

1 **Purified meta-Cresol Purple dye perturbation: how it influences spectrophotometric pH**
2 **measurements**

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17 **Abstract**

18 Ocean acidification, a phenomenon of seawater pH decreasing due to increasing atmospheric
19 CO₂, has a global effect on seawater chemistry, marine biology, and ecosystems. Ocean
20 acidification is a gradual and global long-term process, the study of which demands high-quality
21 pH data. The spectrophotometric technique is capable of generating accurate and precise pH
22 measurements but requires adding an indicator dye that perturbs the sample original pH. While
23 the perturbation is modest in well-buffered seawater, applications of the method in environments
24 with lower buffer capacity such as riverine, estuarine, sea-ice meltwater and lacustrine
25 environments are increasingly common, and uncertainties related to larger potential dye
26 perturbations need further evaluation. In this paper, we assess the effect of purified meta-Cresol
27 Purple (mCP) dye addition on the sample pH and how to correct for this dye perturbation. We
28 conducted numerical simulations by incorporating mCP speciation into the MATLAB CO2SYS
29 program to examine the changes in water sample pH caused by the dye addition and to reveal the

30 dye perturbation mechanisms. Then, laboratory experiments were carried out to verify the
31 simulation results. The simulations suggest that the dye perturbation on sample pH is a result of
32 total alkalinity (TA) contributions from the indicator dye and chemical equilibrium shifts that are
33 related to both the water sample properties (pH, TA, and salinity) and the indicator dye solution
34 properties (pH and solvent matrix). The laboratory experiments supported the simulation results;
35 the same dye solution can lead to different dye perturbations in water samples with different pH,
36 TA, and salinity values. The modeled adjustments agreed well with the empirically determined
37 adjustments for salinities > 5 , but it showed greater errors for lower salinities with disagreements
38 as large as 0.005 pH units. Adjustments are minimized when the pH and salinity of the dye are
39 matched to the sample. When the dye is used over a wide range of salinity, we suggest that it
40 should be prepared in deionized water to minimize the dye perturbation effect on pH in the
41 fresher sample waters with less well-constrained perturbation adjustments. We also suggest that
42 the dye perturbation correction should be based on double dye addition experiments performed
43 over a wide range of pH, TA, and salinity. Otherwise, multiple volume dye addition experiments
44 are recommended for each sample to determine the dye perturbation adjustment. We further
45 create a MATLAB function *dyperturbation.m* that calculates the expected dye perturbation.
46 This function can be used to validate empirically-derived adjustments or in lieu of empirical
47 adjustments if dye addition experiments are unfeasible (e.g., for historical data). This study of
48 dye perturbation evaluation and correction will improve the accuracy of the pH data, necessary
49 for monitoring the long-term anthropogenic-driven changes in the seawater carbonate system.

50

51 **1. Introduction**

52 Atmospheric carbon dioxide (CO₂) has been dramatically increasing since the Industrial
53 Revolution, mainly due to fossil fuel burning and deforestation
54 (<https://www.esrl.noaa.gov/gmd/ccgg/trends/>). The increasing atmospheric CO₂ levels are
55 tempered by the absorption of CO₂ by the ocean and the biosphere. However, oceanic CO₂
56 uptake causes ocean acidification, a seawater pH decreasing phenomenon (Doney et al., 2009).
57 Ocean acidification has a vital effect on the seawater chemistry, marine organisms and global
58 ecosystems (Orr et al., 2005; Doney et al., 2009; Cai et al., 2011; Andersson and Gledhill, 2013;

59 Waldbusser and Salisbury, 2014; Wanninkhof et al., 2015; Kwiatkowski and Orr, 2018;
60 Landschützer et al., 2018). Ocean acidification is a subtle, long-term process, which requires
61 high-quality data to better evaluate and understand it (Orr et al., 2018). The Global Ocean
62 Acidification Observing Network (GOA-ON) suggested that the measurement of pH should
63 achieve an accuracy of 0.02 pH units to assess *Weather* changes, i.e., spatial and short-term
64 variations, while the *Climate* goal, focusing on deciphering decadal trends, requires pH data with
65 an accuracy of 0.003 pH units.

66 The spectrophotometric pH method can be an accurate and precise technique for determining
67 seawater pH (Lai et al., 2016; Liu et al., 2011; Ma et al., 2019; Mosley et al., 2004). This method
68 relies on adding pH-sensitive indicator dye into water samples. Since the dye changes color with
69 pH, the pH of the sample can be determined through absorbance spectra (Clayton and Byrne,
70 1993). There are many pH-sensitive indicator dyes available for spectrophotometric pH
71 measurements, such as meta-Cresol Purple (mCP), thymol blue, and phenol red (Clayton and
72 Byrne, 1993; Liu et al., 2011; Mosley et al., 2004; Patsavas et al., 2013; Soli et al., 2013; Zhang
73 and Byrne, 1996). One of the most commonly used and well-characterized dyes for
74 spectrophotometric pH measurements is mCP. Since Clayton and Byrne (1993) formalized
75 spectrophotometric pH measurements for seawater using mCP, the reproducibility and accuracy
76 of this method have been improved dramatically. The uncertainty sources of spectrophotometric
77 pH measurements include factors such as dye perturbation, dye purity, instrument parameters
78 (wavelength accuracy and absorbance errors), temperature control, pressure (DeGrandpre et al.,
79 2014). The dye purity uncertainty has been addressed for mCP by purification through high-
80 performance liquid chromatography (HPLC), and its physicochemical properties are now well
81 characterized for a wide range of temperature, salinity (S), and pressure (Liu et al., 2011; Müller
82 and Rehder, 2018a; Soli et al., 2013). Beyond these, an automated spectrophotometric system
83 was designed (Carter et al., 2013) to achieve precise, reproducible and fast measurements with
84 consistent temperature and pressure. These methodological improvements have enhanced the
85 repeatability of spectrophotometric pH measurements to 0.0002 pH units (Liu et al., 2011) and
86 decreased its combined uncertainty to 0.005–0.01 pH units (Bockmon and Dickson, 2015; Carter
87 et al., 2013).

88 Despite recent improvements, our understanding of the pH perturbation caused by the dye
89 addition remains limited. The characterization of mCP is performed in strongly buffered
90 solutions, where the dye addition does not cause an appreciable pH perturbation of the buffer
91 solution (Liu et al., 2011). But this is not the case in seawater or brackish samples. To correct the
92 perturbation of sample pH due to dye addition, *The Guide to Best Practices for Ocean CO₂*
93 *Measurements* recommends a simple, empirical perturbation correction procedure by adding
94 double volume indicator (Dickson et al., 2007). This procedure works well for open ocean waters
95 using unpurified mCP. Chierici et al. (1999) also suggested an equation to correct for the dye
96 perturbation effect on seawater sample pH based on both seawater sample pH and S, but this
97 approach has not yet been widely applied to routine pH measurements and was developed before
98 the properties of purified mCP were characterized over the full 0-40 S range. None of the recent
99 research that used purified mCP clearly stated how much the addition of purified mCP would
100 influence the measured sample pH over the full 0- 40 S range.

101 Clearly, there is room for improvement in quantifying the pH measurement uncertainty due to
102 the dye perturbation, identifying the perturbation mechanism, and in establishing a strategy to
103 reduce or even eliminate this source of uncertainty. We comprehensively evaluate the impact of
104 dye perturbation using numerically simulated data and laboratory experiments in this paper. The
105 numerical simulations were performed by adding the mCP acid-base speciation equations into
106 the CO2SYS program (Lewis and Wallace, 1998; Van Heuven et al., 2011; Xu et al., 2017). The
107 simulation results allow us to reveal the controlling chemical mechanisms on dye perturbation.
108 We have also carried out laboratory experiments to verify the numerical simulation. Based on
109 these results, we offer solutions regarding the mCP pH perturbation corrections for
110 spectrophotometric pH measurements.

111 **2. Theory and Method**

112 **2.1 Basic theory**

113 The spectrophotometric pH measurements are based on the absorption spectra of samples with a
114 pH-sensitive indicator dye. The dye, usually mCP, is a diprotic acid with acid and base forms
115 that have distinct absorption peaks. In the pH range of natural seawater, mCP behaves as a
116 monoprotic weak acid, so only its second dissociation constant needs to be taken into

117 consideration (Clayton and Byrne, 1993). In this case, the sample pH (on the total scale)
 118 (Dickson, 1993, 1984) can be estimated from the relative concentration of different mCP species
 119 and its second dissociation constant (Clayton and Byrne, 1993),

$$120 \quad \text{pH} = \text{p}K_2 + \log_{10} \frac{[\text{I}^{2-}]}{[\text{HI}^-]} \quad (\text{Equation 1})$$

121 where $[\text{HI}^-]$ and $[\text{I}^{2-}]$ are the concentrations of the monoprotonated and unprotonated species of
 122 the indicator dye, respectively. The concentration ratio $[\text{HI}^-]/[\text{I}^{2-}]$ can be determined by the
 123 spectrophotometric measurements of the absorbance (A). Therefore, sample pH can be
 124 calculated as (Liu et al., 2011),

$$125 \quad \text{pH} = -\log(K_2 e_2) + \log_{10} \left(\frac{R - e_1}{1 - \frac{e_3 \cdot R}{e_2}} \right) \quad (\text{Equation 2})$$

126 where R is the absorbance ratio of mCP at λ_{578} and λ_{434} ; e_1 , e_2 and e_3 are the molar absorptivity
 127 ratios, $e_1 = \frac{578^{\epsilon_{HI}}}{434^{\epsilon_{HI}}}$, $e_2 = \frac{578^{\epsilon_I}}{434^{\epsilon_{HI}}}$, $e_3 = \frac{434^{\epsilon_I}}{434^{\epsilon_{HI}}}$.

128 To minimize the influence of dye addition to the sample pH, an empirical correction (Figure 1)
 129 was developed to determine the ‘true pH’ that would have been observed under no dye addition
 130 (Clayton and Byrne, 1993; Dickson et al., 2007). The empirical method relies on two
 131 assumptions. First, the change in R is linearly related to the volume of dye added (Figure 1, Step
 132 1). Second, the changes in R per volume (mL) dye added ($\Delta R/\Delta V$) is a simple linear function of
 133 R (Equation 4). Following these two assumptions, the dye perturbation can be evaluated
 134 empirically based on double volume dye addition experiments. Some of the seawater samples are
 135 measured twice with different dye addition volumes to derive a relationship between absorbance
 136 changes and dye addition volume (Figure 1, Step 2). The first measurement is the absorbance
 137 ratio (R_1) of the seawater sample with a single addition of dye (V_1 , mL), and the second
 138 absorbance ratio measurement (R_2) is performed after the second addition of dye to the same
 139 seawater sample (V_2 , mL). The changes in R per milliliter of dye added ($\Delta R/\Delta V = (R_2 - R_1)/(V_2 -$
 140 $V_1)$) can be expressed as a function of R_1 ,

$$141 \quad \frac{\Delta R}{\Delta V} = A + B \cdot R_1 \quad (\text{Equation 3})$$

142 Then this relationship is applied to all the samples to obtain the theoretical R values without dye
143 perturbation, being the empirically corrected absorbance ratio R_{EC} (Clayton and Byrne, 1993;
144 Dickson et al., 2007). For any given sample, the measured R_1 can be corrected to R_{EC} as,

$$145 \quad R_{EC} \equiv R_1 - \frac{\Delta R}{\Delta V} \cdot V_1 = R_1 - (A + B \cdot R_1) \cdot V_1 \quad (\text{Equation 4})$$

146 $\Delta R/\Delta V$ determines how much R_1 should be empirically corrected to R_{EC} (Figure 1, Step 3). The
147 pH value calculated from R_1 via Equation 2 is the seawater pH with the dye perturbation, while
148 the pH value calculated from R_{EC} is considered as the ‘true’ seawater pH (Figure1, Step 4). The
149 difference between the dye-perturbed pH and the ‘true’ pH is ΔpH . Some measurement
150 approaches cannot reliably dispense precise volumes of dye and instead use changes in the
151 absorbance at the mCP isosbestic wavelength as a proxy for the volume of dye added (Carter et
152 al., 2013).

