

1 **Cold surges and dust events: establishing the link between the East Asian Winter Monsoon and the**  
2 **Chinese loess record**

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39 **Abstract**

40 The Chinese loess/palaeosol succession is one of the most comprehensive and intensively studied  
41 archives of Neogene and Quaternary global palaeoclimate events. Its stratigraphic details are widely  
42 recognised to indicate close links to the history and function of the East Asian Winter Monsoon (EAWM)  
43 – one of the most active components of the Earth's climate system. But the formal meteorological links  
44 between the EAWM and dust emission, both in the present day and in the past, have not been established  
45 and with it, the veracity of the loess record as an indicator of the EAWM questioned. Here we show that  
46 present day major dust events over northern China, while largely occurring during spring, are

47 nevertheless 'conditioned' by the strength of the preceding EAWM. We also demonstrate, for the first  
48 time, a close link between the occurrence of dust events and the strength of the EAWM. From these  
49 findings, linked to global-scale climate model simulations, we conclude that the Chinese loess succession  
50 provides a convincing proxy record of the strength of the East Asian Winter Monsoon.

51

## 52 **1. Introduction**

53 The stratigraphy of the Chinese Loess Plateau, comprising inter-bedded loess and palaeosol sequences,  
54 provides an iconic Quaternary terrestrial record of East Asian glacial and interglacial events. The 'Red  
55 Clay' sequences extend the record into the early Neogene, and through this captures both East Asian  
56 'inland aridification' and the on-set of the East Asian monsoon regime. (Ding et al., 1992; Liu and Ding,  
57 1998; Guo et al., 2002; Stevens et al., 2007; An et al., 2014). The Quaternary component of the  
58 succession played a fundamental role in establishing the correlation between the terrestrial and marine  
59 glacial-interglacial records, stressing the loess record's global significance (Heller and Liu, 1984; Kukla,  
60 1987; Ding et al., 2002; Williams, 2014) This correlation carried the implication that the East Asian  
61 glacial-interglacial scale climate shifts, were captured through dust entrainment, transport, deposition and  
62 post-depositional sediment 'modification', which in turn meant that the loess record provides a register of  
63 the function and intensity of the East Asian monsoon (EAM) regime (An et al., 1991; Liu and Ding, 1998;  
64 An et al., 2014) .

65

66 Various loess related proxies have been used to reconstruct monsoon events. The record of East Asian  
67 Summer Monsoon (EASM) variability has been related to the weathering imprint of inter-bedded  
68 palaeosols with, traditionally, some emphasis on magnetic susceptibility (e.g., Heller and Liu, 1982; An  
69 et al., 2014). In the reconstruction of the East Asian Winter Monsoon (EAWM) variations, strong claims

70 have been drawn from changes in the rates of deposition – mass accumulation rates (MAR) - and loess  
71 grain-size changes, not always without difficulties (Stevens et al., 2007). From these studies has emerged:  
72 (i) the general claim of a strengthening of the EAWM during glacial stages with a significant downturn of  
73 EAWM activity during interglacial stages (e.g. Stevens et al., 2007; Hao et al., 2012; An et al., 2014); and  
74 (ii) the recognition of more short term stadial-interstadial to millennial scale events (Porter and An,  
75 1995; Sun et al., 2010, 2011; An et al., 2014).

76

77 Linked to an understanding of the East Asian component of the Asian monsoon winter regime at a range  
78 of time scales, are far-reaching implications, including the possibility of deciphering possible Atlantic–  
79 EAWM and EAWM–Indonesian–Australian Summer Monsoon teleconnections (Sun et al., 2011;  
80 Wyrwoll et al., 2007; Wang et al., 2012; Denniston et al., 2013). The recognition of millennial scale  
81 stronger EAWM events further emphasises the importance of the loess record in pointing to the global  
82 imprint of such events (Porter and An, 1995; Sun et al. 2011; Denniston et al., 2013).

