrations are
301 calculated. The mean and median values of these percentages are 4.51% and 2.98% from
302 isoprene and 24.70% and 25.16% from propylene; the HCHO calculations for these F0AM
303 simulations can be found in supplemental information (Table S4). Based on these results, it is
304 likely that only a small fraction of the total HCHO produced from VOCs in the middle of the
305 PBL came from isoprene. For Spring 2017 in the NYC, LIS, and CT region, anthropogenic
306 VOCs play an important and possibly dominant role in HCHO formation and O₃ production.

307

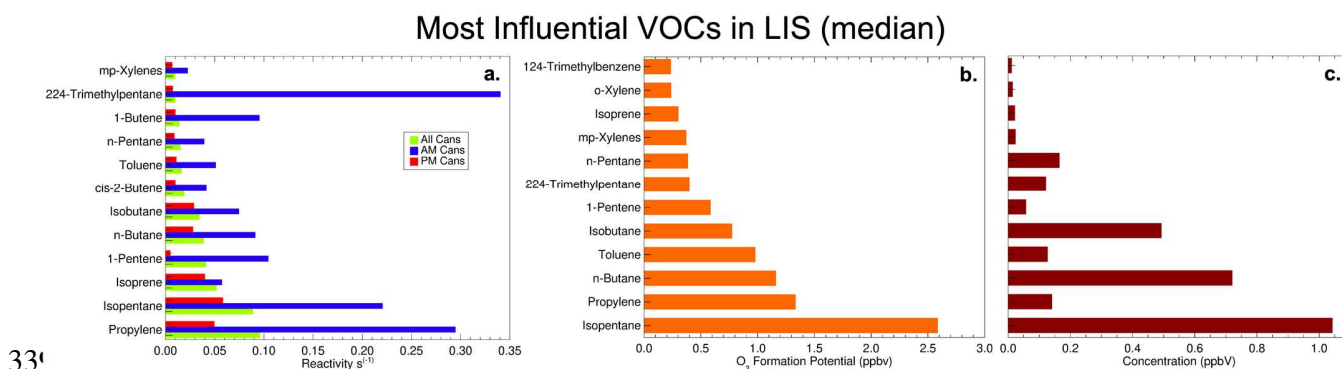
308 **3.2: Reactivity with OH**

309 Reaction of VOCs with OH initiates a series of chemical reactions that form oxidants and
310 lead to photochemical O₃ production as well as secondary organic aerosols. Therefore, rate
311 constants for the reaction of VOCs with OH (k_{OH}) are the primary factor for determining
312 individual contributions of various VOCs to the formation of secondary pollutants [Carter, 1994;
313 Kansal, 2009]. VOCs can also react with O₃, however this reaction is usually slower than the
314 reaction with OH. Nevertheless, O₃ consumption rates are underestimated when solely
315 considering k_{OH} .

316 Rate constant equations for the reaction of OH with each VOC in this analysis are
317 included in Table S1. All 50 VOC species measured are ranked by their OH reactivity during this
318 sampling period, to determine which species are most important for photochemical O₃ formation.
319 In Figure 4a, the 12 VOCs with the highest median OH reactivity values for all cans are shown
320 as green bars. Median OH reactivity values for the listed VOCs are also shown for cans collected
321 in the morning (blue bars) and in the afternoon (red bars), highlighting important diurnal
322 differences. Median OH reactivity is calculated by converting the median concentration of each

323 VOC from ppbv to molecules/cm³, accounting for pressure and temperature, and multiplying this
 324 value by the rate constant k_{OH} . The median and mean observed concentrations for each VOC are
 325 included in Table S1.

326 The morning reactivity of VOCs exceeds afternoon reactivity (Figure 4) due to the larger
 327 abundance of VOCs earlier in the day compared to the afternoon. The morning excess may be
 328 due to 1) a build-up of overnight emissions that are likely anthropogenic VOCs (note: little
 329 biogenic isoprene is emitted at night) resulting in larger morning concentrations prior to peak
 330 solar insolation; 2) photochemical break-down of VOCs throughout the day; or 3) morning
 331 emissions of VOCs being higher than in the afternoon. Of course, all three factors could play a
 332 role. Larger concentrations of VOCs in the morning likely lead to increased abundances of O₃
 333 precursors during the day due to the cascade of associated chemical reactions that form HO₂ and
 334 peroxy radicals upon degradation of VOCs, and lead to a rise in O₃ concentrations later in the
 335 day. Figure 4b shows the median values of OFP for the 12 VOCs with the largest contribution to
 336 ozone formation. The MIR values used for these calculations are included in the supplemental
 337 information (Table S1). The equivalent figure showing the OH reactivity and OFP sorted by
 338 mean value (rather than median) is shown in Figure S5.



