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1 **Variations in the Influence of Diffuse Light on Gross Primary Productivity in Temperate**

2 **Ecosystems**

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17 **Abstract**

18 The carbon storage potential of terrestrial ecosystems depends in part on how 19 atmospheric conditions influence the type and amount of surface radiation available for 20 photosynthesis. Diffuse light, resulting from interactions between incident solar radiation and 21 atmospheric aerosols and clouds, has been postulated to increase carbon uptake in terrestrial 22 ecosystems. However, the magnitude of the diffuse light effect is unclear because existing 23 studies use different methods to derive above-canopy diffuse light conditions. We used site-24 based, above-canopy measurements of diffuse light and gross primary productivity (GPP) from 25 ten temperate ecosystems (including mixed conifer forests, deciduous broadleaf forests, and 26 croplands) to quantify the GPP variation explained by diffuse photosynthetically active radiation 27 (PAR) and to calculate increases in GPP as a function of diffuse light. Our analyses show that 28 diffuse PAR explained up to 41% of variation in GPP in croplands and up to 17% in forests, 29 independent of direct light levels. Carbon enhancement rates in response to diffuse PAR 30 (calculated after accounting for vapor pressure deficit and air temperature) were also higher in 31 croplands (0.011-0.050 μmol CO2 per μmol photons of diffuse PAR) than in forests (0.003-0.018 32 μmol CO2 per μmol photons of diffuse PAR). The amount of variation in GPP and carbon 33 enhancement rate both differed with solar zenith angle and across sites for the same plant 34 functional type. At crop sites, diffuse PAR had the strongest influence and the largest carbon 35 enhancement rate during early mornings and late afternoons when zenith angles were large, with 36 greater enhancement in the afternoons. In forests, diffuse PAR had the strongest influence at 37 small zenith angles, but the largest carbon enhancement rate at large zenith angles, with a trend 38 in ecosystem-specific responses. These results highlight the influence of zenith angle and the

49 **1. Introduction**

50 Forests are estimated to remove up to 27% of human-emitted CO_2 annually (2.6 \pm 0.8 Gt 51 C yr⁻¹), with temperate forests responsible for about half of this uptake globally (Le Quéré et al., 52 2013; Sarmiento et al., 2010). It is uncertain how this amount of carbon uptake will change in 53 the future because forest carbon processes are affected by complex interactions driven by 54 changes in climate and natural- and human-caused shifts in plant species composition and 55 canopy structure. Isolating and quantifying the impacts of individual drivers of land-atmosphere 56 CO2 exchange could improve these calculations of the future terrestrial carbon sink. 57 One important factor influencing photosynthesis and hence forest CO₂ uptake is light 58 availability. Rates of leaf-level $CO₂$ uptake increase with solar radiation until leaves are light 59 saturated (Mercado et al., 2009). This implies that forest $CO₂$ uptake is greater on sunny days 60 when leaves are fully exposed to direct light. However, increases in diffuse light, which is 61 produced when clouds and aerosols interact with and scatter incoming solar radiation, may be 62 even more beneficial than equal increases in direct light. At the ecosystem level, key processes 63 related to photosynthesis, including gross primary productivity (GPP), net ecosystem exchange 64 (NEE), and light-use efficiency (LUE), can increase in magnitude when the proportion of light 65 entering a forest canopy is more diffuse (Gu et al., 1999b; Hollinger et al., 1994; Jenkins et al., 66 2007; Oliphant et al., 2011; Urban et al., 2012; Zhang et al., 2011). In addition, global 67 simulations from 1960-1999 indicate that increases in the proportion of diffuse light reaching 68 plant canopy surfaces may have amplified the global land carbon sink by 24% (Mercado et al., 69 2009).

70 Several mechanisms have been proposed to explain how diffuse light increases 71 ecosystem CO2 uptake and LUE. First, diffuse light can penetrate deeper into a forest canopy

72 and reach lower canopy leaves that would normally be light-limited on clear days when light is 73 mostly direct (Hollinger et al., 1994; Oliphant et al., 2011). Second, the same amount of light is 74 distributed across more leaves when diffuse light is dominant, which can minimize light 75 saturation and photo-inhibition of upper canopy leaves and increase canopy LUE or 76 photosynthesis (Gu et al., 2002; Knohl and Baldocchi, 2008). Third, diffuse light can create 77 conditions favorable for photosynthesis by reducing water and heat stress on plants (Steiner and 78 Chameides, 2005; Urban et al., 2012). Finally, a fourth hypothesis suggests that diffuse light has 79 a higher ratio of blue to red light, which may stimulate photochemical reactions and stomatal 80 opening (Urban et al., 2012).

81 There is no consensus regarding the magnitude of effect that diffuse light has on 82 ecosystem carbon processing. Studies using derived values of diffuse light suggest that LUE is 83 higher when most incident light is diffuse and can result in maximum carbon uptake under 84 moderate cloud cover (Gu et al., 2002; Min and Wang, 2008; Rocha et al., 2004). However, 85 studies using a three-dimensional canopy model and a land surface scheme predict that diffuse 86 radiation will not lead to significant increases in carbon uptake on cloudy days as compared to 87 clear days because of reductions in total shortwave radiation (Alton et al., 2005; Alton et al., 88 2007). If clouds decrease surface radiation enough to lower total canopy photosynthetic activity, 89 this could offset any potential GPP gain resulting from increased LUE under diffuse light 90 conditions (Alton, 2008).

