

1 **Climate-driven variability in *Euphausia pacifica* size distributions off northern California**

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7

8 **Abstract**

9

10 We analyzed an 11-year data set of length measurements for *Euphausia pacifica*, resolved by
11 developmental stages ranging from early furcilia through adults, to characterize temporal variability in
12 size distributions in relation to environmental and climate forcing in coastal waters of the northern
13 California Current. Our analysis reveals that coastal size distributions of this linchpin species are related
14 to temperature and chlorophyll *a* concentration, and indicate that warm climate events disrupt and
15 suppress typical seasonal dynamics, resulting in persistent shifts towards populations dominated by
16 smaller juveniles and adults. The 2014-16 marine heatwave (MHW) resulted in sharp and sustained
17 transition in the size structure of *E. pacifica* off northern California; we captured numerous mature
18 individuals smaller than any previously reported in the literature and larger size classes were
19 conspicuously absent or rare when the MHW most strongly impacted coastal waters. Early life history
20 stages exhibited a contrasting response, shifting instead towards larger size distributions during warm
21 conditions. Advection appears to be the dominant mechanism driving rapid shifts in *E. pacifica* size
22 structure off northern California, which implies alongshore and cross-shelf gradients in life history
23 expression (e.g., changes in size-at-maturation) related primarily to temperature. Variability in food
24 availability (chlorophyll concentration) and physiological responses (e.g., changes in growth rates) likely
25 contribute to seasonal variability, but appear insufficient to account for sharp, climate-driven
26 transitions. Climate-driven shifts in individual and total biomass of this critical and dominant forage
27 species have important ecological implications, particularly for growth and survival of species at higher
28 trophic levels. Given that warm water events are likely to increase with climate change, the effects of
29 the 2014-16 MHW on *E. pacifica* in coastal waters may serve as a harbinger for future change. Detailed
30 information on size structure of euphausiid populations is fundamental to understanding population
31 responses to and ecosystem consequences of climatic and environmental forcing, and has been
32 incorporated in assessments that support ecosystem-based approaches to fisheries management.

33

34 **1. Introduction**

35

36 Within the California Current Ecosystem (CCE) euphausiids, commonly known as krill, play a crucial role
37 linking primary production to higher trophic levels, including numerous ecological and economically
38 important organisms such as salmonids, rockfishes, seabirds, and marine mammals (Tanasichuk, 2002;
39 Brodeur and Pearcy, 1992; Abraham and Sydeman, 2004; Becker et al., 2007; Schoenherr, 1991). Among
40 those species that are not directly dependent on euphausiids, most are within two trophic links of them
41 at some point in their life history (Field. et al., 2006). Changes in euphausiid populations can alter
42 trophic pathways and ecosystem functioning, often with substantive consequences for higher trophic
43 levels (Sydeman et al., 2006; Lindley et al., 2009; Ruzicka et al., 2012; Wells et al., 2012; Jones et al.,
44 2018; Santora et al., 2020). Recognition of their importance to marine ecosystems and fisheries
45 productivity motivated a pre-emptive moratorium on euphausiid harvest in the California Current and
46 has spurred investigations focused on spatial patterns in relation to oceanographic and bathymetric
47 structure (Gómez-Gutiérrez et al., 2005; Ressler et al., 2005; Santora et al., 2011; Santora et al., 2012;
48 Fiechter et al., 2020) and temporal variability in response to climate forcing (Dorman et al., 2015a;
49 Sydeman et al., 2013; Brinton and Townsend, 2003).

50

51 Several studies have documented the response of euphausiid populations to environmental and climatic
52 variability in various regions of the CCE. Variability in the onset and intensity of upwelling affects the
53 abundance of *Euphausia pacifica* and *Thysanoessa spinifera*, species that usually dominate under typical
54 (relatively cool and productive) upwelling conditions (Tanasichuk, 1998; Gómez-Gutiérrez et al., 2005).
55 Delayed onset of upwelling can affect higher trophic levels by negatively impacting euphausiid biomass
56 (Sydeman et al., 2006, Dorman et al., 2011). El Niño events have particularly strong effects on
57 euphausiid populations and communities throughout the CCE and regions to the north and south
58 (Mackas and Galbraith, 2002; González et al., 2000; Brinton and Townsend, 2003). During these events,
59 euphausiid biomass tends to decline, especially for cool water species, and warm-water species are
60 encountered more frequently and with greater abundance (Smith, 1985; Marinovic et al., 2002; Lilly and
61 Ohman, 2018; Bjorkstedt and Robertson, in prep.).

62

63 Population level responses to climate variability are likely to be mirrored by changes in individual
64 characteristics related to environmental conditions that affect development and growth. Euphausiid

65 energetics are strongly influenced by temperature (Lasker, 1966; Fowler et al., 1971; Ross, 1981; Ross
66 1982a; Iguchi and Ikeda, 1995). Growth and development rates of *E. pacifica* are positively related to
67 temperature (Ross, 1982a), yet individuals tend to mature at smaller size under warmer conditions (Ross
68 1982b; Forster et al., 2012). How *E. pacifica* respond to variability in prey resources is less clear. In the
69 northern California Current, *E. pacifica* growth does not exhibit a strong relationship with chlorophyll
70 concentration (Shaw et al., 2010), whereas tight coupling between growth and chlorophyll
71 concentration has been observed elsewhere along the coast of North America (Pinchuk and Hopcroft,
72 2007; McLaskey et al., 2020). Although several studies document selectivity for diatoms and varying
73 degrees of omnivory in *E. pacifica* (e.g., Ohman, 1984; Du and Peterson, 2014; Dilling et al., 1998),
74 effects of food quality or diversity on euphausiid growth rates remain poorly. Complicating matters,
75 many euphausiids, including *E. pacifica*, have the capacity to shrink during periods of starvation, bouts of
76 intense reproductive effort, or exposure to increased temperatures (Nicol et al., 1992; Pinchuk and
77 Hopcroft, 2007; Feinberg et al., 2007; Shaw et al., 2010; Marinovic and Mangel, 1999).

78
79 In this paper, we describe an 11-year data set of length measurements for *Euphausia pacifica*, resolved
80 by developmental stage from early furcilia through adults, collected in coastal waters off northern
81 California, and characterize variability in size-structure in relation to seasonal and interannual variability
82 in temperature and chlorophyll *a* concentration. Specifically, we examine the hypothesis that *E. pacifica*
83 tend to be larger (smaller) during cooler (warmer) conditions. This hypothesis is based on well-
84 documented relationships between temperature and euphausiid growth and development dynamics
85 (e.g., Ross, 1981; Ross, 1982a; Feinberg, 2006) and broad observations that indicate temperature has an
86 important role in determining individual body size and population size distributions (e.g., Atkinson,
87 1994; Daufresne et al., 2009). We also examine the hypothesis that greater (lower) chlorophyll *a*
88 concentration is related to larger (smaller) euphausiids. Including chlorophyll concentration in our
89 analysis reflects the common observation that, all else constant, more food typically supports greater
90 growth. We are particularly interested in the biological response to the 2014-16 marine heatwave
91 (MHW; Hobday et al., 2018; Jacox et al., 2018a), but we also assess responses to lesser warming events
92 (e.g., the 2009-10 El Niño) in contrast to periods marked by cooler conditions and more intense
93 upwelling. *E. pacifica* are a dominant species in the California Current (Brinton and Wyllie, 1976; Dorman
94 et al., 2005), therefore understanding its dynamics in the context of environmental variability is
95 expected to yield important insight to the structure and function of the California Current Ecosystem.
96

97 **1.1 Recent Climate Variability in the California Current**

98 During the past decade the CCE experienced substantial climate variability, including record-breaking
99 warm water anomalies. Except for a brief interruption by the moderate 2009-10 El Niño, basin scale
100 indices reflected relatively cool and productive conditions from 2007 through 2013 (Fig. 1). During
101 winter 2013-14 climate indices affecting the CCE underwent a major phase shift; the Pacific Decadal
102 Oscillation (PDO), North Pacific Gyre Oscillation (NPGO), and Multivariate ENSO Index (MEI) changed
103 phase, indicating a transition to a warmer and less productive North Pacific Ocean. Coincident with and
104 following this climatic shift, the CCE experienced a severe and unprecedented MHW that lasted several
105 years (2014-16) and extended throughout the entire NE Pacific (Hobday et al., 2018; Jacox et al., 2018a).
106 In the northern portion of the California Current, the initial phase of the MHW was marked by the arrival
107 of warm 'blob' waters in fall 2014 (Bond et al., 2015; Peterson et al., 2015). Warm water anomalies
108 persisted into 2016 and were influenced by one of the strongest tropical El Niño events on record during
109 2015-16 (Gentemann et al., 2016; McPhaden, 2015; Jacox et al., 2016). By early 2017, conditions in the
110 central and southern CCE had returned to near normal. However, recovery from the MHW and El Niño
111 was slower in the northern CCE, where warm water anomalies and zooplankton taxa persisted into mid-