83

84 These palaeoclimate interpretations of the significance of inferred winter loess depositional rates and  
85 grain size changes have been recently cast into doubt by claims that the record of dust events in the  
86 Chinese loess succession does not relate to a stronger EAWM (Roe, 2009; Lu et al., 2011). The claims  
87 have now entered the more general literature (Williams, 2014:153) and challenge a vast research effort  
88 with far-reaching implications for our understanding of global scale climate teleconnections and drivers.  
89 Roe’s (2009) claims are based on the fact that present-day dust events occur in spring and hence do not  
90 relate to or indicate winter synoptic states. Roe (op. cit.) recognises that such dust events are driven by  
91 strong winds associated with cyclogenesis and the passage of strong cold fronts, and points to the fact

92 that such events occur as a result of the breakdown of the Siberian High – the ultimate driver of the  
93 EAWM.

94

95 The details of proposed EAWM changes, whether in strength and/or frequency, have generally not been  
96 firmly grounded in a framework of the controlling climate drivers. With associated discussions bringing  
97 with them an element of circularity – high mass accumulation rates (MARs) and ‘coarse’ grain-size  
98 indicate a stronger EAWM, and from a stronger EAWM, high MARs and ‘coarse’ grain-sizes can be  
99 expected. Our objective here is to break this nexus and specifically determine the relationship between  
100 dust events–loess deposition and the strength of the EAWM. We attempt this by focusing on the  
101 controlling climatology of dust events and employ Ocean Atmosphere Global Climate Model simulation  
102 results for selected periods over the last 21,000 years to strengthen our claims.

103

## 104 **2. Cold surges in the climatology of the East Asian winter monsoon**

105 The EAWM dominates the climate of East Asia during the winter months (e.g., Chang et al., 2006) and is  
106 closely associated with the development of a cold core high pressure system over the Siberian-Mongolian  
107 region (Fig. 1). The strength of the EAWM is defined through the associated surface pressure and/or a  
108 consideration of the details of its dynamic controls (Jhun and Lee, 2004; Li and Yang, 2010). During the  
109 EAWM, the Siberian High with its central pressure reaching in excess of 1035 hPa, dominates much of  
110 the Eurasian continent; individual cases of central pressure as high as 1085 hPa have been reported. More  
111 strong northwesterly flows occur at its eastern margins, where the flow separates into one branch directed  
112 eastward into the subtropical western Pacific, and then tending southward in the direction of the South  
113 China Sea. At 500 hPa, a trough (the East Asian Trough) is evident, aligned with the longitudes of Japan.

114 At 200 hPa, the Polar East Asian Jet is prominent, with its maximum located just southeast of Japan. The  
115 Polar Jet is associated with strong baroclinic instability, large vertical wind shear and cold air advection.

116

117 The cold air 'excursions', also described as 'cold surges', are channeled by the trough southwards and are  
118 a characteristic feature of the EAWM (Lau and Chang, 1987) that impact strongly on the winter climate  
119 of eastern China. Their path is in part related to relief controls of the Tibetan Plateau, with their effect  
120 extending to the tropics, where they can lead to the flare-up of convective activity over the Maritime  
121 Continent (Chan and Li, 2004). The recognition of cold-surges as a meteorological feature extends back  
122 to the early part of the last century (Li, 1937; Lu, 1937). They were then termed 'cold waves' and were  
123 recognised as being associated with 'sandstorms' (Lu, 1937). Lu (op.cit.) described an event in March  
124 1936, which could be traced over much of east China extending into the south and associated with a  
125 regionally extensive sandstorm.

126

127 It is now recognised that the frequency of cold-surges can act as a surrogate for the strength of the  
128 EAWM (i.e. surface pressure) (Zhang et al., 1997; Ding and Sikka, 2006). For the period 1979 to 1995,  
129 Zhang et al. (1997) found an average annual occurrence of 13 cold surges per year based on the  
130 NCEP/NCA reanalysis data. Using different criteria, Chen et al. (2004) proposed that about 17 events  
131 occur every winter. Cold surges can last from 5 to 14 days (Zhang et al., 1997), hence forming a  
132 significant component of the winter climate of the southern East Asia – South China Sea region. But cold  
133 surges do not necessarily constitute dust depositional events, as emphasised by Roe (2009). In the present  
134 climate, the frequency of large-scale Asian dust events peaks between March and May (Shao and Dong,  
135 2006), with more than 85% of the annual dust storms occurring during spring (Zhu et al., 2008). The  
136 annual number of dust events originating from the Gobi Desert ranges between 20 and 35 (Sun et al.,

137 2001), and given their mainly spring occurrence, they would appear to be a poor indicator of the strength  
138 of the EAWM.

