

1           ASSESSING THE EVOLUTION OF SOIL MOISTURE AND VEGETATION  
2           CONDITIONS DURING THE 2012 UNITED STATES FLASH DROUGHT

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Abstract

This study examines the evolution of several model-based and satellite-derived drought metrics sensitive to soil moisture and vegetation conditions during the extreme flash drought event that impacted major agricultural areas across the central U.S. during 2012. Standardized anomalies from the remote sensing based Evaporative Stress Index (ESI) and Vegetation Drought Response Index (VegDRI) and soil moisture anomalies from the North American Land Data Assimilation System (NLDAS) are compared to the United States Drought Monitor (USDM), surface meteorological conditions, and crop and soil moisture data compiled by the National Agricultural Statistics Service (NASS).

Overall, the results show that rapid decreases in the ESI and NLDAS anomalies often preceded drought intensification in the USDM by up to 6 weeks depending on the region. Decreases in the ESI tended to occur up to several weeks before deteriorations were observed in the crop condition datasets. The NLDAS soil moisture anomalies were similar to those depicted in the NASS soil moisture datasets; however, some differences were noted in how each model responded to the changing drought conditions. The VegDRI anomalies tracked the evolution of the USDM drought depiction in regions with slow drought development, but lagged the USDM and other drought indicators when conditions were changing rapidly. Comparison to the crop condition datasets revealed that soybean conditions were most similar to ESI anomalies computed over short time periods (2-4 weeks), whereas corn conditions were more closely related to longer-range (8-12 week) ESI anomalies. Crop yield departures were consistent with the drought severity depicted by the ESI and to a lesser extent by the NLDAS and VegDRI datasets.

54 Keywords – Flash drought; drought monitoring; soil moisture; evapotranspiration; crop  
55 impacts; agriculture; satellite data

56

## 57 1. Introduction

58         The 2012 drought that impacted major agricultural areas across the central U.S.  
59 was the worst drought to affect this region since 1988 and had similar magnitude and  
60 spatial extent to the severe droughts that occurred during the 1930s and 1950s (Hoerling  
61 et al. 2014). The almost complete absence of heavy rainfall events during the growing  
62 season, combined with record high temperatures, strong winds, and abundant sunshine,  
63 led to rapid decreases in soil moisture content and the rapid emergence of flash drought  
64 conditions (Lydolph 1964; Mozny et al. 2012; Otkin et al. 2013; Mo and Lettenmeier  
65 2015). According to the U.S. Drought Monitor (USDM; Svoboda et al. 2002), drought  
66 coverage and intensity rapidly increased during June and July in response to the  
67 anomalous weather conditions, with nearly 80% of the contiguous U.S. characterized by  
68 at least abnormally dry conditions by the end of summer. Most of the central U.S.,  
69 including the Corn Belt, experienced severe drought (or worse) conditions at some point  
70 during the growing season (Mallya et al. 2013). Recent modeling studies have shown  
71 that this exceptional drought event was not forced by tropical sea surface temperature  
72 anomalies. Instead, it was associated with natural variations in the weather that led to the  
73 development of a persistent upper-tropospheric ridge that inhibited convection and  
74 caused exceptionally warm temperatures to occur across the region for several months  
75 (Kumar et al. 2013; Wang et al. 2014; Hoerling et al. 2014; Diffenbaugh and Scherer  
76 2013).

77           The 2012 drought was one of the most expensive natural disasters in U.S. history  
78 with Federal crop indemnity payments alone exceeding \$17 billion (USDA 2013). Crop  
79 losses were especially large because the most severe drought conditions occurred during  
80 critical stages of crop development, such as pollination in corn and the grain filling stage  
81 in soybeans. Prior work has shown that even short periods (e.g. several days) of intense  
82 water stress can result in large crop yield reductions (e.g., Meyer et al. 1993; Saini and  
83 Westgate 1999; Calvino et al. 2003; Earl and Davis 2003; Barnabas et al. 2008; Mishra  
84 and Cherkauer 2010; Prasad et al. 2011; Kebede et al. 2012; Hunt et al. 2014). In 2012,  
85 however, severe moisture and heat stress lasted for more than a month across most major  
86 agricultural areas of the country, thereby leading to the lowest corn yields since 1995. If  
87 long-term yield trends are accounted for, the percentage yield loss was one of the largest  
88 on record going back to 1866 (Hoerling et al. 2014; Boyer et al. 2013). The large yield  
89 loss is consistent with a recent study by Lobell et al. (2014) that assessed yield trends  
90 during recent decades for different levels of moisture stress. Their analysis showed that  
91 yield gains have been smallest on a percentage basis for growing seasons in which large  
92 vapor pressure deficits indicative of severe drought conditions occur during critical crop  
93 yield development stages. As drought conditions spread westward during the summer,  
94 ranchers also experienced substantial impacts through a combination of higher feed  
95 prices, a lack of high quality forage, and heat-related animal stress, with many ranchers  
96 forced to either sell or relocate their livestock to other parts of the country (USDA 2012).  
97 The rapid onset of severe drought conditions meant that farmers and ranchers had little  
98 time to prepare for its adverse effects. It is possible, however, that greater use of drought  
99 indicators that respond quickly to changing conditions, such as the satellite-derived

100 Evaporative Stress Index (ESI; Anderson et al. 2007a,b), may promote drought mitigation  
101 efforts during future flash drought events by providing earlier warning of drought  
102 development (Otkin et al. 2014, 2015a, b).

103 High-resolution estimates of soil moisture and vegetation health conditions are  
104 necessary to accurately assess the severity and geographic extent of drought conditions at  
105 spatial and temporal scales sufficient for stakeholders to make informed management  
106 decisions. Moreover, an accurate assessment of current conditions is a prerequisite for  
107 producing useful drought intensification forecasts over monthly to seasonal time scales.  
108 In this paper, the evolution of several drought indicators sensitive to vegetation health  
109 and soil moisture conditions will be examined during the onset and development of the  
110 2012 flash drought. These indicators include the ESI, which uses satellite thermal  
111 infrared observations and a land surface energy balance model to estimate anomalies in  
112 evapotranspiration (ET) and the Vegetation Drought Response Index (VegDRI; Brown et  
113 al. 2008) that uses satellite, land, and climate observations to assess vegetation health  
114 conditions. The evolution of the satellite-derived datasets will be compared to modeled  
115 soil moisture anomalies from the North American Land Data Assimilation System  
116 (NLDAS; Xia, et al. 2012a,b; 2014) and to time series of precipitation and meteorological  
117 conditions. The accuracy of these datasets will be assessed for different locations and  
118 time periods through comparison with USDM drought analyses and county-level crop  
119 and range condition datasets compiled by the United States Department of Agriculture  
120 (USDA) National Agricultural Statistics Service (NASS). Though the NASS datasets are  
121 qualitative, they provide very valuable ground truth of the actual impact of the drought on  
122 agriculture. Each of these datasets is described in Section 2. The overall evolution of the

123 drought and relationships between the drought indicators and crop conditions and yield  
124 are assessed in Section 3, with conclusions presented in Section 4.

125

## 126 2. Data and Methodology

### 127 2.1. Evaporative Stress Index

128 The ESI depicts standardized anomalies in ET fraction ( $ET/ET_{ref}$ ), where ET is  
129 the actual ET flux retrieved under clear-sky conditions and  $ET_{ref}$  is a reference ET flux  
130 based on a Penman-Monteith formulation (Allen et al. 1998). Reference ET is used in  
131 this equation to minimize the impact of non-moisture related drivers of ET, such as the  
132 seasonal cycle in solar radiation, when assessing anomalies in ET. Similarly, the use of  
133 clear-sky ET minimizes impacts of cloud cover on ET variability, again focusing on soil  
134 moisture drivers. The Atmosphere-Land Exchange Inverse (ALEXI) model (Anderson et  
135 al. 1997, 2007a, 2011) is used to estimate the actual ET flux. ALEXI uses a two-source  
136 energy balance model (Norman et al. 1995) and land surface temperature (LST) retrievals  
137 obtained from satellite thermal infrared imagery to compute sensible, latent, and ground  
138 heat fluxes for vegetated and bare soil components of the land surface. The partitioning  
139 of the surface energy fluxes is accomplished using vegetation cover fraction estimates  
140 derived from the MODIS leaf area index product (Myneni et al. 2002). The total surface  
141 energy budget is computed using the observed increase in LST from ~1.5 hr after local  
142 sunrise until 1.5 hr before local noon, with closure of the energy balance equations  
143 achieved using the McNaughton and Spriggs (1986) atmospheric boundary layer growth  
144 model. Lower-tropospheric temperature profiles used by the boundary layer model are  
145 obtained from the Climate Forecast System Reanalysis dataset (Saha et al. 2010). The

146 ALEXI model is run each day over the contiguous U.S. (CONUS) with 4-km horizontal  
147 grid spacing using LST retrievals and insolation estimates derived from the Geostationary  
148 Operational Environmental Satellite (GOES) imager.