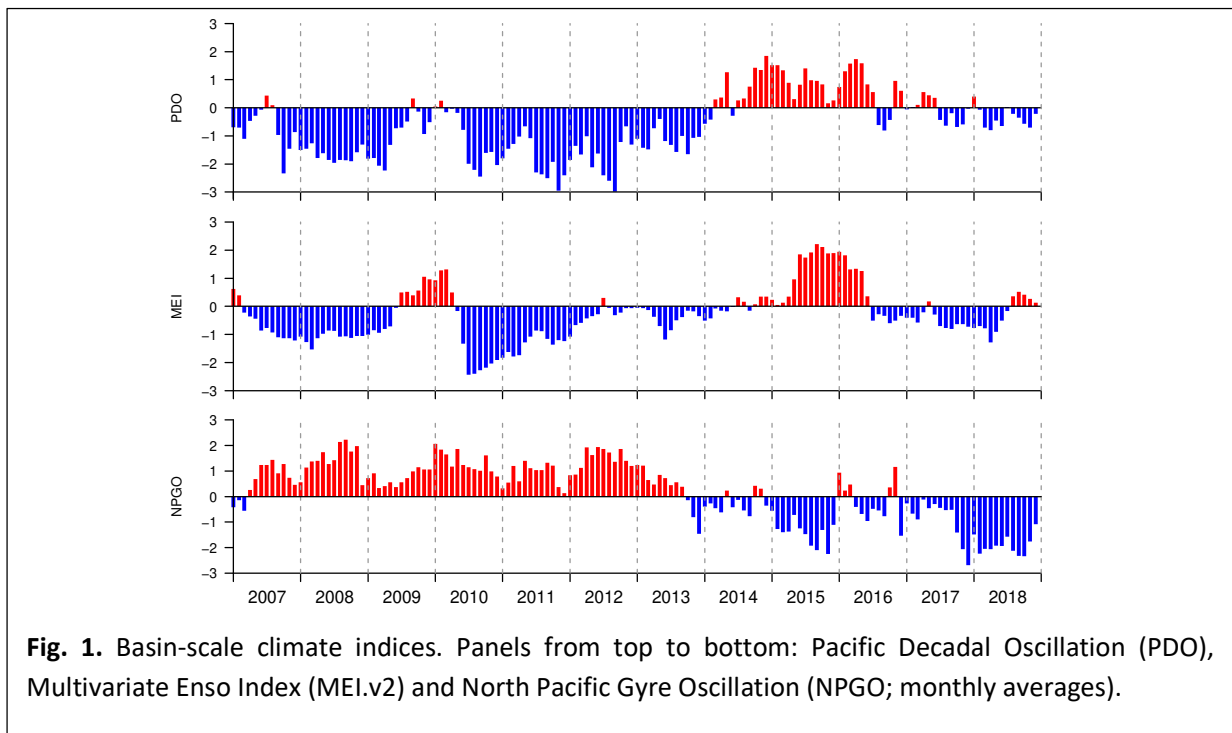
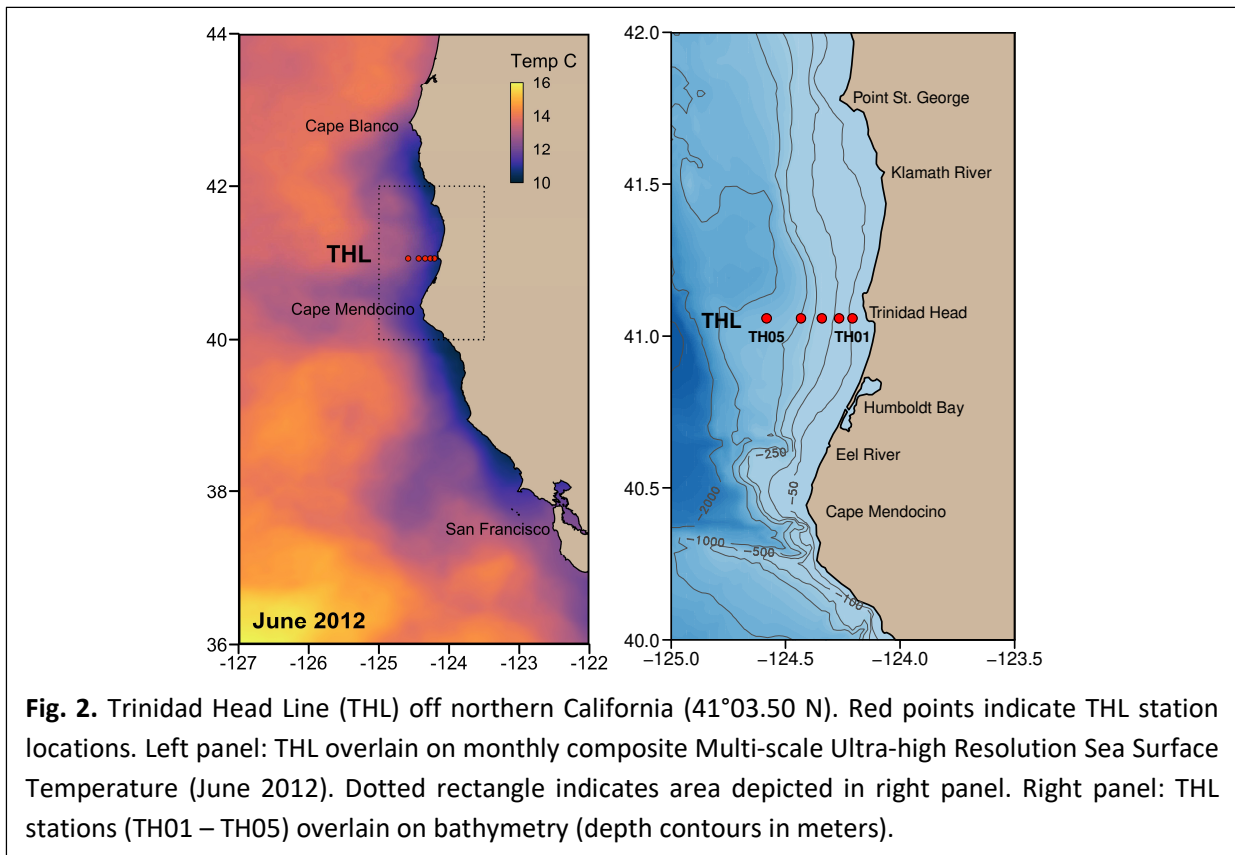


Fig. 1. Basin-scale climate indices. Panels from top to bottom: Pacific Decadal Oscillation (PDO), Multivariate Enso Index (MEI.v2) and North Pacific Gyre Oscillation (NPGO; monthly averages).

112 2017 (Wells et al., 2017; Bjorkstedt and Robertson, in prep.).

113



114 **1.2 Trinidad Head Line**

115 The Trinidad Head Line (THL; 41.05° N; Fig. 2) lies in a zone of transition between the northern California
 116 Current (NCC), which is dominated by strongly seasonal upwelling and downwelling along a generally
 117 linear coast (Hickey and Banas, 2009), and regions to the south where upwelling typically is both less
 118 seasonal and more intense than in the north, and interactions between coastal topography, winds, and
 119 alongshore flow lead to greater complexity and mesoscale structure in coastal circulation (Largier et al.,
 120 1993; Batteen, 1997; Strub and James 2000). The region nominally represented by the Trinidad Head lies
 121 between Cape Blanco and Cape Mendocino, two major headlands that strongly influence coastal
 122 circulation through the generation of mesoscale upwelling jets and eddies (Hayward and Mantyla, 1990;
 123 Barth et al., 2000). Upwelling in this region is most consistent and reaches maximum intensity in the
 124 spring and summer; upwelling can occur during the late fall and winter, but is weaker, and often
 125 countered by southerly winds associated with winter storms (Leising et al., 2015). Circulation in the area
 126 north of Cape Mendocino can be relatively sluggish, even during intense upwelling, owing to the
 127 dominant winds having a substantial onshore component along the westward jutting coastline (Largier

128 et al., 1993; Garcia-Reyes and Largier, 2012). During winter and early spring, coastal waters are
129 impacted by freshwater discharge from major rivers (e.g., the Klamath and Eel) and numerous smaller
130 coastal watersheds.

131

132 **2. Methods**

133

134 **2.1 Trinidad Head Line Sampling**

135 Hydrographic data and plankton samples have been collected along the THL since 2007. The THL
136 currently extends from the inner shelf to the upper slope waters (5 - 37 km offshore) and consists of five
137 stations, three of which lie over the shelf (TH01, 35 m bottom depth; TH02, 75 m; TH03, 140 m) and two
138 offshore of the shelf break (TH04, 460 m; TH05, 780 m) (Fig. 2). Cruises occur at roughly monthly
139 (occasionally biweekly) intervals. Sampling is conducted during 12-hour cruises that, with few
140 exceptions, have been scheduled to sample stations offshore of the shelf during darkness so that
141 zooplankters exhibiting diel vertical migrations will be more susceptible to capture. Shelf stations were
142 often occupied after sundown during the early part of the time series, but are now occupied during late
143 afternoon to minimize navigational challenges and risk of fouling with crab gear.

144

145 **2.1.1 Hydrographic Observations**

146 At each station hydrographic data are collected with a Sea Bird Electronics (SBE) model 19 plus V2 CTD
147 (19 plus until September 2014) cast to a maximum depth of 500 m (150 m until August 2014) or to
148 within a few meters of the seafloor. The CTD is integrated with a SBE ECO-55 water sampler and several
149 auxiliary sensors. The CTD and its auxiliary sensors are calibrated annually or bi-annually by Sea Bird
150 Electronics. CTD casts and zooplankton sampling are performed almost simultaneously. Consequently,
151 collection of hydrographic data occur during late afternoon through midnight. CTD data were initially
152 processed using SBE Data Processing software (version 7.26.4.23) and resulting profiles were projected
153 to 1 m depth intervals. Fluorescence data (WET Labs WET Star WS3S) were used as a proxy for
154 chlorophyll *a* concentration. Depth-by-time Hovmöller plots displaying variability in water column
155 structure were compiled by interpolation across the time series of hydrographic casts for individual
156 stations.