153 **2.2 Numerical simulations of the dye perturbation**

154 The pH perturbation caused by the dye addition can be calculated from the known equilibrium
155 chemistry of seawater and the dye. For our calculation of the dye perturbation, we modified
156 CO2SYS for MATLAB (Lewis and Wallace, 1998; Van Heuven et al., 2011; Xu et al., 2017) to
157 a new function *dyeperturbation.m* (<https://github.com/Sheenaxy/Dyeperturbation>). The CO2SYS
158 program is based on the equilibria of carbonate parameters, other weak acid-base species, and
159 other seawater related chemistry parameters. Any two of the four carbonate parameters, total
160 dissolved inorganic carbon (DIC), total alkalinity (TA), pH, and fugacity or partial pressure of
161 CO_2 ($f\text{CO}_2$ or $p\text{CO}_2$), are chosen as inputs and the other two parameters can be calculated. Other
162 input variables include the choices of equilibrium constants, temperature, pressure, S, the
163 concentration of silicate and phosphate and the pH scale. In this work, we used DIC and TA pair
164 to calculate pH.

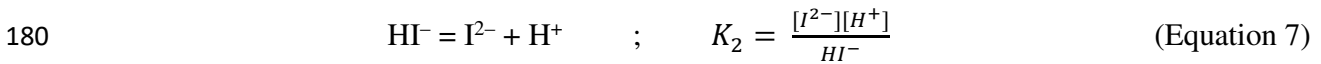
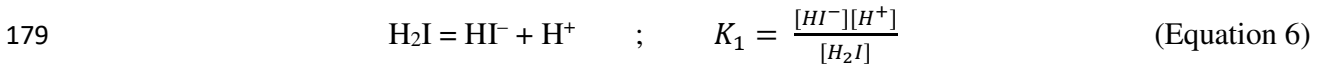
165 In the CO2SYS, DIC is defined as the sum of $[\text{HCO}_3^-]$, $[\text{CO}_3^{2-}]$ and $[\text{H}_2\text{CO}_3]^*$ (by convention
166 $[\text{H}_2\text{CO}_3]^*$ includes both aqueous CO_2 and H_2CO_3). TA is defined as the moles of hydrogen ion
167 equivalent to the excess of proton acceptors over proton donors per kilogram of seawater
168 (Dickson, 1981). The proton acceptors are bases formed from weak acids with $\text{pK} \geq 4.5$ at 25 °C

169 and 0 mol/kg ionic strength, while the proton donors are acids with $pK \leq 4.5$ at 25 °C and 0
 170 mol/kg ionic strength. Thus, TA is defined as,

$$171 \quad TA = CA_{lk} + BA_{lk} + [OH^-] + PA_{lk} + SiAlk - [H^+]_{free} - [HSO_4^-] - [HF^-] \quad (\text{Equation 5})$$

172 where CA_{lk} is the carbonate alkalinity ($CA_{lk} = [HCO_3^-] + 2[CO_3^{2-}]$); BA_{lk} is the borate
 173 alkalinity ($BA_{lk} = [B(OH)_4^-]$); PA_{lk} is the phosphate alkalinity ($PA_{lk} = [HPO_4^{2-}] + 2[PO_4^{3-}] -$
 174 $[H_3PO_4]$); and $SiAlk$ is the silicate alkalinity ($SiAlk = [SiO(OH)_3^-]$). By solving Equation (5)
 175 iteratively we can calculate an equilibrium pH.

176 To test the influence of the dye addition in sample pH, we added the equilibria reactions of mCP
 177 as an additional chemical species into Equation 5. The two-step dissociation equilibria of mCP
 178 are,



181 K_1 is a function of temperature, where $\log K_1 = -782.62/T + 1.1131$ (Liu et al., 2011). When $T =$
 182 298 K , $-\log K_1 = pK_1 = 1.5131$. K_2 is a function of S and temperature. Previous research
 183 considered $p(K_2e_2) = -\log_{10}(K_2e_2)$ as one term to simplify the mCP characterization process (Liu
 184 et al., 2011; Müller and Rehder, 2018b). Here we estimate K_2 as,

$$185 \quad K_2 = \frac{10^{-(p(K_2e_2))}}{e_2} \quad (\text{Equation 8})$$

186 $p(K_2e_2)$ varies from 7.6479 at $S = 35$, $T = 298.15 \text{ K}$ (Liu et al., 2011) to 8.2978 at $S = 0$, $T =$
 187 298.15 K (Lai et al., 2016). Since there is no literature reference for the purified mCP e_2 function
 188 at the full S range ($0 < S < 40$), we linearly interpolated e_2 between $S = 0$ and $S = 40$ at $T =$
 189 298.15 K using $e_2 = 2.306$ for $T = 298.15 \text{ K}$, $S = 0$ (Lai et al., 2016), and $e_2 = 2.22$ for $T = 298.15$
 190 K , $S = 35$ (Clayton and Byrne, 1993) and linearly interpreted them between $S = 0$ and 40 to
 191 simplify the calculation. Note that the e_2 value variation (from $S = 0$ to $S = 35$ at $T = 298.15 \text{ K}$) is
 192 small and has little influence on K_2 value calculated from $p(K_2e_2)$.

193 We define total concentration of indicator dye (TI) as, $TI = [I^{2-}] + [HI^-] + [H_2I]$. Therefore, the
 194 concentration of the different species of the indicator can be rewritten as,

$$195 \quad [I^{2-}] = TI \cdot \frac{K_1 K_2}{[H^+]^2 + [H^+]K_1 + K_1 K_2} \quad (\text{Equation 9})$$

$$196 \quad [HI^-] = TI \cdot \frac{K_1 [H^+]}{[H^+]^2 + [H^+]K_1 + K_1 K_2} \quad (\text{Equation 10})$$

$$197 \quad [H_2I] = TI \cdot \frac{[H^+]^2}{[H^+]^2 + [H^+]K_1 + K_1 K_2} \quad (\text{Equation 11})$$

198 Similar to the carbonate system, the speciation of mCP varies with pH (Figure 2). For a given
 199 pH, we can calculate the concentration of each mCP form.

200 Knowing mCP dissociation constants and the species composition, we can calculate TA
 201 contributed by the indicator addition to the sample. The first dissociation constant K_1 of mCP is
 202 $\sim 1.5 (\leq 4.5)$, and the second dissociation constant is around $7.5 (\geq 4.5)$ so the appropriate zero
 203 level of proton for mCP is HI^- . Thus, the TA of indicator dye solution (TA_I) can be calculated
 204 by,

$$205 \quad TA_I = [I^{2-}]_I - [H_2I]_I + [OH^-] - [H^+]_{\text{free}}. \quad (\text{Equation 12})$$

206 When the indicator solution is added to the seawater sample, it changes the TA of the sample. As
 207 TA mixes conservatively, the TA of a seawater sample with indicator (TA_m) is the mass-
 208 weighted mean TA of the seawater sample (TA_s) and that contributed by the indicator. Thus,
 209 TA_m becomes,

$$210 \quad TA_m = \frac{m_s TA_s + m_I TA_I}{m_s + m_I} = CAlk + BAlk + [OH^-] + PAlk + SiAlk - [H^+]_{\text{free}} - [HSO_4^-] - [HF^-] + IAlk$$

$$211 \quad (\text{Equation 13})$$

212 where $IAlk$ is the indicator dye alkalinity contributed by the mCP in the mixed solution. Note
 213 that $IAlk$ is different from TA_I . Similar to $CAlk$, $BAlk$ and etc., $IAlk$ is defined as,

$$214 \quad IAlk = [I^{2-}]_m - [H_2I]_m = TI_m \cdot \frac{K_1 K_2 - [H^+]^2}{[H^+]^2 + K_1 [H^+] + K_1 K_2} \quad (\text{Equation 14})$$

215 TI_m is the indicator dye concentration in the mixing solution; $[I^2]_m$ (calculated as Equation 9)
216 and $[H_2I]_m$ (calculated as Equation 11) are the concentration of unprotonated and diprotonated
217 mCP species in the mixing solution, respectively. By solving Equation (13) iteratively we can
218 calculate a new equilibrium pH with the indicator dye addition. Note that since all chemical
219 equilibria must be changed with the indicator dye solution addition, all TA components such as
220 CA_{alk} , BA_{alk} , IA_{alk} etc. will change too. We can, therefore, calculate the sample pH after adding
221 the dye using Equation (13) and the original sample pH without dye perturbation using Equation
222 (5). By subtracting the pH before dye addition, the pH changes due to dye addition can be
223 calculated. The theoretical absorbance value R can be calculated from Equation (2) with known
224 pH.

225 In our simulations, we set the pH and the concentration of the dye solution, dye/sample volume
226 ratio to the values in the standard operating procedure (SOP) of *The Guide to Best Practices for*
227 *Ocean CO₂ Measurements* (Dickson et al., 2007). We added 20 μ L of 2.5 mmol/L mCP solution
228 of pH = 8.0 to 15 mL sample (thus final dye concentration in the sample = 3.3 μ mol/L). The SOP
229 did not mention the S of the indicator solution. Here, we set the indicator solution $S = 0$ to match
230 our laboratory experiments. In addition, we simulated the effect of dye perturbation under the
231 dye solution $S = 35$ in Section 4.3.1.

232 The simulation was run with 50000 randomly chosen combinations of DIC and TA to reflect the
233 ranges commonly found in the marine environment. TA is in the range of 300 to 2800 μ mol kg^{-1}
234 (bounding the open ocean TA range using a value appropriate for a low and high TA riverine
235 endmember, respectively (Cai et al., 2008)); DIC is between 240 μ mol kg^{-1} and 4200 μ mol kg^{-1} ;
236 S is within 0–40. Although there is no general relationship among the chosen DIC, TA and S , we
237 constrained the TA/DIC ratio to be within 0.8–1.5 to be more realistic for marine and estuarine
238 environments. All simulations were performed under 25 °C. The selected dissociation constants,
239 temperature, and pressure inputs are listed in Table 1. We also conducted three groups of
240 simulation tests that match our laboratory experiments as described below.

241 **2.3 Laboratory experimental design**

242 Previous research suggested the effect of dye perturbation on sample pH is determined by the
243 sample ionic strength and buffer capacity, yet this has not been experimentally demonstrated

244 (Chierici et al., 1999; Mosley et al., 2004). We designed three groups of laboratory experiments
245 to verify the simulation results on dye perturbation (Figure 3). In the first group of experiments,
246 we used water samples with fixed high S (36.3) and varying TA (about 1100 $\mu\text{mol kg}^{-1}$, 1450
247 $\mu\text{mol kg}^{-1}$, 2000 $\mu\text{mol kg}^{-1}$, 2350 $\mu\text{mol kg}^{-1}$). The second group of experiments was performed
248 using water samples with fixed high TA ($\sim 2350 \mu\text{mol kg}^{-1}$) and varied S (about 0, 4.8, 10, 15.5,
249 25.5, 36.3). The third group of experiments was performed using samples with fixed S-TA
250 relationship ($\text{TA} = 35.634 * \text{S} + 1377.1$; $\text{S} = 0, 12, 24, 36$). In each group, we prepared two types
251 of water and mixed them together to get four different TA-S pairs. Low S water originated from
252 Mississippi and Atchafalaya river waters (both are of high TA, (Guo et al., 2012)). High S water
253 is northern Gulf of Mexico seawater. Water was adjusted by additions of 0.2 mol/L NaOH or 0.2
254 mol/L HCl, which modify TA without significantly modifying S. The water was mixed to get
255 different S or TA values along the gradient connecting the endmembers. Then, we adjusted the
256 pH (and DIC) of these water samples by bubbling pure CO_2 gas or pure N_2 gas. TA was titrated
257 with a ROSSTM combination electrode 8102 (Thermo Fisher Scientific) on a semi-automated
258 open cell titrator (AS-ALK2, Apollo SciTech) and calibrated with certified reference material
259 from Andrew G. Dickson's lab, Scripps Institution of Oceanography.