149 While the ESI ideally includes only clear-sky retrievals of ET, incomplete cloud  
150 screening of the thermal infrared-derived LST inputs can add noise to the ET time series  
151 used in the index computation. These errors are reduced using a temporal smoothing  
152 algorithm that identifies days with ET estimates that differ by more than one standard  
153 deviation from surrounding days within a 14 day moving window. Anderson et al.  
154 (2013) have shown that this method effectively removes cloud-contaminated ET  
155 estimates because abrupt changes in daily ET are more likely to occur because of cloud  
156 effects on surface heating than to rapid changes in soil moisture content. The remaining  
157 clear-sky ET estimates are then composited over longer time periods to achieve more  
158 complete domain coverage.

159 Standardized ET fraction anomalies, expressed as pseudo z-scores normalized to a  
160 mean of 0 and a standard deviation of 1, are computed each week using 2, 4, 8, and 12  
161 week composite periods. The mean ET fraction and standard deviations for each  
162 composite period are computed at each grid point in the CONUS domain using data from  
163 2001-2014. Standardized anomalies are computed as:

$$164 \quad ESI(w, y) = \frac{\langle v(w, y) \rangle - \frac{1}{ny} \sum \langle v(w, y) \rangle}{\sigma(w)} \quad (1)$$

165 where the first term in the numerator is the composite ET fraction for week  $w$  and year  $y$   
166 at a given grid point, the second term is the mean ET fraction for week  $w$  averaged over  
167 all years, and the denominator is the standard deviation. By standardizing the anomalies,

168 this means that negative (positive) values depict below (above) average ET fluxes, which  
169 are typically associated with lower (higher) than average soil moisture content and poorer  
170 (better) than average vegetation health in the absence of other stressors such as disease.

171

## 172 2.2. Vegetation Drought Response Index

173 VegDRI is an empirical method that combines satellite observations of vegetation  
174 health with climate data and other information about the land surface to identify regions  
175 containing drought stressed vegetation. Two climate-based drought indices, including the  
176 Standardized Precipitation Index (SPI; McKee et al. 1993) computed over a 36-week  
177 time period and the self-calibrated Palmer Drought Severity Index (Wells et al. 2004) are  
178 used by VegDRI. Normalized difference vegetation index data from the Advanced Very  
179 High Resolution Radiometer are used to calculate seasonal greenness and start of season  
180 metrics input into VegDRI. Several static biophysical variables describing environmental  
181 characteristics that influence drought stress on vegetation, such as land use/land cover,  
182 soil available water holding capacity, ecoregion type, and irrigation, are also included in  
183 the model. A classification and regression tree analysis is then applied to the historical  
184 information in the database to empirically derive VegDRI analyses each week. VegDRI  
185 output is typically displayed as discrete categories; however, because the underlying data  
186 are continuous, they were converted into standardized anomalies using data from 2000-  
187 2012 to ease comparison with other datasets used in this study. VegDRI data at 1-km  
188 native resolution were aggregated to the 4-km ESI grid. A complete description of the  
189 VegDRI model can be found in Brown et al. (2008) and Tadesse et al. (2015).

190



191 2.3. North American Land Data Assimilation System

192 Modeled soil moisture anomalies were computed using data from several NLDAS  
193 models (Xia et al. 2012a,b), including the Noah (Ek et al. 2003; Barlage et al. 2010; Wei  
194 et al. 2013), Mosaic (Koster and Suarez, 1996), and Variable Infiltration Capacity (VIC;  
195 Liang et al. 1996) models. Each land surface model simulates soil moisture content in  
196 multiple layers using energy and water balance equations. Because the models differ in  
197 their treatment of key processes such as evaporation, drainage, vegetation rooting depth,  
198 and canopy uptake, their soil moisture responses can differ due to local climate, soil, and  
199 vegetation characteristics. Daily soil moisture values from each model and the ensemble  
200 mean of all models (hereafter referred to as NMV\_AVE) were interpolated from the  
201 0.125° resolution NLDAS grid to the 4-km ESI grid using a nearest neighbor approach.  
202 Soil moisture data in the topsoil (0-10 cm) and total column (0-200 cm) layers were  
203 averaged over 2- and 4-week periods, with standardized anomalies for each soil layer  
204 (hereafter referred to as TS and TC, respectively) computed at weekly intervals using  
205 data from 1979-2014. The soil moisture response of each model will be compared to the  
206 ensemble mean and to the other drought indicators.

207

208 2.4. North American Regional Reanalysis

209 The evolution of the near-surface atmospheric conditions was evaluated using  
210 NARR data (Mesinger et al. 2006). Daily averages were computed for 10-m wind speed,  
211 2-m temperature, and 2-m dew point depression using analyses available every 3 hour on  
212 a 32-km resolution grid. The daily averages were then interpolated to the ESI grid using

213 a nearest neighbor approach, with standardized anomalies for 1-week periods computed  
214 at weekly intervals using data from 2000-2014.

215

## 216 2.5. Precipitation datasets

217 Gridded daily precipitation for 1948-2014 obtained from the Climate Prediction  
218 Center's (CPC) 0.25° resolution precipitation analysis (Higgins et al. 2000) was  
219 interpolated to the ESI grid using a nearest neighbor approach and then summed at  
220 weekly intervals to create 1-, 4-, 8-, and 12-wk accumulated precipitation amounts. SPI  
221 values for 4-, 8-, and 12-wk periods were subsequently computed. The SPI is a  
222 standardized variable widely used to identify meteorological drought conditions, with  
223 values less (greater) than zero indicating the observed precipitation was less (more) than  
224 the climatological median precipitation for a given length of time and time of year.

225

## 226 2.6. United States Drought Monitor

227 The USDM is a widely used drought analysis generated each week through expert  
228 synthesis of multiple data sources, including precipitation and soil moisture anomalies,  
229 surface stream flow departures, various drought metrics, crop and range conditions, and  
230 impact reports from local observers. Because it conveys drought information at multiple  
231 time scales and for a wide range of impacts (including socioeconomic), the USDM  
232 should not be considered an absolute measure of drought severity. By using a variety of  
233 data sources, most with high spatial resolution (sub-county), the USDM can depict both  
234 large-scale and localized areas of drought. For this study, weekly USDM analyses in  
235 shapefile format were interpolated to the 4-km ALEXI grid by assigning numerical

236 values to each drought category, with abnormally dry (D0) = 0, moderate drought (D1) =  
237 1, severe drought (D2) = 2, extreme drought (D3) = 3, and exceptional drought (D4) = 4.  
238 When no drought conditions were present the value was set to -1.

239

## 240 2.7. USDA crop and soil moisture datasets

241 The USDA NASS produces publically available state-level soil moisture and crop  
242 condition estimates each week from April-November based on survey data collected from  
243 ~4000 local experts knowledgeable in visually identifying crop status and soil moisture  
244 conditions. For this study, the author signed a confidentiality agreement with the USDA  
245 NASS to access county-level crop condition and soil moisture datasets, where the data  
246 were spatially smoothed to ensure that no individual records or confidential data were  
247 publically released. Health condition estimates ranging from very poor to excellent are  
248 reported for pasture and range and for all major agricultural crops, including corn,  
249 soybeans, cotton, winter wheat, spring wheat, peanuts, barley, oats, and sorghum. In  
250 addition, categorical topsoil and subsoil moisture assessments ranging from very short to  
251 surplus are made each week, with the former (latter) category indicating that the soil  
252 moisture content is much less (greater) than that required for normal crop development.  
253 Numerical values were then assigned to each crop condition (very poor, poor, fair, good,  
254 and excellent) and soil moisture (very short, short, adequate, and surplus) category, with  
255 average crop conditions computed for each county using all reports available during a  
256 given week. These county level datasets were spatially smoothed using a 3x3 grid point  
257 square moving window after first being interpolated to the 4-km ESI grid. Crop and soil

258 moisture anomalies were computed each week by subtracting the mean conditions from  
259 the 2002-2014 period of record.

260 The impact of the drought conditions on the end-of-season crop yield was also  
261 assessed using county level yield statistics compiled by NASS. Corn, soybean, winter  
262 wheat, and spring wheat yields from 2000-2014 were obtained from the NASS Quick  
263 Stats database (<http://quickstats.nass.usda.gov>). A least squares regression line was fit to  
264 the annual yield time series for each county and crop to account for local changes in yield  
265 over time, and then trend-adjusted yield departures were computed for each year.

