1	A Novel Hybrid Artificial Neural Network - Parametric Scheme for Postprocessing
2	Medium-Range Precipitation Forecasts
3	
4	
5	by
6	
7	
8	Mohammadvaghef Ghazvinian <sup>a</sup> , Yu Zhang <sup>a</sup> , Dong-Jun Seo <sup>a</sup> , Minxue He <sup>b</sup> , Nelun Fernando <sup>c</sup>
9	
10	
11	
12	
13	<sup>a</sup> Department of Civil Engineering, The University of Texas at Arlington, Arlington, TX 76019,
14	USA
15	<sup>b</sup> California Department of Water Resources, 1416 9 <sup>th</sup> Street, Sacramento, CA 95814, USA
16	<sup>c</sup> Texas Water Development Board, 1700 North Congress Avenue, Austin, TX 78701, USA
17	
18	
19	
20	
21	Revised Version: March 25, 2021
22	Manuscript submitted to the Advances in Water Resources
23	
24	
25	
26	
27	
28	Corresponding author: Mohammadvaghef Ghazvinian
29 30 31	Department of Civil Engineering, The university of Texas at Arlington, 427 Nedderman Hall-416 Yates Street, Arlington, TX 76019, USA Email address: mohammadvaghef.ghazvinian@mavs.uta.edu

1

### 32 Abstract

Many present-day statistical schemes for postprocessing weather forecasts, in particular 33 precipitation forecasts, rely on calibration using prescribed statistical models to relate forecast 34 35 statistics to distributional parameters. The efficacy of such schemes is often constrained not only by prescribed predictor-predictand relation, but also by arbitrary choices of temporal window 36 37 and lead time range for training. To address this limitation, we propose an end-to-end, computationally efficient hybrid postprocessing scheme capable of producing full predictive 38 distributions of precipitation accumulation without explicit stratification of forecast-observation 39 pairs by forecast lead time and season. The proposed framework uses the censored, shifted 40 gamma distribution (CSGD) as the predictive distribution but uses an artificial neural network 41 (ANN) to estimate the distributional parameters of CSGD through a unified approach. This 42 approach, referred to as ANN-CSGD, allows for simultaneous estimation of distributional 43 parameters over multiple lead times and seasons in a single model by incorporating the latter 44 variables as predictors to the ANN. We test our proposed ANN-CSGD model for postprocessing 45 of ensemble mean forecasts of 24-h precipitation totals over selected river basins in California, at 46 47 one- to seven-day lead times, from the Global Ensemble Forecast System (GEFS). The probabilistic quantitative precipitation forecasts (PQPFs) from the ANN-CSGD, are more skillful 48 overall than those from the benchmark CSGD and the Mixed-type meta-Gaussian distribution 49 (MMGD) models. The ANN-CSGD PQPFs highly improve the performance of those from 50 CSGD in predicting the probability of precipitation (PoP) and are also much sharper and reliable 51 52 at higher precipitation thresholds. We demonstrate how the hybrid approach, by using the entire 53 available training data and its modified formulation, efficiently represents interactions between GEFS forecasts and season/lead times, thus leading to enhanced predictive performance. 54

55 Keywords: Statistical postprocessing; Artificial neural networks; Probabilistic quantitative
56 precipitation forecast; Predictive distribution

### 57 **1. Introduction**

Statistical postprocessing techniques are increasingly used to improve the reliability and skill of 58 real time probabilistic quantitative precipitation forecasts (PQPFs) produced by numerical 59 weather prediction (NWP) models. Broadly speaking, these techniques can be categorized as 60 nonparametric and parametric ones. A prominent example of the former is the Analog approach 61 (Hamill and Whitaker 2006; Hamill et al. 2015). The parametric techniques rely on prescribed 62 parametric forms of conditional (predictive), joint and marginal distributions, and employ 63 various techniques ranging from regression to the method of moments, and their variants, for 64 estimating distributional parameters. Many of the modern parametric approaches fall under the 65 66 broad umbrella of Ensemble Model Output Statistics (EMOS; Gneiting et al. 2005), also known as nonhomogeneous regression. As the name implies, the EMOS approaches use prescribed 67 predictive distributions and relate distributional parameters to ensemble statistics through a set of 68 regression equations (Scheuerer and Hamill 2015; Zhang et al. 2017; Stauffer et al. 2017). 69

The extent to which postprocessing techniques have improved forecast skill has varied in 70 practice (Li et al. 2017; Wilks 2018; Vannitsem et al. 2020). There are several common 71 72 limitations in postprocessing methods adopted to date. Among the frequently cited are the 73 inflexible and subjective way of selecting predictors, structural rigidity that makes it difficult to 74 integrate ancillary predictors, and the ad hoc way of determining spatial-temporal training domains (see related discussions in Rasp and Lerch 2018). The advent of machine learning 75 techniques offers many new opportunities to address these limitations. Relative to the parametric 76 approaches, EMOS techniques included, some of the recent machine learning techniques offer 77

3

flexibility in identifying predictors, in integrating ancillary information, and in capturing 78 79 complex, nonlinear predictor-predictand relationships that are difficult to characterize parametrically (see, e.g., Taillardat et al. 2019). Particularly promising are the various artificial 80 neural networks (ANNs) which have been known for their ability to model nonlinear 81 dependencies. Recent years have seen an explosion of ANN-based prediction paradigms (Liu et 82 al. 2016; Brenowitz and Bretherton 2018; Gentine et al. 2018; Rasp et al. 2018; Chapman et al. 83 84 2019; Cloud et al. 2019; Gagne et al. 2019; Lagerquist et al. 2019). Yet, the use of these techniques in the context of postprocessing remains relatively limited. Rasp and Lerch (2018) is 85 perhaps the first attempt of this nature. The authors explored a hybrid scheme that retains a 86 87 parametric form of the predictive distribution of 2-m temperature but relies on ANNs to estimate the distribution parameters from the ensemble statistics of 2-m temperature as well as ancillary 88 89 variables. Scheuerer et al. (2020), in a similar vein, developed an ANN-based scheme for 90 producing 7-day accumulated PQPFs at subseasonal range (2-4 weeks) from NWP ensemble forecasts, and showed that the PQPFs thus generated broadly outperforms climatology. Other 91 92 studies of note include Bremnes (2020) where ANN was used for postprocessing wind speed forecasts. Collectively, these studies indicate that embedding local information and incorporating 93 ancillary forecast variables can lead to larger improvements in forecast skills. They further 94 suggest that ANN models, contrary to the common perception of being black boxes, can help 95 uncover, and offer physical insights to the meteorological processes that underpin the links 96 between predictors and predictands. 97

98 Inspired by the successes of recent ANN-based postprocessing approaches, and motivated by the 99 broader need for improving the skill of PQPF while circumventing limitations inherent in 100 existing EMOS schemes, we propose a hybrid ANN-nonhomogeneous regression-based scheme 101 capable of postprocessing precipitation forecasts at multiple lead times and seasons in a unified way. The proposed scheme retains the parametric form of the predictive distribution of 102 precipitation proposed by Scheuerer and Hamill (2015) and Baran and Nemoda (2016), but 103 departs from the conventional EMOS by using ANNs to relate NWP forecasts to the 104 distributional parameters. The potential advantages of the proposed scheme, which we will 105 henceforth refer to as ANN-CSGD are three-fold. First, this scheme does not require an explicit 106 107 prescription of predictor-predictand relationships as is currently done in EMOS models - it can 108 discover and integrate arbitrary nonlinear relationships through training. Second, the training of the model can be done using the entire data archive and thereby obviate the need for explicit 109 110 treatment of lead time-based and seasonally varying NWP forecast errors. Third, it can account for seasonal variations in the interaction between NWP forecasts and temporal predictors. 111

112 In this paper we describe and evaluate the proposed scheme which relies only on the ensemble 113 mean of NWP forecasts as the major predictor. The evaluation is conducted for sub-basins within three selected river basins in California. The proposed scheme is applied to postprocess 114 Global Ensemble Forecast System (GEFS; Hamill et al. 2013) precipitation reforecasts along 115 with two benchmark schemes. The first is the single predictor version of the censored, shifted 116 gamma distribution (CSGD; Scheuerer and Hamill 2015). The second is the Mixed-type Mata-117 118 Gaussian Distribution (MMGD; Wu et al. 2011), which has been the standard method in the U.S. National Weather Service (NWS) Hydrologic Ensemble Forecast Service (HEFS; Demargne et 119 al. 2014). Our overarching hypothesis is that the flexibility accorded by the ANN-based model in 120 establishing complex predictor-distributional parameter relationships, in determining temporal 121 training windows, and in lumping forecasts for different lead times, will help the proposed 122 scheme attain superior predictive performance relative to the benchmarks. 123

The reminder of this paper is organized as follows. Section 2 describes the proposed ANN-CSGD scheme as well as the benchmark methods, data, and experimental setup. Section 3 presents the outcomes of the experiments and section 4 summarizes the findings and discusses future possible extensions.