260 The pH of these water samples was measured spectrophotometrically. We prepared a 2.5
261 mmol/L indicator solution by dissolving purified mCP (from Robert H. Byrne's laboratory,
262 University of South Florida) in deionized CO_2 free water. NaOH was added to adjust the
263 indicator solution pH to ~ 8.0 . The instrument setup and analysis procedure followed the design
264 of Carter et al. (2013). We used an Agilent 8453 UV-visible spectroscopy system, a 48K-steps
265 Kloehe pump controlled by a computer, a 10 cm flow-through cuvette with a water jacket, and a
266 thermal controlled circulating water bath as described in Carter et al. (2013). The system was
267 further automated to analyze a set of 8 water samples in sequence in about 40 minutes (AS-pH1,
268 Apollo Scitech). All measurements were conducted at 25°C . We added 20 μL of 2.5 mmol/L
269 purified mCP to 15 mL water sample as suggested by *The Guide to Best Practices for Ocean*
270 *CO_2 Measurements* (Dickson et al., 2007). Measurement stability was checked with 2-amino-2-
271 hydroxymethyl-1,3-propanediol (Tris) buffer solution (Batch #T33, from Andrew Dickson at
272 Scripps Institution of Oceanography, University of California, San Diego). Every sample was
273 measured by adding single-volume dye (20 μL) and double-volume dye (40 μL). Each sample
274 was measured at least three times. We also conducted multiple dye addition experiments (adding

275 10 μL , 20 μL , 30 μL , 40 μL , 50 μL , and 60 μL of indicator) to some of these water samples to
276 test the perturbation correction assumptions.

277 **3. Results**

278 **3.1 Computer simulation results**

279 **3.1.1 Theoretical dye perturbation on the sample pH**

280 We calculated pH from DIC, TA and S inputs with and without adding the indicator dye
281 alkalinity term. Figure 4 shows that ΔpH (the water sample pH after a dye addition minus the
282 water sample pH without the dye) is between 0 and 0.005 pH units for low-pH samples (~7.0–
283 7.5), indicating the dye addition ($S = 0$, $\text{pH} = 8.0$) leads to an increase in the water sample pH.
284 For sample pH higher than 7.5, ΔpH becomes negative, meaning the dye addition decreases the
285 original sample pH. Intuitively, it would make sense if this transitional x-intercept value were
286 located at the pH of the sample (in this case $\text{pH} = 8.0$). However, this is only the case for very-
287 low salinity samples in Figure 4b. This is because when the low-ionic strength dye mixes with
288 the high-ionic strength seawater, the speciation of the dye is shifted in favor of I^{2-} and protons are
289 released (Müller and Rehder, 2018a). In Figure 4a, the absolute ΔpH of low-TA samples (dark
290 blue dots) is larger than those of high-TA samples (red dots), showing the dye addition has a
291 larger influence on the pH of low-TA samples. Figure 4b shows that sample S affects ΔpH
292 values as well. The ΔpH of high-S samples (yellow dots, Figure 4b) changes less than that of
293 low-S samples (dark blue dots, Figure 4b). Because both TA and S affect ΔpH , we made three
294 extra groups of simulations to separate the effects of these two factors. The first group had fixed
295 TA and variable S. The second group had fixed S and variable TA. The last group had a fixed
296 TA-S relationship. The results of these simulations are shown together with the results from the
297 laboratory experiments in Section 3.2.

298 **3.1.2 Testing the assumptions of the empirical dye perturbation correction**

299 As explained in Section 2.1, the empirical dye perturbation correction is based on two
300 assumptions: (i) the absorbance ratio R changes linearly with the volume of dye added (or the
301 change of R with dye addition volume is a constant), and (ii) the absorbance ratio change per mL
302 dye addition ($\Delta R/\Delta V$) is a simple linear function of the absorbance ratio (R_1) with single volume

303 dye addition. We assessed the validity of these assumptions using our new MATLAB program
304 *dyeperturbation.m*.

305 To study the linearity between R and V, we calculated the sample pH with variable volumes of
306 dye additions from 1 μL to 100 μL . Then, the pH values were converted to the absorbance ratio
307 R based on Equation (2). Figure 5 shows the relationship between the dye addition volume and
308 the theoretical R, which is a simple linear relationship, and thus $\Delta R/\Delta V$ is a constant for each
309 DIC-TA pair. However, $\Delta R/\Delta V$ is not the same constant for every DIC-TA pair. When R is
310 smaller than 0.5 (referring to $\text{pH} < 7.4$), R is near constant or only slightly increases with
311 increasing dye volume. If $R > 0.5$, the $\Delta R/\Delta V$ has a negative slope against R. These results show
312 the first assumption is valid given a perfect spectrophotometer. Laboratory multiple dye addition
313 experiments also confirmed this conclusion derived from the simulation results (Figure 5b).

314 To test the second assumption, we evaluated the relationship between $\Delta R/\Delta V$ and R_1 and we
315 calculated the pH before and after single and double dye additions. Equation (2) was solved to
316 acquire R_1 and R_2 , respectively. $\Delta R/\Delta V$ can be determined as the empirical correction method
317 suggests, $\Delta R/\Delta V = (R_2 - R_1)/(V_2 - V_1)$. The slope of the relationship between $\Delta R/\Delta V$ and R_1
318 changes depending on the sample properties. $\Delta R/\Delta V$ decreases faster (slower) with increasing R_1
319 under low (high) TA or S (Figure 6 -8). $\Delta R/\Delta V$ vs R_1 is strongly correlated, but not perfectly
320 linear, which may influence the calculation of the empirical corrected sample pH a little. The
321 $\Delta R/\Delta V$ decreasing rates in different TA or S samples will be discussed in detail in Section 3.2.

322 **3.2 Laboratory experiments results supporting model simulations**

323 To verify the simulation results, we conducted laboratory experiments using natural water from
324 the northern Gulf of Mexico, the Mississippi River and the Atchafalaya River. The endmember
325 waters were mixed to get various S and TA gradients as described in Section 2.3. The results of
326 the laboratory experiments are compared with those from computer simulations.

327 **3.2.1 Fixed S and varying TA group**

328 The $\Delta R/\Delta V$ and ΔpH values agree well between the simulations and the experimental results
329 under fixed high S (Figure 6). The simulation results demonstrate that the $\Delta R/\Delta V$ decreasing
330 rates with R_1 depend on the sample TA: $\Delta R/\Delta V$ of high-TA samples decreases with R_1 less than

331 that of low-TA samples (Figure 6a). The simulation results suggest that dye perturbation on
332 high-TA samples ($\sim 2400 \mu\text{mol kg}^{-1}$) pH is rather small ($-0.0013 - 0.0005$ pH units; Figure 6b).
333 When sample TA is lowest ($\sim 1100 \mu\text{mol kg}^{-1}$), dye perturbation effect on ΔpH is the largest at
334 about -0.0030 to 0.0016 pH units (dark blue line in Figure 6b).

335 In the experimental results, $\Delta\text{R}/\Delta\text{V}$ decreases less with increasing R_1 when sample TA is high
336 ($\sim 2400 \mu\text{mol kg}^{-1}$); and $\Delta\text{R}/\Delta\text{V}$ decreases more rapidly when sample TA is low ($\sim 1100 \mu\text{mol kg}^{-1}$).
337 Based on Figure 6c, we made a simple linear regression of $\Delta\text{R}/\Delta\text{V}$ against R_1 for each TA
338 group and calculated the ΔpH (Figure 6d) using Equations (3) and (4). Experimental results
339 (Figure 6d) are similar to the simulation results (Figure 6b); dye perturbation increases with
340 sample pH increasing and TA decreasing. In Figure 6d, the ΔpH for high-TA samples (~ 2400
341 $\mu\text{mol kg}^{-1}$) is closer to 0 ($-0.001 - 0.000$ pH units) compared to ΔpH for low-TA samples (~ 1100
342 $\mu\text{mol kg}^{-1}$). Laboratory results support the simulation results (Figure 9). Therefore, for high
343 salinity samples, the differences between simulated $\Delta\text{R}/\Delta\text{V}$ - experimental $\Delta\text{R}/\Delta\text{V}$ ($\Delta(\Delta\text{R}/\Delta\text{V})$)
344 is close to 0 with a 0.06 standard deviation. The simulated ΔpH and the experimental ΔpH are
345 very close to each other with average $\Delta(\Delta\text{pH})$ 0.0008 (blue square dots in Figure 9b).

346 **3.2.2 Fixed TA and varying S group**

347 Since Mississippi river water has the same high TA ($\sim 2350 \mu\text{mol kg}^{-1}$) as the northern Gulf of
348 Mexico seawater, we can obtain a series of samples with the same high TA but distinct S by
349 mixing these two types of water in different ratios. Our simulation shows the decrease in $\Delta\text{R}/\Delta\text{V}$
350 with R_1 is minimized (relative to low S sample) when sample S is high (Figure 7a). The
351 experimental results support the simulation results (Figure 7a, 7c and 9a): when $S = 36$, $\Delta\text{R}/\Delta\text{V}$
352 decreases to -0.2 at $R_1 = 1.9$. When $S = 0$, $\Delta\text{R}/\Delta\text{V}$ decreases to -0.8 at $R_1 = 1.5$. Both simulation
353 results (Figure 7b) and experimental results (Figure 7d) show the dye perturbation on sample pH
354 is large in low-S water, while the perturbation is small in high-S water. In terms of ΔpH , the
355 experimental and simulated results match well when water sample $S > 5$ (Figure 9b). However,
356 when sample salinity equals zero and pH is low, the experimental and simulated perturbation
357 results have obvious difference (orange cross dots in red circle, Figure 9b). This could be both
358 the uncertainties in the spectrophotometric pH measurements and in the carbonate constants used

359 in the theoretical calculations at low S ($S < 5$) that may contribute to the mismatch of theory and
360 observation.

361 3.2.3 Fixed TA and S relationship group

362 The last scenario is mixing low-TA and low-S water with high-TA and high-S water in different
363 ratios to obtain a series of TA-S gradient water. This case is representative of natural waters in
364 most estuarine and coastal areas, where river TA is much lower than seawater TA and that TA
365 and S have a linear relationship during mixing. Simulation results (Figure 8a) show that $\Delta R/\Delta V$
366 decrease little with increasing R_1 in highly saline and alkaline water with $S = 36$ and $TA = 2800$
367 $\mu\text{mol kg}^{-1}$, while $\Delta R/\Delta V$ decreases rapidly with R_1 in the fresh riverine water of $S = 0$ and $TA =$
368 $1100 \mu\text{mol kg}^{-1}$. The relationship between $\Delta R/\Delta V$ and R_1 is nearly linear when $R_1 > 0.5$. As for
369 ΔpH , the dye perturbation on sample pH is very small ($-0.0015 - 0.001$ pH units) for $S = 36$ and
370 $TA = 2800 \mu\text{mol kg}^{-1}$, while it is large ($-0.01 - 0.003$ pH units) for low-S and low-TA water
371 (Figure 8b, blue line). The experimental results (Figure 8c and 8d) agree with the simulation
372 results (Figure 8a, 8b and Figure 9). Both of them show that dye perturbation presents little
373 influence on $\Delta R/\Delta V$ and sample pH in high-S and high-TA samples. Except for the water of $S =$
374 0 , the relationship of $\Delta R/\Delta V$ vs. R_1 can be considered as linear (Figure 8c) when $R_1 > 0.5$. The
375 average difference between simulated ΔpH and experimental ΔpH is -0.0006 , indicated by
376 yellow dots in Figure 9b.

377

378 4. Discussion

379 4.1 The mechanism of dye perturbation

380 The addition of mCP can perturb the original sample pH by (i) modification of the water sample
381 TA, and (ii) H^+ redistribution in the solution. The contribution of extra TA can be calculated
382 from the indicator solution properties. We simulated the addition of $20 \mu\text{L}$ (or $50 \mu\text{mol}$) mCP
383 solution (2.5 mmol/L , $S = 0$, $\text{pH} = 8$) to 15 mL high-S seawater sample at $25 \text{ }^\circ\text{C}$. Under these
384 conditions, the added $[\text{H}^+]$, $[\text{OH}^-]$, and the total mole concentration of mCP solution would be
385 diluted 750 times. Thus, the total concentration of mCP in the water sample is about 3.33
386 $\mu\text{mol/L}$. The mCP solution is a mixture of H_2I , I^{2-} and HI^- . The mCP solution ($\text{pH} = 8.0$, $S=0$)

387 mainly contributes ~20 % I^{2-} and ~80 % HI^- to the seawater sample (Figure 2). The added HI^-
388 does not affect the water sample TA, since HI^- is the zero level of proton in the titration. The TA
389 of the indicator dye solution is $TA_I = [I^{2-}]_I - [H_2I]_I + [OH^-] - [H^+]_{free}$. After mixing the indicator
390 dye and the water sample, the sample will have mCP species in solution and a new TA that is the
391 mass-weighted average alkalinity of the indicator solution TA_I and the original water sample TA,
392 $TA_m = (m_s TA_s + m_I TA_I) / (m_s + m_I)$. When the original water sample TA_s is high, the
393 capacity of the sample resisting extra TA addition is high as well. Thus, the effect of indicator
394 addition perturbation is smaller for high-TA samples than for low-TA samples.