157

158 **2.1.2 Zooplankton Collections and Sample Processing**

159 Zooplankton were sampled with a 0.7 m diameter Bongo sampler fitted with 505 and 335 μm mesh that
160 has been dyed dark purple and flowmeters (General Oceanics). At each station, the net was deployed
161 rapidly to a maximum depth of 100 m (or within a few meters of the sea floor) and subsequently
162 retrieved along an oblique profile by maintaining a steady wire retrieval rate (20 m min^{-1}) and adjusting
163 ship speed through the water to maintain a wire angle of $45^\circ (\pm 5^\circ)$ throughout the tow. Nets were
164 gently and thoroughly washed down to concentrate retained plankton in the cod-end and samples were
165 immediately transferred to sample jar(s) for preservation. For this paper, we analyze samples collected
166 with the 505 μm mesh net, and preserved in 5% formalin buffered with a saturated solution of sodium
167 borate in seawater.

168

169 In the laboratory, samples were temporarily transferred to fresh water and quantitatively fractioned
170 with a 1.25L Motodo box plankton splitter (Aquatic Sciences) to produce tractable subsamples for
171 analysis. When samples were split, subsamples consisted of at least 100 euphausiids. Euphausiids were
172 removed under magnification, and subsequently identified to species and stage, and for adults,
173 distinguished by sex. Taxonomic identification was corroborated with online sources (Brinton et al.,
174 1999). Stage designations were determined independently of length by examination of morphological
175 features indicative of developmental stage, again supported by comparison to published descriptions.
176 Specifically, adult euphausiids were distinguished from juveniles on the basis of whether an individual
177 exhibited male or female sexual characteristics, not a critical size threshold. Because *E. pacifica* F4 and
178 F5 furcilia cannot be distinguished based on telson spines, these stages were combined and treated as
179 one stage. For each species-stage combination that occurred in a sample, body length (BL) of up to 25
180 individuals were measured in millimeters from the back of the eye to base of the telson (Shaw et al.,
181 2010). Larger individuals were measured using the microscope's ocular micrometer. Smaller specimens
182 ($< 5 \text{ mm}$) were measured from images captured with a Luminera camera (Infinity 2). Both ocular
183 micrometer and digital measurement were regularly calibrated with a stage micrometer.

184

185 **2.2 Local Oceanographic Conditions (MURS and CUTI)**

186 As indices of regional oceanographic conditions, we obtained monthly sea surface temperatures (SSTs)
187 and temperature anomalies derived from the Jet Propulsion Laboratory's (JPL) Multi-scale Ultra-high
188 Resolution (MUR) SST Analysis (JPL MUR MEaSURES Project., 2010) for 2007-2018 for a grid cell centered
189 at 41.05° N and 125.00° W (<https://coastwatch.pfeg.noaa.gov/erddap/griddap/jplMURSST41mday.html>;

190 <https://coastwatch.pfeg.noaa.gov/erddap/griddap/jplMURSST41anommday.html>). Monthly anomalies
191 are a mean of the daily anomaly product for a given month. Daily anomalies are produced by the
192 Environmental Research Division (ERD) as deviations from a climatology based on the source data (2003-
193 2014). We also obtained daily and monthly values of the Coastal Upwelling Transport Index (CUTI; Jacox
194 et al., 2018b) for 2007-2018 for a grid cell centered at 41° N (accessed 1 June 2019,
195 http://www.mjacox.com/wp-content/uploads/2019/06/CUTI_monthly.csv). Cumulative CUTI was
196 calculated using daily CUTI values from January 1 through December 31.

197

198 **2.3 Data processing and statistical analysis**

199 All data analysis was conducted in R (version 3.5.1; R Core Team, 2018).

200

201 Preliminary inspection of the data indicated that variability in abundance and size-distribution were
202 generally coherent across the transect, especially across station where a particular life history stage was
203 consistently present at high density (Fig. 4). Therefore, except where noted, we present results from
204 analysis of data aggregated by cruise. This resolution is expected to effectively capture responses to
205 climate forcing, while minimizing the influence of covariation among stations within a cruise and among
206 individuals within a station on statistical significance of fitted models. General patterns emerging from
207 our analysis at the resolution of cruises were corroborated by analysis of data aggregated at different
208 temporal resolutions (e.g., by year) or analysis at the scale of individual stations.

209

210 For each life history stage (LHS), we estimated mean length, weighted by areal density at each station.
211 Areal density ($N m^{-2}$) was estimated by expanding the number of individuals identified to account for the
212 fraction of the total sample analyzed, the volume of water filtered, and the depth of the sampled water
213 column. We also developed density-weighted, cruise-specific length distributions. For adult and juvenile
214 life history stages, we characterized multi-modal length distributions (when apparent) by binning size
215 distributions into 1 mm size classes and implementing Hasselblad's (1966) steepest-descent algorithm.

216

217 To assess coherence in size variability across LHS, we estimated correlations in mean length among LHS
218 within cruises using function 'ccf' and 'cor' in package 'corrplot' (version 0.84; Wei et al., 2017).

219 Preliminary examination indicated that correlations among stages tended to be strongest for concurrent
220 measurements (data not shown).

221

222 Based on the uneven spacing of our observations over time, the extent of some gaps (> 2 months in
223 some cases), and a reluctance to obscure variability at short time scales, we elected to analyze data in its
224 original temporal structure rather than to aggregate or interpolate data to evenly spaced time intervals.
225 We examined temporal variability in environmental conditions and mean length at seasonal and
226 interannual time scales. At the seasonal scale, we describe a seasonal climatology for mean cruise
227 temperature, mean cruise $\log_{10}(\text{chlorophyll } a + 1)$, and mean stage-specific length by fitting generalized
228 additive models (GAMs) for each response variable as a function of day of year. GAMs were based on a
229 cyclic cubic spline to favor continuity across the start and end of the annual seasonal cycle. To resolve
230 interannual variability, we fit GAMs for mean environmental conditions and mean length on cruise date.
231 The number of distinct piecewise smooths in the time series GAM was selected by visual inspection to
232 ensure that the fitted curve captured low-frequency variability and lacked substantial excursions from
233 trends in mean length over time. GAMs were fit using functions in package 'mgcv' (Wood, 2017).

234

235 To examine stage-specific relationships between *E. pacifica* length and environmental conditions, we fit
236 linear mixed effects (LME) models, in which the potential for random variability in the slope and
237 intercept of length-environment relationships associated with (calendar) year was assessed, and if
238 present accounted for. Models were fit to relate mean length to temperature and chlorophyll *a*
239 concentration (based on raw fluorescence) averaged over the sampled (max 100 m) water column
240 across the transect. Chlorophyll concentrations were $\log_{10}(n+1)$ transformed prior to averaging. Based
241 on preliminary examination of the data, we considered models that allowed the relationship between
242 length and temperature to vary with year and included chlorophyll concentration as a covariate with
243 effects independent of temperature.

244

245 Optimal model structure, in terms of random and fixed effects, was identified following protocols
246 outlined in Zuur et al. (2009), and using Akaike's Information Criterion corrected for small sample size
247 ('AICc' in 'MuMIn' package, version 1.43.6; Bartoń, 2019) and likelihood-ratio tests ('anova'; R Core
248 Team, 2018) to inform model selection. Models were fit using functions in 'nlme' (version 3.1-137;
249 Pinheiro et al., 2018). Collinearity was assessed and optimal models were validated by graphical
250 evaluation of assumptions (Zuur, 2010). Results from this analysis proved insensitive to the extent of the
251 upper water column selected or to minor adjustments to the stations included in the spatial average of
252 environmental conditions.

253

254 To support interpretation of results from LME models, we generated metrics that express percent
255 change in euphausiid body length associated with change in temperature (dT) and chlorophyll
256 concentration (dC) for an individual of mean length for each LHS at mean temperature and mean
257 chlorophyll. dT scales the slope of the fixed effect of temperature by mean length as a general indicator
258 of percent change in length associated with a change in temperature of 1°C. dC scales the slope of the
259 fixed effect of chlorophyll concentration by the mean length*0.1 as an indicator of percent change in
260 length with a change of 0.1 in $\log_{10}([\text{chl } a]+1)$. We also calculated a metric to capture the range of
261 variability arising from year-to-year variability in the length-temperature relationship (rYT). This metric
262 expresses the range of expected length across years (excluding 2010 due to incomplete seasonal
263 coverage) as a percentage relative to mean length at mean temperature and mean chlorophyll. Positive
264 (negative) rYT indicates larger size during warm (cool) years.

265

266 To complement results emerging from analyses based on mean length and mean environmental
267 conditions, we applied canonical correspondence analysis (CCA) using function 'cca' in package 'vegan'
268 (version 2.5-5; Oksanen et al., 2013) to assess correlation between water column structure and adult
269 size distributions at the scale of individual samples. For each sample, we binned *E. pacifica* adult lengths
270 into 1 mm size classes to generate a station-by-size class matrix. A corresponding temperature-by-
271 station matrix was created by binning temperature in the upper 100 m of the water column to 0.25°C
272 intervals and quantifying the vertical extent (at 1 m resolution) of the water column in each temperature
273 bin. CCA was applied to estimate the strength and structure correlation between the resulting size and
274 temperature matrices. Statistical significance was assessed with a permutation test using function
275 'anova.cca' in package 'vegan' (n = 10,000).