### 128 **2.** Materials and methods

#### 129 2.1. Proposed model

The censored, shifted gamma distribution (CSGD) introduced by Scheuerer and Hamill (2015), has been a popular choice to represent the right skewed, mixed-type dichotomous-continuous nature of the predictive distribution of precipitation (Scheuerer and Hamill 2015; Baran and Nemoda 2016; Zhang et al. 2017; Scheuerer et al. 2020). Let  $F_{k,\theta}$  denote the cumulative distribution function (CDF) of the gamma distribution with shape parameter k > 0 and scale parameter  $\theta > 0$ . The CDF at realized precipitation value y, and quantile functions of CSGD for any  $0 \le p < 1$  are defined by (Scheuerer and Hamill 2015; Baran and Nemoda 2016):

$$F^{0}{}_{k,\theta,\delta}(y) = \begin{cases} F_{k,\theta}(y-\delta), & y \ge 0\\ 0, & y < 0 \end{cases}$$
(1)

$$q_p = max \left[ 0, \delta + F_{k,\theta}^{-1}(p) \right]$$
<sup>(2)</sup>

137 where the additional parameter,  $\delta < 0$  shifts the gamma distribution to the negative values. To 138 form the CSGD, the shifted gamma distribution is left censored at zero by assigning the mass 139 probability  $F_{k,\theta}(-\delta)$  to the origin to account for non-negativity of precipitation amounts. To 140 relate the mean  $\mu = k\theta$ , standard deviation  $\sigma = \sqrt{k}\theta$ , and shift parameter  $\delta$  of predictive CSGDs 141 to the predictors, we propose a fully connected (dense) feed forward neural network where each 142 node receives a linear combination of weighted outputs from nodes in the previous layer, adjusts it by adding a bias quantity, and applies an activation function to the result. Our proposed ANN-

144 CSGD structure (Fig. 1) consists of the following elements:

- Input layer, where covariates are introduced to the network.
- One hidden layer; we use the exponential linear unit (ELU) with α = 1 as the activation
   function to introduce nonlinearity to the network

$$f(x) = \begin{cases} x, & x > 0 \\ \alpha [\exp(x) - 1], & x \le 0 \end{cases}$$
(3)

- 148 ELUs are known to provide more precise and faster learning compared to the other 149 activation functions in deep learning experiments (see, Clevert et al. 2015).
- Layer normalization (Ba et al. 2016) which normalizes each sample output from hidden nodes to maintain the mean and standard deviation of node outputs within each example close to 0 and 1, respectively. Recent studies (see, e.g., Xu et al. 2019) show that Layer normalization helps stabilize the training process by enabling smoother gradients and yields faster training convergence.
- Output layer with a linear activation function. We set three CSGD parameters as functions of the network outputs  $O_i$  to constrain the values of these parameters to reasonable ranges (i.e.,  $\mu, \sigma > 0$  and  $\delta < 0$ ). Therefore, we set  $\delta = -sqrt(O_1^2)$ ,  $\mu =$ exp $(O_2)$ , and  $\sigma = exp(O_3)$ . These additional functions can be interpreted as inverse link functions used in conventional distributional regression or generalized additive models for location, scale, and shape (GAMLSS; Rigby and Stasinopoulos 2005) (see, also, Cannon 2012; Rasp and Lerch 2018).
- We incorporate the ensemble mean forecast, forecast lead time (1 to 7 days), and month of the year of the verifying observations (1 to 12) as predictors to the ANN. Using the latter two predictors enables us to train a single model to postprocess forecasts from multiple lead times

and months. Lead time values are normalized by dividing each quantity by the maximum value 165 (i.e., day/7).To account for seasonal cycle, the cosine 166 we use term  $[cos(2\pi(month-1)/12)]$  to both introduce the cyclical nature of the month of the year to the 167 network and to enforce the network to encode the annual cycle of precipitation over the study 168 area (see, Liu et al. 2018; Scheuerer et al. 2017). 169

We retain the average value of continuous ranked probability score (CRPS) of predictive CSGDs
as the loss function for training the weights and biases of the ANN-CSGD. The ANN is trained
by minimizing the CRPS computed using collocated and coincidental forecast-observation pairs
over training data (see the appendix B for the mathematical definition of CRPS)

$$CRPS = \frac{1}{N} \sum_{i=1}^{N} \operatorname{crps}(F_{k_i \theta_i \delta_i}, y_i)$$
(4)

The analytical expression of CRPS for a paired CSGD predictive distribution and verifying
observation was proposed by Scheuerer and Hamill (2015). Similarly, we implement

$$\operatorname{crps}\left(F_{k_{i}\theta_{i}\delta_{i}}, y_{i}\right) = (y_{i} - \delta_{i})\left[2F_{k_{i},\theta_{i}}(y_{i} - \delta_{i}) - 1\right] - \frac{\theta_{i}k_{i}}{\pi}B\left(\frac{1}{2}, k_{i} + \frac{1}{2}\right)\left[1 - F_{2k_{i},\theta_{i}}(-2\delta_{i})\right] + \theta_{i}k_{i}\left[1 + 2F_{k_{i},\theta_{i}}(-\delta)F_{k_{i}+1,\theta_{i}}(-\delta_{i}) - F_{k_{i},\theta_{i}}(-\delta_{i})^{2} - 2F_{k_{i}+1,\theta_{i}}(y_{i} - \delta_{i})\right] + \delta F_{k,\theta}(-\delta)^{2}$$

$$(5)$$

where B(0,0) is the beta function, and  $(k_i, \theta_i, \delta_i)$  are three parameters of *ith* predictive CSGD with  $y_i$  being the corresponding verifying observation. To minimize the loss function, we use the Adam stochastic gradient descent-based optimization algorithm (Kingma and Ba 2014) and update model parameters based on small batches randomly sampled from the training dataset. One major challenge in applying ANNs is to constrain the complexity of the model while attaining optimal predictions. Overfitting can occur if a very complex structure is used. Several regularization techniques to reduce generalization errors in ANNs are available as reviewed by 183 Goodfellow et al. (2016). Among them, we use early stopping, which is one of the most popular184 and widely used regularization techniques in ANNs.

In our work, we leave 20% of the available training data as the validation set and do not include them in training process. This practice enables us to reduce overfitting by monitoring the average loss value over the validation set while we train the model, and return the best possible training parameters (weights and biases) at the time when the lowest CRPS for the validation set is achieved. We terminate training when no further decrease in validation set loss is seen after 15 iterations through all training batches or the entire training data (epochs), with up to 1000 epochs.

We train ANNs using the previously described process, with all possible combinations ofdifferent settings, using the early stopping technique for the following hyperparameters

- *Number of nodes in the hidden layer:* {5,10,15}
- 195

196

• *Batch size:* {2048, 4096, 8192}

• *Learning rate of the Adam optimization algorithm:* {0.01,0.005}

All networks are trained with the same random number generator (seed) and are evaluated 197 198 based on the average loss value in the validation set. The ANN configuration with the lowest validation loss is chosen for out-of-sample predictions. Individual tested ANNs have  $0{7n + 3}$ 199 trainable parameters where *n* refers to the number of nodes in the hidden layer. We used a simple 200 201 (non-trained) layer as the normalization layer. Our assessments showed that training Layer 202 normalization parameters (beta and gamma) does not yield significant improvement over the non-trained one and possibly increases the risk of overfitting due to the increased number of 203 204 overall network parameters.

205

- 207 2.2. Benchmark models
- 208 2.2.1. CSGD

To generate postprocessed precipitation forecasts at a given location, for each forecast lead time and month of the year, Scheuerer and Hamill (2015), first fit three climatological CSGD parameters ( $\mu_{cl}$ ,  $\sigma_{cl}$  and  $\delta_{cl}$ ) to locally observed training precipitation data using a 91-day temporal window centered around the 15<sup>th</sup> of each month. In the second step these parameters are included in nonlinear, nonhomogeneous regression equations to relate monthly parameters of predictive CSGDs to statistics of spatially smoothed ensemble of forecasts.

In this study, we use the regression equations that incorporate only the ensemble mean:

$$\mu = \mu_{cl}/a_1 \log\{1 + \left[(\exp(a_1) - 1)\left(a_2 + a_3 \overline{f}/\overline{f_{cl}}\right)\right]\}$$
(6)

$$\sigma = a_4 \sigma_{cl} \sqrt{\mu/\mu_{cl}} \tag{7}$$

$$\delta = \delta_{cl} \tag{8}$$

where  $\overline{f}$  and  $\overline{f_{cl}}$  correspond to the raw ensemble mean forecasts and their climatological mean in training data, respectively. In the version of CSGD described in Scheuerer and Hamill (2015), the predictive shift parameter  $\delta$  is kept identical to the climatological shift to ensure that the predictive CSGD reverts to climatology as a limiting case when the forecast becomes less skillful (e.g., at longer lead times) (see related discussion in Scheuerer and Hamill 2015).

The four regression coefficients  $a_1, a_2, a_3, a_4$  are estimated by minimizing the CRPS using the closed form expression proposed by Scheuerer and Hamill (2015) (see sec. 2.1) as a function of CSGD parameters over training data.

Past studies (Scheuerer and Hamill 2015; Baran and Nemoda 2016; Zhang et al. 2017; Baran

and Lerch 2018; Taillardat et al. 2019) show that CSGD method and its variants perform well in

comparison with other modern postprocessing techniques. Recent exploratory analyses (see,
Ghazvinian et al. 2020, Fig. 1) showed that the climatological CSGD shift parameter, derived by
CRPS minimization approach, tends to be inflated and this leads to an underestimation of a
probability of precipitation (PoP). This bias directly affected the performance of predictive
CSGD, primarily in predicting PoP and, to a degree, the predicted magnitude of precipitation.
This was particularly evident at shorter lead times and in rainy seasons where the predictive
distribution of precipitation deviates widely from climatology.