395 The addition of the $HI^- - I^{2-}$ pair can change the sample original pH by H^+ exchange between the
396 sample solution and the dye. The H^+ redistributes among all the chemical species including
397 carbonate, boron, and phosphate as well as the $HI^- - I^{2-}$ pair. Then, a new chemical equilibrium is
398 established, in which the TA and I^{2-} contributed by mCP plays a role. The addition of the free H^+
399 and OH^- initially present in the indicator solution is small enough to be neglected. The exchange
400 of H^+ that occurs among different chemical species and the solution can be significant if the
401 indicator solution encounters a vastly different TA or pH environment. Beyond TA and pH, the
402 exchange and redistribution of H^+ is also influenced by the sample S. When the mCP dye is
403 added to a sample that has very different S from that of the indicator solution, the pK values of
404 mCP shift dramatically and H^+ transfers among different acid-base species. Thus, the dye
405 perturbation can be the result of dissociation coefficients changes. Overall, the influence of the
406 dye perturbation is determined by the S, TA, and pH of the water samples and the mCP solution.

407 **4.2 pH bias caused by the empirical correction method**

408 Our experiments results show that the dye perturbation magnitude is distinct in different S and
409 TA samples. Therefore, the method for empirical dye perturbation correction, which does not
410 take the properties of samples into account, needs careful re-evaluation when applied to water
411 samples spanning a large range of properties. In this section, we assess several common biases
412 that are caused by applying the empirical dye perturbation correction.

413 **4.2.1 One dye perturbation curve for all samples**

414 Applying a unique dye perturbation curve (one linear regression for $\Delta R/\Delta V$ vs. R_1) to a series of
415 samples with different TA and S can lead to biases (Figure 10a, black vs. blue vs. yellow line)
416 because the change of $\Delta R/\Delta V$ vs. R_1 is not constant with changing TA and S. If only one
417 averaged linear regression of $\Delta R/\Delta V$ vs. R_1 (as the black line in Figure 10a indicates) is made for
418 all samples ($S = 0\text{--}40$, $TA = 1100\ \mu\text{mol kg}^{-1} - 2800\ \mu\text{mol kg}^{-1}$), the samples with high S and
419 high TA are about 0 – 0.001 pH overcorrected (yellow arrows in Figure 10a and 10b). On the
420 contrary, the dye perturbation influence on low-S and low-TA samples (as the blue line and
421 arrow indicates in Figure 10a) should be larger than that calculated from the averaged $\Delta R/\Delta V$ vs.
422 R_1 regression line. For a sample with low S and low TA ($S = 0$, $TA = 1100\ \mu\text{mol kg}^{-1}$), the
423 corrected pH from average $\Delta R/\Delta V$ vs. R_1 regression line is about 0.001–0.006 pH units is under-
424 corrected or is smaller than it should be (blue arrows, Figure 10b).

425 **4.2.2 Insufficient double dye addition samples over a wide range of pH**

426 If the double dye addition samples only fall in a very narrow pH range, the dye perturbation
427 correction bias can be very large when the average $\Delta R/\Delta V$ vs. R_1 regression line is extrapolated
428 to extreme pH (Figure 10a, red line). For example, if water samples are from a coastal cruise, the
429 samples may cover a wide S and TA range ($S = 0\text{--}40$, $TA = 1100\text{--}2800\ \mu\text{mol kg}^{-1}$). However,
430 if the double dye addition samples insufficiently cover the pH range of the field samples (only in
431 the range of $R = 1.0 \sim 1.5$, as figure 10a red solid line shows), then $\Delta R/\Delta V$ to R_1 slope will
432 decrease far more than the actual sample slope. The $\Delta R/\Delta V$ of high-S and high-TA samples ($S =$
433 40 , $TA = 2800\ \mu\text{mol kg}^{-1}$) from the extrapolation line (figure 10a, red dash line) is much smaller
434 than its theoretical values at high pH (Figure 10a, red arrow), which can cause an overcorrection
435 of 0.005 pH units (Figure 10b, right red arrow). For the low S and low TA samples ($S = 0$, $TA =$
436 $1100\ \mu\text{mol kg}^{-1}$), the $\Delta R/\Delta V$ from extrapolation is larger than its theoretical value at low pH, and
437 the pH correction is 0.002 pH units smaller than its theoretical value (left red arrows in Figures
438 10a and 10b). It is therefore important to ensure that the lines are fit based to dye-addition
439 experiment data from a set of samples that is representative of the measurements.

440 The above discussions do not include possible human errors that can lead to enormous dye
441 perturbation correction errors. In our initial tests, we added 30 μL stock solution (2.5 mmol/L) to
442 15 mL sample for the sample measurement and 60 μL mCP stock solution for the double dye

443 addition experiment. In the double dye addition test, the final mCP concentration in the sample
444 solution was too high (~10 $\mu\text{mol/L}$), which caused the absorbance peak values in double dye
445 addition experiments to exceed the linear range of the Agilent 8453 instrument. This introduced
446 an erroneous ΔpH correction as large as 0.01, at least twice as larger as the true correction value.
447 Such errors should be avoided in the dye perturbation experiments. The final mCP concentration
448 in the sample should always fall in the range of linear detection range. Since pH data literatures
449 do not generally report their double dye correction information, it is difficult to evaluate the
450 extent of this error. However, we caution field workers to check if their double addition is still
451 within the analytical linear range of the instrument if the experimentally derived double dye
452 correction is suspiciously large. We further note that the negative bias visible in Figure 9a
453 (average $\Delta(\Delta\text{R}/\Delta\text{V}) = -0.0008$) could be attributed to a non-linear response in one wavelength vs.
454 another (i.e., the I^{2-} peak at 578 nm may not increase with increasing dye concentration to the
455 same degree as the HI^- at 434 nm). For most samples, this results in small errors in the empirical
456 adjustments, but for others (e.g., the 3 circled Fixed TA group points in Figure 9b) the impact of
457 these errors on pH can be significant. Alternately, a negative $\Delta(\Delta\text{R}/\Delta\text{V})$ bias could be explained
458 if the dye pH were lowered by the dye gradually absorbing CO_2 .

459 **4.3 Reduction of dye perturbation**

460 **4.3.1 The optimization of dye solution preparation**

461 The preparation of the dye solution is very important because it is directly related to the
462 magnitude of the dye perturbation. The purified mCP solid is in the form of H_2I , which is very
463 hard to dissolve in the water. Adding NaOH to the solvent can help with the dissolution and
464 adjust the mCP indicator pH. Since the natural seawater pH falls in the range of 7.4 – 8.5, the
465 recommended dye pH is about 7.9 ± 0.1 pH (SOP 6b, (Dickson et al., 2007)). However, the
466 amount of added NaOH is different for deionized water and 0.7 mol/L NaCl ($S = 35$) solution. In
467 the 0.7 mol/L NaCl, the ionic strength is high, and the mCP pK_2 value is smaller than it is in the
468 deionized water, (Müller and Rehder, 2018a) which means more I^{2-} is needed to keep the same
469 pH (Equation 1). Therefore, more NaOH should be added to the stock solution to produce more
470 I^{2-} , thus contributing more TA to the water sample.

471 When purified mCP is used for open ocean spectrophotometric seawater pH measurements, it is
472 common to dissolve the indicator in 0.7 mol/L NaCl solution and adjust the pH to approximately
473 8.0. The perturbation effect of this indicator solution is very small for high S-TA seawater
474 sample, and the perturbation is near 0 when the sample pH is close to 8.0 (Figure 11b). However,
475 this indicator solution is not ideal for low S samples (e.g., brackish or riverine water sample
476 (blue dots, Figure 11b)). The shift of pK_2 per unit salinity is the largest at very low ionic
477 strength (Müller and Rehder, 2018b). When this $S = 35$, $pH = 8.0$ mCP indicator is added to the
478 low S water sample, the pK_2 increase is quite significant even though the dye is diluted by a
479 much larger volume of sample. This pK_2 change causes a greater pH perturbation for applying
480 high S stock solutions to low S sample than applying low S stock solution to high S samples
481 (Figure 11, low-S blue dots).

482 In addition, we should be cautious when using the $S = 35$ and $pH = 8.0$ indicator to measure
483 high-S and low-pH seawater samples (e.g., deep seawater measured at $T = 25\text{ }^\circ\text{C}$) because the
484 second assumption of the empirical dye perturbation correction method, $\Delta R/\Delta V$ vs. R_1 having
485 linear relationship, is no longer valid (especially when R close to 0.5, Figure 11a). The non-
486 linearity at low pH results is the result of the R ratio being defined with the absorbance
487 associated with the acidic form of mCP in the denominator. This amplifies R changes at high pH
488 and diminishes R changes at low pH. To avoid the non-linear part of the $\Delta R/\Delta V$ vs. R_1
489 relationship, we can choose to use a $S = 35$ and pH less than 8 indicator dye solution or to use a S
490 $= 0$ and $pH = 8.0$ indicator dye solution (that has a more linear response over the entire pH range,
491 Figure 10a, 11a).

492 Overall, the preparation of the dye solution should be based on the expected sample S , pH , and
493 TA. The goal is to achieve a minimum dye perturbation around the sample original pH. We
494 provide a new MATLAB program to determine the pH and S of the indicator solution that will
495 produce the lowest pH perturbation. An example is provided in the supplementary material.

496 **4.3.2 Practical method for dye perturbation correction**

497 The empirical dye perturbation correction method is limited because it does not take the sample
498 S and TA into consideration when performing the double dye addition experiments. Chierici et
499 al. (1999) took S into consideration, but not TA, when applying the indicator addition

500 perturbation correction. In this paper, both experimental and computational results suggest that
501 the dye perturbation is smaller for high-TA and high-S samples than that for low-TA and low-S
502 samples. Here, we suggest that the selection of double dye addition samples should cover as
503 wide range of S, TA and pH as the collection of water samples analyzed, and the regression line
504 for $\Delta R/\Delta V$ vs. R_1 should be made separately for different S and TA groups. Otherwise, it will be
505 difficult to well define the $\Delta R/\Delta V$ vs. R_1 regression line, which may cause the biases discussed in
506 Section 4.2.

507 Another method to eliminate dye perturbation is an empirical extrapolation of R to zero dye
508 addition for every individual sample (Lai et al., 2016). This method can remove the effect of S
509 and TA influence of the sample. Simulation results (Figure 5a) indicate that for a single sample,
510 R and volume of dye added (concentration) have a simple linear relationship. We verified the
511 simulation results with the laboratory experiments (Figure 5b). Therefore, it is practical to
512 linearly extrapolate the R to zero dye addition for each sample. Previous research suggested that
513 for freshwater or low-S samples, dye perturbation on sample pH should be determined for every
514 sample to obtain the best accuracy due to the low buffer capacity of freshwater (Lai et al., 2016;
515 Moseley, 2004). For seawater, even though the buffer capacity is much higher than that of
516 freshwater, we still suggest performing multiple volume dye addition experiments to all the
517 samples when not enough samples covering different S and TA are available. This approach has
518 the limitations that (1) random errors in individual dye-addition experiment measurements can be
519 erroneously extrapolated unless the experiment is repeated enough times to adequately
520 characterize the dye perturbation response and (2) that this requires more analysis time, dye, and
521 sample.

522 If experiments for dye perturbation correction cannot be performed, we provide a new MATLAB
523 function that can calculate the theoretical dye perturbation. With known sample properties (TA,
524 DIC, S and T), indicator dye properties (pH, S, concentration) and sample/dye mixing volume
525 ratios, the dye perturbation on sample pH can be calculated. The function (*dyeperturbation.m*)
526 details and an example result are provided in the Supplementary Material.

527 **5. Conclusions**

528 The addition of mCP solution to water samples changes the sample TA composition and
529 redistributes hydrogen ion among other chemical equilibrium parameters. The effect of this dye
530 perturbation is distinct for water samples with different TA, S, and pH; and it also depends on
531 the mCP solution preparation. To determine the “true pH” of the water sample without a dye
532 perturbation, a correction for the dye perturbation is needed, which is very important for low-
533 buffered, low-S and low-TA samples. This work not only reveals the mechanism by which the
534 dye affects the carbonate parameters of the seawater samples, but also allows improving the
535 accuracy of spectrophotometric pH measurements. The high-quality data will help with
536 identifying small decadal pH changes and give a better understanding of ocean acidification.