91 Several studies using measurements of diffuse light support the hypothesis that LUE is 92 higher under diffuse light, consistent with studies using derived diffuse light data (Dengel and 93 Grace, 2010; Jenkins et al., 2007). In addition, total carbon uptake can be greater under cloudy, 94 diffuse light conditions compared to clear skies in three forest types (Hollinger et al., 1994; Law

95 et al., 2002). Aerosol-produced diffuse light also leads to an increase in the magnitude of NEE 96 in forests and croplands (Niyogi et al., 2004). Additional observation-based analyses indicate 97 that diffuse light increases carbon uptake when compared to the same level of direct light, but 98 also when total light levels decrease (Hollinger et al., 1994; Urban et al., 2007; Urban et al., 99 2012).

100 The magnitude of the diffuse light effect on terrestrial carbon uptake may depend on 101 ecosystem type or canopy structural characteristics. A regional modeling study suggests that 102 diffuse light can increase net primary productivity (NPP) in mixed and broadleaf forests, but has 103 a negligible effect on croplands (Matsui et al., 2008). Another study using derived diffuse light 104 data suggests that LUE increases with diffuse light, and that differences among ecosystems are 105 potentially dependent on vegetation canopy structure (Zhang et al., 2011). The influences of 106 ecosystem type and vegetation structure are also supported by an observation-based study 107 showing that under diffuse light, $CO₂$ flux into a grassland decreased, but increased by different 108 amounts in croplands depending on the species of crop planted (Niyogi et al., 2004). However, 109 another study using derived diffuse light data found no difference in the effect of patchy clouds 110 on LUE among 23 grassland, prairie, cropland, and forest ecosystems in the Southern Great 111 Plains (Wang et al., 2008). Inconsistencies among these studies may be due to differences in the 112 methods and models used to obtain diffuse light or sky conditions and assess their impacts on 113 ecosystem carbon processing (Gu et al., 2003).

114 Climate modelers have begun incorporating the influence of diffuse light on ecosystem 115 carbon uptake into land surface schemes as more details of canopy structure are added to models 116 (Bonan et al., 2012; Dai et al., 2004; Davin and Seneviratne, 2012). Our study provides insight 117 into the importance of diffuse light on ecosystem carbon processing for improving projections of 118 the terrestrial carbon sink. We seek here to 1) quantify how much variation in ecosystem GPP is 119 explained by diffuse light, independent of direct radiation levels, 2) compare the influence of 120 diffuse light on GPP among temperate ecosystems differing in canopy structure and species 121 composition, and 3) determine the strength of diffuse light enhancement of GPP while 122 accounting for its correlation with zenith angle, vapor pressure deficit (VPD), and air 123 temperature. Unlike many previous studies (Alton, 2008; Butt et al., 2010; Gu et al., 1999b; Min 124 and Wang, 2008; Zhang et al., 2010), we drive our analyses only with direct field measurements 125 of diffuse light, rather than with derived values from radiation partitioning models, which may be 126 biased by incorrect representations of clouds and aerosols. Finally, our paper highlights the 127 changes in the diffuse light effect across the diurnal cycle and the role of time of day on the 128 diffuse light enhancement in terrestrial ecosystems, which will be important in scaling this effect 129 from individual sites to the globe.

130 **2. Materials and Methods**

131 *2.1 Data Sources*

132 All analyzed data were collected and processed by investigators participating in the 133 AmeriFlux program (http://ameriflux.lbl.gov/), a network of meteorological towers in the United 134 States (U.S.) that measures net fluxes of water vapor and CO₂ between the land surface and the 135 atmosphere and corresponding meteorological, soil, and vegetation conditions (Baldocchi, 2003). 136 Data collection, analysis, and metadata are standardized, reviewed, and quality controlled by 137 AmeriFlux for all sites. GPP is calculated by subtracting the modeled ecosystem respiration 138 from observed NEE. Respiration is modeled empirically based on NEE observations during the 139 night, when GPP is assumed to be zero. We focus our study on GPP instead of another measure 140 of carbon processing because it describes ecosystem CO₂ uptake, is affected directly by

141 radiation, and is the first step in processing atmospheric $CO₂$ into long-term storage in 142 ecosystems.

143 *2.2 Site Selection*

144 We selected temperate AmeriFlux sites within the contiguous U.S. with at least three 145 years of Level 2 (processed and quality controlled) NEE and GPP. Among these, we specifically 146 selected sites that contain equipment to measure above-canopy total and diffuse 147 photosynthetically active radiation (PAR, 400-700 nm) and report at least three years of diffuse 148 PAR values to AmeriFlux. For the University of Michigan Biological Station (UMBS), we 149 obtained updated total and diffuse PAR data from site coordinators that were not yet available on 150 the AmeriFlux website at the time of our analyses. After separating sites with crop rotations by 151 species, there were sufficient data for ten sites covering three ecosystem types, including mixed 152 forest (Howland Logged, Howland N Fertilized, Howland Reference), deciduous broadleaf forest 153 (Morgan Monroe and UMBS), and cropland (Mead Irrigated Maize, Mead Irrigated Rotation: 154 Maize, Mead Irrigated Rotation: Soybean, Mead Rainfed Rotation: Maize, Mead Rainfed 155 Rotation: Soybean). Site characteristics and data availability are listed in Table 1.