139

### 140 **3. Methods**

141 To establish the relationship between cold air outbreaks in winter (i.e. the strength of the EAWM), and  
142 dust activities in spring, we analysed cold surge events recorded at stations maintained by the China  
143 Meteorological Administration in North China (north of 34°N) over the 41-year time period between  
144 1961 and 2001. Initially, the data-sets employed in this study are from the daily mean surface temperature  
145 recorded at 833 stations across China by the China Meteorological Administration. In order to focus on  
146 northern China, this study examines only the stations that are north of 34°N. Then, all records of station  
147 records north of 34°N were screened for missing data and stations with too many missing values were  
148 removed. A station year is considered 'missing' if more than 1% of the days are missing. After screening,  
149 a total of 280 stations with 'non-missing' years during the study period could be retained. Here, a cold  
150 surge event at a single station is defined if the event-temperature drop  $\Delta T$  exceeds 10°C and the event  
151 minimum temperature  $T_{\min}$  is lower than its 10-day climate mean by -5°C. By averaging and then  
152 normalising the time series for the 280 stations during winter, a cold wave frequency index (CWFI) was  
153 derived. Wei and Lin (2009) provide further details of the calculations involved and Wang and Ding  
154 (2006) and Ma et al. (2008) provide examples of the applications of the criteria derived.

155

156 Corresponding dust event frequencies were obtained from synoptic data from 162 weather stations in  
157 Northern China (data available from the China Meteorological Administration). For these stations, days  
158 with dust events (also generally described as "dust weather"), including dust storms, blowing dust and  
159 dust in suspension, were registered. For an individual station, dust frequency is defined as the total

160 number of days between March and May in which a dust event is observed. Zhou et al. (2006) provide a  
161 detailed introduction to these data-sets.

162

163

### 164 **3. Results**

165 The correspondence between the dust event frequency and CWFI series (Fig. 2) yields a correlation  
166 coefficient of 0.73. The time series show a matching decreasing trend with time, and on detrending both,  
167 the positive correlation of 0.4 is retained. These data clearly suggest that a high winter cold surge  
168 frequency in northern China is accompanied by a high dust event frequency in the following spring.

169

170 In order to provide more details on the likely relationship between dust and cold-surge events, ‘strong’  
171 ( $>0.7$ ) and ‘weak’ ( $<-0.7$ ) CWFI years were selected and composite analyses of the corresponding dust  
172 events in the following spring were carried out. The composite analyses confirm a significant  
173 correspondence between the winter cold surges and spring dust events. The results clearly show that the  
174 CWFI and dust frequency are positively correlated, making it evident that when the winter cold surge  
175 frequency in northern China is high, then the following spring dust event frequency is also high (Fig. 3).

176

### 177 **4. Discussion**

178 The circulation controls of matching high (low) dust/strong (weak) CWFI events are highlighted by the  
179 difference in the spring 500 hPa geopotential height and spring sea level pressure between ‘strong’ and  
180 ‘weak’ CWFI years (Fig. 4). In comparing the two composites, the deepening of the Aleutian Low is  
181 prominent and in strong CWFI years, the East Asian Trough is on average, 20 gpm lower, with surface  
182 pressure differing by 2 hPa. During a strong CWFI year (high dust frequency year), the Siberian High is



183 relatively strong; the Aleutian Low is relatively deep with its western part having a stronger northerly  
184 wind component; the Polar High is strong; the European Shallow Trough is relatively deep; the East  
185 Asian Trough is deep and the Ural Ridge is strong. A conceptual overview of the three-dimensional  
186 structure of the EAWM during strong EAWM years is given in Figure 5. The outlined circulation controls  
187 can be placed into the more general context of hemisphere-scale conditions (Gong et al., 2001; Wei and  
188 Lin, 2009; Park et al., 2011; Chang and Lu, 2012). Gong et al. (2006) established a correlation between  
189 the inter-annual variation of dust storm frequency over northern China and the Arctic Oscillation (AO). A  
190 negative phase of the AO is associated with a deeper East Asian Trough and a strong Siberian High with  
191 associated cold surges in northeast Asia (Gong et al., 2001; Wu and Wang, 2002) – i.e. a strong EAWM.  
192 This reconstruction emphasises that during strong EAWM years, the Siberian High has a relatively high  
193 surface pressure, with the Aleutian Low relatively deep and with its western part having a strong  
194 northerly wind component. The East Asian Trough is deeper with a strong Polar Jet evident in the upper  
195 troposphere. These configurations lead to stronger winter cold air outbreaks followed by dust events in  
196 the following spring.