537 The empirical method of dye perturbation correction based upon a single regression is limited
538 and may cause biases on corrected pH (biases of ~0.001–0.005 pH units). We suggest taking TA
539 and S of the analyzed samples into consideration when selecting the samples to perform the
540 double dye addition for the empirical method of dye perturbation correction method. If the
541 number of samples is low or the samples do not cover a wide range of TA, S and pH, a better
542 way to eliminate the dye perturbation is studying the dye perturbation through multivolume dye
543 additions to each sample, ideally repeated multiple times to minimize the effects of random
544 errors. We also provide a new MATLAB function that helps calculate theoretical dye
545 perturbations, which can be used as a sanity check for dye perturbation adjustments or use it in
546 lieu of an empirical adjustment. The function is validated by empirical measurements, but
547 considerable uncertainties remain in low salinity sample due to uncertainties in carbonate
548 chemistry and mCP speciation coefficients.

549

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561 materials or equipment identified are necessarily the best available for the purpose.

562

563 Table 1. Constant choices for CO2SYS calculation.

| Parameters | Choices |
|---|---------------------------|
| Carbonate Dissociation Constants K_1 , K_2 | (Millero et al., 2010) |
| Dissociation constants of HSO_4^- , KSO_4 | (Dickson, 1990a) |
| Dissociation constants of F, KF | (Dickson and Riley, 1979) |
| Dissociation constants of B, KB | (Dickson, 1990b) |
| Sulfate Concentration | (Morris and Riley, 1966) |
| Fluoride Concentration | (Riley, 1965) |
| Boron Concentration | (Uppström, 1974) |

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567 **References:**

568 Andersson, A.J., Gledhill, D., 2013. Ocean Acidification and Coral Reefs: Effects on
569 Breakdown, Dissolution, and Net Ecosystem Calcification. *Ann. Rev. Mar. Sci.* 5, 321–348.
570 <https://doi.org/10.1146/annurev-marine-121211-172241>

571 Bockmon, E.E., Dickson, A.G., 2015. An inter-laboratory comparison assessing the quality of
572 seawater carbon dioxide measurements. *Mar. Chem.* 171, 36–43.
573 <https://doi.org/10.1016/j.marchem.2015.02.002>

574 Cai, W.-J., Wang, Y., Hodson, R.E., 1998. Acid-Base Properties of Dissolved Organic Matter in
575 the Estuarine Waters of Georgia, USA. *Geochim. Cosmochim. Acta* 62, 473–483.
576 [https://doi.org/10.1016/S0016-7037\(97\)00363-3](https://doi.org/10.1016/S0016-7037(97)00363-3)

577 Cai, W.J., Guo, X., Chen, C.T.A., Dai, M., Zhang, L., Zhai, W., Lohrenz, S.E., Yin, K.,
578 Harrison, P.J., Wang, Y., 2008. A comparative overview of weathering intensity and
579 HCO₃⁻ flux in the world's major rivers with emphasis on the Changjiang, Huanghe,
580 Zhujiang (Pearl) and Mississippi Rivers. *Cont. Shelf Res.* 28, 1538–1549.
581 <https://doi.org/10.1016/j.csr.2007.10.014>

582 Cai, W.J., Hu, X., Huang, W.J., Murrell, M.C., Lehrter, J.C., Lohrenz, S.E., Chou, W.C., Zhai,
583 W., Hollibaugh, J.T., Wang, Y., Zhao, P., Guo, X., Gundersen, K., Dai, M., Gong, G.C.,
584 2011. Acidification of subsurface coastal waters enhanced by eutrophication. *Nat. Geosci.*
585 4, 766–770. <https://doi.org/10.1038/ngeo1297>

586 Carter, B.R., Radich, J.A., Doyle, H.L., Dickson, A.G., 2013. An automated system for
587 spectrophotometric seawater pH measurements. *Limnol. Oceanogr. Methods* 11, 16–27.
588 <https://doi.org/10.4319/lom.2013.11.16>

589 Chierici, M., Fransson, A., Anderson, L.G., 1999. Influence of m-cresol purple indicator
590 additions on the pH of seawater samples: Correction factors evaluated from a chemical
591 speciation model. *Mar. Chem.* 65, 281–290. [https://doi.org/10.1016/S0304-4203\(99\)00020-](https://doi.org/10.1016/S0304-4203(99)00020-1)
592 1

- 593 Clayton, T.D., Byrne, R.H., 1993. Spectrophotometric seawater pH measurements : total
594 hydrogen results. *Deep. Res.* 40, 2115–2129. [https://doi.org/10.1016/0967-0637\(93\)90048-](https://doi.org/10.1016/0967-0637(93)90048-8)
595 8
- 596 DeGrandpre, M.D., Spaulding, R.S., Newton, J.O., Jaqueth, E.J., Hamblock, S.E., Umansky,
597 A.A., Harris, K.E., 2014. Considerations for the measurement of spectrophotometric pH for
598 ocean acidification and other studies. *Limnol. Oceanogr. Methods* 12, 830–839.
599 <https://doi.org/10.4319/lom.2014.12.830>
- 600 Dickson, A.G., 1993. The measurement of sea water pH. *Mar. Chem.* 44, 131–142.
601 [https://doi.org/10.1016/0304-4203\(93\)90198-W](https://doi.org/10.1016/0304-4203(93)90198-W)
- 602 Dickson, A.G., 1990a. Standard potential of the reaction: $\text{AgCl(s)} + 1/2\text{H}_2(\text{g}) = \text{Ag(s)} + \text{HCl(aq)}$,
603 and the standard acidity constant of the ion HSO_4^- in synthetic sea water from 273.15
604 to 318.15 K. *J. Chem. Thermodyn.* [https://doi.org/10.1016/0021-9614\(90\)90074-Z](https://doi.org/10.1016/0021-9614(90)90074-Z)
- 605 Dickson, A.G., 1990b. Thermodynamics of the dissociation of boric acid in synthetic seawater
606 from 273.15 to 318.15 K. *Deep Sea Res. Part A, Oceanogr. Res. Pap.*
607 [https://doi.org/10.1016/0198-0149\(90\)90004-F](https://doi.org/10.1016/0198-0149(90)90004-F)
- 608 Dickson, A.G., 1984. pH scales and proton-transfer reactions in saline media such as sea water.
609 *Geochim. Cosmochim. Acta.* [https://doi.org/10.1016/0016-7037\(84\)90225-4](https://doi.org/10.1016/0016-7037(84)90225-4)
- 610 Dickson, A.G., 1981. An exact definition of total alkalinity and a procedure for the estimation of
611 alkalinity and total inorganic carbon from titration data. *Deep Sea Res. Part A, Oceanogr.*
612 *Res. Pap.* [https://doi.org/10.1016/0198-0149\(81\)90121-7](https://doi.org/10.1016/0198-0149(81)90121-7)
- 613 Dickson, A.G., Riley, J.P., 1979. The estimation of acid dissociation constants in seawater media
614 from potentiometric titrations with strong base. I. The ionic product of water - K_w . *Mar.*
615 *Chem.* [https://doi.org/10.1016/0304-4203\(79\)90001-X](https://doi.org/10.1016/0304-4203(79)90001-X)
- 616 Dickson, A.G., Sabine, C.L., Christian, J.R., 2007. Guide to best practices for ocean CO₂
617 measurements. North Pacific Marine Science Organization.
- 618 Doney, S.C., Fabry, V.J., Feely, R.A., Kleypas, J.A., 2009. Ocean Acidification: The Other CO₂

619 Problem. *Ann. Rev. Mar. Sci.* 1, 169–192.
620 <https://doi.org/10.1146/annurev.marine.010908.163834>

621 Guo, X., Cai, W., Huang, W., Wang, Y., Chen, F., Murrell, M.C., Lohrenz, S.E., Jiang, L., Dai,
622 M., Hartmann, J., Lin, Q., Culp, R., 2012. Carbon dynamics and community production in
623 the Mississippi River plume 57, 1–17. <https://doi.org/10.4319/lo.2012.57.1.0001>

624 Kwiatkowski, L., Orr, J.C., 2018. Diverging seasonal extremes for ocean acidification during the
625 twenty-first century. *Nat. Clim. Chang.* 8, 141–145. [https://doi.org/10.1038/s41558-017-](https://doi.org/10.1038/s41558-017-0054-0)
626 0054-0

627 Lai, C.Z., DeGrandpre, M.D., Wasser, B.D., Brandon, T.A., Clucas, D.S., Jaqueth, E.J., Benson,
628 Z.D., Beatty, C.M., Spaulding, R.S., 2016. Spectrophotometric measurement of freshwater
629 pH with purified meta-cresol purple and phenol red. *Limnol. Oceanogr. Methods* 14, 864–
630 873. <https://doi.org/10.1002/lom3.10137>

631 Landschützer, P., Gruber, N., Bakker, D.C.E., Stemmler, I., Six, K.D., 2018. Strengthening
632 seasonal marine CO₂ variations due to increasing atmospheric CO₂. *Nat. Clim. Chang.* 8.
633 <https://doi.org/10.1038/s41558-017-0057-x>

634 Lewis, E., Wallace, D., 1998. Program developed for CO₂ system calculations. Ornl/Cdiac-105.
635 <https://doi.org/4735>

636 Liu, X., Patsavas, M.C., Byrne, R.H., 2011. Purification and characterization of meta-cresol
637 purple for spectrophotometric seawater pH measurements. *Environ. Sci. Technol.* 45, 4862–
638 4868. <https://doi.org/10.1021/es200665d>

639 Ma, J., Shu, H., Yang, B., Byrne, R.H., Yuan, D., 2019. Spectrophotometric determination of pH
640 and carbonate ion concentrations in seawater: Choices, constraints and consequences. *Anal.*
641 *Chim. Acta* 1081, 18–31. <https://doi.org/10.1016/j.aca.2019.06.024>

642 Millero, F.J., Graham, T.B., Huang, F., Bustos-Serrano, H., Pierrot, D., 2006. Dissociation
643 constants of carbonic acid in seawater as a function of salinity and temperature. *Mar. Chem.*
644 <https://doi.org/10.1016/j.marchem.2005.12.001>

645 Morris, A.W., Riley, J.P., 1966. The bromide/chlorinity and sulphate/chlorinity ratio in sea
646 water. *Deep. Res. Oceanogr. Abstr.* [https://doi.org/10.1016/0011-7471\(66\)90601-2](https://doi.org/10.1016/0011-7471(66)90601-2)

647 Mosley, L.M., Husheer, S.L.G., Hunter, K.A., 2004. Spectrophotometric pH measurement in
648 estuaries using thymol blue and m-cresol purple. *Mar. Chem.* 91, 175–186.
649 <https://doi.org/10.1016/j.marchem.2004.06.008>

650 Müller, J.D., Rehder, G., 2018a. Metrology of pH Measurements in Brackish Waters—Part 2:
651 Experimental Characterization of Purified meta-Cresol Purple for Spectrophotometric pHT
652 Measurements. *Front. Mar. Sci.* 5, 1–9. <https://doi.org/10.3389/fmars.2018.00177>

653 Müller, J.D., Rehder, G., 2018b. Metrology of pH Measurements in Brackish Waters—Part 2:
654 Experimental Characterization of Purified meta-Cresol Purple for Spectrophotometric pHT
655 Measurements. *Front. Mar. Sci.* 5, 1–9. <https://doi.org/10.3389/fmars.2018.00177>

656 Orr, J.C., Epitalon, J.M., Dickson, A.G., Gattuso, J.P., 2018. Routine uncertainty propagation for
657 the marine carbon dioxide system. *Mar. Chem.* 207, 84–107.
658 <https://doi.org/10.1016/j.marchem.2018.10.006>

659 Orr, J.C., Fabry, V.J., Aumont, O., Bopp, L., Doney, S.C., Feely, R.A., Gnanadesikan, A.,
660 Gruber, N., Ishida, A., Joos, F., Key, R.M., Lindsay, K., Maier-Reimer, E., Matear, R.,
661 Monfray, P., Mouchet, A., Najjar, R.G., Plattner, G.-K., Rodgers, K.B., Sabine, C.L.,
662 Sarmiento, J.L., Schlitzer, R., Slater, R.D., Totterdell, I.J., Weirig, M.-F., Yamanaka, Y.,
663 Yool, A., 2005. Anthropogenic ocean acidification over the twenty-first century and its
664 impact on calcifying organisms. *Nature* 437, 681–686. <https://doi.org/10.1038/nature04095>