340 **Figure 4:** (a) VOCs ranked by their reactivity with OH. Values are for the median reactivity of all VOCs. The green
 341 bars represent all observations collected during the flight campaign, the blue bars represent observations collected
 342 during the morning hours, and red bars represent observations collected during the afternoon hours. (b) All VOC

343 observations are ranked by their Ozone Formation Potential (OFP) and sorted by the median OFP. (c) The
344 corresponding median VOC concentration for each VOC species is shown.

345 Isoprene is the 3rd most important VOC according to OH reactivity (Figure 4a) but 10th
346 according to OFP (Figure 4b). Isoprene has a high reactivity and MIR value indicating that under
347 the right circumstances, isoprene can contribute significantly to the production of O₃ (Table S1).
348 In our can samples, the concentration of isoprene is much lower than that of many other VOCs,
349 so the contribution of isoprene to O₃ formation is relatively small (Figure 4b) despite this large
350 reactivity. Based on these observations collected during the spring in the NYC/CT/LIS region, it
351 is likely that strong biogenic sources of isoprene such as oak trees are not yet fully leafed and
352 therefore are not emitting enough isoprene to compete with anthropogenic VOC sources. This
353 means the relative concentration of biogenic isoprene is low compared to other anthropogenic
354 VOCs.

355 As discussed above, our FOAM box-model calculations showed that only ~5% of the
356 HCHO produced in the model is directly related to isoprene concentrations, while 25-30% is due
357 to propylene. This finding means that regardless of the concentration of HCHO, the dominant
358 atmospheric source appears to be the oxidation of propylene and likely other anthropogenic
359 VOCs, rather than isoprene. Without direct measurement of HCHO, a complete understanding of
360 VOC oxidation in the atmosphere is not possible. Therefore, follow-up studies using in situ
361 aircraft observations of HCHO are necessary to provide context for the VOC observations and
362 provide a means to assess the validity of VOC representation in chemical mechanisms used in
363 models.

364 Atmospheric lifetimes of VOCs are critical components of their respective contributions
365 to local O₃ formation. Propylene, mainly emitted from polypropylene production, petrochemical
366 facilities and diesel exhaust [Buzcu and Fraser, 2006; Hocking, 2005; Liu et al., 2008], has a

367 daytime atmospheric lifetime of ~5 hours [Atkinson, 2000; Warneke *et al.*, 2004]. Isopentane has
368 a longer lifetime of ~1 day and is emitted mainly from gasoline evaporation and engine exhaust
369 [Buzcu and Fraser, 2006; Liu *et al.*, 2008]. Isoprene which is biogenically produced, and HCHO
370 mainly generated from secondary reactions in the atmosphere but also produced from
371 petrochemical facilities and vehicle emissions [Lin *et al.*, 2012; Parrish *et al.*, 2012], have
372 atmospheric lifetimes of ~1-2 hours [Atkinson, 2000; Warneke *et al.*, 2004].

373 As the day proceeds and air is convected and advected off the surface, isoprene is quickly
374 removed by reaction with OH resulting in longer-lived anthropogenic VOCs becoming more
375 important contributors to the local and transported O₃ formation. Isoprene also reacts quickly
376 with NO₃ at night, resulting in consistent lifetimes for daytime and nighttime chemistry.
377 Propylene has a longer nighttime lifetime of ~16 hours as its reaction with NO₃ is much slower
378 than that of isoprene. This allows for propylene to buildup up in the atmosphere at night while
379 isoprene is removed [Warneke *et al.*, 2004]. Therefore, as isoprene is quickly removed during all
380 times of the day, longer-lived anthropogenic VOCs can become more important contributors to
381 local and transported O₃ formation.