Site (SiteID)	Lat, Lon (°)	Years of Data	Canopy Height (m)	Vegetation Community	Management	LAI $(m2 m-2)$	Climatic Annual Precipitation (mm)	Mean Growing Season Temperature ^a $(^{\circ}C)$	Mean Growing Season VPDa (kPa)
Howland Logged $(US-Ho3)$	45.207, -68.725 45.204, -68.740 45.209, -68.747	2006-2008	20 _p	Dominated by red spruce (Picea rubens) and eastern hemlock (Tsuga canadensis). Also contains balsam fir (Abies balsamea), white pine (Pinus strobus), white cedar (Thuja occidentalis), red maple (Acer rubrum), and paper birch (Betula papyrifera) ^c .	Selected logging and harvest $(2001)^{b}$	2.1 to $~10$	1000 ^b	16.7	0.83
Howland Reference $(US-Ho1)$		2006-2008			Minimal disturbance since 1900s ^d	$\sim 6^{b}$		17.6	0.87
Howland N Fertilized $(US-Ho2)$		2006-2009			N addition $(2001 - 2005)^{d,e}$	$\sim 6^{b}$		16.5	0.82
Mead Irrigated Maize $(US-Ne1)$	41.165, -96.476	2001-2012	2.9 ^f	Maize (Zea mays)	Center-pivot irrigationf	5.7 ^e	887f	27.0	1.33
Mead Irrigated Rotation: Maize $(US-Ne2)$	41.164, -96.470	2001, 2003, 2005, 2007, 2009-2012	2.9 ^e 1.0 ^e	Maize (Z. mays)	Center-pivot irrigationf	5.3 ^e		26.2	1.14
Mead Irrigated Rotation: Soybean $(US-Ne2)$		2002, 2004, 2006, 2008		Soybean (Glycine max)		4.9 ^e			
Mead Rainfed Rotation: Maize $(US-Ne3)$	41.179, -96.439	2001, 2003, 2005, 2007, 2009, 2011	2.6 ^e	Maize (Z. mays)	Naturally rainfed ⁹	4.2 ^e 3.8 ^e		26.7	1.39
Mead Rainfed Rotation: Soybean $(US-Ne3)$		2002, 2004, 2006, 2008, 2010, 2012	0.9 ^e	Soybean (G. max)					
Morgan Monroe (US-MMS)	39.323, -86.413	2007-2010	27 ^h	Dominated by sugar maple (A. saccharum), tulip poplar (Liriodendron tulipifera), sassafras (Sassafras albidum), white oak (Quercus alba),	None	5^{i}	1012^{j}	24.3	1.12

156 Table 1: AmeriFlux site information and ecosystem characteristics

^aValues calculated from AmeriFlux data, ^bScott et al., 2004, ^cHollinger et al., 2004, ^dAmeriFlux website, ^epersonal communication

158 with site investigator, ^fYan et al., 2012, ^gVerma et al., 2005, ^hDragoni et al., 2011, ⁱOliphant et al., 2011, ^jCurtis et al., 2002, ^kGough et

159 al., 2013

160 *2.3 Definition of Analysis Period*

161 To determine the maximum effect of diffuse light on GPP, we limited our period of 162 analysis to the portion of the year when ecosystems are most productive. We used a carbon-flux 163 phenology approach, where NEE is the defining variable for phenological transitions and the 164 peak-growing season is the time period when NEE is at its maximum magnitude (Garrity et al., 165 2011). To do this, we first calculated 5-day NEE means for each site and year. Climate, 166 vegetation composition, and inter-annual weather variability lead to phenological variation 167 among sites (Richardson et al., 2013). Therefore, we adjusted our definition for the beginning 168 and end of the peak-growing season to uniformly capture a representative portion of the NEE 169 peak across sites and years. We defined the start of the growing season as the first day when the 170 5-day NEE average was within 90% of the year's fourth highest 5-day NEE average. The 171 fourth-highest value was used to account for any extreme NEE values that may have occurred 172 because of anomalous weather conditions. We set the end of season as the last day within 75% 173 of the year's fourth-highest 5-day NEE average. The cutoff for the start of the peak-growing 174 season is higher than the cutoff for the end of the season because canopy leaf-out and growth 175 initiation typically occur quickly in seasonal sites, whereas canopy phenological changes are 176 slower at the end of the season. While this approach cannot detect the exact beginning and end 177 of the season, the criteria we used provide a uniform method for defining the period during 178 which plants were at full seasonal growth and activity at our sites. We included only daytime 179 values by excluding points with total PAR values \leq 20 µmol m⁻² s⁻¹, assuming such low radiation 180 levels are characteristic for nighttime.

181 *2.4 Data Analysis*

182 For each site, we combined all available peak-growing season daytime data and removed 183 observations with negative measurements of diffuse PAR, direct PAR, or GPP, as these were 184 likely sensor errors or marginal weather conditions (e.g., rain events). We also excluded data 185 points with missing air temperature and VPD. We divided the remaining data into nine 186 categorical groups based on solar zenith angle and the time of observation. We chose to bin by 187 zenith angle to account for the effect of the sun's position on the amount of direct and diffuse 188 PAR above a canopy, differences in radiation penetration through the canopy, and changes in 189 plant hydraulics throughout the day. Zenith angle was calculated as the following: 190 cos $\varphi = \sin \phi \sin \delta + \cos \phi \cos \delta \cos [15(t - t_0)]$ (1) 191 where φ is the zenith angle, ϕ is the latitude, δ is the solar declination angle, *t* is time, and *t*⁰ is the 192 time of solar noon (Campbell and Norman, 1998). Given the latitudes of the sites, we defined 193 mornings to begin at zenith angles between 76-100°, noon to occur at the minimum calculated 194 zenith angles of 16-30°, and the end of daylight to occur around 76-100°. 195 The effect of diffuse PAR on GPP may depend on total light conditions. For example, 196 little scattering occurs under clear skies, which results in low diffuse and high direct PAR levels. 197 As a result, small increases in diffuse PAR are unlikely to have a strong impact on canopy 198 photosynthesis due to large amounts of direct PAR available for photosynthesis. If direct PAR 199 levels are low, however, such as on cloudy days or during the morning and evening, the increase 200 in diffuse PAR will have a larger effect because canopy leaves are below light-saturation. To 201 calculate direct PAR, we subtracted the observed diffuse PAR from the observed total PAR. 202 Because GPP and PAR are known to have a strong relationship that can be empirically described