233 2.2.2. MMGD

The MMGD (Herr and Krzysztofowicz 2005; Wu et al. 2011) was developed by the U.S. NWS 234 235 as a component of the Meteorological Ensemble Forecast Processor (MEFP) of the operational HEFS (Demargne et al. 2014). This mechanism is routinely used to generate calibrated PQPF 236 from single-valued precipitation forecasts (ensemble mean) at river basin scales and at temporal 237 238 aggregation scales ranging from 6-h to 3-months and for lead times up to 9-months (Wu et al. 2018; Demargne et al. 2014). In contrast to the CSGD, where PoP and the probability of 239 240 magnitude of precipitation are estimated using the same predictive distribution, MMGD uses a Bayesian approach to break down the predictive distribution to explicitly account for the 241 dichotomous-continuous nature of precipitation. 242

Let *X* and *Y* denote the random variables of a single-valued quantitative precipitation forecast and the observed precipitation amount, respectively. The conditional distributions of observed precipitation, given a current forecast of no precipitation and positive precipitation, are given as follows (details of this derivation can be found in Wu et al. 2011; Ghazvinian et al. 2020):

$$F_{Y|X}(y|x, x = 0) = P(Y = 0|X = 0) + P(0 < Y \le y|X = 0)$$

$$= a + (1 - a)G_Y(y)$$
(9)

$$F_{Y|X}(y|x, x > 0) = P(Y \le y|X = x, X > 0)$$

$$= c(x) + (1 - c(x))D_{Y|X}(y|x)$$
(10)

where a and c(x) represent mass probabilities of observed precipitation being equal to zero, and 247  $G_Y(y) =$ 248 are combined with the continuous conditional distributions  $P(Y \le y | X = 0, Y > 0)$  and  $D_{Y|X}(y | x) = P(Y \le y | X = x, X > 0, Y > 0)$  to construct the 249 predictive distributions. To estimate  $D_{Y|X}(y|x)$ , its marginal continuous variates [X|X > 0, Y >250 0] and [Y|X > 0, Y > 0] undergo normal quantile transformation (NQT), yielding standard 251 normal variates  $U = \Phi^{-1}[D_X(x)]$  and  $V = \Phi^{-1}[D_Y(y)]$ . Following the meta-Gaussian 252 distribution theorem of Kelly and Krzysztofowicz (1997),  $D_{Y|X}(y|x)$  assumes the following 253 form 254

$$D_{Y|X(y|x)} = \Phi\left[\frac{\Phi^{-1}[D_Y(y)] - \rho \Phi^{-1}[D_X(x)]}{\sqrt{1 - \rho^2}}\right]$$
(11)

where  $\Phi()$  and  $\Phi^{-1}()$  denote the standard normal CDF and quantile function of standard normal distribution, respectively; and  $\rho$  is the Pearson's product correlation coefficient between *U* and *V*.

The performance of MMGD has been evaluated in a number of studies (see, e.g., Wu et al. 258 2011; Brown et al. 2014a; Demargne et al. 2014; Kim et al. 2018; Seo et al. 2015; Ghazvinian et 259 al. 2019). While conclusions indicate that overall, MMGD produces reliable POPFs and is 260 capable of preserving the skill in the raw forecast, its PQPFs underestimate heavy-to-extreme 261 precipitation amounts (low reliability for higher thresholds). The latter finding was also 262 corroborated by Zhang et al. (2017), where the authors compared the performances of MMGD 263 264 and CSGD over the Mid-Atlantic region in U.S. Their results pointed to the superior performance of CSGD. In that study, CSGD's ability to ingest additional ensemble statistics as 265

predictors was shown to play a key role in its outperformance. Further performance comparisons by Ghazvinian et al. (2020), which relied on only the ensemble mean predictor and were conducted over the American River Basin in California, pointed to the clear outperformance of MMGD, particularly in predicting PoP. The authors confirmed that the use of a two-part scheme helped improve the representation of the predictive distribution.

We select MMGD as the second reference model to further address these discrepancies in the findings of previous studies. This enables us to determine whether our unified ANN-CSGD model improves upon the operational paradigm (MMGD), especially in situations where CSGD underperforms the latter, and helps us identify possible factors that contribute to the differential performance of the three schemes.

276 2.3. Data and experimental setup

The experiments focus on 24-h mean areal precipitation (MAP) totals over sub-basins of three major river basins in the service area of the NWS California-Nevada River Forecast Center (CNRFC; https://www.cnrfc.noaa.gov).

We use ensemble mean precipitation forecasts from January 1985 through December 2016 (32 280 years) for lead times 1 to 7 days. These data were obtained from the Global Ensemble Forecast 281 System (GEFS; version 10) reforecast dataset (Hamill et al. 2013) and were processed by the 282 283 CNRFC at 1-degree spatial resolution and 6-h accumulation intervals issued daily at 00 universal time (UTC). As ground truth, we use the basin MAP data generated by the CNRFC. The MAP 284 data were created using the so-called Mountain Mapper tool, which relies on the Parameter-285 elevation Regressions on Independent Slopes Model (PRISM; Daly et al. 2008) to group gauges 286 and interpolate gauge reports onto the domain of each watershed. The CNRFC MAP series are at 287 6-h increments and are available for the period between October 1948 and September 2017. The 288

MAP data were temporally aggregated to 24-h accumulation and paired with coincidentalreforecasts.

Postprocessing experiments are performed over sub-basins in the American River Basin (NFDC1, FOLC1), the Russian River Basin (WSDC1, GUEC1), and the Eel River Basin (DOSC1, FTSC1) (Fig. 2), and separately for upper/lower elevation zones when applicable. Subbasin names and corresponding NWS IDs are presented in Table 1. The CNRFC runs HEFS routinely to produce postprocessed PQPFs and ensemble streamflow forecasts for many of the sub-basins.

For each river basin, we selected one headwater and one downstream sub-basin for the hindcast experiment to examine the potential elevation dependence in forecast skills. The selected basins have been recognized for their importance in water resources management and flood control, as noted in past hydrometeorological forecast postprocessing/verification studies (see, e.g., Wu et al. 2011; Brown et al. 2012; Seo et al. 2015; He et al. 2016; Scheuerer et al. 2017; Ghazvinian et al. 2020).

The climate of the region is characterized by very dry summers, with most of its annual 303 precipitation falling during the cool season (October - April), and the highest monthly averaged 304 precipitation typically recorded in January. The American River originates from the Tahoe and 305 El Dorado national forests of the Sierra Nevada and is one of the major water supply sources for 306 307 California. Streamflow in the American River is mainly (2/3) supplied from wintertime rainfall 308 and snowmelt runoff, with a small portion (1/3) from spring to early summer snowmelt runoff (Dettinger et al. 2014). On the other hand, the Russian, and Eel River Basins are coastal basins 309 where snowmelt runoff is much less important (Scheuerer et al. 2017). To be consistent with the 310

CNRFC operations, we use the nearest neighbor interpolation (Brown et al. 2014a; Seo et al.
2015; Ghazvinian et al. 2020) to pair forecasts-observations.

For generating PQPFs and evaluating the performances of ANN-CSGD relative to the two 313 benchmark models, we adopt an 8-fold cross validation approach. In this approach, for a given 314 basin, we divide the data to 8 consecutive 4-year length folds. Predictions for each fold are 315 produced using each postprocessing mechanism trained with the data of remaining 7 folds (28 316 317 years). Postprocessed out of sample forecasts from all models are verified against observations in 318 individual months of the year in verification years and separately for each sub-basin and forecast lead time. This leads to 32 years of verified forecasts for each sub-basin and lead time. While the 319 320 ANN-CSGD uses the entire available training data (i.e., covering all lead times and seasons) for training and hyperparameter tuning, the benchmark models are trained using subsamples 321 322 representing each forecast lead time and a month/season of the year. To gain insights on how 323 increasing the length of training record and using different seasonal windows for training can affect the predictions of benchmark models, we train each model with different training window 324 sizes and regulations. A summary of training schemes for ANN-CSGD and benchmark models is 325 provided as follows: 326

Unified approach (*ANN-CSGD*) uses forecast-observation pairs of all months and lead times of training years for training and hyperparameter tuning, resulting in a training sample size of up to 7 lead times × 28 years × 365 days = 71540, 20% of which is dedicated for hyperparameter tuning and not used in training.

MMGD and CSGD with 61 days and 91 days training windows (*MMGD-61, CSGD-61*)
 and (*MMGD-91, CSGD-91*) use 61 and 91 training days around the 15<sup>th</sup> of each month
 across training years for generating PQPF for out of sample data of that month, yielding

training sample size up to 28 years × 61 days = 1708 and 28 years × 91 days = 2548 for each lead time and month, respectively. 61 days and 91 days training windows have been used in several past studies (e.g., Hamill et al. 2015; Scheuerer and Hamill 2015; Scheuerer et al. 2017; 2018; Wu et al. 2018).

- MMGD seasonal training scheme (*MMGD-seasonal*), where forecasts in out of sample data from the cool (October-April) and dry (May-September) seasons are postprocessed by a model trained using the data in each season. Thus, a single model is trained for each season and each lead time.
- CSGD seasonal training scheme (*CSGD-seasonal*) (Scheuerer et al. 2020) where the climatological CSGD parameters ( $\mu_{cl}, \sigma_{cl}$  and  $\delta_{cl}$ ) as well as the climatological mean forecast  $\overline{f_{cl}}$  are derived using a 61-day window around the 15<sup>th</sup> of each month, but the same regression coefficients are used across cool and dry seasons to increase the training sample size.

The latter two training schemes yield a sample size of up to 5942 and 4284 for the cool and dryseasons, respectively.