276

277 We calculated individual biomass by first converting body length (BL; mm) to total length (TL; mm) and
278 subsequently calculating dry weight (DW, μg) and biomass (B; μgC) per the following relationships

279
$$\text{TL} = 1.218 * \text{BL} + 0.2807$$

280
$$\text{DW} = 0.795 * \text{TL}^{3.239}$$

281
$$\text{B} = -1.985 + 0.401 * \text{DW}$$

282 following (Shaw et al., 2010; Feinberg et al., 2007; Ross, 1982a). Biomass (divided by 1000) was then
283 expanded to areal density (mgC m^{-2}) by summing over size-distributions and applying cruise-specific
284 density estimates.

285

286 **3. Results**

287

288 Our data set includes measurements of over 30,000 individual *E. pacifica*, spanning F1 – F7 furcilia,
 289 juvenile, and adult life history stages (Table 1) from the five core stations (TH01-TH05) over the course
 290 of 119 cruises conducted between November 2007 and December 2018 and associated hydrographic
 291 data. *E. pacifica* represent approximately 80% (by numerical density) of the euphausiid assemblage
 292 along the THL. The data set includes observations from a handful of cruises that were cut short by
 293 deteriorating weather, so that only three or four stations were occupied. Four cruises were conducted
 294 on an unusual schedule so that offshore stations were collected during daylight hours. We exclude data
 295 from these cruises in statistical analyses as preliminary examination of these data suggested a strong
 296 day-night bias in abundance, presumably related to well-documented diel vertical migration behaviors
 297 (Brinton, 1965; Brinton and Townsend, 2003). We lack samples from cruises during summer 2010 due to
 298 the loss-at-sea of our Bongo net.

Stage	This study		Boden (1950)	Ross (1981)		Feinberg et al. (2006)		Ross (1981)	
	Range of Mean Length	Range of Mean Biomass	Range of Length	8°C		10.5°C		12°C	
				Development Time	Duration	Development Time	Duration	Development Time	Duration
Ad	8.46 - 17.27	662 - 6412	*8.75 - 17.78	--	--	--	--	--	--
Jv	3.98 - 11.06	62 - 1549	*4.65 - 7.93	68	--	58.4	--	45	--
F7	3.40 - 5.79	37 - 200	4.07 - 5.06	60	8	55.4	3	40	5
F6	3.25 - 4.53	32 - 93	3.82 - 4.65	55	5	51	4.4	36.5	3.5
F4/5	2.93 - 4.34	23 - 81	3.17 - 4.23	50	5	43.4	7.6	33	3.5
F3	2.57 - 3.62	15 - 46	2.67 - 3.99	38.5	11.5	32.4	11	27.5	5.5
F2	2.19 - 3.44	9 - 39	2.59 - 3.09	31	7.5	26.7	5.7	23	4.5
F1	1.53 - 2.93	2 - 23	2.18 - 2.35	22	9	21.3	5.4	17.5	5.5

Adult lifespan: 1+ years (Ross, 1982; Feinberg et al., 2006); * from Brinton and Wyllie, 1976

Table 1. Stage-specific range of mean lengths (mm) and biomass ($\mu\text{g C ind}^{-1}$) of *Euphausia pacifica* adults (Ad), juveniles (Jv), and F7 – F1 furcilia measured in this study and in Boden (1950) and Brinton and Wyllie (1976). Median development time (days to stage) and stage duration (days) at 8, 10.5, and 12 °C (from Ross, 1981 and Feinberg et al., 2006) provided as context for interpretation of results.

299

300 **3.1 Local Oceanographic Conditions**

301 A comprehensive analysis of variability in hydrographic structure along the Trinidad Head Line is beyond
 302 the scope of this paper and will be developed elsewhere. For the present analysis, we describe
 303 hydrographic conditions at the outer-shelf station (TH03; Fig. 2) as representative of the state of the
 304 ocean along the Trinidad Head Line. Observations at this station (and elsewhere along the transect)
 305 capture annual transitions between seasonal upwelling and downwelling, including variability in the

306 timing of these transitions, as well as the local signature of larger scale processes captured in basin-scale
307 oceanographic climate indices spanning the THL record (Fig. 3).

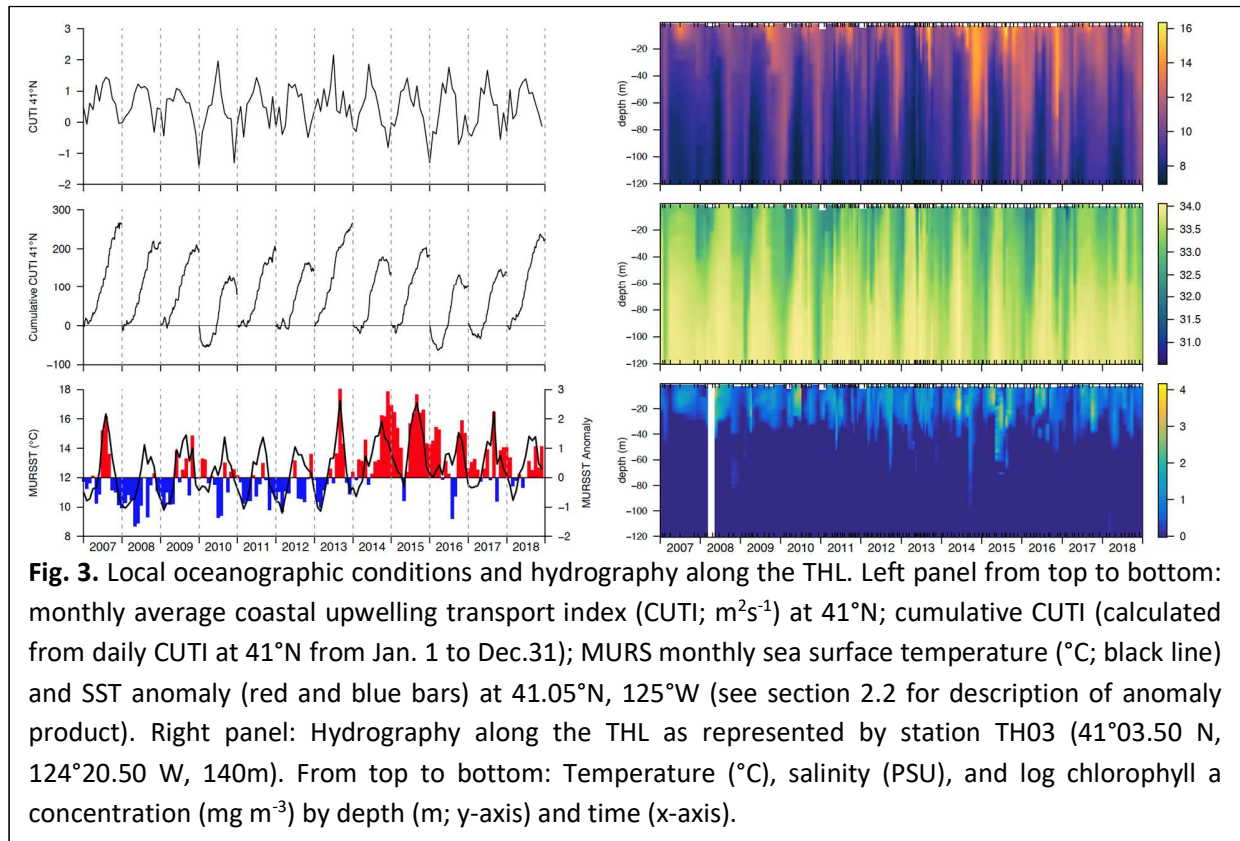


Fig. 3. Local oceanographic conditions and hydrography along the THL. Left panel from top to bottom: monthly average coastal upwelling transport index (CUTI; m^2s^{-1}) at 41°N ; cumulative CUTI (calculated from daily CUTI at 41°N from Jan. 1 to Dec.31); MURS monthly sea surface temperature ($^\circ\text{C}$; black line) and SST anomaly (red and blue bars) at 41.05°N , 125°W (see section 2.2 for description of anomaly product). Right panel: Hydrography along the THL as represented by station TH03 ($41^\circ03.50\text{ N}$, $124^\circ20.50\text{ W}$, 140m). From top to bottom: Temperature ($^\circ\text{C}$), salinity (PSU), and log chlorophyll a concentration (mg m^{-3}) by depth (m; y-axis) and time (x-axis).

308
309 Several features of the THL hydrographic record and local oceanographic conditions warrant attention
310 as important context for interpreting variability in euphausiid length over time (Fig. 3). Conditions along
311 the Trinidad Head Line were unusually cool during the short period of record prior to the 2009-10 El
312 Niño. Although moderate in strength, the 2009-10 El Niño caused warming and freshening over much of
313 the water column in response to changes in wind forcing (Fig. 3). Following the 2009-10 El Niño, the MEI
314 transitioned to La Niña conditions (2010-2012; Fig. 1). Along the THL, this period was generally
315 characterized by the juxtaposition of cool spring conditions driven by strong upwelling followed by rapid
316 warming and freshening of the upper water column as wind forcing weakened. A minor shift to positive
317 MEI during 2012 manifested locally in delayed and weak upwelling. Additionally, strong downwelling in
318 spring 2012 resulted in relatively fresh water and low chlorophyll values along the THL (Fig. 3). The first
319 phase of the 2014-16 MHW was marked by the arrival of the warm 'blob' in coastal waters (Bond et al.,
320 2015; Chao et al., 2017) and the rapid development of strong warming extending to at least 60 m depth.