197

198 The Siberian High forms in response to strong radiative cooling in the lower troposphere, confining it to  
199 the lower levels of the troposphere – below 500 hPa (Panagiotopoulos et al., 2005). Consequently, given  
200 these controls and the correspondingly shallow nature of the Siberian High, it is to be expected that it was  
201 significantly enhanced during glacial stages. To support our claim, we use results from a transient  
202 simulation of the last 21,000 years (TraCE-21K – Liu et al., 2009; He et al., 2013) with the coupled  
203 atmosphere–ocean Community Climate System Model Version 3 (CCSM3). Fig. 6 shows the simulated  
204 sea surface pressures over the East Asian region for selected time periods during the last 21,000 years.  
205 The model results indicate a strengthening of the Siberian High during glacial stages, showing

206 progressive weakening of the Siberian High by 12 ka, with this trend continuing into the Holocene. From  
207 the model results and the present day relationship between the strength of the Siberian High, cold surges  
208 and dust events, the inference of an increase in the frequency of cold surge events during glacial stages  
209 follows.

210

211 Given the clear association of dust events and the strength of the Siberian High (i.e. the EAWM), it  
212 follows that glacial (interglacial) stages were characterised by high (low) loess mass accumulation rates  
213 (MAR), possibly associated with coarser (finer) loess grain-sizes (e.g. Kohfeld and Harrison, 2003;  
214 Stevens et al., 2007; Sun et al., 2011; Kang et al., 2013; An et al., 2014). In providing a more secure  
215 chronology, Kang et al. (2015) have been able to give more refinement to these claims. They demonstrate  
216 that despite significant variations, all sites considered by them show increased MARs during ~26 to 19 ka,  
217 with notable peaks in some stratigraphic successions during 23 to 19 ka. Generally lower MARs were  
218 recognised for the time period, 19 to 12 ka. These results show a general correspondence with the strength  
219 of the model reconstruction of the Siberian High (Fig. 6), with decreasing strength of the Siberian High  
220 corresponding to lower MARs. In this it should not be overlooked that land surface characteristics can be  
221 a control on dust entrainment and that this may be reflected in MARs. For instance, the distribution of  
222 frozen ground at the Last Glacial Maximum (Liu and Jiang, 2016) is likely to have had an impact on  
223 entrainment rates and hence the 'matching' MARs of some loess successions.

224

225 While our model results point to a strengthening of the Siberian High during glacial stages and hence a  
226 stronger EAWM, the degree to which this can be expected to be apparent in loess grain-size variations  
227 may not always be straightforward. It is generally recognised in the Chinese loess literature that grain-size  
228 characteristics are related to source controls, transport paths and wind strength (e.g., Pye, 1987, cited in

229 An et al., 2014) But despite these complexities, a strong claim exists for a link between the grain-size  
230 characteristics of Chinese loess and the strength of the EAWM (summary in An et al., 2014: 54-57). The  
231 potential importance of such claims is highlighted by the recent findings of Sun et al. (2011). Using the  
232 Gulang and Jinguang loess successions, they argue that grain-size variations over the last ~60,000 years  
233 reflect changes in the strength of the EAWM and draw the conclusion that Atlantic meridional  
234 overturning imprints itself on the EAWM through changes in the westerly circulation that ‘transmits’ this  
235 signal from the North Atlantic to the EAWM region (Sun et al., 2011). These are very important  
236 inferences that stress the need for focused field and climate modeling work, specifically addressing the  
237 questions of the controls of loess grain-size. A need that is further emphasised by the recent Nie et al.  
238 (2015) claim of the link between Yellow River sediment sources and the development of the Chinese  
239 Loess Plateau.