203 by a rectangular hyperbola, we used the non-linear regression function in the R program (R

204 Development Core Team, 2012) to fit the following relationship:

$$
GPP_{fitted} = (\alpha \gamma \ PAR_{dir}) / (\gamma + \alpha \ PAR_{dir}) \tag{2}
$$

206 where *GPPfitted* is the value of GPP predicted by total PAR using a rectangular hyperbola model 207 (Eq. 2), α is the canopy quantum efficiency, γ is the canopy photosynthetic potential, and *PARdir* 208 is direct PAR (Gu et al., 2002). The α and γ are the fitted parameters and are solved iteratively. 209 We used the initial conditions of 0.044 µmol CO₂ per µmol photons and 23.7 µmol CO₂ m⁻² s⁻¹ 210 for α and γ , respectively (Ruimy et al., 1995). The resulting empirical relationships for each site 211 are presented in Appendix 1.

212 To remove the confounding effect of direct PAR, we first calculated the residuals 213 between observed GPP and *GPPfitted*. We then compared those residuals against diffuse PAR for 214 ten sites and nine zenith angle bins. For each zenith angle category, we estimated the variation 215 in GPP residuals that can be explained by diffuse PAR alone using the following simple linear 216 regression:

$$
GPP_r = GPP - GPP_{fitted} = \beta_0 + \beta_1 PAR_{diff} + \varepsilon
$$
\n(3)

218 and a combination of diffuse PAR, VPD, and air temperature using the following multiple linear 219 regression:

$$
GPP_r = GPP - GPP_{fitted} = \beta_0 + \beta_1 PAR_{diff} + \beta_2 VPD + \beta_3 T_a + \varepsilon
$$
\n(4)

221 where *GPPr* represents the residuals between the observed GPP and *GPPfitted* and *PARdiff* is 222 diffuse PAR. *T_a* is air temperature measured at the eddy covariance tower and β_0 , β_1 , β_2 and β_3 223 are the fitted parameters estimating the model intercept and the linear slopes of the effects of 224 diffuse PAR ,VPD, and air temperature at each solar zenith bin, respectively. The *ε* is the error 225 term.

226 ANOVA comparisons between the simple (diffuse PAR only) and multiple linear 227 regressions (including VPD and air temperature) showed that the multiple linear regression 228 model (Eq. 4) was significantly better ($p \le 0.05$) than the simple regression model, with the 229 exception of nine site/bin combinations. We did not include interactions in the multiple linear 230 regression because ANOVA tests indicated that the interaction terms did not improve the model 231 consistently, and improvements to the residual sum of squares averaged only 3.5% in cases 232 where interaction terms were significant. We also accounted for multiple testing over solar 233 zenith angle bins and different sites by using the Bonferroni correction to calculate a new critical 234 p-value. Light-response curves could not be fit to all scenarios, reducing the final number of 235 comparisons to 83. Thus, for the simple and multiple linear regression comparisons, we consider 236 a relationship significant if $p \le 6.02 \times 10^{-4} (= 0.05/83)$.

237 **3. Results**

238 *3.1 Relationship between diffuse PAR and GPP*

239 We found significant positive relationships between diffuse PAR and *GPPr* throughout 240 the day, except in a few cases where diffuse PAR was not a significant predictor of *GPPr* (Fig. 1, 241 Fig. 2, black bars). Exceptions to these relationships occurred mainly at the Mead crop sites 242 during mid-day and to a lesser extent at the UMBS forest during early mornings and late 243 afternoons (Fig. 2, black bars). In addition, a rectangular hyperbola could not be fit to the direct 244 PAR and GPP data in the afternoon at large zenith angles at the Mead sites and Morgan Monroe 245 (Appendix 1). Overall, the linear fits between diffuse PAR and *GPPr* indicate that across sites 246 and zenith angles, diffuse PAR explains 3-22% of variation in *GPPr* in the morning and 3-41% 247 of variation in *GPPr* in the afternoon (Fig. 2, black bars).

Fig 1. Simple linear regressions (Eq. 3) between diffuse PAR and GPP_r for observations around 10:00 -14:00 standard time (zenith angles from 16-30°, other zenith angle bins not shown). Regression lines are only plotted for models with $p < 6.02 \times 10^{-4}$ (Bonferroni-corrected critical value).

Fig 2: Proportions of variation in GPP_r explained by environmental variables. Solid bars represent R² values from simple linear regressions that include only the effect of diffuse PAR (Eq. 3). The total height of the bars (solid and white together) represents the R^2 from multiple linear regressions that include effects of air temperature (T_a) and vapor pressure deficit (VPD) with diffuse PAR (Eq. 4). Only R² values with p < 6.02 x 10⁻⁴ (Bonferroni-corrected critical value) are plotted. The minimum calculated zenith angle for these sites was ~16°.