### 349 **3. Results**

In this section we present verification results using different metrics (see appendix B for mathematical definitions and details). We first use the continuous ranked probability skill score (CRPSS) to assess the overall predictive performance of PQPFs from ANN-CSGD relative to those from the benchmark models with different training scenarios. Subsequently, we analyze ANN-CSGD's performance relative to the benchmark models with a 61-day training window, using Brier skill score (BSS), reliability diagrams, and mean squared error skill score (MSESS).

356 3.1. Overall predictive performance of PQPFs

Fig. 3 compares CRPSS of PQPFs from ANN-CSGD and those from the benchmark models with different training scenarios and for the three river basins. The results are computed using cross validated-forecasts from all months and are aggregated over sub-basins of each river basin with *MMGD-61* as the reference forecast. To assess whether differences in predictive performances shown are statistically significant, we perform one-sided Diebold-Mariano test (Diebold and Mariano 1995) for all possible pairs of model comparisons (see appendix B for details). These results are provided in tables S1–S3 in the supplemental material to this article.

Overall, ANN-CSGD generates the most skillful PQPFs across lead times. In the American 364 River (Fig. 3a), ANN-CSGD outperforms its baseline CSGD with different training scenarios by 365 366 a wide margin. The improvement upon each CSGD scheme is statistically significant at all lead times. Nevertheless, performance differences between ANN-CSGD and each of MMGDs are not 367 statistically significant. In the Russian River Basin (Fig. 3b), ANN-CSGD significantly 368 369 outperforms each of benchmark models in a large number of cases. In the Eel River Basin (Fig. 3c), ANN-CSGD outperforms both MMGDs and CSGDs, though its difference with MMGD-61 370 is not statistically significant. It is apparent that the relative performance of MMGD and CSGD 371 varies by river basin and at different lead times. Except for the American River Basin, where 372 most differences are not statistically significant, the seasonal version of MMGD trails behind 373 374 those calibrated with 61- and 91-day moving windows.

For all three river basins, the performance differences of CSGD-61 and CSGD-91 are not statistically significant across the lead times. Interestingly, unlike MMGD-seasonal, CSGDseasonal tends to considerably improve its performance at longer lead times and for all river basins. The training strategy used in CSGD-seasonal was recently introduced by Scheuerer et al. (2020) in their subseasonal forecast scheme (+ 2 week ahead). This scheme presumes that NWP forecast error characteristics change on a season scale when the forecast has very limited skill. Our result confirms the hypothesis that performance is enhanced through the use of wider seasonal windows. Expanding the seasonal window potentially reduces the risk of overfitting of nonlinear CSGD regression model coefficients at longer lead times when the signal to noise ratio is rather poor.

The results corroborate our postulation that different temporal data pooling methods for training 385 386 statistical postprocessing models exert influences on the accuracy of postprocessed PQPFs. The use of MMGD as an alternative scheme serves to further illustrate the significance of ANN-387 CSGD model. EMOS methods such as CSGD are deemed inflexible in that the response variable 388 389 in these models is assumed to follow a single unimodal parametric distribution (see, e.g., Taillardat et al. 2016; Wu et al. 2019; Baran and Lerch 2018), which potentially limits their 390 performance. As such, why does ANN-CSGD retain its superior performance relative to CSGD 391 392 across lead times and study basins while both use the same predictive distribution? This is most likely due to the fact that ANN-CSGD uses the entire training dataset and encodes nonlinear lead 393 time- and seasonal-error dependencies in forecasts in an adaptable manner. Thus, it can preserve 394 the skill of raw forecast, particularly at longer lead times, where postprocessing via CSGD-395 seasonal offers marginal benefit, or even degrades forecast skill. Another advantage of the 396 397 proposed scheme is that it reduces the risk of overfitting due to the early stopping algorithm 398 implemented as a part of its training.

399 3.2. Brier skill score and reliability

Fig. 4 shows the results of BSS for three thresholds > 0.25, > 30 and 60 mm/24h and for the three river basins. While both ANN-CSGD and CSGD underperform MMGD in predicting events > 0.25 mm (i.e., PoP), ANN-CSGD, interestingly, conspicuously outperforms CSGD

(Figs. 4a-c). As pointed out by Ghazvinian et al. (2020), CSGD performs poorly in predicting the 403 PoP due to its reliance on the climatological shift parameter (see also sec. 2.2.1 for further 404 details). When the forecast is very skillful, the predictive CSGD departs from climatology, so 405 does the optimal shift parameter. At longer forecast lead times, the forecast skill declines and the 406 predictive CSGD tends to approach the unconditional climatological one. This feature is 407 reflected in the improvement in CSGD's performance across the lead times. ANN-CSGD, on the 408 409 other hand, directly estimates the shift parameter of the predictive CSGD as an arbitrary function 410 of predictors, thus eliminating the need for a climatological shift parameter. This results in large and statistically significant improvements relative to the CSGD in predicting the PoP. As for the 411 412 outperformance of MMGD relative to the ANN-CSGD, we hypothesize that the flexible two-part structure of MMGD is likely a major contributor. A detailed discussion on this matter can be 413 414 found in Ghazvinian et al. (2020).

415 At the middle threshold of 30 mm/day, ANN-CSGD outperforms both schemes in the American River Basin (Fig. 4d). In the Russian River Basin and the Eel River Basin (Figs. 4e 416 and f), the relative performance of ANN-CSGD and CSGD is mixed but both manage to 417 outperform MMGD, except at Day 7 in the Russian River basin where CSGD slightly 418 underperforms, though it is not statistically significant (not shown here). At the highest 419 threshold, namely 60 mm/day (Figs 4g-i), ANN-CSGD outperforms all other schemes. CSGD 420 mostly outperforms MMGD in the Russian River Basin (Fig. 4h) but underperforms the latter in 421 American River and Eel River basins (Fig. 4g, i). 422

To compare the calibration of PQPFs produced through each scheme, we plot reliability diagrams for the same events and evaluate the contribution of reliability and resolution to the Brier score (Figs. 5,6, and 7). To attain a large enough sample size to better study larger thresholds, we lump cross-validate forecasts at all lead times, and divide forecast probabilities
[0,1] into 15 evenly distributed probability categories to discern the differential performance of
schemes under higher probability categories. The major findings for each river basin are
summarized as follows:

American River Basin: In predicting positive precipitation events (> 0.25 mm/day) 430 431 (Figs. 5a-c), ANN-CSGD's outperformance relative to CSGD is attributed to improvements in both reliability (lower REL) and resolution (higher RES). ANN-432 CSGD mitigates to a great extent the underforecast issue of CSGD. ANN-CSGD 433 generates PQPFs that are more reliable than MMGD but are characterized with lower 434 resolution, yielding an overall inferior predictive performance. At higher thresholds 435 (Figs. 5d-i), ANN-CSGD clearly outperforms both CSGD and MMGD in terms of both 436 437 reliability and resolution. As shown in the histograms embedded in each subplot, ANN-CSGD generates PQPFs that are able to issue high probabilities in predicting mid-to-438 heavy precipitation with higher frequencies, and this points to improved sharpness 439 (Figs. 5f and i). 440

441 Russian River Basin: Similar to the American River Basin, at the lowest threshold (Figs. 6a-c), ANN-CSGD produces forecasts with higher reliability (lower REL) than 442 MMGD but with lower resolution and overall lower predictive skill (higher BS). In >443 30 mm/day ANN-CSGD performs better than CSGD in terms of both reliability and 444 resolution (Figs. 6e, f). At the highest threshold (Figs. 6h, i), the lack of reliability in 445 ANN-CSGD PQPFs relative to those from CSGD is compensated by the higher 446 resolution, and this leads to a superior predictive performance of the former as 447 448 evidenced by the lower BS. MMGD at both thresholds (Figs. 6d, g) produces less reliable PQPFs with lowest sharpness. At the 30 mm/day threshold (Fig. 6d), MMGD
PQPFs' resolution is somewhat higher but is compensated by lower reliability.

Eel River Basin: At the lowest threshold (> 0.25mm/day) (Figs. 7a-c), the relative performance of schemes is quite similar to that for the other two river basins, with ANN-CSGD outperforming MMGD in terms of reliability but not resolution. At higher thresholds (Figs. 7d-i), PQPFs from ANN-CSGD are more reliable and sharper and, overall, more skillful (lowest BS). Though at the highest threshold (i.e., > 60 mm/day), the former exhibit slightly lower resolution than those from MMGD, but this is compensated by superior reliability.

# 458 3.3. Evaluation of deterministic forecasts

459 Finally, we compute mean squared error skill score (MSESS) to evaluate the performance of 460 the distribution mean of PQPF produced using each scheme relative to the GEFS ensemble mean forecast (Fig. 8). These results are accompanied by the results of the Diebold-Mariano test based 461 on the squared error of mean PQPFs (see Tables S4–S6 in the supplemental material). The 462 relative performance varies among the river basins. For the American River Basin (Fig. 8a), all 463 postprocessed PQPFs outperform the GEFS ensemble mean in terms of MSESS. ANN-CSGD 464 PQPFs perform favorably against MMGD PQPFs for all three river basins (the performance 465 differences are not statistically significant). For both the Russian and Eel River Basins (Figs. 8b 466 and c), MSESS values are generally lower relative to those for the American River Basin. This, 467 as we posit, is attributable to location-dependent biases in the GEFS ensemble mean forecast. 468 For example, GEFS is more skillful in the Russian and Eel River Basins according to the MSESS 469 results relative to climatological forecasts (the results are shown in Fig. S1 of supplemental 470 471 materials). For the Russian River Basin (Fig. 8b), underperformance of postprocessed PQPF 472 relative to the GEFS ensemble mean is seen; however, the performance differences are not 473 statistically significant. Unlike the benchmarks, mean PQPF from ANN-CSGD for Russian 474 River Basin significantly outperforms GEFS ensemble mean forecast in all lead times. For both 475 the Russian and Eel River Basins (Figs. 8b, c), ANN-CSGD tends to outperform the other two 476 schemes, though the performance differentials are not statistically significant when comparing 477 with MMGD.