665 Patsavas, M.C., Byrne, R.H., Liu, X., 2013. Purification of meta-cresol purple and cresol red by
666 flash chromatography: Procedures for ensuring accurate spectrophotometric seawater pH
667 measurements. *Mar. Chem.* 150, 19–24. <https://doi.org/10.1016/j.marchem.2013.01.004>

668 Riley, J.P., 1965. The occurrence of anomalously high fluoride concentrations in the North
669 Atlantic. *Deep. Res. Oceanogr. Abstr.* [https://doi.org/10.1016/0011-7471\(65\)90027-6](https://doi.org/10.1016/0011-7471(65)90027-6)

670 Screening, T., Lecture, C., n.d. Nakamura_Harned_Cells 36–41.

671 Soli, A.L., Pav, B.J., Byrne, R.H., 2013. The effect of pressure on meta-Cresol Purple
672 protonation and absorbance characteristics for spectrophotometric pH measurements in
673 seawater. *Mar. Chem.* 157, 162–169. <https://doi.org/10.1016/j.marchem.2013.09.003>

674 Uppström, L.R., 1974. The boron/chlorinity ratio of deep-sea water from the Pacific Ocean.
675 *Deep. Res. Oceanogr. Abstr.* [https://doi.org/10.1016/0011-7471\(74\)90074-6](https://doi.org/10.1016/0011-7471(74)90074-6)

676 Van Heuven, S., Pierrot, D., Rae, J.W.B., Lewis, E., Wallace, D.W.R., 2011. MATLAB Program
677 Developed for CO2 System Calculations. ORNL/CDIAC-105b.
678 https://doi.org/10.3334/CDIAC/OTG.CO2SYS_MATLAB_V1.1

679 Waldbusser, G.G., Salisbury, J.E., 2014. Ocean Acidification in the Coastal Zone from an
680 Organism’s Perspective: Multiple System Parameters, Frequency Domains, and Habitats.
681 *Ann. Rev. Mar. Sci.* 6, 221–247. <https://doi.org/10.1146/annurev-marine-121211-172238>

682 Xu, Y.Y., Pierrot, D., Cai, W.J., 2017. Ocean carbonate system computation for anoxic waters
683 using an updated CO2SYS program. *Mar. Chem.* 195, 90–93.
684 <https://doi.org/10.1016/j.marchem.2017.07.002>

685 Zhang, H., Byrne, R.H., 1996. Spectrophotometric pH measurements of surface seawater at in-
686 situ conditions: Absorbance and protonation behavior of thymol blue. *Mar. Chem.* 52, 17–
687 25. [https://doi.org/10.1016/0304-4203\(95\)00076-3](https://doi.org/10.1016/0304-4203(95)00076-3)

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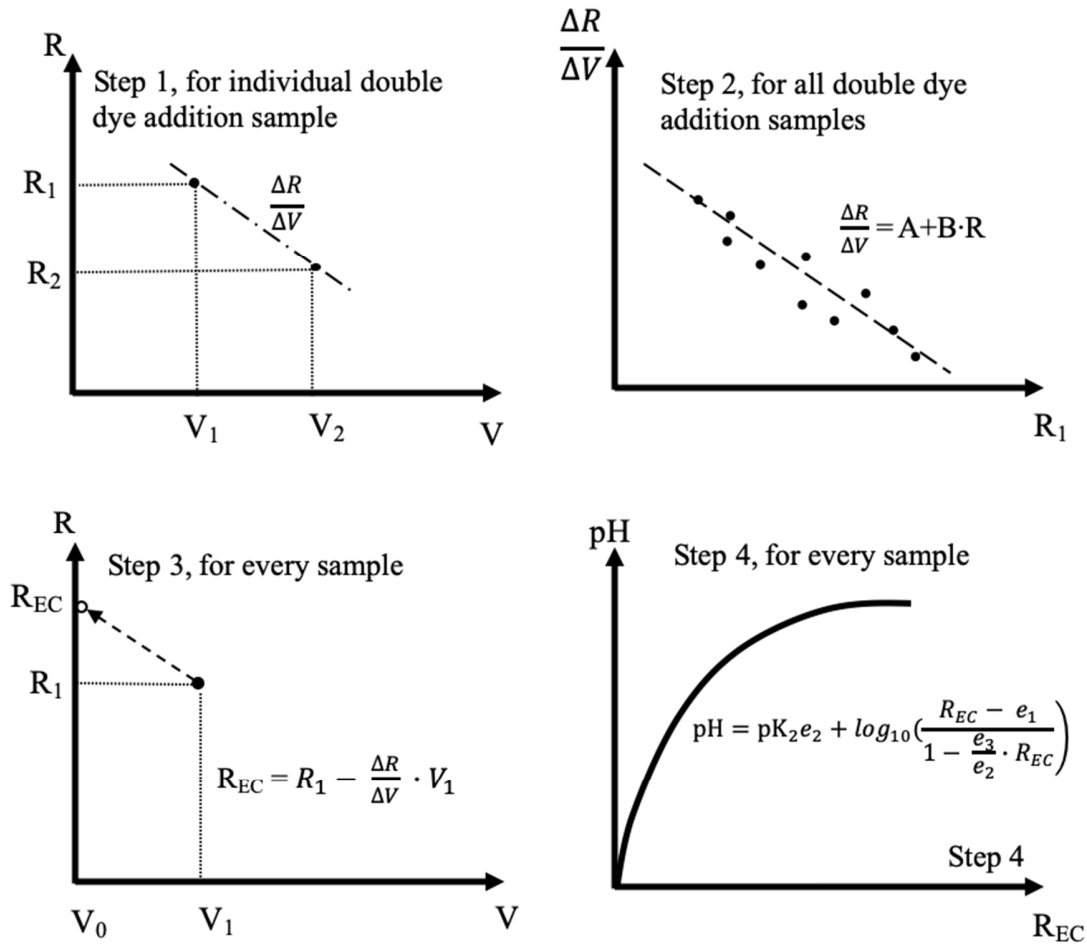


Figure 1. Four steps of empirical dye perturbation. Step 1, calculate $\frac{\Delta R}{\Delta V}$ from limited number of double dye addition experiments. Step 2, plot $\frac{\Delta R}{\Delta V}$ against R_1 to get a regression linear relationship. The first 2 steps are the empirical dye perturbation assumptions. Step 3, apply extrapolated $\frac{\Delta R}{\Delta V}$ to every sample to get R_{EC} . Step 4, use R_{EC} to calculate sample 'true' pH without dye addition.

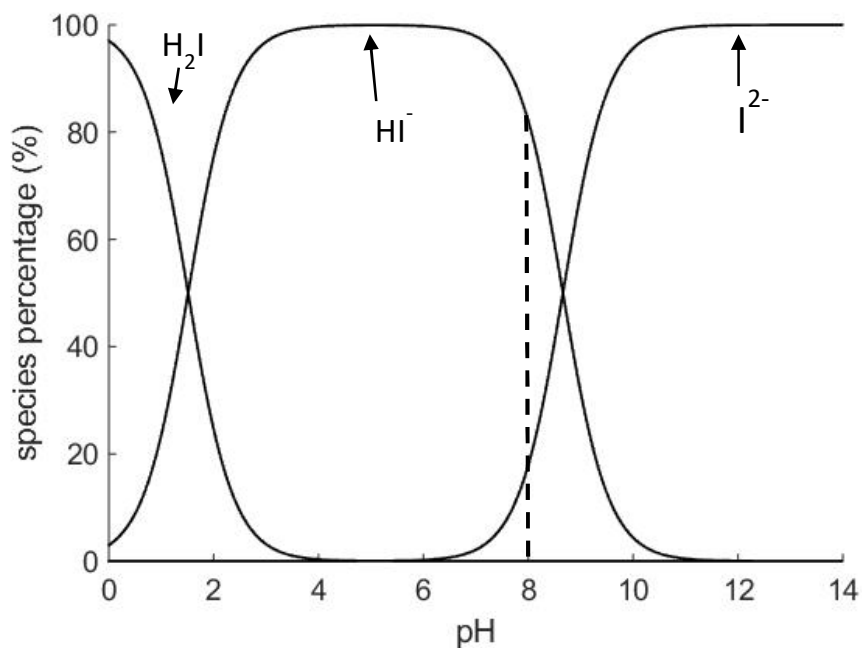


Figure 2. Bjerrum plot showing the species distribution of purified mCP as a function of pH ($T = 298.15K$, $S = 0$). The mCP indicator solution is usually adjusted to $pH = 8.0$ (the dash line), where $HI^- = 80\%$, and $I^{2-} = 20\%$.

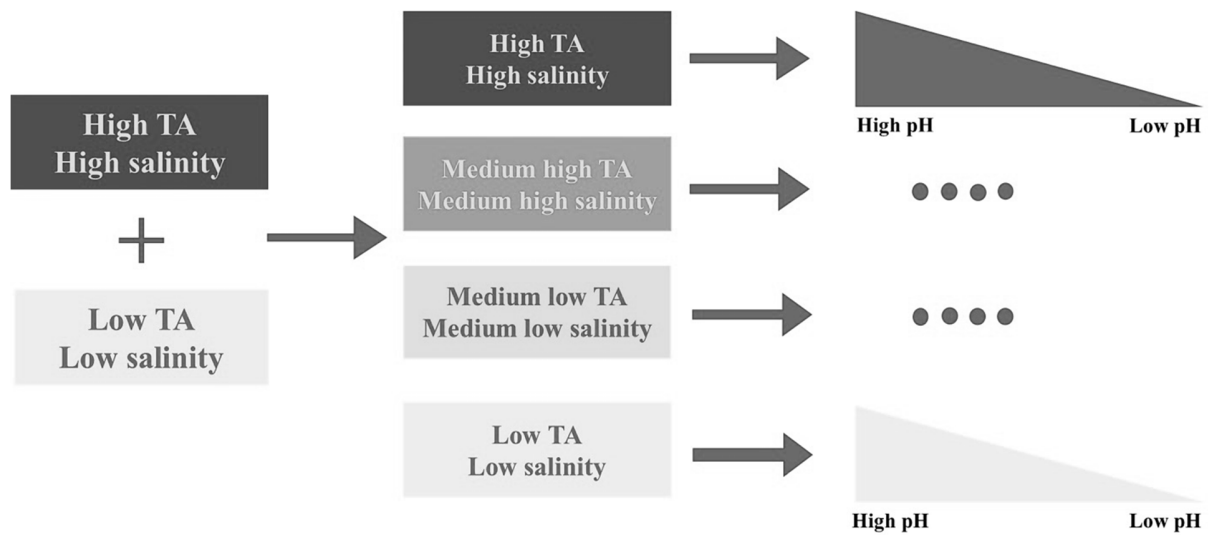


Figure 3. The laboratory experiment design of fixed S and TA relationship. High TA-S water mixed with low TA-S water to get 4 types of water with different TA and salinity. Each type of water was adjusted to obtain a pH gradient by dissolving/removing CO₂.