382

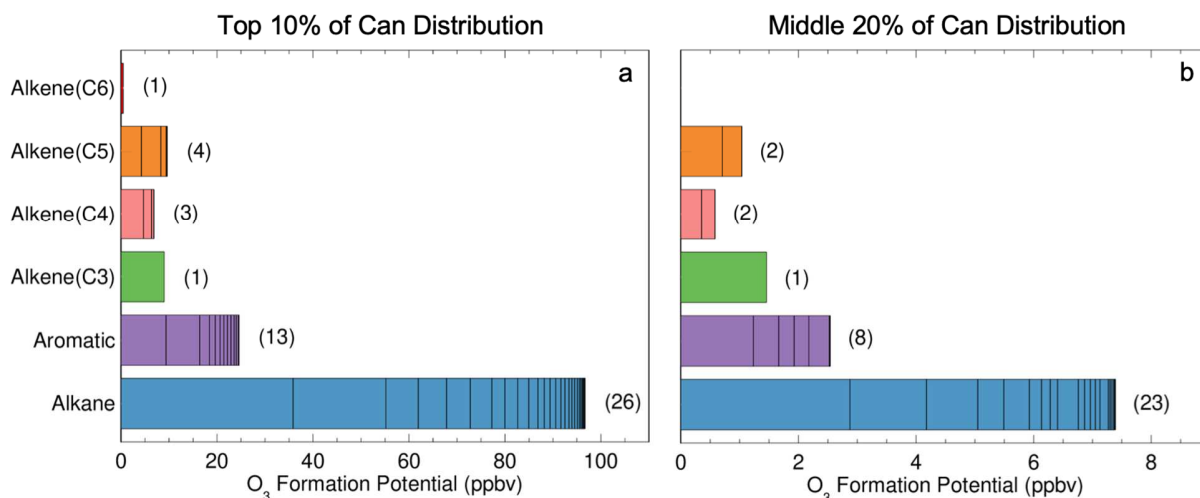
383 **3.3: Ozone Formation Potential**

384 All VOCs measured during this campaign were also classified by their chemical
385 classification category, and the resulting OFP for each category is shown in Figure 5. This
386 analysis does not include HCHO as it was not measured. The measurements of each VOC from
387 the 35 cans collected are first sorted by mixing ratio (ppbv), from highest to lowest. The average
388 of the top 10% of measurements for each VOC is found, representing an average of the largest
389 values measured during this campaign. The OFP for each VOC is then calculated using the top

390 10% average described above, and classified into one of the following categories: alkane,
391 aromatic, alkene (c3), alkene (c4), alkene (c5), alkene (c6), as shown in Figure 5a. This process
392 is repeated using the average of the middle 20% of the can measurements, and shown in Figure
393 5b. The bars in Figure 5 show the O₃ formation potential due to the various VOC categories for
394 the worst case (top 10%) and median case (middle 20%) for May 17 and 18, 2017 in the
395 NYC/LIS region. The values reported in Figure 5 are similar to Figure 12 of *Kleinman et al.*
396 [2005], who examined the reactivity associated with each VOC category for multiple
397 metropolitan regions, including NYC. For the RAMMPP observations, the flight tracks are more
398 directly downwind of NYC along LIS and more VOCs are measured than reported by *Kleinman*
399 *et al.* [2005] from the NARSTO-NE campaign conducted in 1996 [*Solomon et al.*, 2000].

400 The bar for each VOC group in Figure 5 has black lines indicating the OFPs for the
401 individual contributing species in the category (numerical values are listed in the supplementary
402 material Table S5), and the total number of contributing species to the group is shown in
403 parentheses at the end of the bar. The OFP for the alkane group is comparatively large because
404 26 alkanes were measured (over half of all VOCs measured). Isopentane is the most prominent
405 VOC in this category. Propylene is the only VOC in the alkene (c3) category and accounts for
406 significant amounts of O₃ formation potential, while isoprene is in the alkene (c5) category and
407 makes a relatively small contribution to OFP (see Table S5 for numerical values).

NYC Long Island Sound Can Data Grouped by VOC Category



408

409 **Figure 5:** Ozone formation potential for the categories of hydrocarbons for the (a) top 10% of measurements and the
410 (b) middle 20% of measurements for all VOCs measured throughout the RAMMPP 2017 flight campaign. Isoprene
411 is the first section of the C5 bars in (a) and the second section in (b). Analysis is similar to Figure 12 from *Kleinman*
412 *et al.* [2005]. OFP values for each VOC are listed in Table S5.

413 The results described above provide evidence that anthropogenic VOCs are playing a
414 very important role in the photochemical production of O₃ over CT and LIS. However, given the
415 sample size and short duration of these flights, further *in situ* observations of air both at the
416 surface and aloft, at various times throughout the O₃ season, along with measurements of HCHO,
417 are needed to confirm this result is representative of general conditions in the NYC metropolitan
418 region. Continued monitoring of important species like NO_x, O₃, and VOCs in the LIS region
419 would improve confidence in the conclusion that implementing emission controls that target the
420 most important anthropogenically produced VOCs would improve air quality in the
421 NYC/LIS/CT region.