321 This transition, although appreciable, was preceded by surface warming in 2013, during which, intense
322 spring upwelling gave way to the strongest surface temperature anomaly in the recent local record
323 (Gentemann et al., 2016), and early 2014 when unusually warm conditions were observed in advance of
324 the more extensive (deeper) warm 'blob' signal (Leising et al., 2015). Following a brief period of
325 upwelling-driven cooling in early 2015, weak upwelling followed by sustained downwelling and
326 poleward flow associated with the strong 2015-16 El Niño influenced warm water anomalies; winter and
327 spring temperatures remained at least 1°C warmer than average at the surface and had a strong,
328 persistent signature in the upper 60-70 m of the water column (Fig. 3). Warm conditions persisted into
329 late 2016 and early 2017, but by late 2017 and into 2018 unusually warm conditions were again
330 restricted to the upper water column. Warm water anomalies abated further during winter 2018, as
331 relatively weak storm activity and downwelling allowed relatively cool, salty waters to persist over the
332 shelf.

333

334 **3.2 Cross-shelf patterns in stage-specific distributions and size structure**

335 Stage-specific cross-shelf distributions of *E. pacifica* exhibited a consistent ontogenetic pattern marked
336 by a shift from the broad dispersion of early furcilia stages (F1 and F2) across the transect towards the
337 more concentrated distribution of adults and juveniles over the outer shelf and offshore of the shelf
338 break (Fig. 4). Stage-specific size distributions of immature life history stages were generally consistent
339 across the transect in aggregate (Fig. 4) and within cruises (correlations of mean size among stations
340 across the time series ranged from 0.71 to 0.99, yielding slope estimates of 0.74 to 1.16 from linear
341 models forced through the origin (data not shown). Size distributions of adult *E. pacifica* were consistent
342 at offshore stations where adults were most abundant. The few adult *E. pacifica* captured at inshore
343 stations are an exception to this pattern, as these individuals tended to be larger than those captured
344 further offshore. Adult length distributions frequently included two or more distinct modes. When this
345 occurred, further inspection indicated that these modes were in general consistently represented across
346 stations and did not arise from aggregation of spatially distinct modes in the size distribution. In general,
347 stage-specific size distributions overlapped, but mean size consistently differed across stages over time.

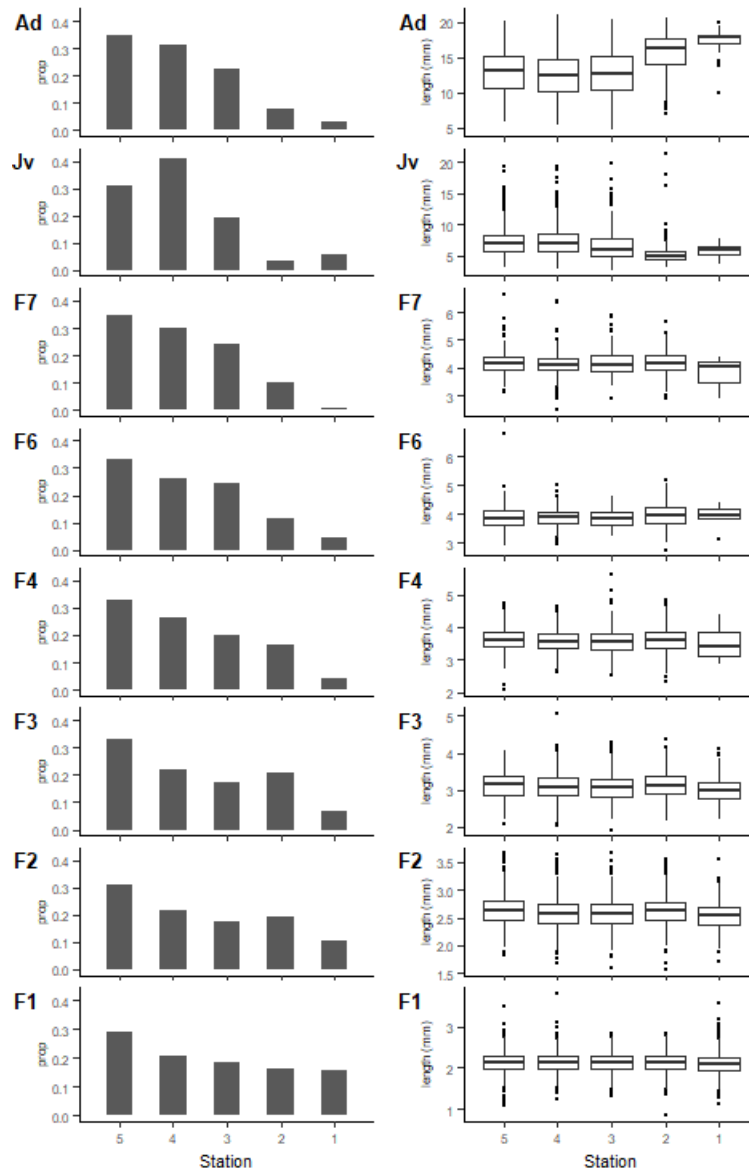


Fig. 4. Stage-specific cross-shelf distributions (left-hand column) and size-distributions (right-hand column) of *Euphausia pacifica* aggregated over all cruises. From top to bottom: adults (Ad), juveniles (Jv) and F7-F1 furcilia stages. Cross-shelf distributions are represented as relative abundance (y-axis) based on areal density ($N\ m^{-2}$) at stations (x-axis) TH05 (offshore) – TH01 (nearshore). Box represents the interquartile range (IQR) bounded by the 25th and 75th quartiles. Whiskers extend 1.5 IQR from the median.

348

349 **3.3 Coherence in Stage-Specific Length Distributions**

350 Mean length of adult males and females was significantly correlated ($r^2 = 0.90$, $p < 0.001$; Fig. A.1).

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We observed substantial overlap between stage-specific size distributions (Fig. 4 and Table 1), yet differences in mean size across stages (including males and females) was maintained over time. Mean length of juveniles and adults was significantly correlated ($r^2 = 0.34$, p -value < 0.001 ; Table 2), as were mean lengths across furcilia stages (Table 2).

3.3.2 Furcilia

Mean length was positively correlated among all furcilia stages ($r^2 = 0.23 - 0.72$; $p < 0.05 - 0.001$; Table 2), and correlations tended to be strongest across sequential stages. Among furcilia stages, only F1 showed a significant correlation in mean size with adults (Table 2).

Stage	A	J	F7	F6	F4/5	F3	F2	F1
A		0.34***	-0.08	0.08	-0.01	0.06	0.13	0.23*
J	0.54		-0.11	-0.03	-0.11	-0.09	-0.10	0.13
F7	-0.45	-0.41		0.62***	0.56***	0.55***	0.23*	0.41***
F6	0.60	-0.13	0.72		0.68***	0.55***	0.26*	0.41***
F4/5	-0.04	-0.47	0.67	0.62		0.72***	0.30**	0.43***
F3	0.55	-0.54	0.84	0.65	0.95		0.42***	0.52***
F2	1.34	-0.54	0.37	0.31	0.42	0.43		0.38***
F1	0.02	0.02	0.21	0.27	0.27	0.43	0.33	

Table 2. Correlations (top panel) and slope estimates (bottom panel) of *Euphausia pacifica* mean body length by cruise among life history stages (furcilia F1-F7, J = juveniles, A = adults). Slopes were estimated from linear models of mean length for younger stages on older stage (e.g., F1 regressed on F2). Significance of correlation: * $p < 0.05$, ** $p < 0.01$, *** $p < 0.001$.

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3.4 Temporal variability in stage-specific size distributions

In the following sections, we describe temporal variability in stage-specific size distributions at seasonal and interannual scales in the context of corresponding variability in temperature and $\log_{10}(n+1)$ -transformed chlorophyll *a* concentration, averaged over stations and depths within a cruise.

3.4.1 Seasonal patterns

Seasonal climatology of mean temperature conformed to characteristic patterns of cooling during spring and summer associated with upwelling and equatorward transport and warming during fall and winter

371 (Fig. 5). This pattern held for temperature during years influenced by warming events, during which
372 much of the divergence in temperature from climatological conditions occurred during fall and winter.
373 Winter temperatures remained low during winter under cool climate conditions, with intense upwelling-
374 driven cooling noted only during spring 2013.

375
376 Chlorophyll *a* concentration also exhibits a generally consistent seasonal pattern, with seasonal maxima
377 occurring during the summer months (Fig. 5). Substantial blooms occurred during spring and summer in
378 both cool and warm years.

379

380 **3.4.1.1 Adults and Juveniles**

381 Adult *E. pacifica* captured along the THL tended to be larger during summer and smaller during winter
382 across all years (Fig. 5). Seasonal (climatological) variability in adult size was more strongly (inversely)
383 coherent with seasonal patterns in mean temperature than with climatological chlorophyll *a*
384 concentration. Offsets from the mean seasonal pattern were linked to interannual climate variability:
385 mean size tended to exceed seasonal means throughout the year during cool conditions, and
386 conversely, to be smaller than expected throughout the year during warm events, and in particular
387 during the 2014-2016 MHW.