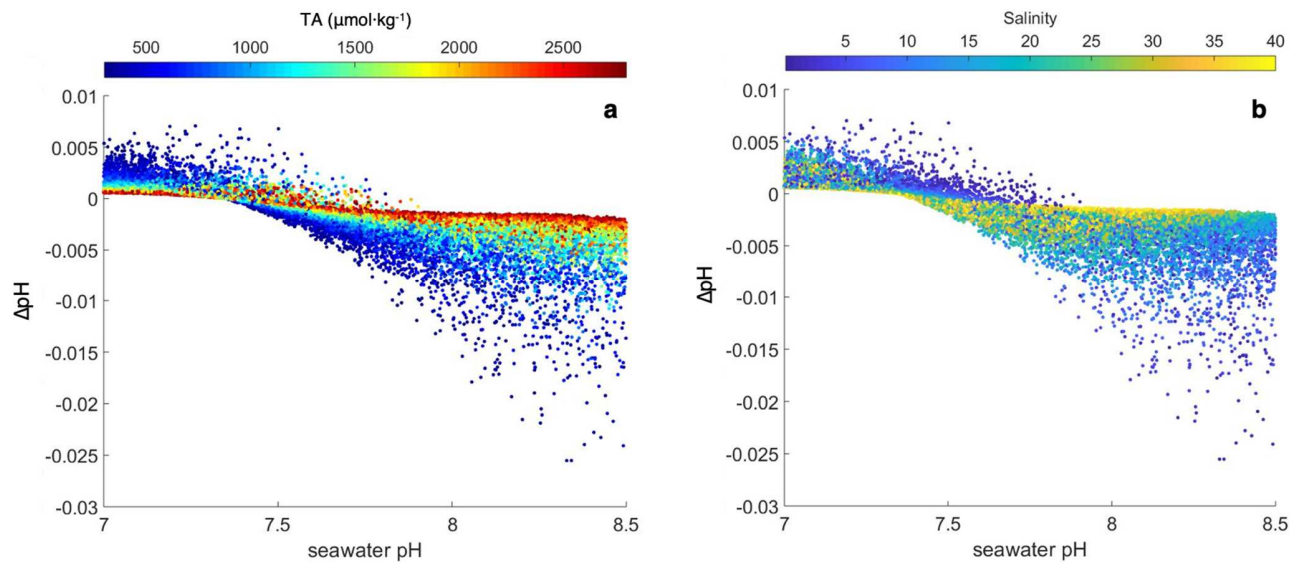


Figure 4. Dye perturbation simulation on sample pH using 50000 random combinations of DIC, TA, and S with a dye stock of $S = 0$ and $\text{pH} = 8$. The change in pH caused by the dye addition is shown on the Y axis as ΔpH (seawater pH with dye addition – seawater pH) as a function of sample pH in absence of the dye (X axis). Figure 3a visualizes the variation in ΔpH with TA while 3b shows the variation in ΔpH with S. Both panels indicate the influence of TA and salinity on dye perturbation.

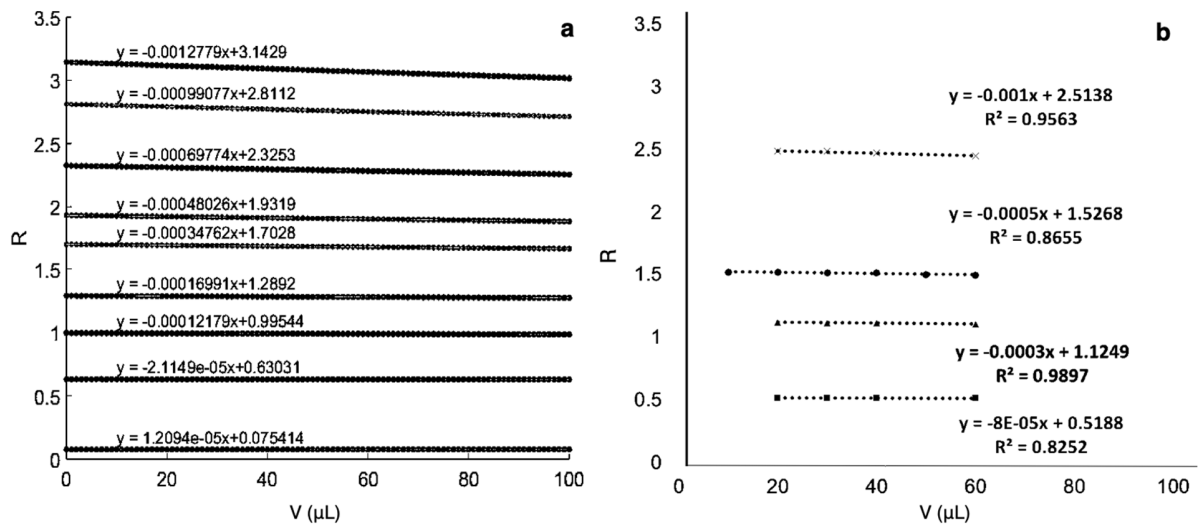


Figure 5. The absorbance ratio (R) as a linear function of dye volume (V). Subplot (a) is simulation result. Each line represents a different combination of sample DIC, TA and S, thus a different pH. Subplot (b) is laboratory result, supporting the simulation results.

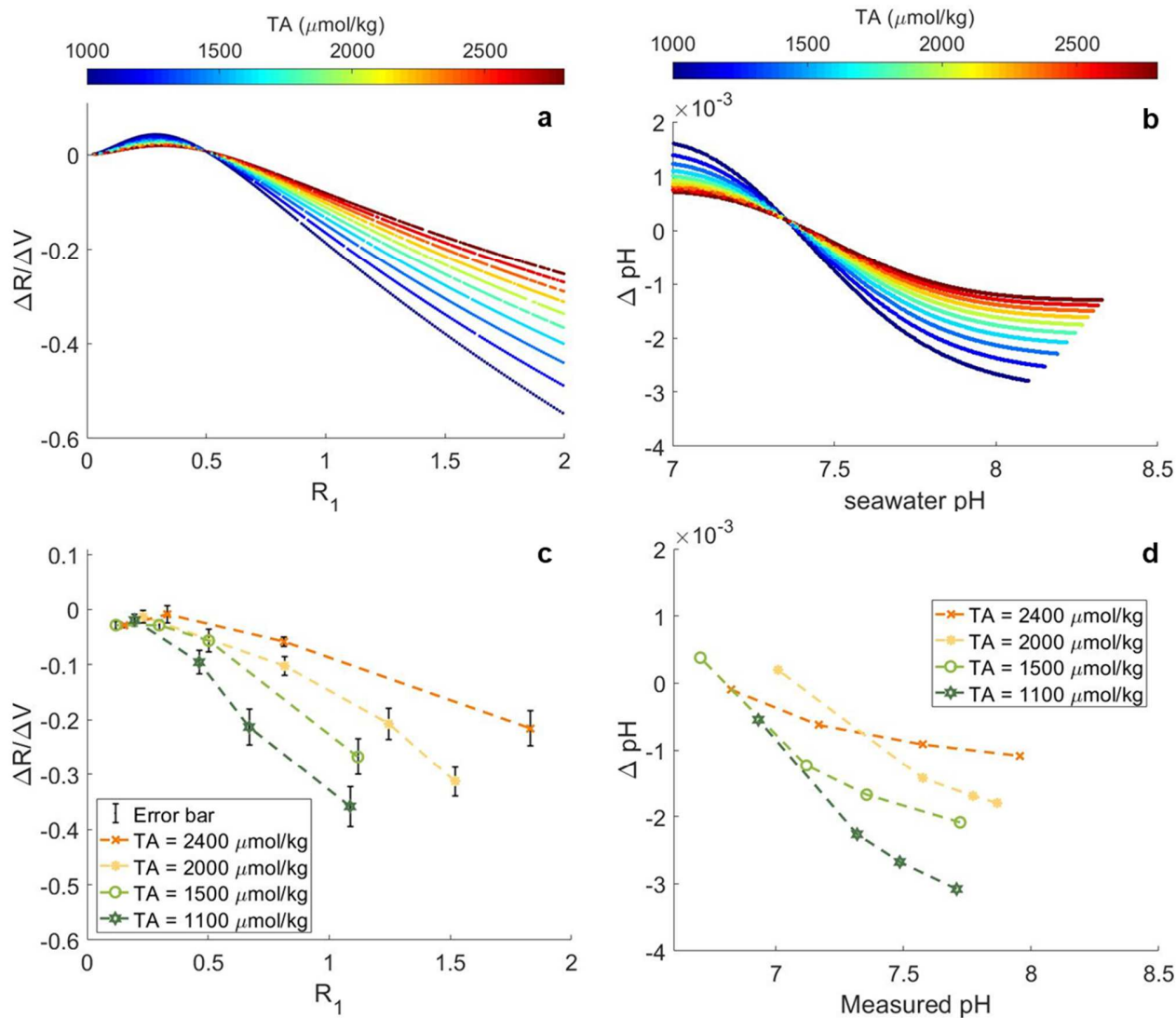


Figure 6. Simulation (a and b) and experimental (c and d) results of dye perturbation on a sample with fixed high salinity ($S = 36$) and variable TA. The high S-TA is Gulf of Mexico Seawater, while the high S – low TA water is the Gulf of Mexico Seawater with HCl addition. Subplots a and c show the relationship between $\Delta R/\Delta V$ and R_1 , which depends on the sample TA. Subplots b and d show the ΔpH (= pH with dye perturbation – pH without dye perturbation) changes of the water sample due to the indicator addition.

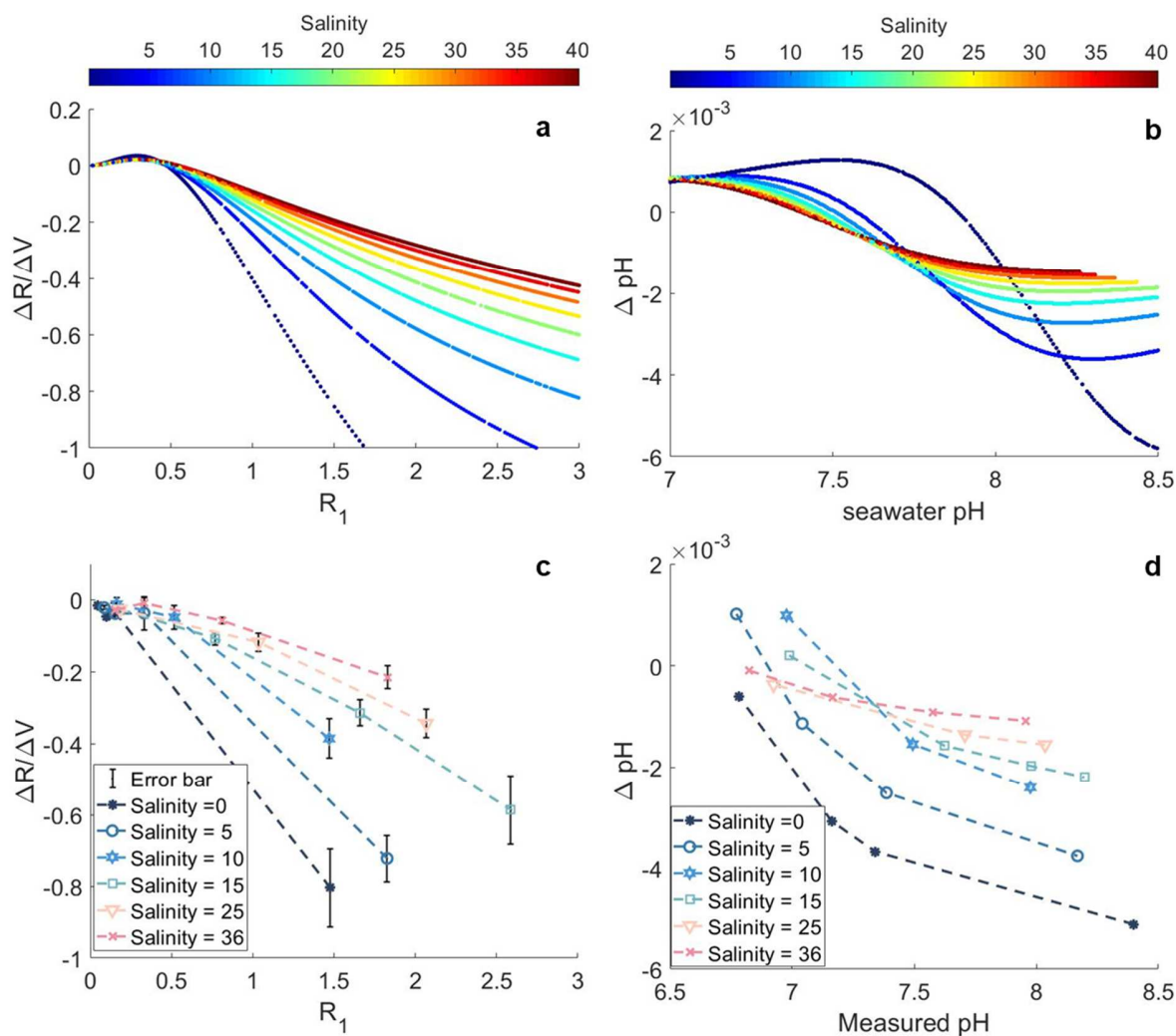


Figure 7. Simulation (a and b) and experimental (c and d) results of dye perturbation on the sample with fixed high TA ($\text{TA} = 2350 \mu\text{mol kg}^{-1}$) and variable S. The high S-TA is Gulf of Mexico seawater, while the low S – high TA is from Mississippi river water. Subplots a and c show that the decreasing rate of $\Delta R/\Delta R \sim R_1$ depends on the sample S. Subplots b and d represent the ΔpH (= pH with dye perturbation – pH without dye perturbation) changes of the water sample due to the indicator addition.