422

423 **4: Conclusions**

424 Aircraft measurements of VOCs obtained during May 17 and 18, 2017 over the New
425 York City (NYC) metropolitan area, Long Island Sound (LIS), and Connecticut (CT) suggest

426 anthropogenic VOCs contribute significantly to photochemical O₃ production. A total of 35
427 canisters were collected and 50 VOC species were quantified. Our analysis indicates the relative
428 importance of anthropogenic VOCs like isopentane and propylene (propene) to the formation of
429 O₃ is larger than that of isoprene, a mainly biogenic VOC. This finding may be due to the short
430 lifetime of isoprene (~1-2 hours) compared to many anthropogenic VOCs and/or minimal
431 isoprene emissions in the NYC metro region at this time of year, resulting in the greater relative
432 impact of longer-lived anthropogenic VOCs to O₃ formation.

433 Since formaldehyde (HCHO) was not measured during these flights, we use a 0-D box
434 model to estimate the percentage of HCHO produced by propylene and isoprene based on their
435 observed concentrations. Only ~5% of the HCHO produced in the model is due to the breakdown
436 of isoprene, while 25-30% is due to propylene degradation. This result shows that the likely
437 sources of HCHO in the atmosphere for this flight campaign are predominantly anthropogenic in
438 nature, and not from isoprene.

439 Air circulating aloft within the Planetary Boundary Layer (PBL) cannot quickly replenish
440 with VOCs and other pollutants emitted from the surface. This isolation can create an O₃
441 photochemical production regime that differs from the surface, complicating our understanding
442 of O₃ production within the PBL. Conditions like this may exist over bodies of water and during
443 early spring and late fall months when biogenic VOC sources are significantly reduced.
444 Therefore, the longer-lived anthropogenic VOCs and other anthropogenic pollutants remain in
445 these air parcels and are carried downwind contributing to poor air quality in relatively distant
446 regions, particularly when transported over bodies of water such as LIS.

447 The data collected during this measurement campaign also provide evidence for the
448 diurnal variation of the concentration of mid-PBL VOCs. When the VOCs are sorted by time of

449 day, morning VOC reactivity (blue bars in Figure 4a) is much greater than afternoon reactivity
450 (red bars in Figure 4a). This finding indicates that in the NYC metropolitan region and areas
451 downwind, VOCs emitted overnight and in early morning hours when photochemistry is inactive
452 and isoprene is scavenged by NO_3 [Warneke *et al.*, 2004] may contribute to a build-up of
453 anthropogenic VOCs near the surface. As the sun rises and photochemistry resumes, the high
454 morning concentrations of VOCs will form HO_2 and peroxy radicals, which contribute to high
455 afternoon concentrations of O_3 . As atmospheric conditions evolve from VOC-sensitive in the
456 morning hours to NO_x -sensitive in the afternoon [Chen and Brune, 2012; Mao *et al.*, 2010], the
457 atmosphere likely transitions through conditions where the MIR and OFP values discussed in
458 this work are applicable, producing the maximum amount of O_3 . Therefore, controlling the
459 emission of anthropogenic VOCs, especially in the early morning and overnight hours, may help
460 reduce the maximum concentrations of O_3 that are usually reached in the afternoon.

461 Previous studies indicate near-surface air over large bodies of water like LIS or Lake
462 Michigan (downwind from Chicago) is likely VOC-sensitive due to stagnant meteorology,
463 heavily polluted air, and suppressed PBL heights [Sillman *et al.*, 1993]. Therefore, surface
464 measurements of VOCs and NO_x may not fully explain photochemical O_3 production for
465 conditions aloft (250 to ~1500 m) [Kleinman, 2000; Sillman, 1999] in air parcels depleted in
466 isoprene over these bodies of water. The O_3 production from anthropogenic VOCs could have
467 important implications for air quality attainment strategies for municipalities downwind of
468 metropolitan areas such as NYC, where the vertical mixing of the boundary layer can bring
469 VOCs and any O_3 produced in the pollution plume down to the surface [Sillman, 1999; 2002;
470 Sillman *et al.*, 1993]. Importantly, surface monitoring may not measure the VOCs most critical to
471 air quality aloft. As industry develops, new technologies are implemented, and societal activity

472 changes, the atmospheric composition of anthropogenic VOCs will likely evolve. Therefore,
473 continued monitoring of VOCs at the surface and aloft is critical to understanding the
474 complicated photochemistry that leads to O₃ production in areas designated as in non-attainment
475 of the NAAQS O₃ standard like the NYC region.