478

#### 479 **4. Discussion and conclusions**

We propose a unified, univariate, hybrid neural network-parametric PQPF postprocessing 480 scheme capable of producing postprocessed forecasts for lead times at least up to 7 days 481 (medium-range). This scheme retains the use of parametric predictive distribution, but employs 482 483 ANN to estimate distribution parameters from forecast-observation pairs. The predictors explored in this study include ensemble mean forecast, forecast lead time, and month of the year, 484 whereas the predictands are three parameters of the predictive censored, shifted gamma 485 distribution (CSGD). The ANN-CSGD model parameters were obtained by minimizing a loss 486 function that is the closed-form expression of CRPS for CSGD (Scheuerer and Hamill 2015), 487 488 with the Adam stochastic gradient descent algorithm (Kingma and Ba 2014) as the optimization approach. To test the performance of our model, we conducted cross-validation experiments to 489 generate medium-range (lead times 1-7 days) daily accumulated PQPFs over selected river 490 491 basins in the service area of the CNRFC. We used two benchmarking processing schemes in this study, namely the CSGD EMOS (Scheuerer and Hamill 2015) with a single-predictor 492 formulation and the NWS operational postprocessor mixed-type Meta-Gaussian distribution 493 (MMGD). These benchmark models were calibrated based on different seasonal data pooling 494

scenarios to investigate the possible impacts of training window size and strategies on theperformance of postprocessed PQPFs.

Verification results showed that ANN-CSGD, in general, outperform the baseline CSGD and 497 MMGD in terms of overall calibration, and significantly so in some cases. Interestingly, ANN-498 CSGD mainly impacts (improves) BSS of PQPF from CSGD at the lowest threshold, which has 499 disproportionate impacts on CRPSS. ANN-CSGD manages to address the CSGD's poor 500 501 performance in predicting PoP as noted in Ghazvinian et al. (2020). While the ANN-CSGD 502 performance comparison results are mixed in predicting 30 mm/day thresholds, it outperforms both benchmark models in predicting large-extreme events (> 60mm/day). On average, the 503 504 proposed method generates high probability forecasts for heavy precipitation more frequently 505 than benchmarks as assessed by sharpness histograms (higher sharpness). This is particularly useful to CNRFC's operational precipitation and flood forecasting practice and, thus, could 506 507 benefit real-time reservoir operations (e.g., determining reservoir release schedules) in California. In its current practice, CNRFC relies on HEFS to produce ensemble PQPFs from 508 509 NWP precipitation forecasts and then generates ensemble streamflow forecasts, which are used to guide real-time flood management and control practices. The MMGD model, embedded in 510 HEFS, has shown to systematically underestimate heavy precipitation amounts, leading to 511 negative biases in subsequent flood forecasts (Demargne et al. 2014; Brown et al. 2014b). The 512 superior performance of the proposed ANN-CSGD on heavy precipitation estimation makes it a 513 viable tool to address limitations in the forecast skills for extreme precipitation and floods. These 514 improvements in forecasts will, in turn, serve to aid real-time reservoir operations and flood risk 515 management. 516

517 In contrast to the CSGD version of Scheuerer and Hamill (2015), the proposed method directly estimates predictive CSGD's shift parameter given each set of predictors. In doing so, it 518 circumvents the need of invoking climatology, and thereby alleviates the bias issue in estimating 519 the PoP in the existing CSGD scheme. Furthermore, the use of ANN allows for representations 520 of complex interactions between three predictive CSGD parameters. Together, these new 521 features help the scheme produce sharper (narrower) predictive distributions than the benchmark 522 523 CSGD. Moreover, ANN-CSGD is able to use much larger training data with extra high forecast-524 observation values, and efficiently translate this to predictive skill at the highest threshold.

The new scheme also has a distinct practical advantage in that it eliminates the need for more 525 526 computationally expensive and operationally labor-intensive approach used in most 527 contemporary statistical postprocessing schemes. Whereas the benchmark models need to be retrained for every forecasting lead time and month/season, ANN-CSGD does not, and it can 528 529 simultaneously utilize forecast-observation pairs across all lead times, months, and seasons. Our results support our hypothesis that the fixed size seasonal window training schemes for current 530 postprocessing methods may not be sufficient for generating consistently skillful POPFs across 531 all lead times. In other words, the performance of existing schemes may be improved by 532 identifying an optimal seasonal training window specific for each lead time, depending on the 533 534 study area and the statistical model at hand. For example, it was shown that a seasonal CSGD tended to improve the performance benchmark 61-day and 91-day CSGDs at longer lead times 535 but not in shorter lead times. ANN-CSGD, on the other hand, automatically adapts to the 536 changes in raw forecasts-observations errors along with all lead times and seasons, and hence, is 537 capable of producing PQPFs with consistently higher skills. 538

A major limitation of nonhomogeneous regression or GAMLSS techniques is that their performance is dependent on the robustness of user-prescribed regression relationships. Moreover, they are typically limited in digesting ordinal temporal covariates such as those used in the ANN-CSGD model. The proposed model, by contrast, can freely learn to characterize arbitrary nonlinear predictor-distribution parameters relationships and among-predictors interactions efficiently.

A well-known challenge in training ANN models is model configuration (hyperparameter 545 546 tuning) to achieve the best validation score. Generally, it is very difficult to find the best possible ANN configuration in a very large parameter space. As pointed out by Scher (2018), there is a 547 548 trade-off between robustness, which depends on the depth and thoroughness of grid search, and computational expenses. For example, our initial assessment showed that maintaining the 549 architecture but expanding the number of layers does not significantly improve the model 550 551 performance. Other regularization techniques such as L1 could be used in combination with early stopping to further reduce generalization errors. However, these techniques could require deeper 552 search for hyperparameters and, therefore, increase computational complexity. We also 553 experimented with training embedding layers with different sizes {2,3,4,5,6,7} to project discrete 554 lead times onto a larger vector of inputs but only found very marginal improvements in the 555 556 validation score. Therefore, we decided not to include embedding layers in our final model.

In future work, we aim to extend the current approach to create a spatially adaptable scheme for postprocessing medium-range ensemble precipitation forecasts on a gridded basis. We expect to achieve this by incorporating geographical information into the network as shown by Scheuerer et al. (2020) in their subseasonal forecasting approach. For example, entire ensemble members or their statistics at a grid point, in addition to those from a specific radius of surrounding grid points, can be direct inputs to the model as the predictors. Such a model potentially eliminates
the need for generating a local superensemble to address the issue of displacement errors in
gridded precipitation forecasts.

Additionally, the current study focuses on 24-hour accumulated precipitation. In operations, 565 CNRFC produces 6-hourly PQPFs and updates their forecasts every 6 hours during major storm 566 events. To align with CNRFC operations, we also plan to explore the performance of the 567 568 proposed ANN-CSGD in generating 6-hourly PQPFs in our future work. Finally, stacked 569 convolution or Long Short-Term Memory (LSTM) layers applied on top of embedding vectors, appear to be very effective in object detection (Krizhevsky et al. 2012), in computer vision, and 570 571 in Natural Language Processing (Collobert et al. 2011), including Machine Translation and Question Answering (Devlin et al. 2018). We envision investigating similar techniques to 572 possibly improve the skill of postprocessed forecast at longer lead times. 573

574

### 575 Acknowledgements

The authors thank the editor and reviewers for their valuable comments that helped improve the article. The first author was financially supported by the faculty startup fund for Dr. Yu Zhang provided by UT Arlington, NOAA Grant NA18OAR4590370-01, Texas Water Development Board Contract No. 1800012276, and NSF grant 1909367. These supports are duly acknowledged here. The authors would also like to thank Michael Scheuerer at Norwegian Computing Center (NR) whose comments and suggestions led to the development of the scheme, and CNRFC staff for providing the forecast and analysis archive.

583

584

# Appendix A

### Implementation details

We implemented our ANN codes in python (Python Software Foundation 2018) using Google's deep learning platform, Tensorflow (Abadi et al. 2016) and Keras API (Chollet et al. 2015). For fitting CSGD climatological and predictive distributions, R (R Core Team 2018) scripts provided by Dr. Michael Scheuerer were used. To calibrate NWS postprocessor, mixedtype meta-Gaussian distribution (MMGD), a research version, very similar to the operational one was implemented in R.

591

### Appendix B

### 592

# Verification metrics used in this study

## 593 A. Mean squared error skill score (MSESS)

The mean squared error skill score (MSESS; Jolliffe and Stephenson 2003) measures the reduction in mean squared error (MSE) of deterministic forecast (mean PQPF/ensemble mean) and verifying observations relative to the reference forecast.

$$MSESS = 1 - \frac{1}{n} \sum_{i=1}^{n} (\bar{x}_i - y_i)^2 / \frac{1}{n} \sum_{i=1}^{n} (\bar{x}_i^{ref} - y_i)^2$$
(A1)

597 Positive values of MSESS indicates improvement in skill of deterministic forecast relative to the598 reference forecast.