266

### 267 3. Results

268

#### 269 3.1. Large-scale drought analysis

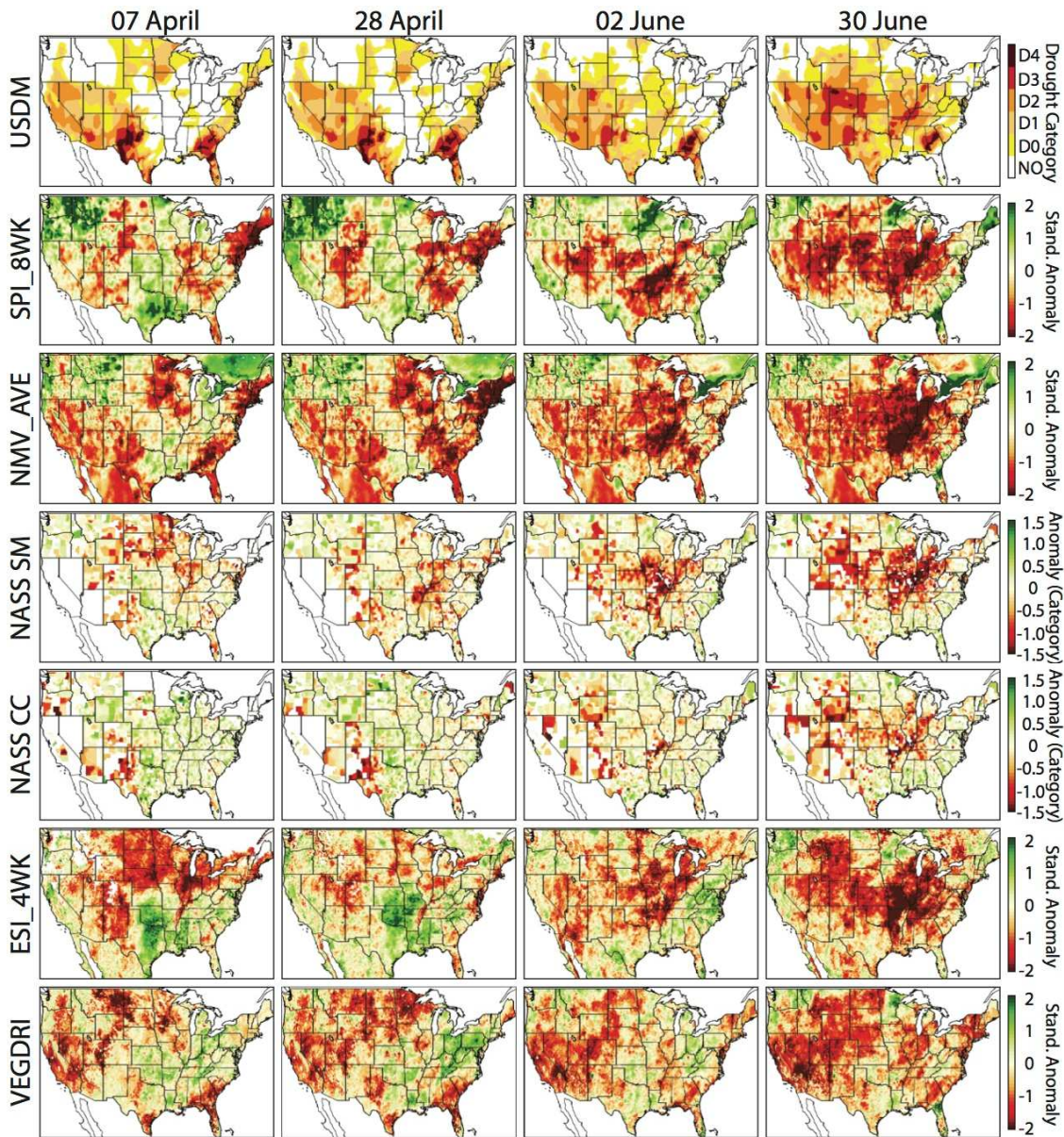
270 This section examines the overall evolution of conditions across the U.S. during  
271 the 2012 drought event from drought onset during late spring through drought maturation  
272 during the summer and the northwestward progression of the core drought area during the  
273 fall. Figures 1 and 2 show the evolution of the USDM, SPI\_8WK, NMV\_AVE topsoil  
274 moisture, NASS topsoil moisture and crop condition, ESI\_4WK, and VegDRI datasets at  
275 monthly intervals from 07 April to 28 October. The time period lengths for each variable  
276 were chosen to minimize differences in their response time to the anomalous conditions.  
277 For example, compared to the 8-wk SPI, a shorter 4-wk time period was used to compute  
278 the ESI and NMV\_AVE anomalies because vegetation and soil moisture tend to respond  
279 to rainfall anomalies occurring over longer time periods. The VegDRI data, however,

280 will represent drought on a slightly longer time scale due to its use of long-term climate  
281 variables such as the 36-week SPI.

282         On 07 April, drought conditions were present across the southwestern and north-  
283 central U.S. and along the East Coast, with the worst conditions located in Georgia and  
284 west Texas according to the USDM. Overall, these drought areas were well captured by  
285 the NMV\_AVE and ESI\_4WK anomalies; however, there were some differences in their  
286 spatial extent and magnitude. For example, negative ESI\_4WK anomalies cover a much  
287 larger area of the northern U.S. These negative ET anomalies developed in response to a  
288 prolonged period of record heat during March (Blunden and Arndt 2013) and indicate  
289 that the newly emerged vegetation became moisture stressed because their shallow roots  
290 were unable to access sufficient subsoil moisture once the top few cm of the soil profile  
291 became dry. Thus, the ESI\_4WK anomalies across this part of the country are indicative  
292 of short-term dryness at this time. Their large spatial extent, however, is consistent with  
293 the widespread negative topsoil moisture anomalies reported in the NASS dataset. The  
294 VegDRI analysis also depicts drought in many parts of the country, including the  
295 southwestern and southeastern U.S.; however, it does not depict drought over New  
296 Mexico and Texas or over New England because the vegetation signal is considered too  
297 weak at this time of the year.

298         By 28 April, dry conditions were becoming more widespread across the eastern  
299 U.S. according to the SPI\_8WK, NMV\_AVE, and NASS soil moisture datasets;  
300 however, only minor changes were made to the USDM analysis. The ESI\_4WK dataset  
301 contains a large area of positive anomalies across the south-central U.S. within a region  
302 of above average rainfall. These anomalies indicate that the vegetation was growing

303 rapidly in response to the favorable conditions, which is supported by the positive NASS  
 304 crop condition anomalies. The VegDRI dataset also contains positive anomalies across



305 Fig. 1 Temporal evolution of the United States Drought Monitor (USD  
 306 Drought Monitor (USD) drought depiction, and 8-week Standardized Precipitation Index (SPI), 4-week modeled total  
 307 column soil moisture (NMV\_AVE), National Agricultural Statistics Service (NASS) topsoil  
 moisture and crop condition, 4-week Evaporative Stress Index (ESI), and Vegetation  
 Drought Response Index (VegDRI) anomalies from 07 April until 30 June 2012.

306 this part of the country; however, they are smaller than the ESI\_4WK anomalies. Across  
 307 the north central U.S., the ESI\_4WK anomalies had become less extreme, possibly

308 because the vegetation had developed a deeper root structure that could access more soil  
309 moisture and support higher ET rates. Farther to the east, large precipitation deficits from  
310 Pennsylvania to southern New England led to the development of large topsoil moisture  
311 anomalies in the NMV\_AVE dataset. The ESI\_4WK anomalies were near normal except  
312 for areas along the Atlantic Coast, whereas the VegDRI anomalies were mostly positive  
313 across the region.

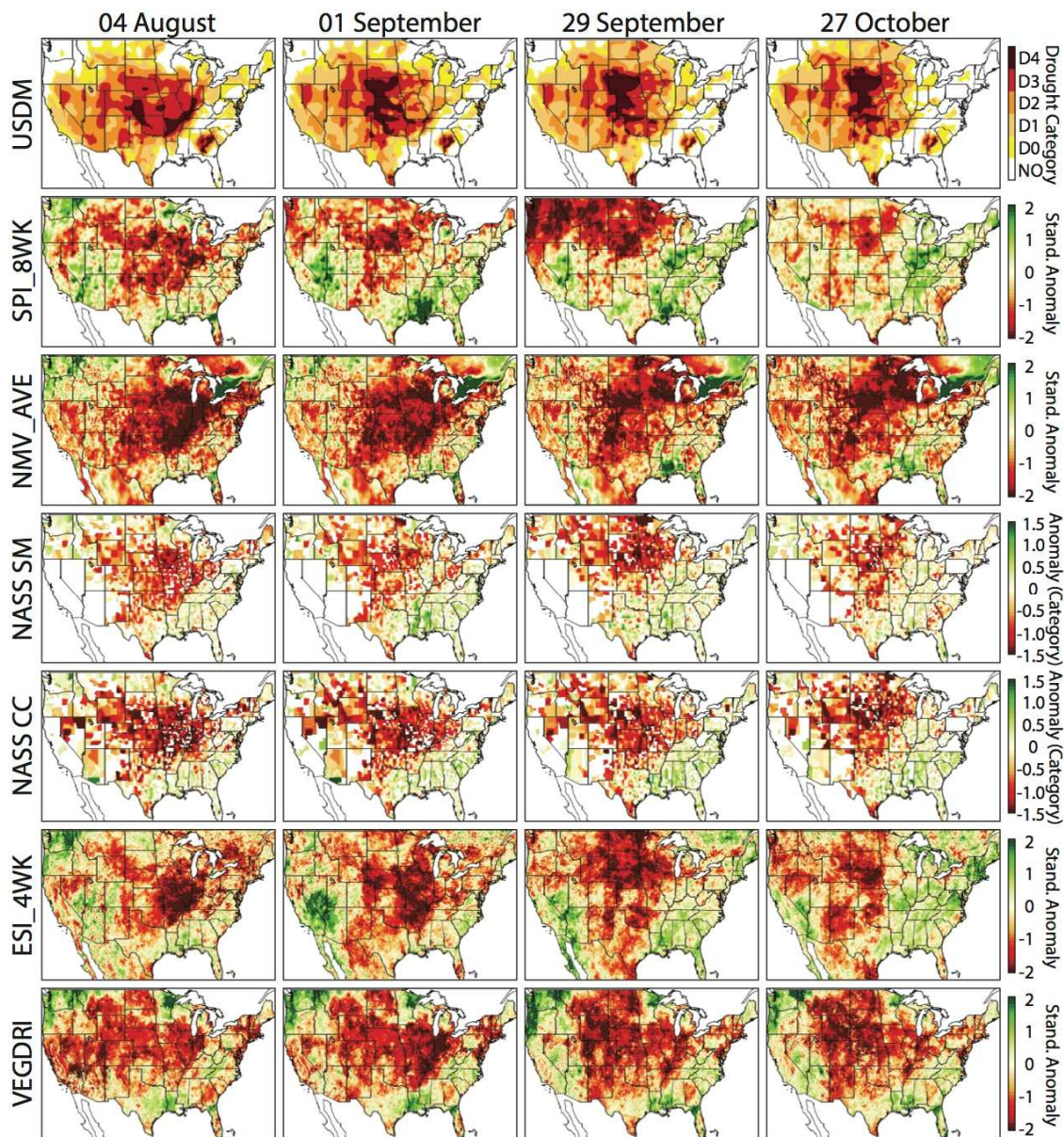
314 By 02 June, large negative SPI\_8WK anomalies had developed across most of the  
315 southern U.S., with especially large rainfall deficits located in the south central U.S. The  
316 rapid transition from positive to negative anomalies is also evident in the ESI\_4WK and  
317 NMV\_AVE datasets, which now contain large negative anomalies across most of the  
318 central U.S. The NASS datasets indicate that the topsoil moisture content and to a lesser  
319 extent the crop conditions were below average across most of the central U.S. The worst  
320 soil moisture conditions were located in the mid-Mississippi River valley where some  
321 areas experienced up to a 2-category increase in drought severity during the previous five  
322 weeks according to the USDM. Unlike the other datasets, the VegDRI anomalies mostly  
323 remained positive or only became slightly negative across the central U.S. The delayed  
324 response of this metric to the rapidly worsening conditions likely results from its use of  
325 long-term climate variables such as the 36-wk SPI that change more slowly than fast  
326 response drought indicators such as the ESI.