388

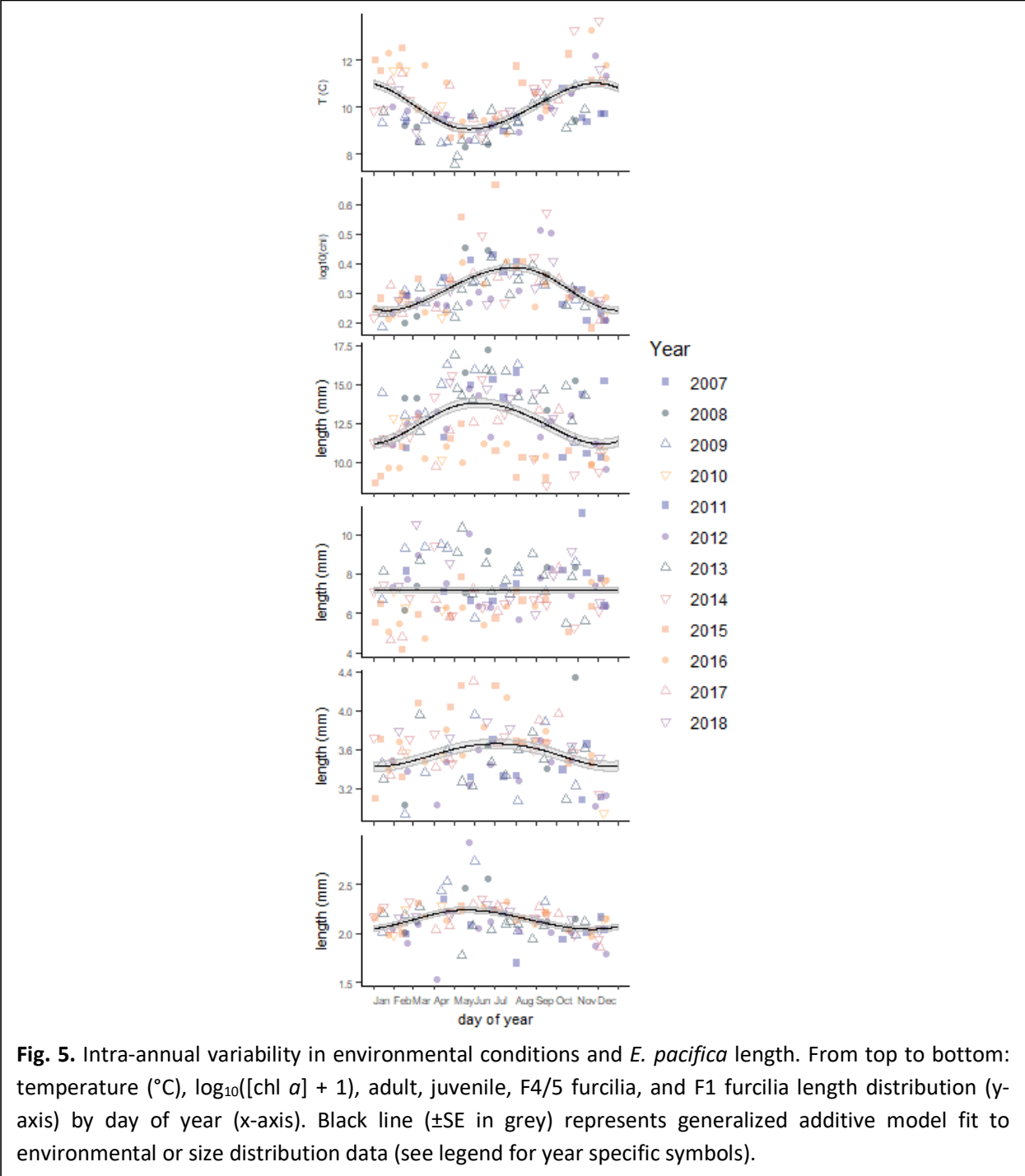
389 We did not observe a consistent seasonal pattern in mean size of juvenile *E. pacifica*. Rather, the timing
390 of seasonal maxima and minima varied substantially across years (Fig. 5). As observed for adults,
391 however, juvenile *E. pacifica* tended to be smaller than average throughout warm years and larger
392 during cool years.

393

394 **3.4.1.2 Furcilia**

395 Intermediate furcilia stages exhibited an opposite pattern with respect to the offset between mean size
396 and seasonal variability in size: individuals tended to be larger under warmer than usual conditions and
397 smaller during cool years (Fig. 5). Seasonal trends in mean size of these stages tend to lag temperature
398 climatology and to overlap more with the seasonal pattern in chlorophyll *a* concentration. The earliest
399 furcilia stage, F1, did not exhibit consistent offsets from seasonal mean size in relation to climate
400 variability.

401



402 **3.4.2 Interannual trends**

403 Interannual variability in mean temperature is entirely consistent with patterns described above (3.1
404 Local Oceanographic Conditions), with clear indication of sustained warming associated with the 2009-
405 10 El Niño and the 2014-16 MHW, and cool periods during the early part of the record (late 2007-2009)
406 and 2011 through mid-2014 (Fig. 6).

407
408 Low-frequency variability in chlorophyll *a* concentration was most apparent in the magnitude of
409 seasonal maxima (Fig. 6). Several substantial blooms were observed over the course of the time series,
410 often coinciding with (or closely following) evidence of relaxation from upwelling. Chlorophyll *a*
411 concentration remained low during years directly influenced by El Niño events (2010 and 2016) or
412 reduced upwelling (2017).

413

414 **3.4.2.1 Adults and Juveniles**

415 The pattern of low-frequency variability in mean length of adult *E. pacifica* is very nearly inverse that
416 observed for mean temperature (Fig. 6). From 2007 to mid-2009, under consistently cool conditions,
417 large (>15 mm) adults were present year-round and dominated the population. In mid-2009, as the
418 2009-10 El Niño began to influence the region, the size structure of the population shifted; large
419 euphausiids were present but smaller (<10 mm) adults made up a substantial fraction of the population.
420 Although we lack data to characterize any trends in size during the summer 2010, the tendency for small
421 (<10 mm) and large (>15 mm) size classes to be simultaneously present year-round continued into 2011
422 and on through the first half of 2014. During this period, large adults tended to be dominant during
423 summer months while small adults were more abundant in winter. One exception to this pattern
424 occurred in 2012, when the maximum size of adults was lower and modal structure of the size
425 distribution were less distinct. Mean adult length declined sharply with the arrival of warm 'blob' waters
426 in mid-2014, individuals larger than 12 mm effectively disappeared from our samples. Larger adults (> 15
427 mm) were briefly abundant during a period of intense upwelling in early 2015, but again disappeared as
428 the 2015-16 El Niño renewed unusually warm conditions along the coast. During this period, the typical
429 seasonal pattern of increasing adult size during spring and summer was diminished in scope, such that
430 mean adult length remained at or below the time series mean (12.9 mm ± 3.03 SD) and on several
431 occasions was as low as 60-70% of the time series mean. Beginning in 2017, large (>15 mm) adults were
432 again present in substantial numbers, and the lower bound of adult sizes increased. However, the upper
433 range and mean adult size distributions had not yet recovered to values observed prior to 2014.

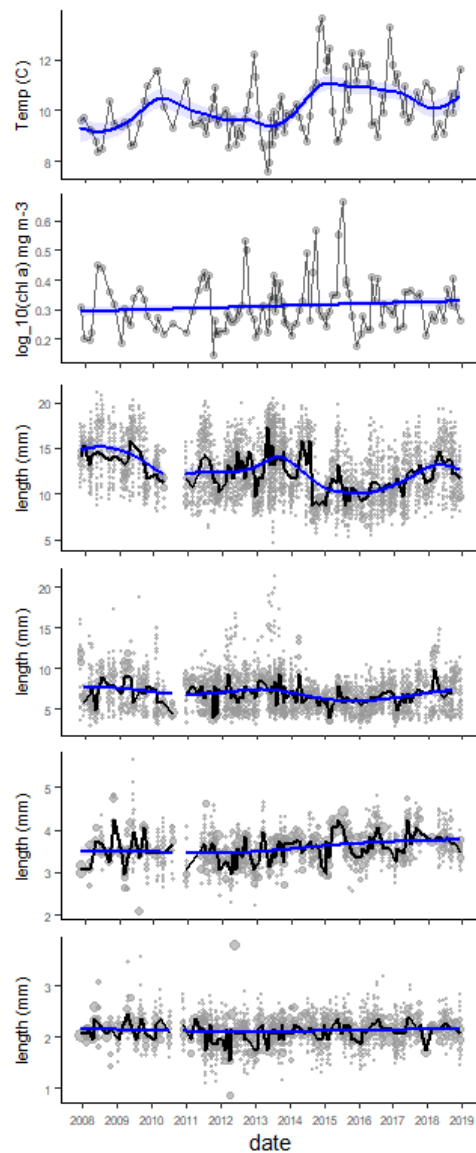


Fig. 6. Interannual variability in environmental conditions and *E. pacifica* length distributions. From top to bottom: temperature (°C), $\log_{10}([\text{chl } a]+1)$, adult, juvenile, F4/5 furcilia, and F1 furcilia length distribution based on body length measurements (mm; y-axis) by cruise date (x-axis). Black line tracks mean length by cruise. Blue line indicates fit of generalized additive model (tuned to capture interannual trends) fit to mean of cruise-specific length distributions (gray points scale with stage-specific log-transformed density; N m^{-2}).

434

435 Low-frequency variability in juvenile length generally corresponds to patterns observed for adults (Fig.

436 6). However, the decline in juvenile size begins in late-2013, prior to the sharp decline in adult mean

437 length. The few unusually large juveniles detected in our samples were all captured in 2013 or during
438 the cool years preceding the 2009-10 El Niño.