### 599 B. Brier skill score (BSS)

The Brier score (BS; Brier 1950) is equivalent to mean squared error of probabilistic forecast
exceeding a given threshold over *n* pairs of forecast and observations

$$BS(\tau) = \frac{1}{n} \sum_{i=1}^{n} [F_i(\tau) - I\{y_i \ge \tau\}]^2$$
(A2)

where  $F_i(\tau)$  is the probability of probabilistic forecast exceeding the threshold value  $\tau$ , and **I**(.) is the indicator (step) function that takes the value 1 if the *i*th verifying observation exceeds the threshold value and 0 otherwise. BS is negatively oriented and ranges from zero to one. Toassess the improvement in BS relative a reference forecast, we compute Brier skill score

$$BSS = 1 - BS/BS_{ref} \tag{A3}$$

Positive values of BSS indicate improvement of BS over that of reference forecast. Brier score can be decomposed to three terms: *reliability or Type-I conditional bias, resolution,* and *uncertainty* (Murphy 1973; Wilks 2011)

$$BS(\tau) = Reliability(\tau) - Resolution(\tau) + Uncertainity(\tau)$$
(A4)  
$$= \frac{1}{n} \sum_{i=1}^{K} N_i [F_i(\tau) - \bar{o}_i(\tau)]^2 - \frac{1}{n} \sum_{i=1}^{K} N_i [\bar{o}_i(\tau) - \bar{o}(\tau)]^2 + \bar{o}(\tau) [1 - \bar{o}(\tau)]$$

where *K* indicated the number of categories, forecast are aggregated to, *N* is the number of cases in each category,  $\bar{o}_i(\tau)$  is the average climatological probability (ACP) exceeding the threshold  $\tau$ in that category and  $\bar{o}(\tau)$  is the overall ACP. It should be noted that uncertainty term as seen is independent of the forecast source. Probabilistic forecasts with lower/higher reliability/resolution values are desirable.

### 614 *C. Continuous ranked probability score (CRPS)*

The continuous ranked probability score (CRPS; Matheson and Winkler 1976) measures the integral of squared differences between the cumulative distribution function (CDF) of probabilistic forecast and verifying observation. It is a popular metric to assess the overall predictive performance of probabilistic forecasts (sharpness and reliability; see Gneiting et al. 2007 for further details). CRPS averaged over the sample of forecast-observations with size of *n* is given by

$$CRPS = \frac{1}{n} \sum_{i=1}^{n} \int_{-\infty}^{\infty} [F_i(x) - I\{y_i \le x\}]^2 dx$$
(A5)

where  $F_i$  denotes the CDF of PQPF at the ith forecast instance and  $y_i$  is the verifying observation. I (.) is the indicator (step) function which takes the value of 1 if  $x \ge y_i$  and 0 elsewhere. Continuous ranked probability skill score (CRPSS) is routinely used to assess the performance of probabilistic forecast relative to a reference forecast

$$CRPSS = 1 - CRPS/CRPS_{ref} \tag{A6}$$

### 625 D. Reliability diagrams and sharpness histograms

626 The reliability and resolution of a probabilistic forecast for exceeding some specific thresholds  $(\tau)$  can be assessed graphically using reliability diagrams. The reliability diagram consists of a 627 plot of the average values of forecast probabilities exceeding  $\tau$ , against that of observed relative 628 629 frequencies over each defined probability category. In a reliable probabilistic forecast, the reliability diagram should be a close 1:1line. Interested readers are referred to Brocker and Smith 630 (2007) and Wilks (2011) for details on how to interpret the deficiencies in probabilistic forecasts 631 using reliability diagrams. To assess the sharpness of PQPF for specific thresholds, we use 632 sharpness histograms to investigate the frequency of forecast probabilities for different 633 probability bins. Note, a sharp forecast is characterized by higher frequencies for the forecast 634 635 probabilities close to either 0 or 1.

### 636 E. The Diebold-Mariano test

637 To assess statistical significance of verification score differences between two forecast 638 methods, we use the Diebold-Mariano statistical test of the null hypothesis of equal predictive 639 performance (Diebold and Mariano 1995). Let  $\Delta = S_{F1} - S_{F2}$  denote the vector of verification

score S differences from two competing forecast methods  $F_1$  and  $F_2$  over verification sample 640 with length  $n, \overline{\Delta} = 1/n \sum_{i=1}^{n} \Delta_i$ , and  $\hat{\sigma}_{\Delta}$  a suitable estimator of asymptotic standard deviation of 641  $\Delta$ . Under standard regularity conditions, the test statistic  $t_n = \sqrt{n} \frac{\overline{\Delta}}{\widehat{\sigma}_A}$  asymptotically follows a 642 standard Gaussian distribution under the null hypothesis of no difference in predictive 643 performances of two competing forecast methods. Following the past studies (Baran and Lerch 644 2016, 2018; Rasp and Lerch 2018)  $\hat{\sigma}_{\Delta}$  can be estimated by square root of sample autocovariance 645 up to lag k - 1 for the k step-ahead forecasts to account for temporal dependencies in forecast 646 errors. We use one-sided Diebold-Mariano tests. The alternative hypothesis is that forecast 647 method  $F_2$  underperforms forecast method  $F_1$  and the statistical significancy of the test's statistic 648 can be assessed by obtaining corresponding *p*-value. we perform the tests based on both CRPS 649 and squared error of mean PQPF (on a limited basis) and for each lead time and separately for 650 each river basin. To address spatial dependence of forecast errors, scores are averaged across 651 sub-basins in each river basins (M. Scheuerer 2021, personal communication). Further, we adjust 652 the test results by accounting for test multiplicity (i.e., simultaneously analyzing test results of 653 multiple lead times) using false discovery rate (FDR) method (Benjamini and Hochberg 1995) 654 by controlling the FDR at the level  $\alpha_{FDR} = 0.05$ . Note that, this procedure was discussed by Wilks 655 (2016) in spatial context where test results are interpreted simultaneously across multiple grid 656 points but also was suggested to be applied whenever the results of simultaneous several 657 658 hypothesis tests are reported or interpreted.

659

### 660 **References**

Abadi, M., and Coauthors, 2016: Tensorflow: A system for largescale machine learning. Proc.
 USENIX 12th Symp. On Operating Systems Design and Implementation, Savannah, GA,

663

Advanced

# Systems

Association,

265 - 283,

664 https://www.usenix.org/system/files/conference/osdi16/osdi16-abadi.pdf.

Computing

- Ba, J. L., J. R. Kiros, and G. E. Hinton, 2016: Layer normalization. arXiv preprint
  arXiv:1607.06450, https://arxiv.org/abs/1607.06450.
- Baran, S., and D. Nemoda, 2016: Censored and shifted gamma distribution based EMOS model
  for probabilistic quantitative precipitation forecasting. Environmetrics, 27, 280–292,
  https://doi.org/10.1002/env.2391.
- Baran, S., and S. Lerch, 2016: Mixture EMOS model for calibrating ensemble forecasts of wind
- 671 speed. Environmetrics, 27, 116–130, <u>https://doi.org/10.1002/env.2380</u>.
- Baran, S., and S. Lerch, 2018: Combining predictive distributions for statistical post-processing
  of ensemble forecasts. Int. J. Forecast., 34, 477–496,
  https://doi.org/10.1016/j.ijforecast.2018.01.005.
- Benjamini, Y., and Y. Hochberg, 1995: Controlling the false discovery rate: A practical and
  powerful approach to multiple testing. J. Roy. Stat. Soc. B, 57, 289–300,
  https://doi.org/10.1111/J.2517-6161.1995.TB02031.X.
- Bremnes, J. B., 2020: Ensemble postprocessing using quantile function regression based on
  neural networks and Bernstein polynomials. Mon. Wea. Rev.,148, 403–414,
  https://doi.org/10.1175/MWR-D-19-0227.1.
- Brenowitz, N. D., and C. S. Bretherton, 2018: Prognostic validation of a neural network unified
  physics parameterization. Geophys. Res. Lett., 45, 6289–6298,
  https://doi.org/10.1029/2018GL078510.

31

- Brier, G. W., 1950: Verification of forecasts expressed in terms of probability. Mon. Wea. Rev.,
  78, 1–3, https://doi.org/10.1175/1520-0493(1950)078<0001:VOFEIT>2.0.CO;2.
- Bröcker, J. and L.A. Smith, 2007: Increasing the Reliability of Reliability Diagrams. Wea.
  Forecasting, 22, 651–661, https://doi.org/10.1175/WAF993.1.
- Brown, J. D., D. Seo, and J. Du, 2012: Verification of Precipitation Forecasts from NCEP's
  Short-Range Ensemble Forecast (SREF) System with Reference to Ensemble Streamflow
  Prediction Using Lumped Hydrologic Models. J. Hydrometeor., 13, 808–836,
  https://doi.org/10.1175/JHM-D-11-036.1.
- Brown, J. D., L. Wu, M. He, S. Regonda, H. Lee, and D.J. Seo, 2014a: Verification of
  temperature, precipitation, and streamflow forecasts from the NOAA/NWS Hydrologic
  Ensemble Forecast Service (HEFS): 1. Experimental design and forcing verification. Hydrol,
  519, 2869–2889, https://doi.org/10.1016/j.jhydrol.2014.05.028.
- Brown, J. D., M. He, S. Regonda, L. Wu, H. Lee, and D.J. Seo, 2014b: Verification of
  temperature, precipitation, and streamflow forecasts from the NOAA/NWS Hydrologic
  Ensemble Forecast Service (HEFS): 2. Streamflow verification. Hydrol, 519, 2869–2889,
  https://doi.org/10.1016/j.jhydrol.2014.05.030.
- Cannon AJ., 2012: Neural networks for probabilistic environmental prediction: conditional
  density estimation network creation and evaluation (CaDENCE) in R. Comput.
  Geosci.41:126–35, https://doi.org/10.1016/j.cageo.2011.08.023.
- 703 Chapman, W. E., Subramanian, A. C., Delle Monache, L., Xie, S. P., and F. M. Ralph, 2019:
- 704 Improving atmospheric river forecasts with machine learning. Geophysical Research Letters,
- 705 46, 10,627–10,635. https://doi.org/10.1029/2019GL083662.