476 While many important VOCs were measured during LISTOS, the observations not
477 exhaustive. The data from the May 17 and 18 2017 flights did not include ethylene, another
478 important and abundant VOC with both anthropogenic and biogenic sources that contributes
479 significantly to HCHO formation [*Franco et al.*, 2022]. Ethylene is measured at surface PAMS
480 monitors, however the first concentrations reported in 2017 are for May 31, almost two weeks
481 after the data discussed in this paper was collected. Flights conducted as part of the LISTOS field
482 mission in 2019 report measurements of ethylene and HCHO but analysis of these cans is
483 beyond the scope of this paper. As is evidenced by *McDonald et al.* [2018] and *Coggon et al.*
484 [2021], future field missions should expand measurement capabilities beyond the well-monitored
485 primary VOCs to also include Volatile Chemical Products (VCPs) and terpenes. These missing
486 classes of VOCs have been shown to impact O₃ production in urban regions, specifically the
487 New York metropolitan region and are critical to improving modeling capabilities in these
488 regions. Gathering direct measurements of these species will provide comparison opportunities
489 for large air quality models when emission inventories are further speciated to account for these
490 VOCs. Moving forward, the air quality community must continue using a multifaceted approach,
491 including satellite observations such as those to be provided by the upcoming Tropospheric
492 Emissions: Monitoring of Pollution (TEMPO) mission [*Zoogman et al.*, 2017], surface
493 monitoring, and aircraft measurements of various VOCs and VCPs to understand the
494 complexities of atmospheric O₃ formation.

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512 All original data related to this paper can be downloaded from the NASA LISTOS archive
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514 summary of the data used in this analysis are provided in the supplementary information.
515

516 **References**

- 517 AirNow (Accessed 2018), AirNow, edited, p. <https://www.airnow.gov/>.
518
- 519 Atkinson, R. (2000), Atmospheric chemistry of VOCs and NO_x, *Atmospheric Environment*, 34(12-14), 2063-2101,
520 doi:Doi.10.1016/S1352-2310(99)00460-4.
521
- 522 Baker, A. K., A. J. Beyersdorf, L. A. Doezema, A. Katzenstein, S. Meinardi, I. J. Simpson, D. R. Blake, and F.
523 Sherwood Rowland (2008), Measurements of nonmethane hydrocarbons in 28 United States cities,
524 *Atmospheric Environment*, 42(1), 170-182, doi:10.1016/j.atmosenv.2007.09.007.
525
- 526 Buzcu, B., and P. Fraser (2006), Source Identification and Apportionment of Volatile Organic Compounds in
527 Houston, TX, *Atmospheric Environment*, 40, 2385-2400.
528
- 529 Caicedo, V., R. Delgado, W. T. Luke, X. Ren, P. Kelley, P. R. Stratton, R. R. Dickerson, T. A. Berkoff, and G.
530 Gronoff (2021), Observations of bay-breeze and ozone events over a marine site during the OWLETS-2
531 campaign, *Atmospheric Environment*, 263, doi:10.1016/j.atmosenv.2021.118669.
532
- 533 Carter, W. P. L. (1994), Development of Ozone Reactivity Scales for Volatile Organic Compounds, *Journal of Air*
534 *& Waste Management Association*, 44(7), 881-899, doi:10.1080/1073161X.1994.10467290.
535
- 536 Carter, W. P. L. (2010), Updated Maximum Incremental Reactivity Scale and Hydrocarbon Bin Reactivities for
537 Regulatory Applications, *California Air Resources Board Contract 07-339*,
538 <https://www.arb.ca.gov/regact/2009/mir2009/mir2010.pdf>.
539
- 540 Chen, S., and W. H. Brune (2012), Global sensitivity analysis of ozone production and O₃-NO_x-VOC limitation
541 based on field data, *Atmospheric Environment*, 55, 288-296, doi:10.1016/j.atmosenv.2012.03.061.
542
- 543 Chen, X., et al. (2019), On the sources and sinks of atmospheric VOCs: an integrated analysis of recent aircraft
544 campaigns over North America, *Atmospheric Chemistry and Physics*, 19(14), 9097-9123, doi:10.5194/acp-
545 19-9097-2019.
546
- 547 Coggon, M. M., et al. (2021), Volatile chemical product emissions enhance ozone and modulate urban chemistry,
548 *Proc Natl Acad Sci U S A*, 118(32), doi:10.1073/pnas.2026653118.
549
- 550 EPA (1970), The Clean Air Act, edited, Title 42, Chapter 85 United States Congress §7401 et seq.,
551 <https://www.epa.gov/laws-regulations/summary-clean-air-act>.
552
- 553 EPA (1990), Clean Air Act Amendments of 1990, in *Public Law*, edited.
554
- 555 EPA (2016a), NAAQS Table, *Criteria Air Pollutants*, <https://www.epa.gov/criteria-air-pollutants/naaqs-table>.
556
- 557 EPA (2016b), Technical Assistance Document for the National Air Toxics Trends Stations Program, *Revision 3*,
558 [https://www3.epa.gov/ttnamti1/files/ambient/airtox/NATTS%20TAD%20Revision%203_FINAL%22Oct](https://www3.epa.gov/ttnamti1/files/ambient/airtox/NATTS%20TAD%20Revision%203_FINAL%22Oct%20ober%202016.pdf)
559 [ober%202016.pdf](https://www3.epa.gov/ttnamti1/files/ambient/airtox/NATTS%20TAD%20Revision%203_FINAL%22Oct%20ober%202016.pdf).
560
- 561 Federal_Register (2022a), Determinations of Attainment by the Attainment Date, Extensions of the Attainment
562 Date, and Reclassification of Areas Classified as Marginal for the 2015 Ozone National Ambient Air
563 Quality Standards, *Federal Register*, 87(68097).
564
- 565 Federal_Register (2022b), Determinations of Attainment by the Attainment Date, Extensions of the Attainment
566 Date, and Reclassification of Areas Classified as Serious for the 2008 Ozone National Ambient Air Quality
567 Standards, *Federal Register*, 87(60926).
568
- 569 Franco, B., L. Clarisse, M. Van Damme, J. Hadji-Lazaro, C. Clerbaux, and P. F. Coheur (2022), Ethylene industrial
570 emitters seen from space, *Nat Commun*, 13(1), 6452, doi:10.1038/s41467-022-34098-8.