240

## 241 **5. Conclusions**

242 Given that the strength of the EAWM is positively related to the occurrence of dust events, it follows that  
243 the details of the Chinese loess record serve as valid proxy indicators of the strength of the EAWM. In  
244 this claim, we use the term ‘strength’ of the EAWM in the strict sense, i.e., related to the frequency of  
245 cold surges, which is determined by the northern hemispheric circulation in northeast Asia, with the  
246 strength of the Siberian High, as measured by its central pressure, being a key driver. Accepting this  
247 inference and combining it with the recent advances in loess dating techniques, recognition of regional  
248 depositional patterns, ‘switching’ loess sources and climate modeling, brings with it the potential for a  
249 comprehensive reconstruction and understanding of the dynamic palaeoclimatology of the EAWM over  
250 long time scales and through this, ‘capturing’ wider global scale climate teleconnections and drivers.

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395

## List of Figures

Figure 1: December/January/February sealevel pressure (mb) – 2000 to 2010 ( NCEP/NCAR Reanalysis).

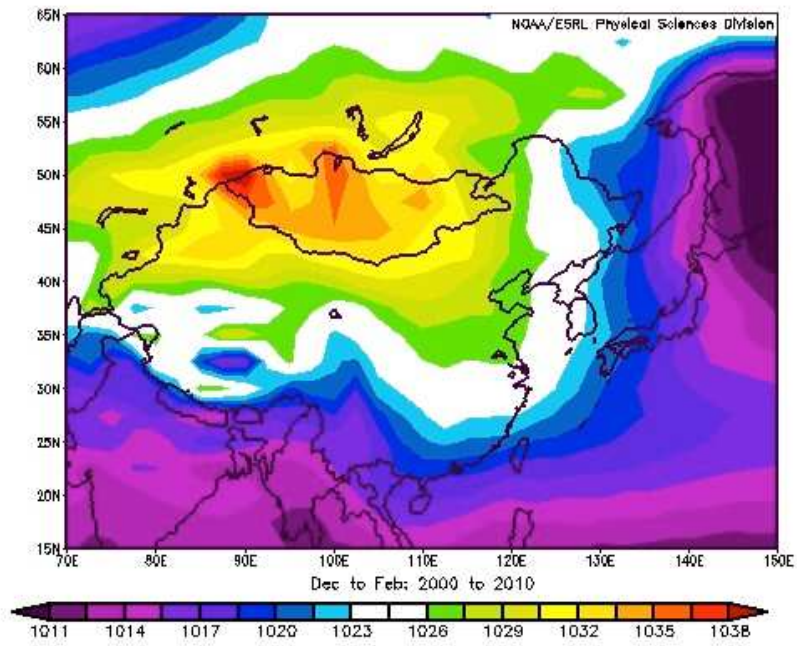
Figure 2: Normalized winter Cold Wave Frequency Index (CWFI; red solid line; red marks; unitless), normalized spring dust event frequencies (blue solid line; blue marks; unitless), and their linear trends (dash lines) for the period 1961-2001.

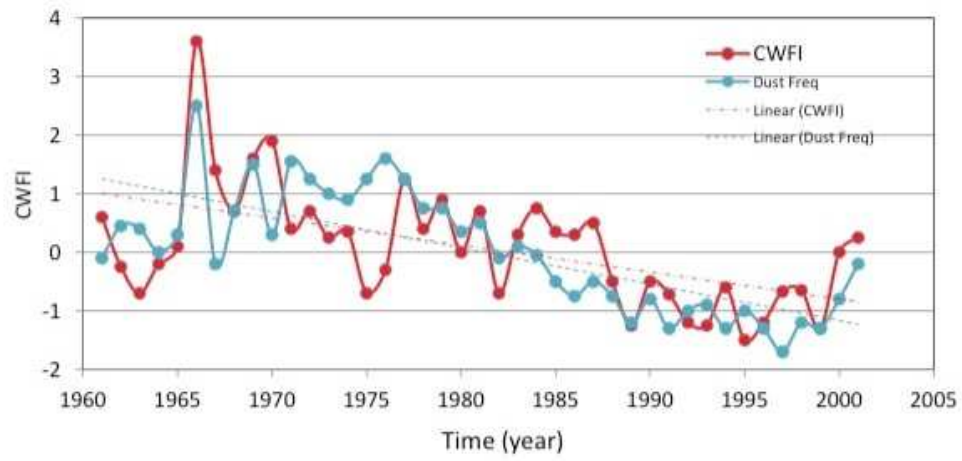
Figure 3: Correlation between CWFI and spring dust event frequencies – shaded area significant at 95%, 99%, and 99.9% confidence level, respectively: (a) raw data, (b) detrended data.