256 *3.2 Diffuse PAR cross-correlation with VPD and air temperature*

257 Concomitant with changes in the partitioning of PAR into direct and diffuse streams, 258 clouds and aerosols change surface VPD and air temperature. These two environmental factors 259 influence stomatal conductance and photosynthesis, and thus affect rates of ecosystem GPP. 260 When including the effects of these two variables on *GPPr* with diffuse PAR (Eq. 4), the amount 261 of variation in *GPPr* explained increases up to an additional 31% during mornings and up to 32% 262 during afternoons (Fig 2, white bars). This increase with VPD and air temperature is greatest 263 across the most zenith angles at the Howland sites, where the multiple linear regression increases 264 explanatory power of *GPPr* by an additional 9-27% and 11-30% in the mornings and afternoons, 265 respectively. VPD and air temperature also account for a relatively larger fraction of the 266 variation of Mead Rainfed Rotation: Soybean *GPP^r* during the mid-day. Although we expected 267 an increase in explanatory power with more variables in the regression, the increase in the 268 explanation of *GPPr* with the addition of these correlated environmental variables is small for the 269 deciduous forests (Morgan Monroe and UMBS). This suggests that the effect of diffuse PAR at 270 the deciduous forests is due to changes in light availability and not from indirect effects driven 271 by the cross-correlation between diffuse PAR and other environmental conditions. Overall, the 272 multiple linear regressions indicate that diffuse PAR is a significant predictor of *GPPr* (except 273 for the sites and zenith angle bins noted in Table 2). In addition, VPD and air temperature could 274 not account for significant amounts of *GPPr* variation under some conditions (Table 2).

275 Table 2: Parameter estimate values from relationships between GPP_r and diffuse PAR, vapor pressure deficit (VPD), and air temperature (T_a). All 276 β_i estimate values (Eq. 4) have p < 6.02 x 10⁻⁴ (Bonferroni-276 β_i estimate values (Eq. 4) have p < 6.02 x 10⁻⁴ (Bonferroni-corrected critical value), except for those designated as NS.

277 *3.3 Magnitude of the effects of diffuse PAR on GPP^r*

278 Howland Forest Reference, Morgan Monroe, and UMBS have not undergone any 279 experimental manipulation (e.g., selective logging, N addition). At these sites, the sign of the 280 significant parameter estimates indicate that in mornings and afternoons, *GPPr* increased with 281 diffuse PAR (Table 2). The predicted increases in GPP_r in the morning were calculated to be 282 0.004-0.010, 0.008-0.011, and 0.010-0.018 μmol CO2 per μmol photons of diffuse PAR at 283 Howland Forest Reference, Morgan Monroe, and UMBS, respectively (Fig. 3). In the afternoon, 284 the increases in *GPPr* were similar in magnitude, and ranged from 0.005-0.011, 0.008-0.009, and 285 0.009-0.018 μmol CO2 per μmol photons of diffuse PAR at Howland Forest Reference, Morgan 286 Monroe, and UMBS, respectively (Fig. 3). 287 The effect of diffuse PAR on rates of *GPPr* varied among forest sites. UMBS had the 288 largest increases in *GPPr* with increases in diffuse PAR, and Howland Forest Reference had the 289 smallest increases in *GPPr*. In addition, the calculated increases in *GPPr* with diffuse PAR 290 appear to depend on zenith angle at two of the sites. At UMBS, the influence of diffuse PAR on 291 *GPP_r* is greatest in the early morning and late afternoon (zenith angles 61-75 $^{\circ}$) and decreases at 292 mid-day (zenith angles 16-45°). At Howland Forest Reference, the response to zenith angle 293 differs and the influence of diffuse PAR on *GPPr* generally increases as the day continues and is 294 highest in the late afternoon (zenith angles 76-100°). However, at Morgan Monroe, the influence 295 of diffuse PAR on *GPPr* did not vary with zenith angle. When we compare across these 296 ecosystems, deciduous forests (UMBS, Morgan Monroe) appear to differ from the mixed conifer 297 forest, particularly in the morning, with differences diminishing in the afternoon.

Fig 3. Diurnal patterns in diffuse PAR β estimates for unmanaged forests across zenith angles from a multiple linear regression that includes VPD and air temperature as covariates (Eq. 4). Error bars indicate one standard error. Only β estimates with p < 6.02 x 10⁻⁴ (Bonferroni-corrected critical value) are plotted.

298

299 At Howland Forest, one site underwent selective logging while a second site was

300 fertilized with 18 kg N/ha on a 21-hectare plot centered around the eddy covariance tower in five

301 to six applications per growing season from 2001-2005 (David Dail, personal communication,

302 2013). Analysis of data at these manipulated sites indicates that the magnitude of increase in

- 303 *GPPr* with diffuse PAR was similar to that of the un-manipulated Howland forest (Fig. 4).
- 304 Differences among forest treatments are not apparent in the morning. In the afternoon, however,
- 305 we observe a trend where diffuse PAR leads to the biggest *GPPr* increase in the forest fertilized
- 306 with N and the smallest change in *GPPr* in the forest that has been selectively logged.

Fig 4: Diurnal patterns in diffuse PAR β estimate values for Howland Forest sites across zenith angles from a multiple linear regression that includes VPD and air temperature as covariates (Eq. 4). Error bars indicate one standard error. Only values with $p < 6.02 \times 10^{-4}$ (Bonferroni-corrected critical value) are plotted.

307

308 At the Mead Irrigated Rotation and Mead Rainfed Rotation sites, soybean and maize are 309 planted in different years, allowing us to examine variations in the effect of diffuse PAR on 310 *GPPr* between crop types (Fig. 5). The increases in *GPPr* for maize were calculated to be 0.011- 311 0.022 μmol CO₂ per μmol photons in the morning and 0.021-0.050 μmol CO₂ per μmol photons 312 in the afternoon. For soybean, the increases in GPP_r in the morning were 0.014-0.021 µmol CO_2 313 per umol photons and in the afternoon were 0.011 -0.028 umol CO₂ per umol photons. Diffuse 314 PAR led to increases in *GPPr* at large zenith angles, but had no effect on *GPPr* at small zenith 315 angles for both crop species (values are only plotted in Fig. 5 if they are significant). In addition, 316 we observed no difference in the magnitude of the effect of diffuse PAR on *GPPr* between 317 soybean and maize in the morning. However, we did observe a greater effect of diffuse PAR on

318 *GPPr* for maize than soybean in the afternoon for zenith angles 46-75°. Irrigation did not appear

Fig 5: Diurnal patterns in diffuse PAR *β* estimate values for Mead crop sites across zenith angles from a multiple linear regression that includes VPD and air temperature as covariates (Eq. 4). Error bars indicate one standard error. Only β values with $p < 6.02 \times 10^{-4}$ (Bonferroni-corrected critical value) are plotted.