571
572 Han, Z. S., J. E. González-Cruz, H. N. Liu, D. Melecio-Vázquez, H. Gamarro, Y. H. Wu, F. Moshary, and R.
573 Bornstein (2022), Observed sea breeze life cycle in and around NYC: Impacts on UHI and ozone patterns,
574 *Urban Climate*, 42, doi:10.1016/j.uclim.2022.101109.
575

576 Harkey, M., T. Holloway, E. J. Kim, K. R. Baker, and B. Henderson (2021), Satellite Formaldehyde to Support
577 Model Evaluation, *J Geophys Res Atmos*, 126(4), doi:10.1029/2020jd032881.
578

579 He, H., et al. (2013), Trends in emissions and concentrations of air pollutants in the lower troposphere in the
580 Baltimore/Washington airshed from 1997 to 2011, *Atmospheric Chemistry and Physics*, 13(15), 7859-7874,
581 doi:10.5194/acp-13-7859-2013.
582

583 Hembeck, L., H. He, T. P. Vinciguerra, T. P. Canty, R. R. Dickerson, R. J. Salawitch, and C. Loughner (2019),
584 Measured and modelled ozone photochemical production in the Baltimore-Washington airshed,
585 *Atmospheric Environment: X*, 2, doi:10.1016/j.aeaoa.2019.100017.
586

587 Hocking, M. B. (2005), *Handbook of Chemical Technology and Pollution Control*, 830 pp., Academic Press,
588 doi:10.1016/B978-0-12-088796-5.X5000-5.
589

590 Kansal, A. (2009), Sources and Reactivity of NMHCs and VOCs in the Atmosphere: A Review, *Journal of*
591 *Hazardous Materials*, 166, 17-26, doi:10.1016/j.jhazmat.2008.11.048.
592

593 Kleinman, L. I. (2000), Ozone process insights from field experiments - part II: Observation-based analysis for
594 ozone production, *Atmospheric Environment*, 34(12-14), 2023-2033, doi:Doi 10.1016/S1352-
595 2310(99)00457-4.
596

597 Kleinman, L. I., P. H. Daum, Y. N. Lee, L. J. Nunnermacker, S. R. Springston, J. Weinstein-Lloyd, and J. Rudolph
598 (2005), A comparative study of ozone production in five U.S. metropolitan areas, *110(D2)*,
599 doi:10.1029/2004jd005096.
600

601 Li, H., et al. (2020), Terpenes and their oxidation products in the French Landes forest: insights from Vocus PTR-
602 TOF measurements, *Atmospheric Chemistry and Physics*, 20(4), 1941-1959, doi:10.5194/acp-20-1941-
603 2020.
604

605 Lin, Y. C., J. J. Schwab, K. L. Demerjian, M.-S. Bae, W.-N. Chen, Y. Sun, Q. Zhang, H.-M. Hung, and J. Perry
606 (2012), Summertime formaldehyde observations in New York City: Ambient levels, sources and its
607 contribution to HOx radicals, *Journal of Geophysical Research: Atmospheres*, 117(D8), n/a-n/a,
608 doi:10.1029/2011jd016504.
609

610 Liu, Y., M. Shao, L. Fu, S. Lu, L. Zeng, and D. Tang (2008), Source Profiles of Volatile Organic Compounds (VOCs)
611 Measured in China: Part 1, *Atmospheric Environment*, 42, 6247-6260.
612