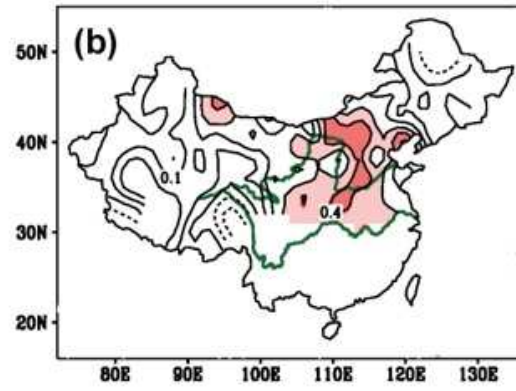
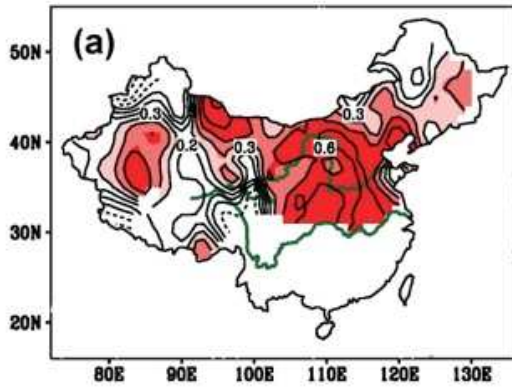
Figure 4: Differences between strong (high dust frequency) and weak (low dust frequency) CWFI years: (a) spring 500 hPa geopotential height (unit: gpm), (b) spring sea level pressure (unit: hPa).

Figure 5: Conceptual representation of the synoptic expression of strong EAWM years promoting cold surges and the incidence of dust events over northern China.

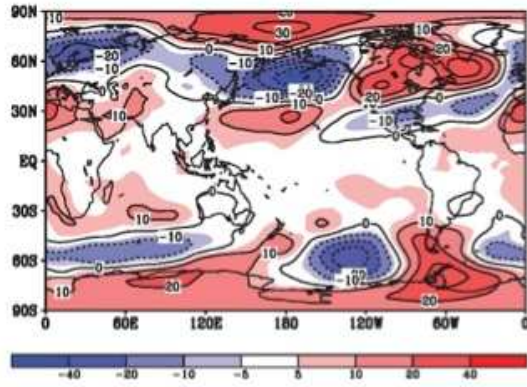
Figure 6: Winter (DJF) model sea level pressure for selected time periods in the last 21 000 years .