- Chollet, F., and Coauthors, 2015: Keras: The Python Deep Learning library. Accessed 2019,
  https://keras.io.
- 708 Clevert, D. A., T. Unterthiner, and S. Hochreiter, 2015: Fast and accurate deep network learning
- by exponential linear units (ELUs). Int. Conf. on Learning Representations, San Juan, Puerto
  Rico, ICLR, 1–14, https://arxiv.org/abs/1511.07289.
- Cloud, K. A., B. J. Reich, C. M. Rozoff, S. Alessandrini, W. E. Lewis, and L. Delle Monache,
  2019: A feed forward neural network based on model output statistics for short-term
  hurricane intensity prediction. Wea. Forecasting, 34, 985–997, https://doi.org/10.1175/WAFD-18-0173.1.
- Collobert, R., J. Weston, L. Bottou, M. Karlen, K. Kavukcuoglu and P. Kuksa, 2011: Natural
  Language Processing (Almost) from Scratch. Journal of Machine Learning Research,
- 717 12:2493-2537. Available online at: https://www.jmlr.org /papers/volume12 /collobert11a
  718 /collobert11a.pdf.
- 719 Daly, C., M. Halbleib, J. I. Smith, W. P. Gibson, M. K. Doggett, G.H.Taylor, J. Curtis, and P. P.
- 720 Pasteris, 2008: Physiographically sensitive mapping of climatological temperature and
- precipitation across the conterminous United States. Int. J. Climatol., 28, 2031–2064,
- 722 https://doi.org/10.1002/joc.1688.
- Demargne, J., and Coauthors, 2014: The science of NOAA's operational Hydrologic Ensemble
  Forecast Service. Bull. Amer. Meteor. Soc., 95, 79–98, https://doi.org/10.1175/BAMS-D-1200081.1.

726	Dettinger, M. D., D. R. Cayan, M. K. Meyer, and A. E. Jeton, 2014: Simulated hydrologic
727	responses to climate variations and change in the Merced, Carson, and American river basins,
728	Sierra Nevada, California, 1900–2099, Clim. Change, 62, 283–317.

- Devlin, J., M. W. Chang, K. Lee, and K. Toutanova, 2018: Bert: Pre-training of deep
  bidirectional transformers for language understanding. arXiv preprint arXiv:1810.04805.
  https://arxiv.org/abs/1810.04805.
- Diebold, F. X., and R. S. Mariano, 1995: Comparing predictive accuracy. J. Bus. Econ. Stat., 13,
  253–263, https://doi.org/10.1080/07350015.1995.10524599.
- Gagne II, D. J., S. E. Haupt, D. W. Nychka, and G. Thompson, 2019: Interpretable Deep
  Learning for Spatial Analysis of Severe Hailstorms. Mon. Wea. Rev., 147, 2827–2845,
  https://doi.org/10.1175/MWR-D-18-0316.1.
- Gentine, P., M. Pritchard, S. Rasp, G. Reinaudi, and G. Yacalis, 2018: Could machine learning
  break the convection parameterization deadlock? Geo-phys. Res. Lett., 45, 5742–5751,
  https://doi.org/10.1029/2018GL078202.
- Ghazvinian, M., Y. Zhang, and D. J. Seo, 2020: A Nonhomogeneous Regression-Based
  Statistical Postprocessing Scheme for Generating Probabilistic Quantitative Precipitation
  Forecast. J. Hydrometeor., 21, 2275–2291, https://doi.org/10.1175/JHM-D-20-0019.1.
- Ghazvinian, M., Seo, D. J., and Y. Zhang, 2019: Improving Medium-range Probabilistic
  Quantitative Precipitation Forecast for Heavy-to-extreme Events through the Conditional
  Bias-penalized Regression. In AGU Fall Meeting 2019. AGU.
  https://agu.confex.com/agu/fm19/meetingapp.cgi/Paper/517742.

34

- Gneiting, T., A.E. Raftery, A.H. Westveld, and T. Goldman, 2005: Calibrated Probabilistic
  Forecasting Using Ensemble Model Output Statistics and Minimum CRPS Estimation. Mon.
  Wea. Rev., 133, 1098–1118, https://doi.org/10.1175/MWR2904.1.
- Gneiting, T., F. Balabdaoui, and A. E. Raftery, 2007: Probabilistic forecasts, calibration and
  sharpness. J. Roy. Stat. Soc., 69B, 243–268, https://doi.org/10.1111/j.14679868.2007.00587.x.
- 753 Goodfellow, I., Y. Bengio, and A. Courville, 2016: Deep Learning. MIT Press, 775 pp.
- Hamill, T.M., G.T. Bates, J.S. Whitaker, D.R. Murray, M. Fiorino, T.J. Galarneau, Y. Zhu, and
- 755 W. Lapenta, 2013: NOAA's Second-Generation Global Medium-Range Ensemble Reforecast
- Dataset. Bull. Amer. Meteor. Soc., 94, 1553–1565, https://doi.org/10.1175/BAMS-D-1200014.1.
- Hamill, T. M., M. Scheuerer, and G. T. Bates, 2015: Analog probabilistic precipitation forecasts
- vsing GEFS reforecasts and climatology-calibrated precipitation analyses. Mon. Wea. Rev.,
- 760 143, 3300–3309, https://doi.org/10.1175/MWR-D-15-0004.1.
- 761 Hamill, T. M., and J. S. Whitaker, 2006: Probabilistic quantitative precipitation forecasts based
- on reforecast analogs: Theory and application. Mon. Wea. Rev., 134, 3209–3229,
- 763 https://doi.org/10.1175/MWR3237.1.
- He, M., and Coauthors, 2016: Verification of ensemble water supply forecasts for Sierra Nevada
  watersheds. Hydrology, 3, 35, https://doi.org/10.3390/hydrology3040035.
- 766 Herr, H. D., and R. Krzysztofowicz, 2005: Generic probability distribution of rainfall in space:
- 767 The bivariate model. J. Hydrol., 306, 234–263, https://doi.org/10.1016/j.jhydrol.2004.09.011.

768	Kelly, K.	S.,	and R	Krzysztofow	icz, 1997	': A	bivariate	meta-Gaussian	density	for	use	in
769	hydrol	ogy.	Stocha	stic Hydrol. Hy	draul., 1	l, 17	–31, https	://doi.org/10.100	07/BF024	284	23.	

- Kim, S., and Coauthors, 2018: Assessing the Skill of Medium-Range Ensemble Precipitation and
- Streamflow Forecasts from the Hydrologic Ensemble Forecast Service (HEFS) for the Upper
  Trinity River Basin in North Texas. J. Hydrometeor., 19, 1467–1483,
  https://doi.org/10.1175/JHM-D-18-0027.1.
- Kingma, D. P., and J. Ba, 2014: Adam: A method for stochastic optimization. Third Int. Conf.
  for Learning Representations, San Diego, CA, ICLR, 1–15, https://arxiv.org/abs/1412.6980.
- Krizhevsky, A., I. Sutskever, and G. E. Hinton, 2012: Imagenet classification with deep
  convolutional neural networks. In Advances in neural information processing systems (pp.
  1097-1105). Available online at: http://www.csri.utoronto.ca/~hinton/absps/imagenet.pdf.
- Lagerquist, R., A. McGovern, and D. J. Gagne II, 2019: Deep learning for spatially explicit
  prediction of synoptic-scale fronts. Wea. Forecasting, 34, 1137–1160,
  https://doi.org/10.1175/WAF-D-18-0183.1.
- Li, W., Q. Duan, C. Miao, A. Ye, W. Gong, and Z. Di, 2017: A review on statistical
  postprocessing methods for hydrometeorological ensemble forecasting. WIREs Water, 4:
  e1246, https://doi.org/10.1002/wat2.1246.
- Liu, Y., E. Racah, J. Correa, A. Khosrowshahi, D. Lavers, K. Kunkel, M. Wehner, and W.
  Collins, 2016: Application of deep convolutional neural networks for detecting extreme
  weather in climate datasets.arXiv.org, https://arxiv.org/abs/1605.01156.