613 Loughner, C. P., et al. (2014), Impact of Bay-Breeze Circulations on Surface Air Quality and Boundary Layer
614 Export, *Journal of Applied Meteorology and Climatology*, 53(7), 1697-1713, doi:10.1175/Jamc-D-13-
615 0323.1.
616

617 Luecken, D. J., W. T. Hutzell, M. L. Strum, and G. A. Pouliot (2012), Regional sources of atmospheric
618 formaldehyde and acetaldehyde, and implications for atmospheric modeling, *Atmospheric Environment*, 47,
619 477-490, doi:10.1016/j.atmosenv.2011.10.005.
620

621 Luecken, D. J., S. L. Napelenok, M. Strum, R. Scheffe, and S. Phillips (2018), Sensitivity of Ambient Atmospheric
622 Formaldehyde and Ozone to Precursor Species and Source Types Across the United States, *Environ Sci*
623 *Technol*, 52(8), 4668-4675, doi:10.1021/acs.est.7b05509.
624

625 Lv, S.-Y., Q.-Y. Liu, Y.-X. Zhao, M.-Q. Zhang, L.-X. Jiang, and S.-G. He (2019), Formaldehyde Generation in
626 Photooxidation of Isoprene on Iron Oxide Nanoclusters, *The Journal of Physical Chemistry C*, 123(8),
627 5120-5127, doi:10.1021/acs.jpcc.8b12471.
628

629 MaineDEP (2017), Standard Operating Procedure for the Determination of Volatile Organic Compounds (VOCs) in
630 Air Collected in Specially-Prepared Canisters and Analyzed by Gas Chromatography/Mass Spectrometry
631 (GC/MS), (*available upon request*), 52.
632

633 Mao, J., et al. (2010), Atmospheric oxidation capacity in the summer of Houston 2006: Comparison with summer
634 measurements in other metropolitan studies, *Atmospheric Environment*, 44(33), 4107-4115,
635 doi:10.1016/j.atmosenv.2009.01.013.
636

637 Mazzuca, G. M., K. E. Pickering, R. D. Clark, C. P. Loughner, A. Fried, D. C. S. Zweers, A. J. Weinheimer, and R.
638 R. Dickerson (2017), Use of tethered sonde and aircraft profiles to study the impact of mesoscale and
639 microscale meteorology on air quality, *Atmospheric Environment*, 149, 55-69,
640 doi:10.1016/j.atmosenv.2016.10.025.
641

642 McDonald, B. C., et al. (2018), Volatile chemical products emerging as largest petrochemical source of urban
643 organic emissions, *Science*, 359(6377), 760-764, doi:10.1126/science.aag0524.
644

645 Parrish, D. D., et al. (2012), Primary and secondary sources of formaldehyde in urban atmospheres: Houston Texas
646 region, *Atmospheric Chemistry and Physics*, 12(7), 3273-3288, doi:10.5194/acp-12-3273-2012.
647

648 Ren, X., et al. (2018), Methane Emissions From the Baltimore-Washington Area Based on Airborne Observations:
649 Comparison to Emissions Inventories, *Journal of Geophysical Research: Atmospheres*, 123(16), 8869-
650 8882, doi:doi:10.1029/2018JD028851.
651

652 Schwantes, R. H., et al. (2020), Comprehensive isoprene and terpene gas-phase chemistry improves simulated
653 surface ozone in the southeastern US, *Atmospheric Chemistry and Physics*, 20(6), 3739-3776,
654 doi:10.5194/acp-20-3739-2020.
655

656 Seltzer, K. M., E. Pennington, V. Rao, B. N. Murphy, M. Strum, K. K. Isaacs, and H. O. T. Pye (2021), Reactive
657 organic carbon emissions from volatile chemical products, *Atmos Chem Phys*, 21(6), 5079-5100,
658 doi:10.5194/acp-21-5079-2021.
659

660 Sillman, S. (1999), The relation between ozone, NO_x and hydrocarbons in urban and polluted rural environments,
661 *Atmospheric Environment*, 33(12), 1821-1845, doi:Doi 10.1016/S1352-2310(98)00345-8.
662

663 Sillman, S. (2002), Some theoretical results concerning O₃-NO_x-VOC chemistry and NO_x-VOC indicators, *Journal*
664 *of Geophysical Research*, 107(D22), doi:10.1029/2001jd001123.
665

666 Sillman, S., P. J. Samson, and J. M. Masters (1993), Ozone production in urban plumes transported over water:
667 Photochemical model and case studies in the northeastern and midwestern United States, *Journal of*
668 *Geophysical Research*, 98(D7), doi:10.1029/93jd00159.
669