327 Conditions continued to rapidly deteriorate across most of the central U.S. during  
328 June in response to the onset of very hot temperatures and the continuation of well below  
329 normal rainfall. By 30 June, large negative NASS topsoil moisture anomalies extended  
330 from the central Rockies eastward across the entire Corn Belt. Crop and range conditions

331 were also beginning to rapidly deteriorate as the vegetation was increasingly unable to  
332 cope with the adverse weather and soil moisture conditions. Overall, the ESI\_4WK and  
333 NMV\_AVE datasets accurately represent the spatial extent of the drought; however, both  
334 depict more severe drought than the USDM in several locations. This is consistent with  
335 prior work (e.g. Otkin et al. 2013) that has shown that the USDM tends to respond too  
336 slowly to rapidly changing conditions. Both datasets indicate that extreme drought had  
337 developed within regions characterized by especially large rainfall deficits along the mid-  
338 Mississippi River valley. The VegDRI anomalies have also decreased within this region,  
339 but remain too small compared to the other datasets. VegDRI performance is better in  
340 the western U.S. where it depicts widespread severe drought conditions.

341         After enduring the hottest July on record and receiving below normal rainfall  
342 (Diffenbaugh and Sherer 2013), extreme to exceptional drought conditions (D3-D4 in the  
343 USDM) encompassed most of the central U.S. by the beginning of August (Fig. 2).  
344 According to the USDM, more than 80% of the U.S. was characterized by at least  
345 abnormally dry conditions at the peak of the drought on 24 July (not shown). Many  
346 locations had experienced flash drought during the previous two months as conditions  
347 rapidly transitioned from being drought free to the two worst (D3 and D4) drought  
348 categories in the USDM. Very large negative NMV\_AVE and ESI\_4WK anomalies  
349 were present within the core drought regions characterized by the largest SPI\_8WK  
350 anomalies. The spatial extent and magnitude of these anomalies are consistent with the  
351 very poor crop conditions present across most of the central U.S. Though the VegDRI  
352 anomalies had also decreased across this part of the country, their magnitude was still  
353 much smaller than the other drought indicators.





354 Fig. 2 Same as Fig. 1 except from 04 August until 27 October 2012.

355 Several episodes of beneficial rainfall during August led to some improvements to  
 356 the drought depiction by 01 September along the eastern periphery of the core drought  
 357 region from Arkansas to Michigan. The wetter conditions in the east combined with the  
 358 continuation of hot, dry weather in the west led to a westward shift of the core drought  
 359 region to the central High Plains. Although the NASS topsoil moisture conditions had  
 360 improved slightly within the eastern Corn Belt, the crops were so badly damaged by this

361 time that only minor gains are evident in the crop condition and ESI\_4WK datasets. The  
362 VegDRI anomalies accurately captured the spatial extent of the severe drought conditions  
363 from the Rocky Mountains eastward to the Mississippi River valley. Further to the west,  
364 unusually heavy rainfall across the Desert Southwest led to very large positive ESI\_4WK  
365 anomalies indicative of much higher than normal ET rates. Some improvements were  
366 also evident in the VegDRI data and to a lesser extent in the NMV\_AVE topsoil moisture  
367 anomalies. The USDM drought depiction improved by one category in most places, but  
368 remained high to reflect the impact of long-term dryness across the region.

369 By 29 September, very dry conditions had developed from the Pacific Northwest  
370 to the Upper Midwest, with SPI\_8WK anomalies  $< -2$  in many locations. These large  
371 rainfall departures combined with warmer than average temperatures led to a northward  
372 expansion of drought conditions into the north central U.S. and further intensification of  
373 the extreme drought over the central High Plains. The worsening drought conditions are  
374 evidenced by the increased spatial extent of large negative anomalies in the ESI\_4WK,  
375 NMV\_AVE, and VegDRI datasets across the north central U.S. and a concurrent increase  
376 in large negative NASS topsoil moisture and crop condition anomalies across this region.

377 Finally, by the end of October, the core drought area had become entrenched over  
378 the central High Plains from the Texas panhandle northward to western South Dakota.  
379 Very dry conditions are evident in each dataset across this part of the country. Continued  
380 wet weather across the eastern U.S., however, led to further improvements to the USDM  
381 drought depiction along the Mississippi River valley. The ESI\_4WK anomalies capture  
382 the improving conditions in the eastern U.S. as indicated by the return to normal or above

383 normal crop condition and soil moisture anomalies in the NASS datasets. Though not as  
384 large, improvements were also evident in the VegDRI and NMV\_AVE datasets.

385

### 386 3.2. Regional drought analysis

387 In this section, the drought evolution will be examined more closely for locations  
388 that experienced severe drought conditions during different parts of the growing season  
389 and are characterized by different climate regimes and agricultural interests. Unlike the  
390 previous section, anomaly time series will be shown at weekly intervals for each variable  
391 and will be assessed separately for each crop type and NLDAS model, and for anomalies  
392 computed over different time periods. The data will be displayed using a visualization  
393 method developed in prior studies (e.g. Otkin et al. 2013) as shown in Fig. 3. The USDM  
394 is displayed in the first column, with weekly rainfall totals and 1-wk anomalies in surface  
395 temperature, dew point depression, and wind speed shown in the next three columns. SPI  
396 values for 4- and 12-wk periods are shown next, followed by anomalies in the NASS  
397 topsoil moisture, subsoil moisture, range, corn, soybeans, and winter wheat conditions.  
398 After that, anomalies are shown for VegDRI and for the ESI computed over 2-, 4-, 8-,  
399 and 12-wk periods. The last sixteen columns show anomalies in topsoil and total column  
400 soil moisture content computed over 2- and 4-wk periods for the Noah, Mosaic, and VIC  
401 models and also for their ensemble mean.

402

#### 403 3.2.1. West-central Missouri

404 Figure 3 shows weekly values for each variable averaged using all grid points in  
405 west-central Missouri (CPC climate division 3). At the beginning of March, abnormally

406 dry conditions were present across the region as signified by negative anomalies in most  
407 datasets. This dryness was partially alleviated by several heavy rainfall events from the  
408 end of March to the first week of May that led to large positive SPI and ESI anomalies at  
409 all time scales. The positive ESI anomalies indicate that the ET had greatly increased in  
410 response to the heavy rainfall and warm temperatures, which is consistent with the above  
411 average range and winter wheat conditions in the NASS dataset. Though improvements  
412 were also evident in the VegDRI and NLDAS datasets, these changes were modest and  
413 most of the anomalies remained negative even though short-term conditions had  
414 improved.

415         After receiving beneficial rainfall during the spring, very dry weather returned to  
416 the region during May and coincided with a prolonged hot and windy spell that caused  
417 soil moisture and crop conditions to rapidly deteriorate. Temporal changes in the short-  
418 range ESI anomalies (2-8 weeks) were exceptionally large at the beginning of the flash  
419 drought event and closely mirrored the observed crop condition changes. The VegDRI  
420 and 12-wk ESI anomalies also decreased, but at a slower rate than the shorter-range ESI  
421 anomalies. Each of the NLDAS models also exhibited rapid decreases in soil moisture at  
422 the end of May that were consistent with changes in the NASS soil moisture dataset and  
423 preceded their appearance in the ESI by one week. These changes first appeared in the  
424 TS moisture and shorter 2-wk composites before appearing in the TC and 4-wk soil  
425 moisture anomalies. As drought conditions intensified during the summer, most of the  
426 drought indicators continued to deteriorate except for the VIC soil moisture anomalies,  
427 which were more sensitive to small rainfall events. The USDM drought severity lagged  
428 the other drought metrics by several weeks during the entire event.