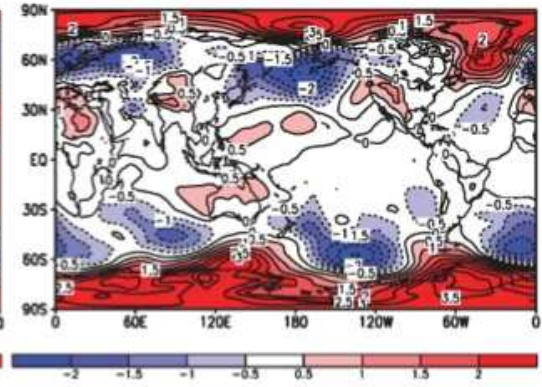


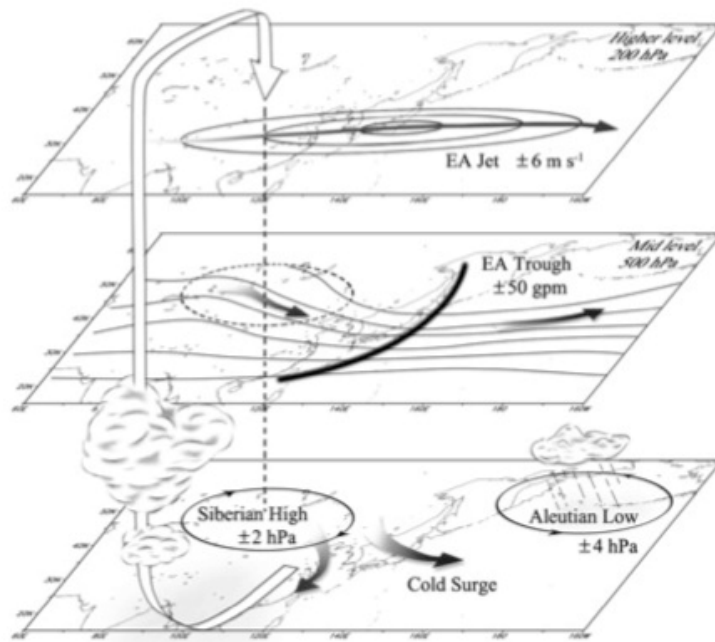


(a) 500 hPa geopotential height (unit: gpm)

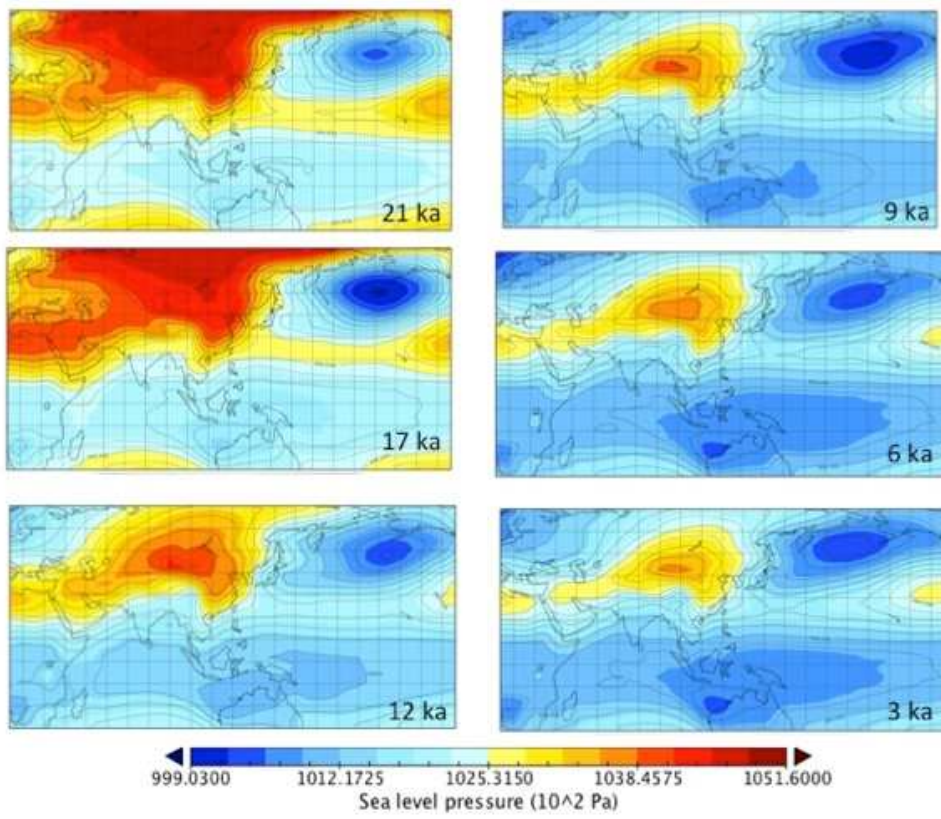


(b) sea level pressure (unit: hPa)









- The Chinese loess/palaeosol succession is one of the most comprehensive and intensively studied archives of Neogene and Quaternary global palaeoclimate events.
- The succession is widely claimed to be linked to the strength of the East Asian Winter Monsoon (EAWM), an inference that has been recently challenged.
- Here we show that present day major dust events over northern China are 'conditioned' by the strength of the preceding EAWM.
- These findings are related to global-scale palaeoclimate model simulations that show a progressive weakening of the EAWM over the last c. 21,000 years.
- From these findings we confirm that the Chinese loess succession provides a convincing proxy record of the strength of the East Asian Winter Monsoon.