670 Solomon, P., E. Cowling, G. Hidy, and C. Furiness (2000), Comparison of scientific findings from major ozone field
671 studies in North America and Europe, *Atmospheric Environment*, 34(12-14), 1885-1920,
672 doi:10.1016/s1352-2310(99)00453-7.
673

674 Stauffer, R. M., A. M. Thompson, D. K. Martins, R. D. Clark, D. L. Goldberg, C. P. Loughner, R. Delgado, R. R.
675 Dickerson, J. W. Stehr, and M. A. Tzortziou (2015), Bay breeze influence on surface ozone at Edgewood,
676 MD during July 2011, *J Atmos Chem*, 72(3-4), 335-353, doi:10.1007/s10874-012-9241-6.
677

678 Stein, A. F., R. R. Draxler, G. D. Rolph, B. J. B. Stunder, M. D. Cohen, and F. Ngan (2015), NOAA's HYSPLIT
679 Atmospheric Transport and Dispersion Modeling System, *Bulletin of the American Meteorological Society*,
680 96(12), 2059-2077, doi:10.1175/bams-d-14-00110.1.

681
682 Tang, J. H., L. Y. Chan, C. Y. Chan, Y. S. Li, C. C. Chang, S. C. Liu, D. Wu, and Y. D. Li (2007), Characteristics
683 and diurnal variations of NMHCs at urban, suburban, and rural sites in the Pearl River Delta and a remote
684 site in South China, *Atmospheric Environment*, 41(38), 8620-8632, doi:10.1016/j.atmosenv.2007.07.029.
685
686 Taubman, B. F., L. T. Marufu, C. A. Piety, B. G. Doddridge, J. W. Stehr, and R. R. Dickerson (2004), Airborn
687 Characterization of the Chemical, Optical, and Meteorological Properties and Origins of a Combined
688 Ozone-Haze Episode over the Eastern United States, *Journal of Atmospheric Sciences*, 61, 1781-1793.
689
690 Torres-Vazquez, A., J. Pleim, R. Gilliam, and G. Pouliot (2022), Performance Evaluation of the Meteorology and
691 Air Quality Conditions From Multiscale WRF-CMAQ Simulations for the Long Island Sound
692 Tropospheric Ozone Study (LISTOS), *Journal of Geophysical Research: Atmospheres*, 127(5),
693 doi:10.1029/2021jd035890.
694
695 Venecek, M. A., W. P. L. Carter, and M. J. Kleeman (2018), Updating the SAPRC Maximum Incremental
696 Reactivity (MIR) scale for the United States from 1988 to 2010, *J Air Waste Manag Assoc*, 68(12), 1301-
697 1316, doi:10.1080/10962247.2018.1498410.
698
699 Warneke, C., et al. (2004), Comparison of daytime and nighttime oxidation of biogenic and anthropogenic VOCs
700 along the New England coast in summer during New England Air Quality Study 2002, *Journal of*
701 *Geophysical Research*, 109(D10), doi:10.1029/2003jd004424.
702
703 Wolfe, G. M., et al. (2016a), Formaldehyde production from isoprene oxidation across NOx regimes, *Atmos Chem*
704 *Phys*, 16(4), 2597-2610, doi:10.5194/acp-16-2597-2016.
705
706 Wolfe, G. M., M. R. Marvin, S. J. Roberts, K. R. Travis, and J. Liao (2016b), The Framework for 0-D Atmospheric
707 Modeling (FOAM) v3.1, *Geoscientific Model Development*, 9(9), 3309-3319, doi:10.5194/gmd-9-3309-
708 2016.
709
710 Wu, Y., et al. (2021), Synergistic aircraft and ground observations of transported wildfire smoke and its impact on
711 air quality in New York City during the summer 2018 LISTOS campaign, *Sci Total Environ*, 773, 145030,
712 doi:10.1016/j.scitotenv.2021.145030.
713
714 Zhu, L., et al. (2016), Observing atmospheric formaldehyde (HCHO) from space: validation and intercomparison of
715 six retrievals from four satellites (OMI, GOME2A, GOME2B, OMPS) with SEAC(4)RS aircraft
716 observations over the Southeast US, *Atmos Chem Phys*, 16(21), 13477-13490, doi:10.5194/acp-16-13477-
717 2016.
718
719 Zoogman, P., et al. (2017), Tropospheric Emissions: Monitoring of Pollution (TEMPO), *J Quant Spectrosc Radiat*
720 *Transf*, 186, 17-39, doi:10.1016/j.jqsrt.2016.05.008.
721
722