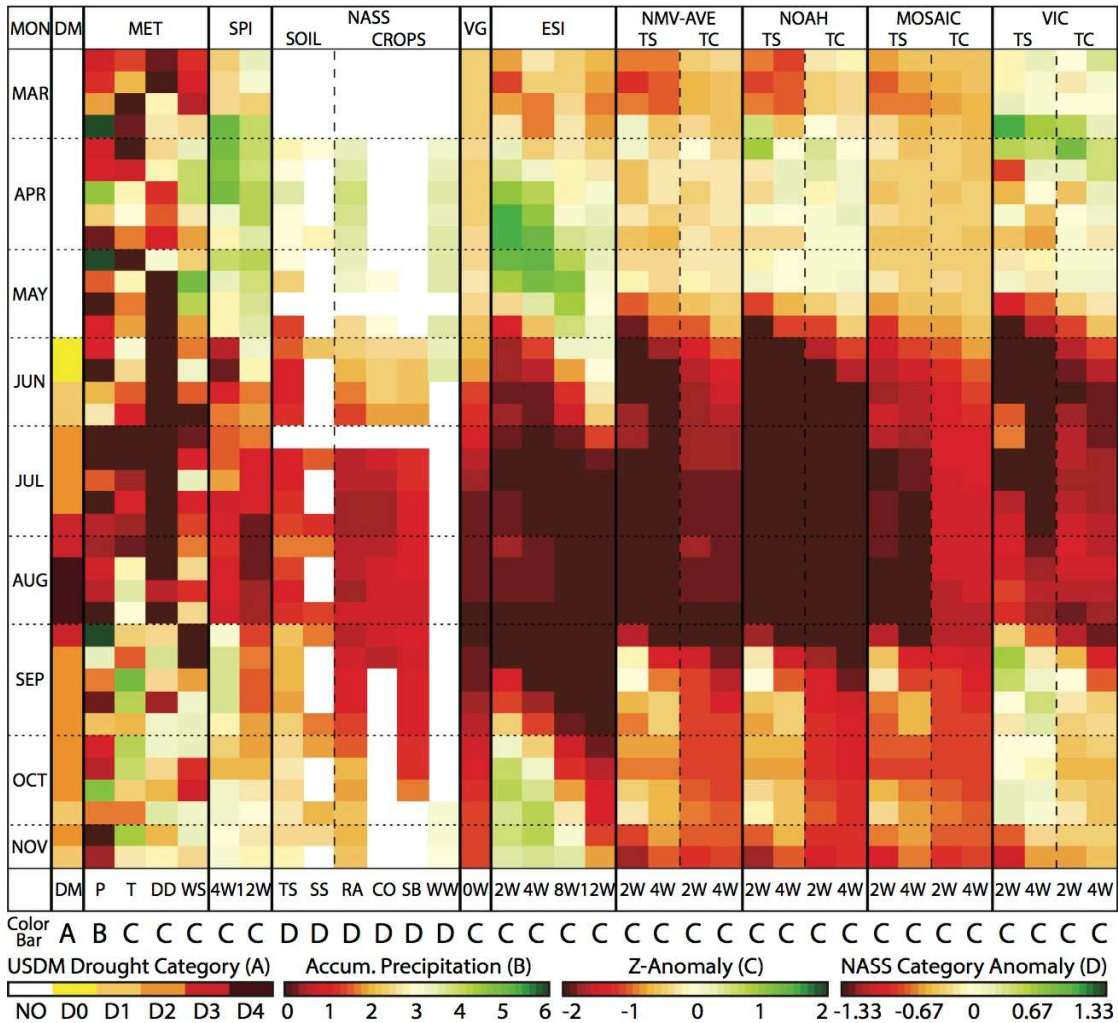


Fig 3. Drought evolution across west-central Missouri during 2012. The weekly USDM drought category is shown in column 1. The weekly rainfall (cm) is shown in column 2, with z-anomalies for 1-week average surface temperature, dew point depression, and wind speed shown in columns 3-5. Note that the z-anomaly color bar is reversed for each of these variables so that positive anomalies indicative of enhanced drying are shown in red and brown colors. The SPI values for 4- and 12-week time periods are shown in columns 6-7. NASS topsoil and subsoil moisture anomalies are shown in columns 8-9, with crop condition anomalies for range, corn, soybeans, and winter wheat shown in columns 10-13. Standardized VegDRI anomalies are shown in column 14. ESI z-anomalies for 2-, 4-, 8-, and 12-week composite periods are shown in columns 15-18, with 2- and 4-week z-anomalies for topsoil (0-10 cm) and total column (0-2 m) soil moisture for the NMV ensemble average and for the Noah, Mosaic, and VIC models shown in columns 19-22, 23-26, 27-30, and 31-34, respectively.

429

430            Heavy rainfall at the beginning of September led to rapid improvements in the TS  
 431 and TC moisture in all of the NLDAS models and the reappearance of positive 4-wk SPI  
 432 values. The ESI anomalies, however, remained at or near their lowest values for the year,

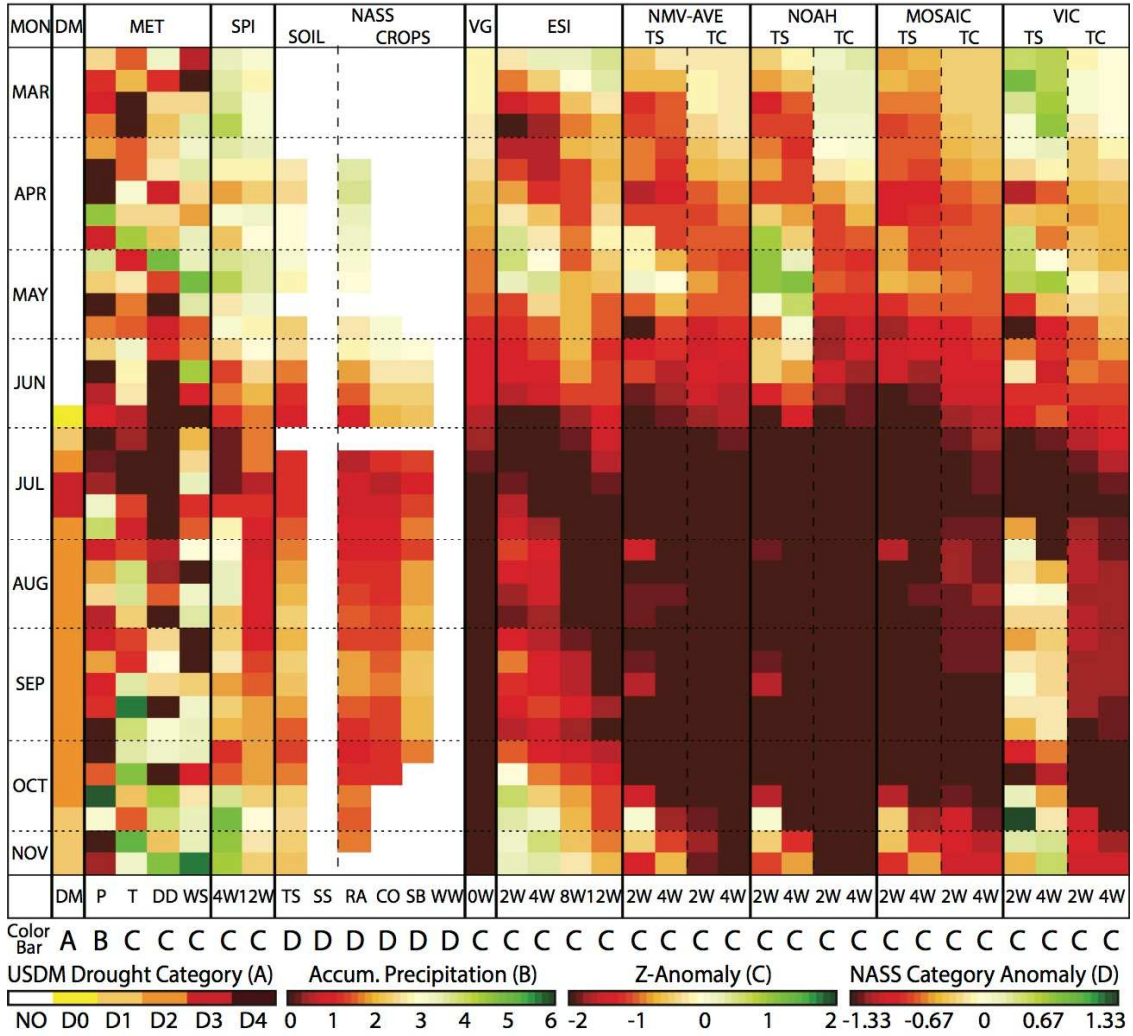
433 and did not exhibit substantial improvements until several weeks later. This behavior  
434 indicates that the vegetation was initially dormant or so badly damaged that it could not  
435 immediately respond to the improving conditions. It took a prolonged period of cool, wet  
436 conditions before the vegetation could recover enough to transpire at higher than normal  
437 rates during October. The initial lack of improvement in the ESI is consistent with trends  
438 in the NASS crop condition datasets. The 2-category improvement in the USDM at the  
439 beginning of September was more representative of the above normal rainfall than it was  
440 of improving vegetation conditions. The VegDRI anomalies reached their lowest values  
441 during the peak of the drought at the end of August and then slowly recovered during the  
442 fall. Consistent with its use of long-term climate variables, its evolution more closely  
443 matched the NLDAS ensemble model average TC soil moisture anomalies during the  
444 drought event; however, differences were larger with respect to the individual models.