439

440 **3.4.2.2 Furcilia**

441 Greater cruise-to-cruise variability in size distributions for furcilia limited our ability to characterize low-
442 frequency trends; most fitted GAMs indicated a slow increase in size over time, sometimes marked by a
443 subtle inflection coincident with the onset of the 2014-2016 MHW (Fig. 6). These patterns corroborate
444 evidence from visual inspection and comparison of seasonal patterns that furcilia tended to be smaller
445 for a given stage during cool years and larger during periods that span or follow the warm water
446 anomalies (2014 through 2017). Although small in absolute terms, the observed changes in mean length
447 represent sustained shifts on the order of 5-8% across periods.

448

449 **3.5 Length-Environment Relationships**

450 **3.5.1 Linear Mixed Effects Models**

451 All life history stages exhibited a decline in mean length with increasing temperature, and for most life
452 history stages this relationship proved to be statistically significant (Table 3). Scaled to stage-specific
453 mean lengths, size-temperature slopes translate to reductions in length on the order of 1.09-5.68% °C⁻¹
454 for immature life history stages and 7.5% °C⁻¹ for adults. For most life history stages, this pattern is
455 associated with seasonal variability: individuals tended to be smaller during warmer (non-upwelling)
456 seasons, and conversely, larger individuals were captured during cooler (upwelling) seasons (Table 3 and
457 Fig. 7; see Fig. A.2 for comprehensive LHS models).

458

459 For most life history stages, the optimal LME model included year-to-year random variability in the
460 intercept of the length-temperature relationship, corresponding to an annual offset of the seasonal
461 pattern (Table A.1).

462

463 The pattern is more complex for adults, for which the optimal model included random variation in both
464 intercept and slope of the length-temperature relationship (Table 3). In this case, year-to-year variability
465 captured a clear shift from steeply sloped responses (with correspondingly higher intercepts) occurring
466 during cool years to length-temperature relationships with shallower slopes and lower intercepts during
467 warm years, such that mean length appears to converge on a common range of smaller sizes with
468 increasing temperature (Fig. 7). The optimal model for variation in the upper bound of adult size

stage (random structure)	Fixed								Full				
	ΔAICc			slope (SE)					r ²	dT	dC	r ²	rYT
	L ~ T + log ₁₀ (C+1)	L ~ T	L ~ log ₁₀ (C + 1)	T	log ₁₀ (C + 1)	r ²	dT	dC					
Adult (rSI)	0	0.77		-0.97 (0.17) ***	2.97 (1.62)	0.34	-7.50	2.29	0.62	-22.02			
Juvenile (rl)	2.16	0		-0.39 (0.11) ***	-0.19 (1.40)	0.12	-5.68	-0.28	0.21	-11.51			
F7 (rl)	0	2.32		-0.09 (0.03) **	0.88 (0.42) *	0.16	-2.24	2.12	0.18	1.51			
F6 (rl)	1.13		0	-0.03 (0.03)	0.96 (0.33) **	0.12	-0.68	2.46	0.11	0.02			
F4/5 (rl)	0			-0.08 (0.03) **	1.06 (0.32) **	0.20	-2.29	2.94	0.44	9.61			
F3 (rl)	0			-0.07 (0.02) ***	0.56 (0.24) *	0.17	-2.30	1.82	0.49	13.72			
F2 (rl)	0.3		0	-0.03 (0.02)	0.64 (0.35) *	0.11	-1.09	2.59	0.20	5.28			
F1 (rl)	0.86	0		-0.08 (0.02) ***	0.22 (0.19)	0.24	-3.60	1.04	0.38	6.95			
A ₁₀ (rSI)	2.25	0		-0.75 (0.19) ***	-0.20 (1.71)	0.19	-5.04	-0.14	0.53	-30.40			
A ₉₀ (rl)	0	4.97		-1.01 (0.16) ***	5.40 (1.94) **	0.36	-10.16	5.44	0.50	-16.27			

Table 3. Summary of stage-specific relationships between mean length and temperature and log₁₀(n + 1)-transformed chlorophyll as characterized by linear (Fixed) and linear mixed-effects (Full) models with year as a random effect on intercept (random structure = rl) or slope and intercept (rSI). ΔAICc values shown for models with various fixed structures. Slope and standard error (SE) of coefficients for model with both environmental covariates (L ~ T + log₁₀ ([chl a])) shown. Asterisks included in slope estimates indicate level of significance (* < 0.05, ** < 0.01, *** < 0.001). r² values under 'fixed' and 'full' indicate variance explained by the fixed effect of T + log₁₀ ([chl a]+1) and by the full LME model, respectively. dT scales the slope of the fixed model by the mean length as a general indicator of percent change in length associated with a change in temperature of 1°C. dC scales the slope of the fixed model by the mean length*0.1 as an indicator of percent change in length with a change of 0.1 in log₁₀([chl a]+1). rYT represents the range of difference in lengths across years as a percentage relative to mean length at mean(T) and mean(C). Grey cells indicate models did not have substantial support. A₁₀ and A₉₀ stage represent lower (10th quantile) and upper (90th quantile) bounds of adult size distributions, respectively.

469 distributions (using the 90th quantile as a robust metric) indicates a strong correlation with temperature,
470 with year-to-year offsets from common slope (Table 3 and Fig. 8). In contrast, the optimal model for
471 variation in the lower bound of adult size distributions (10th quantile) includes year-to-year variability in
472 both slope and intercept of the length-temperature relationship, and exhibits a pattern of steeper
473 slopes during cool years and shallower (even positive) slopes during warm years similar to that observed
474 for variability in mean adult size.

475

476 With the exception of the juvenile and F1 stages, all life history stages exhibited positive relationships
477 between mean length and (log₁₀(n+1)-transformed) chlorophyll, and in many cases, model selection
478 gave preference to models that included both temperature and chlorophyll a as explanatory variables
479 (Table 3). Statistically significant length-chlorophyll relationships were detected only for furcilia stages
480 (excepting F1). For two furcilia stages (F2 and F6), a model relating mean size to chlorophyll

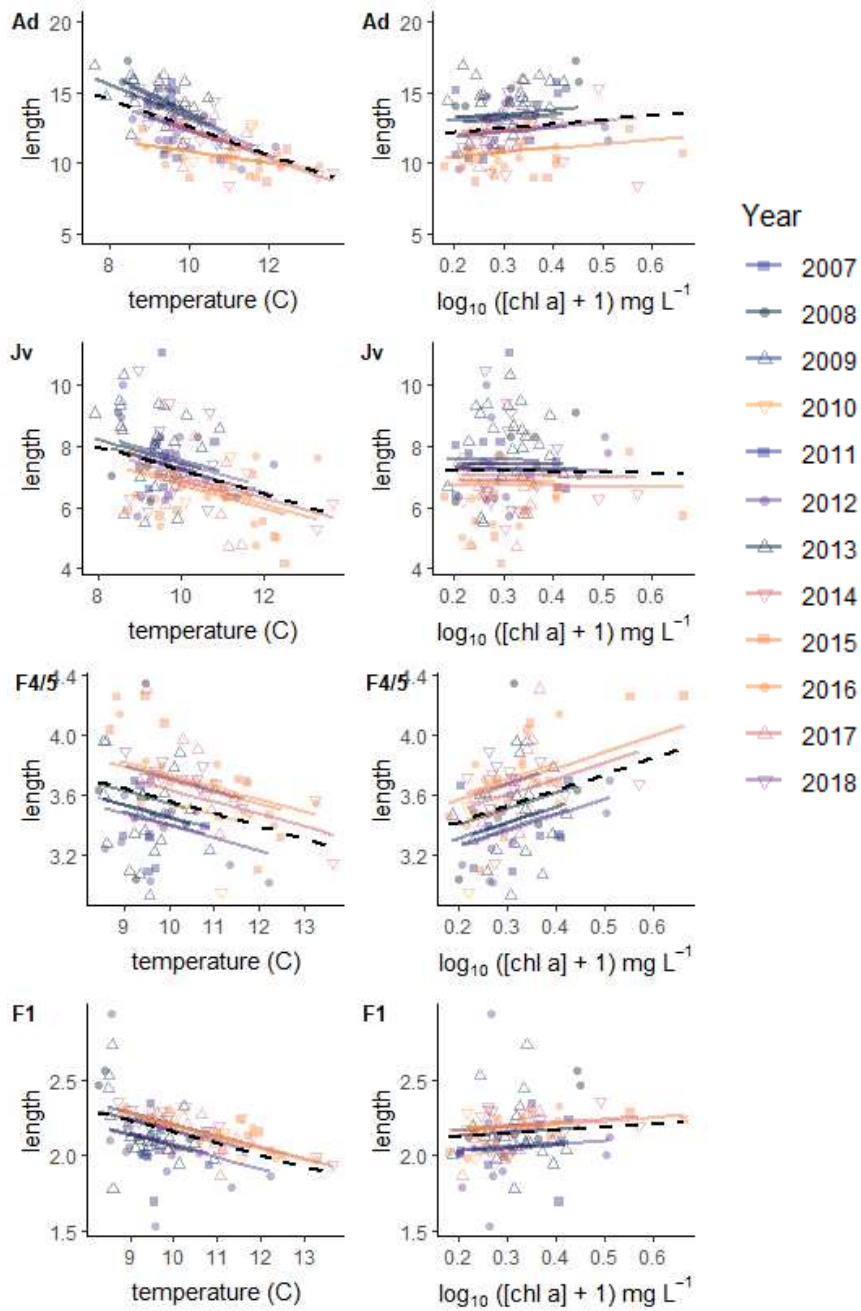
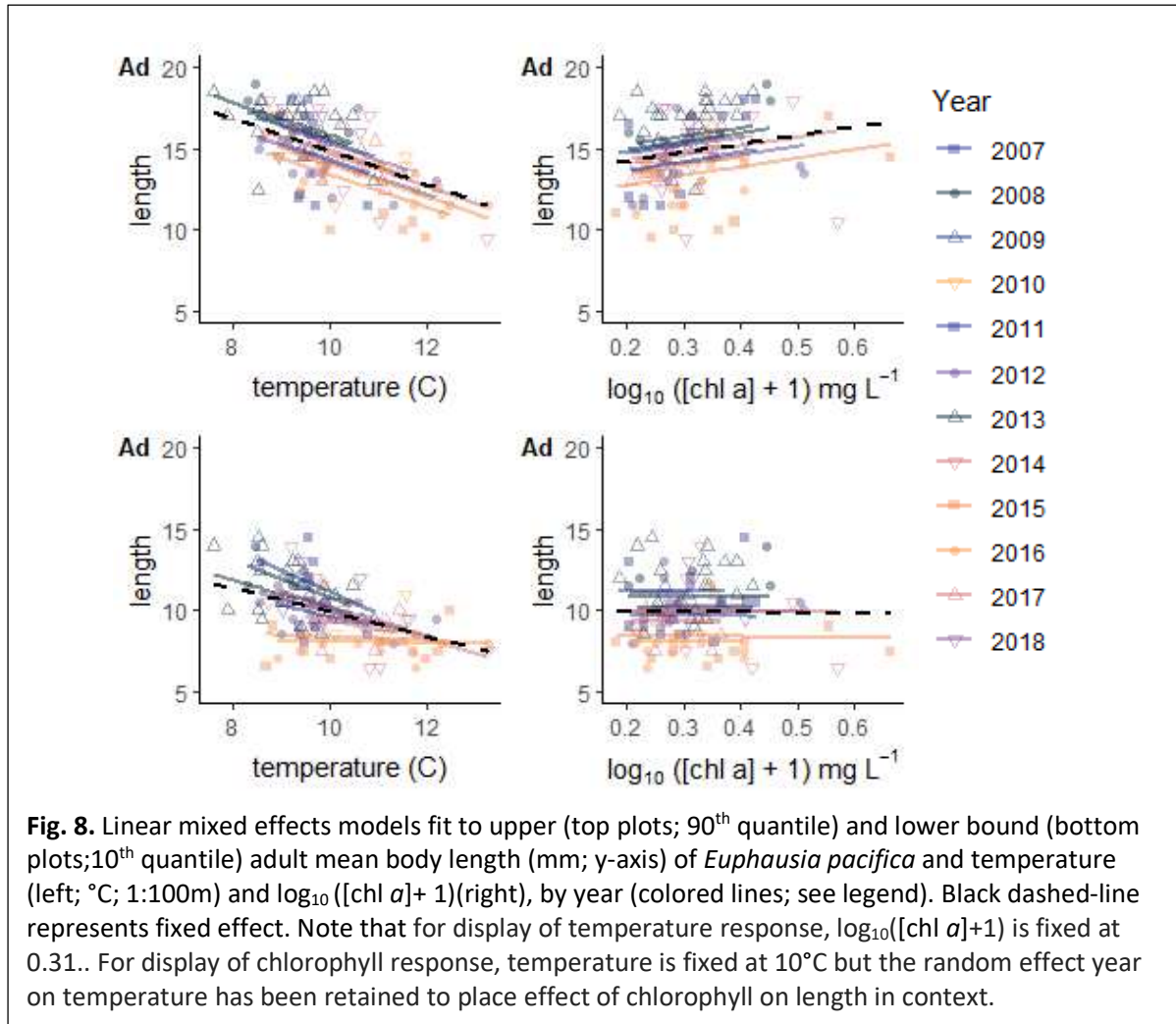