320

321 **4. Discussion**

322 Diffuse light influences Earth's climate by changing the amount and character of light 323 available for photosynthesis, and thus, indirectly controls atmospheric $CO₂$ (Mercado et al., 324 2009). Depending on future anthropogenic emissions and their effects on atmospheric aerosols 325 and clouds, the influence of diffuse light on the terrestrial carbon sink may increase. A more 326 quantitative and mechanistic understanding of the link between diffuse light and land carbon 327 uptake in different ecosystems would allow us to model how changes in diffuse light influence

329 afforestation, and conversion of natural systems to cropland) continues (Arora and Boer, 2010). 330 Past research has identified a positive correlation between diffuse light and ecosystem 331 carbon uptake. However, this result may be due to a cross-correlation with total light 332 availability, where diffuse light could more strongly influence photosynthesis when total light 333 levels are low on overcast days as compared to high light levels on clear days (Gu et al., 1999b; 334 Oliphant et al., 2011; Zhang et al., 2010). The method we use in this paper addresses this 335 confounding factor by removing the effect of direct light on ecosystem carbon uptake before 336 calculating the rate of additional carbon uptake from diffuse light. Importantly, we tested for this 337 potential independent effect using only direct field measurements of diffuse light, as opposed to 338 deriving diffuse light levels with radiation partitioning models that make assumptions about 339 aerosol and cloud conditions over terrestrial ecosystems. Our analysis of ten temperate 340 ecosystems indicates that diffuse PAR correlates positively with *GPPr* and this relationship is 341 independent of direct PAR levels. Specifically, diffuse PAR independently explained up to 22% 342 of the variation in *GPPr* in mornings and up to 41% of the variation in *GPPr* in afternoons. 343 Prior research shows that morning and afternoon responses to diffuse light can differ for 344 the same zenith angles (Alton et al., 2005) and that in multiple ecosystems, rates of carbon 345 enhancement vary across zenith angles (Bai et al., 2012; Zhang et al., 2010). However, to our 346 knowledge, no other studies have investigated full diurnal patterns of diffuse light enhancement. 347 We accomplished this by separating data according to zenith angle and time of day. Our results 348 indicate that in forests, the proportion of variation in *GPPr* explained by diffuse PAR (evaluated 349 through R^2) is greatest at mid-day, and decreases as the sun moves closer to the horizon. The 350 opposite pattern occurs at crop sites, where diffuse PAR did not predict *GPPr* at small zenith

328 atmospheric and terrestrial carbon stocks, particularly as land-use change (e.g., deforestation,

351 angles (mid-day), but did correlate with variation in *GPPr* at larger zenith angles (morning and 352 afternoon). When we examined the magnitude of increase in *GPPr* in response to diffuse PAR 353 (*β1*), the greatest increases were at larger zenith angles in crop sites (0.028 - 0.050 μmol CO2 per 354 μmol photons at 61-75° in the afternoon). In forests, however, diffuse PAR had the strongest 355 influence (\mathbb{R}^2) on *GPP_r* at small zenith angles when the sun is overhead (mid-day), but the largest 356 carbon enhancement rate (*β1*) at larger zenith angles (early mornings and late afternoons) when 357 the sun is closer to the horizon.

358 In addition, some sites show a trend in an asymmetrical diurnal cycle of diffuse light 359 enhancement, most notably in the crop sites. Although increases in *GPPr* with diffuse PAR at 360 forest sites appear to be similar in magnitude throughout the day, some of the zenith angle bins 361 differed between the morning and afternoon. For example, the largest difference in carbon 362 enhancement rates from a morning zenith angle bin to the same bin in the afternoon were 0.005 363 μmol CO2 per μmol photons for mixed conifer forests, 0.003 μmol CO2 per μmol photons for 364 deciduous forests, 0.017 μmol CO₂ per μmol photons for soy, and 0.028 μmol CO₂ per μmol 365 photons for maize, though changes were usually within the standard error of the measurements. 366 The response of *GPPr* to diffuse light may differ in the morning and afternoon because 367 environmental conditions influencing photosynthesis also vary during the day. For example, 368 time lags between the effects of diurnal cycles of radiation and VPD on evapotranspiration 369 (Zhang et al., 2014), stronger hydraulic stresses in the afternoon (Matheny et al., 2014), and 370 morning and afternoon differences in leaf surface wetness that affect stomatal conductance 371 (Misson et al., 2005) might explain the increased importance of diffuse light in the afternoon. 372 These results can be used to evaluate ecosystem and global land surface models by testing if they 373 capture the diurnal patterns we identified.