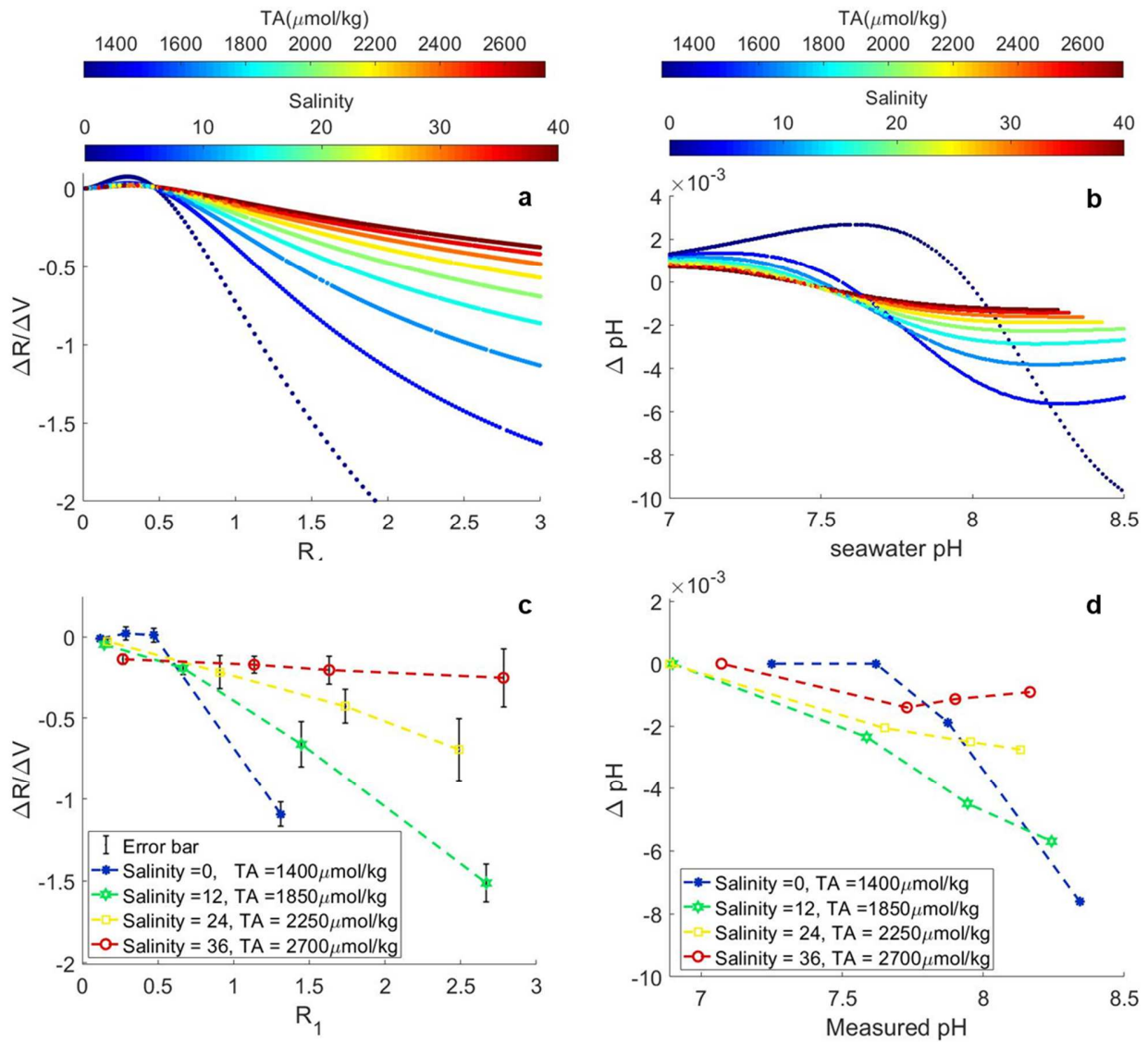


Figure 8. Simulation (a and b) and experimental results (c and d) of dye perturbation on the sample with fixed TA – S relationship. The high S-TA is Gulf of Mexico Seawater, while the low S – low TA is the Atchafalaya river water with HCl addition. Panels a and c shows that the decreasing rate of $\Delta R/\Delta V$ and R_1 decrease with TA and salinity increasing. Panels b and d shows the effect of dye perturbation on pH is smaller for high TA and salinity samples.

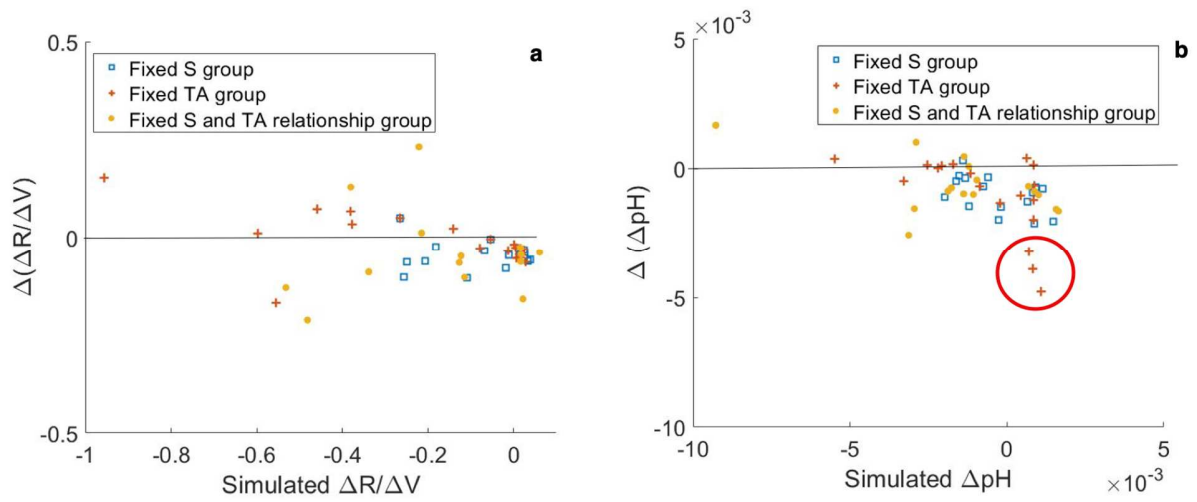


Figure 9. Comparison of simulated vs. experimental $\Delta R/\Delta V$ (a) and ΔpH (b). Y axis is the difference between experimental value and simulated value. Outliers in the red circle are the low salinity samples with fixed high TA. $\Delta(\Delta\text{pH})$ and its related uncertainties are limited by the accuracy and precision of spectrophotometric pH measurement and the carbonate chemical constants at low salinity ($S < 5$).

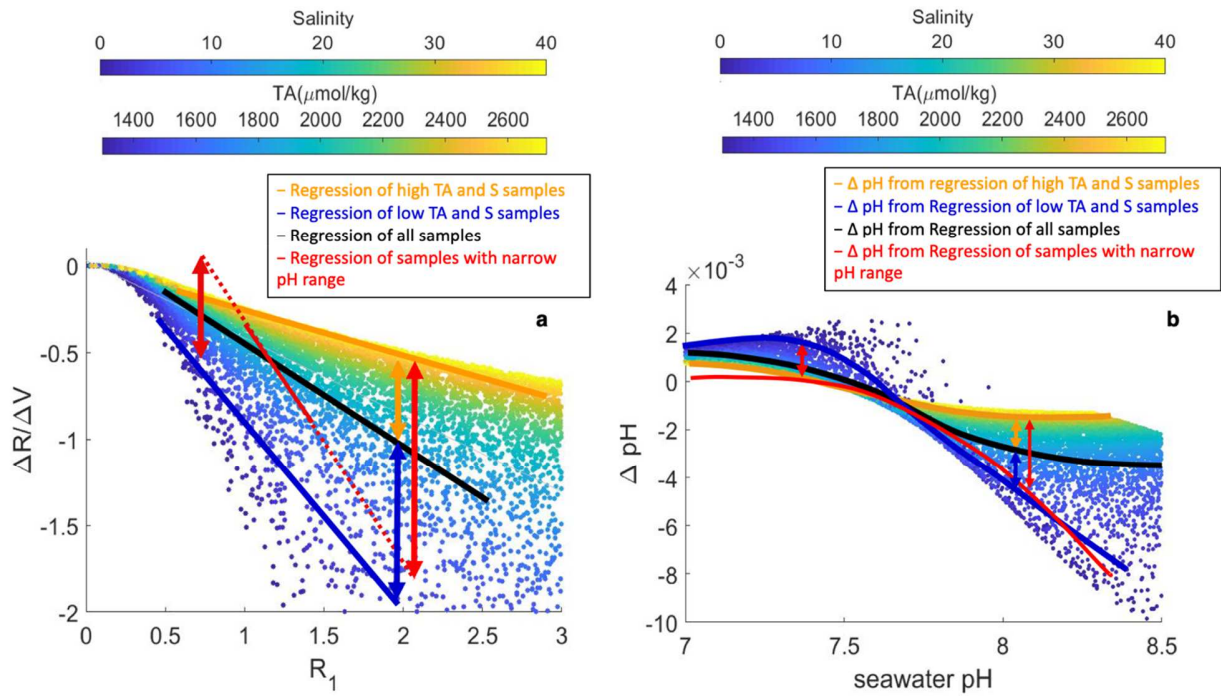


Figure 10. Possible pH biases caused by the empirical dye perturbation. Subplot a represents the distribution of $\Delta R/\Delta V$ versus R_1 . Subplot b shows the dye perturbation on pH (ΔpH). Black and red arrows indicate the differences of regression lines (subplot a) and the differences of ΔpH (subplot b).

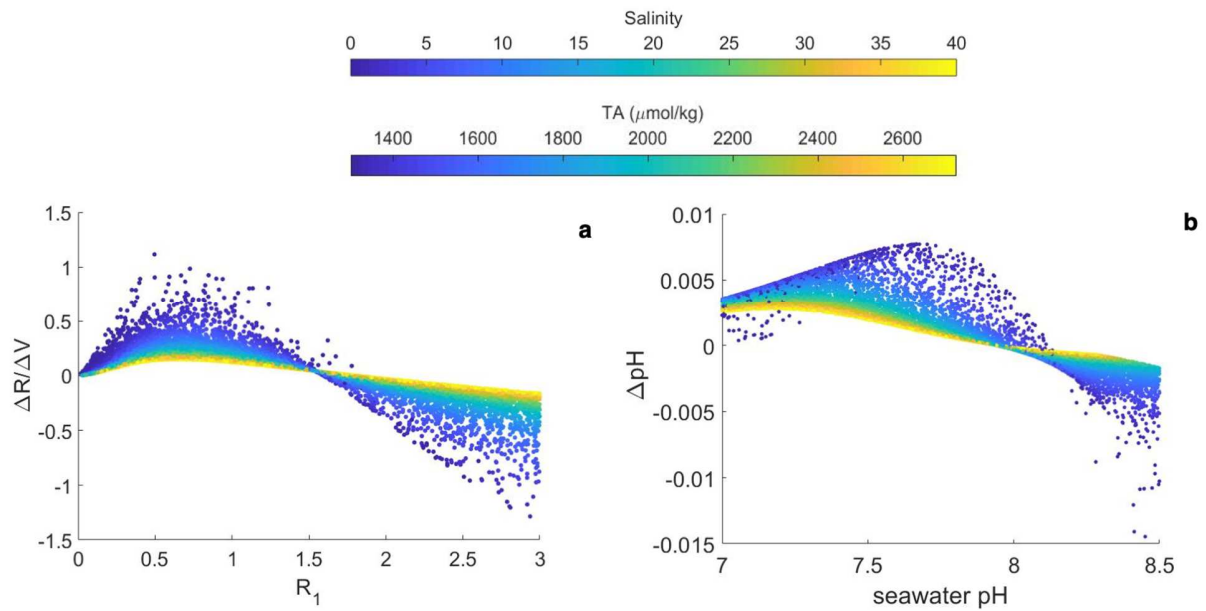


Figure 11. The dye perturbation on sample pH with dye $S = 35$, dye $\text{pH} = 8.0$. The influence of the dye perturbation is smaller for high S-TA water than for low S-TA water. Subplot a shows $\Delta R/\Delta V$ vs R_1 . $\Delta R/\Delta V$ vs. R_1 has a linear relationship when $R_1 > 0.5$. Subplot b shows that when sample salinity and pH are close to the dye, ΔpH is close to 0.