445

### 446 3.2.2. South-central Wisconsin

447 This section describes the evolution of the drought over south-central Wisconsin  
448 (CPC climate division 8). Inspection of Fig. 4 shows that record warmth during March  
449 led to negative anomalies in most datasets despite the slightly above normal rainfall. The  
450 spread in the NLDAS soil moisture anomalies was very large at this time. TS moisture  
451 anomalies were positive in the VIC model, but negative in the Noah and Mosaic models,  
452 whereas the TC moisture anomalies were positive in the Noah model but negative in the  
453 other models. Large model differences are also evident later in the spring during a period  
454 of moderate rainfall that greatly improved TS moisture conditions, especially in the Noah  
455 and VIC models, but led to only minor improvements in the TC soil moisture. The short-

456 range ESI anomalies became slightly positive during May in response to the improved TS  
 457 moisture conditions. The negative VegDRI anomalies present at the beginning of March  
 458 continued to increase during the spring due to long-term dryness across the region.



459 Fig 4. Same as Fig. 3, except for south central Wisconsin.

460 Extreme weather conditions characterized by well below normal rainfall, record  
 461 high temperatures, large dew point depressions, and unusually strong winds developed  
 462 across the region during June and the first half of July. Vegetation conditions as indicted  
 463 by the ESI and NASS datasets rapidly deteriorated during this time period because of the  
 464 increased evaporative demand and the already short soil moisture conditions. Pasture and

465 range conditions were the first to deteriorate, followed in subsequent weeks by decreases  
466 in corn and soybean conditions. Almost all of the satellite and model-based drought  
467 indicators depicted extreme drought conditions by the beginning of July. Though the  
468 USDM drought severity increased an impressive four categories in four weeks during this  
469 flash drought, its period of rapid intensification was delayed by up to 4 weeks compared  
470 to the other datasets, especially those computed over 2- and 4-wk time periods. The  
471 earlier onset of the large negative anomalies in the ESI and modeled soil moisture  
472 datasets, however, is consistent with the large negative anomalies in the NASS datasets.

473         Heavy rainfall during the last two weeks of July allowed conditions to improve  
474 slightly, with the 4-wk SPI returning to normal. Large differences are again evident in  
475 the NLDAS datasets, with the VIC model showing much larger improvements, especially  
476 in TS moisture, that lasted throughout the late summer and fall recovery period. Given  
477 that each of these models had similar anomalies preceding the first rainfall, their different  
478 responses are likely due to differences in their infiltration and runoff rates. Compared to  
479 the NASS TS moisture dataset, the VIC model is likely too wet, whereas the Noah and  
480 Mosaic models are too dry. The VegDRI anomalies became very large in July and then  
481 remained strongly negative during the rest of the growing season even as conditions were  
482 slowly improving in the other datasets. The delayed VegDRI response was likely due to  
483 its use of the 36-wk SPI because this variable remained strongly negative during this time  
484 period. The ESI anomalies displayed different behavior depending on the composite  
485 period length, with the short-term composites showing minor improvements after the first  
486 rainfall, whereas the long-range composites did not improve until September. Overall,  
487 changes in corn conditions were closely related to the long-range (8- and 12-wk) ESI



488 anomalies, whereas the soybean conditions tracked changes in the shorter 2- and 4-wk  
489 anomalies. This behavior is consistent with prior work by Peng et al. (2014) using other  
490 vegetation indices.

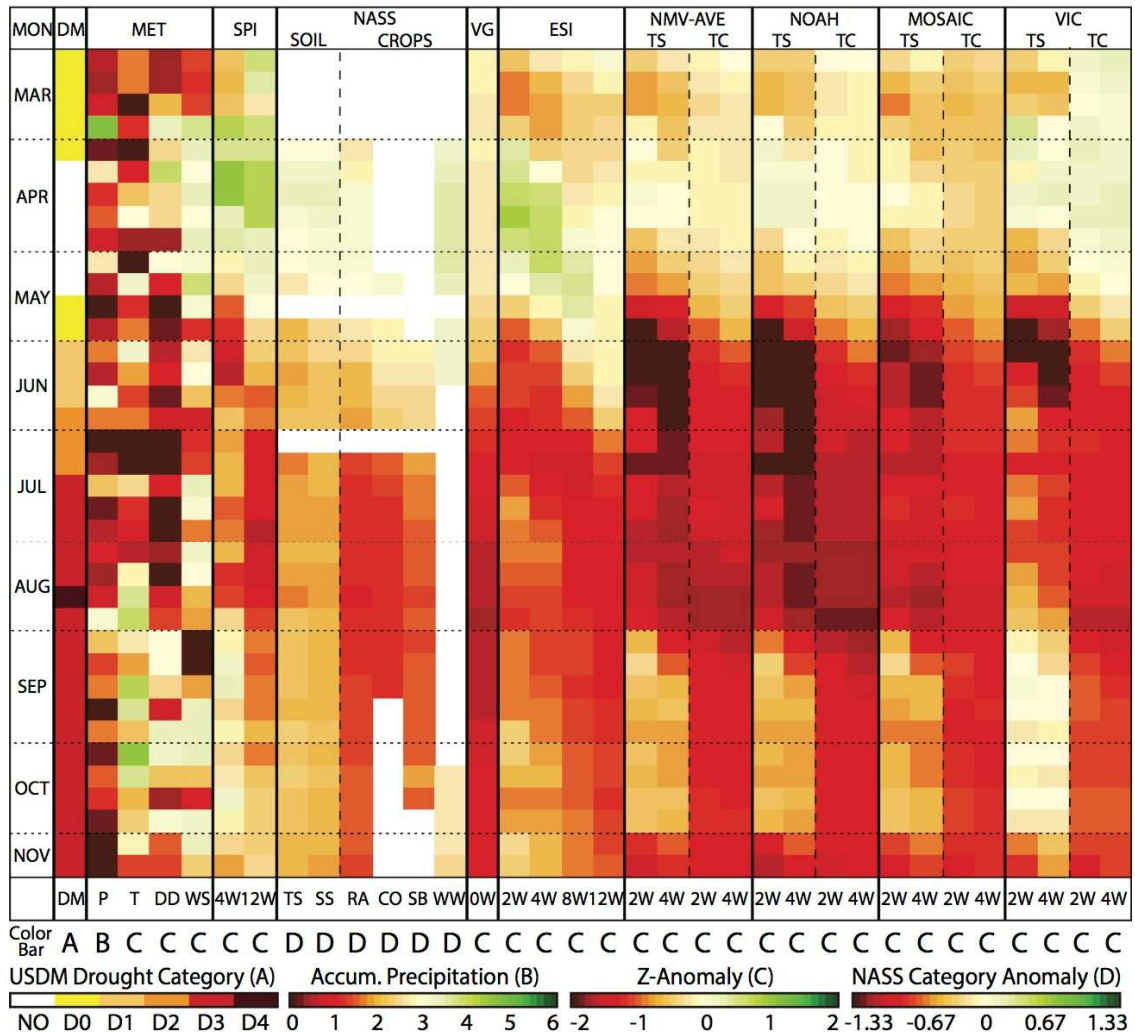
491

### 492 3.2.3. Northwestern Kansas

493 The evolution of the drought conditions over northwestern Kansas (CPC climate  
494 division 1) is shown in Fig. 5. At the beginning of March, most datasets indicated near to  
495 slightly drier than normal conditions. Beneficial rainfall starting at the end of March led  
496 to positive SPI anomalies and above normal soil moisture and crop conditions according  
497 to the NASS datasets. Though there are some differences in magnitude, the evolution of  
498 the modeled TS and TC moisture anomalies are similar in each NLDAS model. The ESI  
499 anomalies remained negative longer than the other datasets during the first part of April  
500 presumably because the vegetation had not yet emerged or was still too small to take full  
501 advantage of the increased soil moisture content. The VegDRI anomalies were near zero  
502 initially before slowly increasing as spring transitioned into summer.

503 Drought conditions began to rapidly intensify during May and June in response to  
504 a prolonged period of hot, windy, and dry weather that quickly depleted the TS moisture  
505 according to the NASS dataset and each of the NLDAS models. Rapid decreases initially  
506 occurred in the NLDAS TS anomalies before becoming evident in the ESI anomalies two  
507 weeks later and the VegDRI dataset after that. Periodic small rainfall events starting in  
508 July led to minor improvements in the 2- and 4-wk ESI composites and the NLDAS TS  
509 moisture anomalies; however, the longer-range ESI composites and NLDAS TC moisture  
510 anomalies continued to decrease during the summer as long-term rainfall deficits

511 continued to accumulate. The negative soil moisture anomalies were largest for the Noah  
 512 model and smallest for the VIC model, similar to the results described in previous  
 513 sections. The VegDRI anomalies continued to become more negative during the fall and  
 514 corresponded well with changes in the NLDAS modeled TC moisture content.