374 Our results indicate that there are ecosystem-specific responses of carbon uptake to 375 diffuse light. The observed differences between crops and forests are consistent with Niyogi et 376 al., 2004 who used measured diffuse shortwave data to show that a crop site with a corn and 377 soybean rotation was more sensitive to increases in aerosol-produced diffuse light than broadleaf 378 and mixed conifer forests. Previous studies have hypothesized that differences in canopy 379 structure among forests, grasslands, and croplands are responsible for differential responses of 380 these ecosystems to diffuse light (Gu et al., 1999a; Niyogi et al., 2004; Oliphant et al., 2011). 381 However, they have not reported site-level canopy architectural measurements to test this 382 potential modifier of land carbon uptake because they are difficult to collect and describe. 383 There are several hypotheses explaining why canopy structure may modify the effect of 384 diffuse light on ecosystem carbon uptake. Canopy gaps, which interact with the angle of 385 incident light, may influence how much light is distributed vertically through a canopy 386 (Hutchison et al., 1980). For example, on clear days in a 30-m tall tulip poplar forest, the 387 amount of radiation reaching the mid- and lower-parts of the canopy is lowest at large zenith 388 angles (Hutchison et al., 1980). The authors attributed this to the low level of total radiation and 389 reduced canopy gaps when the sun is near the horizon. Our analysis of UMBS gap fraction data 390 derived from LAI-2000 measurements shows that as gap fraction decreases, carbon uptake with 391 diffuse light increases (Fig. 6). Because gap fraction here is the ratio of below-canopy PAR to 392 above-canopy PAR, this indicates greater light extinction at larger zenith angles. Greater light 393 extinction in the canopy may increase light scattering, which could expose more leaves to diffuse 394 light. Thus, the response of GPP to diffuse light may be greater at larger zenith angles because 395 of more complete canopy participation in photosynthesis. However, more gap fraction data and 396 canopy light profiles from across sites and collected with uniform methods are needed to test this

397 idea, particularly in crop ecosystems. This would allow us to identify why crops and forests

398 respond differently to diffuse PAR.

Fig. 6. The relationship at UMBS (data from 2007-2011) between a) gap fraction and zenith angle and b) diffuse PAR β (carbon enhancement rate) and zenith angles (same data as shown in Fig. 3). Error bars indicate one standard error.

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400 Second, the distribution of photosynthetic tissues within a canopy depends on the plant 401 community at each site and may contribute to observed differences between crops and forests. 402 Forests have more stratified layers of vegetation and are much taller than crops. This means that 403 leaf area index (LAI) in a forest is distributed over a larger volume than in crop sites. When the 404 sun is overhead, forest canopies shade leaves at lower layers and diffuse light has a greater 405 potential of reaching leaves near the bottom of the forest canopy as compared to direct light. 406 Thus, the opportunity for diffuse light to reach more leaves in the canopy is greater when the sun 407 is overhead (larger \mathbb{R}^2). This explanation is supported by a study in a Norway spruce forest,

408 which showed that needles deeper in the canopy contribute more to overall net ecosystem 409 production on cloudy days than on sunny days (Urban et al., 2012). However, the relative 410 increases in GPP (*β1*) may be smaller than those at crop sites because forest canopies are denser, 411 which increases self-shading. On the other hand, crops are planted to minimize self-shading 412 when the sun is overhead. In addition, *β1* may be higher at crop sites than at forests because 413 multi-directional diffuse light at large zenith angles may reach deeper into crop canopies more 414 effectively than direct light and increase light availability for crop stems, which are more 415 photosynthetic than tree trunks.

416 Modeling studies have shown that species-dependent canopy characteristics, such as leaf 417 clumping, LAI, and leaf inclination angle can affect the influence of diffuse PAR on carbon 418 processing in ecosystems (Alton, 2008; Gu et al., 2002; Knohl and Baldocchi, 2008). This could 419 be due to the penumbral effect, which occurs when the position and types of leaves (e.g., 420 broadleaf and conifer) alter the amount and distribution of light to lower-level leaves (Denholm, 421 1981; Way and Pearcy, 2012). Although the arrangement of leaves in tall canopies with small 422 leaves (e.g., forests) can increase shading of lower canopy leaves, it also increases the probability 423 that leaves and branches scatter light, resulting in more distribution of light in the canopy. 424 However, in shorter canopies with larger leaves (e.g., maize), there is less plant material that can 425 scatter light and these sites may be more dependent on incident diffuse light. This may explain 426 the higher carbon enhancement rates observed at crop sites compared to forests.

427 A few studies have measured how the distribution of light through plant canopies 428 changes under diffuse light, but they are limited in their ability to test the influence of canopy 429 structure on carbon enhancement from diffuse light because they have been conducted in a single 430 ecosystem (Urban et al., 2012; Williams et al., 2014). Because site-level measurements of

431 canopy structure are difficult to obtain, support for the mechanisms through which specific 432 characteristics of canopy structure (e.g., leaf area distribution, leaf clumping) change ecosystem 433 carbon uptake under diffuse light conditions has thus far depended on model assumptions (Alton 434 et al., 2007; Knohl and Baldocchi, 2008). To test whether canopy structural differences in 435 height, canopy gaps, or leaf distribution within a canopy facilitate a diffuse light enhancement, a 436 uniform method of collecting canopy structural data is needed. Methods are available for 437 capturing some of this information, including light detection and ranging (LIDAR) remote 438 sensing (Hardiman et al., 2013). However, no standardized method of collecting data has been 439 applied among sites to allow for inter-site comparisons of canopy structure. Future research 440 should consider collecting data on canopy gaps, leaf distribution, and vertical light distribution to 441 provide datasets that can be used to test whether gaps or leaf distribution within a canopy lead to 442 an enhanced carbon uptake because of increased light distribution. Without this mechanistic 443 connection, modelers cannot determine whether this missing biosphere-atmosphere connection 444 results in a significant under- or over-prediction of the future terrestrial carbon sink. As 445 scientists collect these canopy structural data, we suggest making these data publically available 446 so they can be used to better interpret patterns seen using eddy covariance data.