- Liu, Y., P. Di, S. Chen, and J. DaMassa, 2018: Relationships of Rainy Season Precipitation and
- 789 Temperature to Climate Indices in California: Long-Term Variability and Extreme Events. J.
- 790 Climate, 31, 1921–1942, https://doi.org/10.1175/JCLI-D-17-0376.1.
- Matheson, J. E., and R. L. Winkler, 1976: Scoring rules for continuous probability distributions.
  Manage. Sci., 22, 1087–1096, https://doi.org/10.1287/mnsc.22.10.1087.
- Murphy, A. H., 1973: A New Vector Partition of the Probability Score. J. Appl. Meteor., 12,
  595–600, https://doi.org/10.1175/1520-0450(1973)012<0595:ANVPOT>2.0.CO;2.
- Python Software Foundation, 2018: Python Language Reference, version 3.7. Available at
  http://www.python.org.
- R Core Team, 2017: R: A language and environment for statistical computing. R Foundation for
  Statistical Computing, Vienna, Austria, https://www.R-project.org/.
- Rasp, S. and S. Lerch, 2018: Neural Networks for Postprocessing Ensemble Weather Forecasts.
  Mon. Wea. Rev., 146, 3885–3900, https://doi.org/10.1175/MWR-D-18-0187.1.
- Rasp, S., M. S. Pritchard, and P. Gentine, 2018: Deep learning to represent subgrid processes in
  climate models. Proc. Natl. Acad. Sci. USA,115, 9684–9689,
  https://doi.org/10.1073/pnas.1810286115.
- Rigby, R. A. and D. M. Stasinopoulos, 2005: Generalized additive models for location, scale and
  shape. Journal of the Royal Statistical Society: Series C (Applied Statistics), 54: 507-554,
  https://doi.org/10.1111/j.1467-9876.2005.00510.x.
- Scher, S., 2018: Toward data-driven weather and climate fore-casting: Approximating a simple
  general circulation model with deep learning. Geophys. Res. Lett.,45, 12 616–12 622,
  https://doi.org/10.1029/2018GL080704.

- Scheuerer, M., and T. M. Hamill, 2015: Statistical postprocessing of ensemble precipitation
  forecasts by fitting censored, shifted gamma distributions. Mon. Wea. Rev., 143, 4578–4596,
  https://doi.org/10.1175/MWR-D-15-0061.1.
- Scheuerer, M., T. M. Hamill, B. Whitin, M. He, and A. Henkel, 2017: A method for preferential
  selection of dates in the Schaake shuffle approach to constructing spatiotemporal forecast
  fields of temperature and precipitation. Water Resour. Res., 53, 3029–3046,
  https://doi.org/10.1002/2016WR020133.
- 817 Scheuerer, M. and T.M. Hamill, 2018: Generating Calibrated Ensembles of Physically Realistic,
- High-Resolution Precipitation Forecast Fields Based on GEFS Model Output. J.
  Hydrometeor., 19, 1651–1670, https://doi.org/10.1175/JHM-D-18-0067.1.
- Scheuerer, M., M. B. Switanek, R. P. Worsnop, and T. M. Hamill, 2020: Using Artificial Neural
  Networks for Generating Probabilistic Subseasonal Precipitation Forecasts over California.
  Mon. Wea. Rev., 148, 3489–3506, https://doi.org/10.1175/MWR-D-20-0096.1.
- 823 Seo, D.-J., and Coauthors, 2015: On improving ensemble forecasting of extreme precipitation
- using the NWS Meteorological Ensemble Forecast Processor (MEFP). 2015 Fall Meeting,
  San Francisco, CA, Amer. Geophys. Union, Abstract H51P-08,
  https://agu.confex.com/agu/fm15/meetingapp.cgi/Paper/81958.
- Stauffer, R., N. Umlauf, J.W. Messner, G.J. Mayr, and A. Zeileis, 2017: Ensemble
  Postprocessing of Daily Precipitation Sums Over Complex Terrain Using Censored HighResolution Standardized Anomalies. Mon. Wea. Rev., 145, 955–969,
  https://doi.org/10.1175/MWR-D-16-0260.1.

- Taillardat, M., O. Mestre, M. Zamo, and P. Naveau, 2016: Calibrated ensemble forecasts using
  quantile regression forests and ensemble model output statistics. Mon. Wea. Rev., 144,
  2375–2393, https://doi.org/10.1175/MWR-D-15-0260.1.
- Taillardat, M., A. Fougères, P. Naveau, and O. Mestre, 2019: Forest-Based and Semiparametric
  Methods for the Postprocessing of Rainfall Ensemble Forecasting. Wea. Forecasting, 34,
  617–634, https://doi.org/10.1175/WAF-D-18-0149.1.
- Vannitsem, S., and Coauthors, 2020: Statistical Postprocessing for Weather Forecasts -- Review,
  Challenges and Avenues in a Big Data World. arXiv preprint arXiv:2004.06582,
  https://arxiv.org/abs/2004.06582.
- Wilks, D. S., 2011: Statistical Methods in the Atmospheric Sciences.3rd ed. International
  Geophysics Series, Vol. 100, Elsevier Academic Press, 704 pp.
- Wilks, D. S., 2016: "The stippling shows statistically significant grid points": How research
  results are routinely overstated and over-interpreted, and what to do about it. Bull. Amer.
  Meteor. Soc.,97, 2263–2273, https://doi.org/10.1175/BAMS-D-15-00267.1.
- Wilks, D. S., 2018: Univariate ensemble postprocessing. In S. Vannitsem, D. S. Wilks, & J.
  Messner (Eds.), Statistical postprocessing of ensemble forecasts (pp.49–89).
  https://doi.org/10.1016/B978-0-12-812372-0.00003-0.
- Wu, L., D.J. Seo, J. Demargne, J. Brown, S. Cong, and J. Schaake, 2011: Generation of
  ensemble precipitation forecast from single-valued quantitative precipitation forecast for
  hydrologic ensemble prediction. J. Hydrol., 399, 281–298,
  https://doi.org/10.1016/j.jhydrol.2011.01.013.

852	Wu, L., Y. Zhang, T. Adams, H. Lee, Y. Liu, and J. Schaake, 2018: Comparative Evaluation of
853	Three Schaake Shuffle Schemes in Postprocessing GEFS Precipitation Ensemble Forecasts.
854	J. Hydrometeor., 19, 575–598, https://doi.org/10.1175/JHM-D-17-0054.1.

- Wu, Y., X. Yang, X. Zhang, and Q. Kuang, 2019: Mixture probabilistic model for precipitation
  ensemble forecasting. Q J R Meteorol Soc. 2019; 145: 3516– 3534.
  https://doi.org/10.1002/qj.3637.
- Xu, J., X. Sun, Z. Zhang, G. Zhao, and J. Lin, 2019: Understanding and improving
  layernormalization. *arXiv preprint arXiv:1911.07013*, https://arxiv.org/abs/1911.07013.
- 860 Zhang, Y., L. Wu, M. Scheuerer, J. Schaake, and C. Kongoli, 2017: Comparison of Probabilistic
- 861 Quantitative Precipitation Forecasts from Two Postprocessing Mechanisms. J. Hydrometeor.,
- 862 18, 2873–2891, https://doi.org/10.1175/JHM-D-16-0293.1\_



**Fig. 1.** Schematic of the ANN-CSGD structure. We illustrate hidden layer with 5 nodes for the sake of demonstration. Three parameters of predictive CSGDs are considered as additional functions of ANN outputs to constrain the values of these parameters to reasonable ranges.



Fig. 2. Location map of the study basins as well as basins in the service area of CNRFC within the State of California.



**Fig. 3.** CRPSS for ANN-CSGD and benchmark postprocessing models with different training scenarios (61-day, 91-day, and seasonal window). Displayed are cross-validated CRPSS computed by pooling CRPS values across study sub-basins in each river basins and all months as a function of lead time. MMGD PQPFs with 61-day training window serve as the reference.



**Fig. 4.** Brier skill score (BSS) results for PQPFs from ANN-CSGD and CSGD and for three different thresholds: > 0.25, 30 and 60 mm, averaged over study sub-basins in each river basin and shown as a function of lead time, with MMGD-61 as the reference.



Reliability diagrams, American River Basin, Lead time: 1–7 days

**Fig. 5**. Reliability diagrams for the three thresholds (> 0.25, 30 and 60 mm) and for sub-basins in the American River Basin were computed based on observations and cross-validated postprocessed forecasts pooled across study sub-basins and all forecast lead times. Brier score (BS), Reliability (REL) and Resolution (RES) values are shown in each panel. The insert histograms show the frequencies for each of 15 forecast probability bins in log10 scale for better visibility and the bars show 90% bootstrap confidence intervals of observed frequencies for estimated forecast probabilities. Benchmark models are trained using 61-day window centered around the 15<sup>th</sup> of each month.



Reliability diagrams, Russian River Basin, Lead time: 1-7 days

Fig. 6. As in Fig. 5 except for the Russian River Basin.



Reliability diagrams, Eel River Basin, Lead time: 1-7 days

Fig. 7. As in Fig. 5 except for the Eel River Basin.



**Fig. 8.** As in Fig. 3 except for MSESS with benchmark models trained using 61-day window. GEFS ensemble mean forecast is considered as the reference.

	sub-basin ID	sub-basin name
American River Basin		
	NFDC1HUF	North Fork American River-North Fork Dam (upper)
	NFDC1HLF	North Fork American River-North Fork Dam (lower)
	FOLC1LOF	American River-Folsom Lake
Russian River Basin		
	WSDC1HOF	Dry Creek - Lake Sonoma
	GUEC1LOF	Russian River - Guerneville
Eel River Basin		
	DOSC1HUF	Middle Fork Eel River-Dos Rios (upper)
	DOSC1HLF	Middle Fork Eel River-Dos Rios (lower)
	FTSC1LUF	Eel River-Fort Seward (upper)
	FTSC1LLF	Eel River-Fort Seward (lower)

Table 1. Names and NWS IDs of study sub-basins of each river basin.