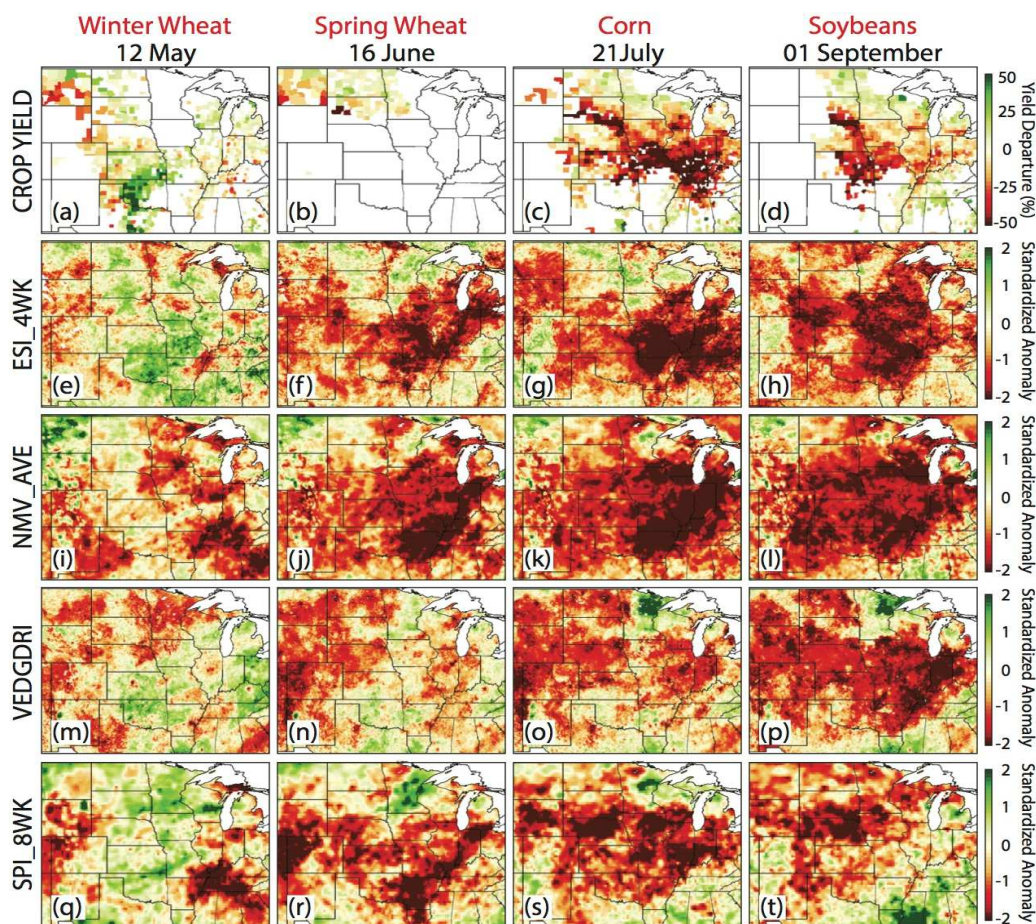
515 Fig. 5. Same as Fig. 3, except for northwestern Kansas.

516

517 3.3. Crop yield analysis

518 In this section, we assess the impact of the severe flash drought conditions on the  
 519 end-of-season yield for major agricultural crops grown across the central U.S. Figure 6

520 shows the trend-adjusted yield departures for corn, soybeans, winter wheat, and spring  
 521 wheat during 2012, along with ESI\_4WK, NMV\_AVE TC soil moisture, VegDRI, and  
 522 SPI\_8WK anomalies during critical times for yield production in each crop. The yield  
 523 departures are expressed as percentages above and below the 2000-2014 yield trend for  
 524 each county to account for local differences in average crop yield and yield trends.



525 Fig. 6. Trend-adjusted yield departures (%) for 2012 for (a) winter wheat, (b) spring wheat, (c) corn,  
 and (d) soybeans for each county computed with respect to the 2000-2014 base line period. ESI  
 4-week standardized anomalies for (e) 12 May, (f) 16 June, (g) 21 July, and (h) 01 September. (i-l)  
 Same as (e-h) except for 4-week NMV\_AVE total column soil moisture standardized anomalies.  
 (m-p) Same as (e-h) except for VegDRI standardized anomalies. (q-t) Same as (e-h) except for  
 8-week SPI standardized anomalies.

526 Overall, it is evident that winter wheat yields were well above average across the  
 527 primary wheat-growing areas in the south-central U.S., most notably in parts of western  
 528 Oklahoma and eastern Kansas where yields were 50% higher than normal. Wheat yields

529 were high in this part of the country because the warm and wet conditions during spring  
530 provided ideal growing conditions that allowed the crop to mature before severe drought  
531 conditions developed by mid-summer. Areas further to the west and north remained in  
532 drought during the spring and early summer; thus, their yields tended to be below normal.  
533 Spring wheat yields over the northern Plains were below normal over South Dakota and  
534 Montana in areas affected by drought; however, they were slightly above normal over  
535 North Dakota where conditions were more favorable.

536         One of the most critical periods for wheat yield production occurs between the  
537 booting and soft dough stages during late spring for winter wheat and early summer for  
538 spring wheat (Hanks and Rasmussen 1982). Overall, for winter wheat, there is a strong  
539 relationship between above average yield over Oklahoma and southeastern Kansas and  
540 positive ESI anomalies on 12 May, with negative ESI anomalies over the High Plains and  
541 the eastern Corn Belt where yields were below average. For spring wheat, the ESI also  
542 contains large negative anomalies in regions with below average yield, such as over most  
543 of Montana and western South Dakota. A strong correspondence also exists between the  
544 VegDRI anomalies and wheat yield departures across most of the central U.S. The  
545 NMV\_AVE anomalies, however, exhibit a weaker relationship to the final yield for both  
546 crops. For example, the NMV\_AVE anomalies are mostly negative across the southern  
547 Plains on 12 May where winter wheat yields were well above average but were mostly  
548 positive across Montana on 16 June where spring wheat yields were below normal.

549         The extreme drought conditions had a much larger impact on corn and soybean  
550 yields across the Midwest. Corn yields were below normal across most of the Corn Belt,  
551 with less than half of normal yield observed in the region extending from South Dakota

552 southeastward across Missouri and the lower Ohio River valley. Soybean yields were  
553 also below normal in most locations, especially across the western Corn Belt and central  
554 Plains where yields were at least 25% below the long-term trend. The different locations  
555 of the largest corn and soybean yield losses are consistent with the evolution of the most  
556 severe drought conditions during the growing season. For example, the largest corn yield  
557 reductions occurred where excessive heat in July combined with the largest precipitation  
558 deficits. July is the most important month for determining corn yield because excessive  
559 heat during that month can significantly decrease pollination efficiency during the critical  
560 silking and tasseling stages (Lobell et al. 2013; Shafiie-Jood et al. 2014). For soybeans,  
561 however, the most important development stages occur during the second half of summer  
562 when soybean pods develop and the seeds still have time to increase in size if the plants  
563 receive adequate rainfall. This meant that soybean yield losses were less severe east of  
564 the Mississippi River because of heavy rainfall during August and September, but were  
565 larger to the west as the core drought region shifted westward during the summer.

566         Comparison of the drought indices on 21 July reveals that the spatial pattern in the  
567 ESI anomalies most accurately corresponds to the observed corn yield departures across  
568 most of the Corn Belt, including the much below average yield from Missouri to southern  
569 Indiana and the above average yield over Minnesota and North Dakota. The NMV\_AVE  
570 anomalies were also strongly negative across the central and eastern Corn Belt; however,  
571 the large anomalies extended too far to the north into areas that actually had near to above  
572 average corn yields. Though VegDRI also exhibits negative anomalies in most locations,  
573 its correspondence to the final corn yield is much weaker than the other datasets because  
574 of its slow response to the rapidly changing conditions experienced during this drought.

575 Its performance improved for soybeans, with negative anomalies and a spatial pattern that  
576 more closely matches those depicted by the SPI\_8WK, ESI and NMV\_AVE datasets  
577 during the bean filling stage (e.g., 01 September).

578 To further assess relationships between the various drought indices and the 2012  
579 crop yields, correlations were computed between the county-level trend-adjusted crop  
580 yield departures and the ESI\_4WK, SPI\_8WK, VEGDRI, and NMV\_AVE TC anomalies  
581 at weekly intervals during the growing season (Fig. 7). The drought monitoring datasets  
582 for a given week were averaged to the individual county level prior to computing the  
583 correlations. Table 1 provides a list of the states used to compute the correlations for  
584 each crop. The correlations typically increase for each crop as the growing season  
585 progresses and reach peak values near critical stages of yield development (shaded areas  
586 in Fig. 7). For most crops, the ESI\_4WK data exhibited the strongest correlations to  
587 yield anomalies during these critical stages, most notably for corn and wheat. Given the  
588 importance of rainfall for yield production, the SPI\_8WK correlations were also strong,  
589 but were weaker than those computed using the ESI\_4WK data except for soybeans. The  
590 stronger correlations exhibited by the ESI\_4WK variable demonstrates that although  
591 rainfall departures are important for yield production, it is also necessary to consider  
592 other drivers of drought such as hot temperatures when assessing agricultural drought  
593 severity and potential impact on yield. Correlations with VEGDRI were generally  
594 weaker than the other variables during the spring and early summer due to its slow  
595 response to the rapidly changing conditions, but increased as drought conditions became  
596 entrenched across the region, with its maximum correlations obtained near the end of the  
597 growing season. Finally, although the NMV\_AVE correlations were relatively strong for

598 corn, they were weaker for the other crops and were even negative for spring wheat.  
 599 Further research is necessary to determine why the modeled soil moisture anomalies had  
 600 such a weak relationship to crop yields during this extreme flash drought event.