447 We also observed differences in diffuse light effects among sites described as the same 448 forest type (e.g., Morgan Monroe and UMBS). This argues for the consideration of site-specific 449 responses to diffuse light because plant community composition of individual forest types (or 450 ecosystems) determine unique canopy structures that can drive how strongly canopy gaps, leaf 451 distribution, and penumbral qualities influence the effect of diffuse light on ecosystem carbon 452 uptake. In particular, there were differences in afternoon carbon enhancement rates between the 453 fertilized and formerly logged Howland Forest sites, which only differ in disturbance activity.

454 Differences in nutrient availability for plants may explain why the N fertilized site correlated 455 more strongly with diffuse light than the logged site. After two years of fertilization, foliage was 456 one of the most N-enriched ecosystem pools (Dail et al., 2009). Increased soil N availability 457 could lead to an increase in leaf N, which correlates with higher concentrations of RuBisCO and 458 chlorophyll (Evans, 1989), implying an interaction between diffuse light and nutrient levels.

459 The effect of diffuse light on carbon uptake between maize and soybean also differed. 460 This may be due to species differences in canopy structure as discussed above, but could also be 461 due to the different photosynthetic pathways soy (C_3) and maize (C_4) use. Maize had a greater 462 increase in carbon uptake with diffuse light than soy did, potentially because C4 plants have a 463 higher light saturation point (Greenwald et al., 2006). Because maize would be farther away 464 from light saturation than soy, an increase in diffuse light (after accounting for cross-correlation 465 with direct light) would bring maize closer to light saturation and thus, increase photosynthesis. 466 In addition, C_4 plants are better adapted to warmer environments, which may cause 467 environmental conditions, such as temperature and water availability, to change crop responses 468 to diffuse light.

469 Finally, our results show that other environmental drivers that co-vary with diffuse PAR 470 also contribute to *GPPr* at some sites. In mixed conifer forests (e.g., the Howland sites), VPD, 471 air temperature, and diffuse PAR together account for substantially more variation in *GPP^r* than 472 diffuse PAR itself does, implying a lesser role for radiation and a larger one for conditions that 473 improve stomatal conductance under cloudy conditions at mixed conifer forests. In contrast, 474 VPD and air temperature, within the ranges of values characteristic of measurement periods at 475 the sites studied here, appear to have small effects on *GPPr* in the broadleaf forests. This implies 476 that the diffuse PAR effect at the broadleaf forests is due to the effect of scattered light itself. At

477 the mixed conifer forests, the peak growing season temperature ranges from 16.5-17.6°C while 478 the temperature is 21.2-24.3°C in the broadleaf forests. Comparing these site temperatures to the 479 optimum temperature range of temperate deciduous trees (20-25°C) and evergreen coniferous 480 trees (10-25°C), broadleaf forests are closer to their optimum temperature range (Larcher, 2003). 481 Considering that photosynthesis varies non-linearly with temperature, the same per unit change 482 in temperature for a cooler site will lead to greater changes in GPP than in a warmer site. 483 Increases in VPD in water-limited situations, on the other hand, should cause photosynthesis to 484 drop because stomata will close to conserve water. However, VPD is actually lower in the 485 mixed forests than in the deciduous broadleaf forests, implying that air temperature is a stronger 486 driver of GPP than is VPD under our study's field conditions.

487 **5. Conclusions**

488 Field measurements show that diffuse PAR accounts for a substantial amount of variation 489 in GPP once the quantity of direct PAR is removed. The observed changes in the diffuse PAR 490 effect on *GPPr* vary across zenith angles, ecosystem types, and plant functional groups, 491 highlighting additional ways that ecosystem structural characteristics and the diurnal cycle 492 influence ecosystem carbon cycling. In addition, observed site-level variation suggests that 493 grouping forests together in regional or global models as the same plant functional type, without 494 considering species composition or canopy structure, may lead to inaccuracies in assessing the 495 impacts of radiation partitioning on modeled surface carbon fluxes.

496 To robustly extend these results, direct measurements of diffuse PAR and ecosystem flux 497 data are needed from a wider range of ecosystems. Furthermore, research that can evaluate 498 mechanisms (e.g, canopy gaps, leaf distribution, and species-specific characteristics) driving 499 terrestrial carbon enhancement under diffuse light will remain stagnant without consistent field

500 measurements of canopy structure at sites with diffuse light and eddy covariance measurements. 501 The incorporation of standard methods for measuring canopy structure and within-canopy light 502 distribution and the availability of these data in common formats from across networks of eddy 503 covariance towers (e.g., AmeriFlux, NEON) would enable the development of better predictive 504 models of carbon exchange in relation to direct and diffuse solar radiation.

505 The interactions between diffuse light and ecosystem productivity may be of increasing 506 importance as the community composition of our terrestrial ecosystems continues to change 507 because of human land use change, natural ecological succession, and climate change. Thus, a 508 more refined understanding of how diffuse PAR modifies atmosphere-land carbon cycling and 509 subsequent representations of this relationship in models will likely advance our understanding 510 of how human management of ecosystems will influence the land carbon sink as well as improve 511 future calculations of atmospheric CO₂ concentrations for global climate projections.

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532 **Appendices**

534 Appendix 1: Values of α and **γ** predicted by best-fit rectangular hyperbolas describing the response of

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538 535 GPP to direct PAR. The α represents the quantum yield and **γ** represents the maximum GPP value. All α 536 and *γ* values listed have $p < 6.02 \times 10^{-4}$ (Bonferroni-corrected critical value), except for those in italics, 537 which have p < 0.01 and those in **bold**, which were not significant because p > 0.05. NS indicates we were unable to fit a light response curve.

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