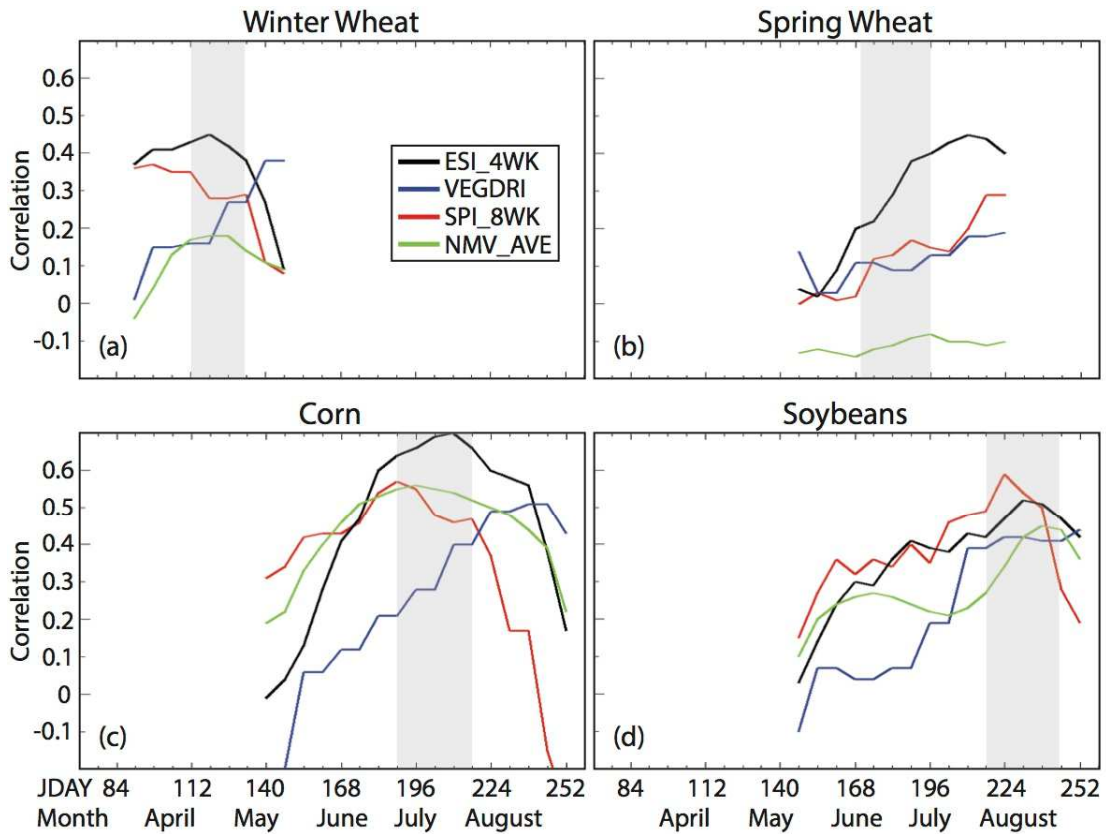


Fig. 7. Time series of correlations between county-level trend-adjusted crop yield departures (%) for (a) winter wheat, (b) spring wheat, (c) corn, and (d) soybeans and ESI\_4WK (black), VEGDRI (blue), SPI\_8WK (red), and NMV\_AVE (green) anomalies at weekly intervals during 2012. The gray-shaded regions in each panel indicate critical development periods for each crop. The time series are only plotted during the growing season for a given crop.

601  
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 606

607 Table 1. States used to compute the yield correlations for winter wheat, spring wheat,  
 608 corn, and soybeans. The correlations were computed using all counties within these  
 609 states that reported crop yields during 2012.

Crop	States
Winter Wheat	Colorado, Illinois, Indiana, Kansas, Kentucky, Michigan, Missouri, Montana, Nebraska, North Dakota, Ohio, Oklahoma, South Dakota, Tennessee, Texas, and Wisconsin
Spring Wheat	Minnesota, Montana, North Dakota, and South Dakota
Corn	Illinois, Indiana, Iowa, Kansas, Kentucky, Michigan, Minnesota, Missouri, Nebraska, North Dakota, Ohio, South Dakota, and Wisconsin
Soybeans	Illinois, Indiana, Iowa, Kansas, Kentucky, Michigan, Minnesota, Missouri, Nebraska, North Dakota, Ohio, South Dakota, and Wisconsin

610

#### 611 4. Conclusions and Discussion

612 This study examined the evolution of several drought indicators sensitive to soil  
 613 moisture and vegetation conditions during the extreme flash drought event that impacted  
 614 most of the U.S., including some of the world's most productive farmland, during 2012.  
 615 The evolution of two satellite-based drought indicators, the ESI and VegDRI, was  
 616 compared to modeled soil moisture anomalies from NLDAS and to observed soil  
 617 moisture and crop conditions compiled by the USDA NASS. The modeled soil moisture  
 618 anomalies were assessed separately for the Noah, Mosaic, and VIC models in the  
 619 NLDAS system, and also for their ensemble mean. The response of each of these  
 620 datasets was compared to observed meteorological conditions and assessed at both  
 621 national and regional scales.

622 Overall, the results showed that rapid temporal changes in the NLDAS and ESI  
 623 datasets often preceded periods of rapid drought intensification in the USDM. In most  
 624 locations, dry conditions initially appeared in the NLDAS TS moisture anomalies before



625 appearing in the ESI and NLDAS TC soil moisture anomalies in subsequent weeks. This  
626 sequence occurs because except for early in the growing season when root depths are still  
627 shallow, vegetation will be able to access soil moisture over more than just the top 10 cm  
628 of the column, which means that ET can remain high even as the TS moisture decreases.  
629 For agricultural drought detection, however, the heightened sensitivity of the TS moisture  
630 to rainfall can lead to false alarms when dry spells are short-lived. Thus, when assessing  
631 agricultural drought severity, it is advantageous to use drought indices that are sensitive  
632 to vegetation, yet able to respond quickly to changing conditions. Decreases in the short-  
633 range (2- and 4-wk) ESI anomalies preceded observed changes in crop conditions by up  
634 to one month in the regional analyses, which is consistent with prior studies by Otkin et  
635 al. (2013, 2014). The NLDAS anomalies were typically similar to concurrent anomalies  
636 in the NASS TS and subsoil moisture datasets. The VegDRI anomalies were most  
637 similar to the TC soil moisture anomalies because that method uses longer-term climate  
638 indicators in addition to remotely sensed vegetation health estimates to assess drought  
639 severity. VegDRI anomalies tended to match the evolution of the USDM in regions with  
640 slow drought development, but lagged the USDM and other drought indicators when  
641 conditions were changing rapidly, making it less suitable as a flash drought early warning  
642 tool. For early warning during rapid onset drought events, it is important to use drought  
643 metrics that are able to capture rapid changes in precipitation, soil moisture, and  
644 vegetation conditions.

645         Comparison of the NLDAS soil moisture anomalies revealed large differences in  
646 behavior for each of the models assessed during this study. The Noah model consistently  
647 depicted the largest soil moisture anomalies, whereas moisture deficits in the VIC model

648 were often less severe because of its greater sensitivity to small rainfall events. In many  
649 situations, the larger improvements depicted by the VIC model were reasonable based on  
650 changes in the NASS soil moisture datasets; however, sometimes these improvements  
651 were too large. More detailed process studies are necessary to identify the reasons for the  
652 model differences and to determine if they also occur during less extreme drought events.

653 A detailed assessment of the NASS crop conditions for three regions revealed that  
654 range conditions were typically the first to deteriorate as drought severity increased  
655 followed thereafter by decreases in corn and soybean conditions. Comparison to the ESI  
656 anomalies showed that soybean conditions were most similar to the short-range (2 and 4  
657 week) ESI composites, whereas corn conditions more closely followed changes in the  
658 longer 8- and 12-wk ESI anomalies. This behavior suggests that crop-specific drought  
659 indices could be developed using ESI anomalies computed over different time periods  
660 that are optimized to depict conditions experienced by each crop. More research is  
661 required to assess this possibility.

662 Crop yield departures were also assessed using county-level yield data. Winter  
663 wheat yields were generally above average because that crop matured before the most  
664 severe drought conditions developed; however, significant yield losses occurred for both  
665 corn and soybeans. Corn yield losses were largest across those regions that experienced  
666 both extreme heat and dry weather during the pollination stage in July. Soybean losses  
667 were largest across the western Corn Belt because of the extreme drought conditions that  
668 developed there during the second half of summer when seed growth occurs. These yield  
669 losses were consistent with the drought severity depicted by the ESI and SPI and to a

670 lesser extent by the NLDAS and VegDRI datasets. These results demonstrate the utility  
671 of county-level crop information for ground truth assessment of drought indices.

672

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681

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