Fig. 7. Linear mixed effects models fit to stage-specific mean body length (mm; y-axis) of *Euphausia pacifica* and temperature (left; °C; 1:100m) and $\log_{10}([\text{chl } a] + 1)$ (right), by year (colored lines; see legend). Black dashed-line represents fixed effect. Life history stages from top to bottom: adults (Ad), juveniles (Jv), F4/5 and F1 furchilia). Note that for display of temperature response, $\log_{10}([\text{chl } a] + 1)$ is fixed at 0.31. For display of chlorophyll response, temperature is fixed at 10°C but the random effect year on temperature has been retained to place effect of chlorophyll on length in context.

481 concentration was identified as being most parsimonious; in both cases, a model that also included
 482 temperature had substantial support (ΔAICc of 0.30 and 1.13 for F2 and F6, respectively; Table 3). For F1

483 furcilia, chlorophyll was excluded from the best model (based on temperature), but retained in a robust
484 alternative model ($\Delta AICc = 0.86$; Table 3).

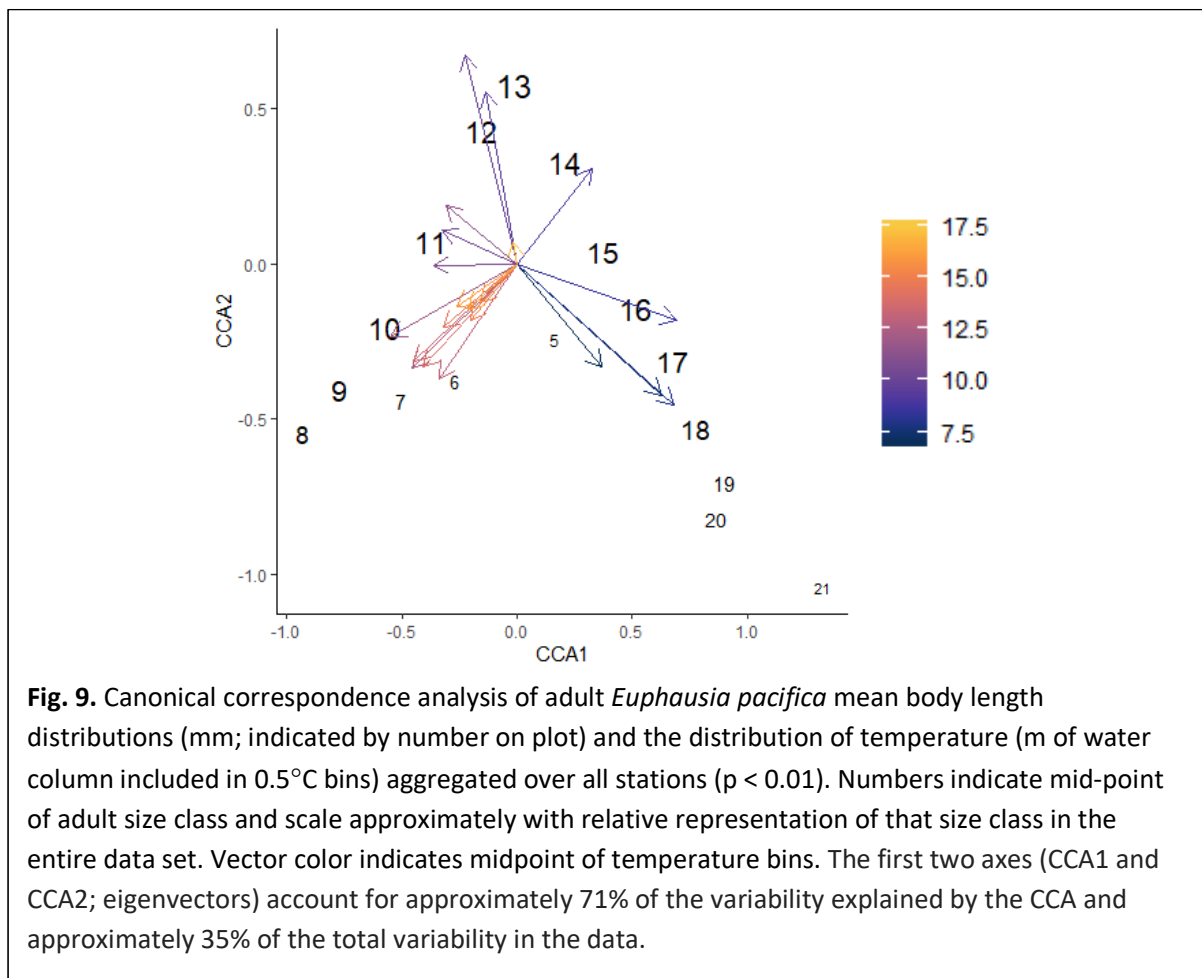


485

486 3.5.2 Canonical Correspondence Analysis

487 Size distributions of adult *E. pacifica* were strongly correlated with water column structure (Fig. 9, $p <$
488 0.01), with small- to large- size classes of adult *E. pacifica* mapping neatly onto the relative proportion of
489 the sampled water column occupied by warm to cold waters, respectively. For example, 16-17 mm
490 adults are strongly associated with the presence of 7-8°C water, while 9-10 mm adults co-occur (in time)
491 with waters from 13-16°C. The presence of 12-13 mm adults typically coincides with the presence of
492 10°C water, regardless of whether, for a particular cruise, 10°C water is relatively cold and deep and 12-
493 13 mm adult euphausiids are among the largest present, or the converse, that 10°C water is relatively

494 warm and near the surface and 12-13 mm adults are on the small side. Approximately 49% of the
495 variability in size distributions can be attributed to water column structure through this analysis, with
496 the first two axes (eigenvectors) accounting for approximately 71% of this explanatory power (i.e.,
497 approximately 35% of the total variability in size distributions). Note that the correlation captures
498 concurrent shifts in the distribution of environmental conditions and euphausiid size structure, and does



499 not imply explicit co-location of euphausiids of a particular size with water of a particular temperature.

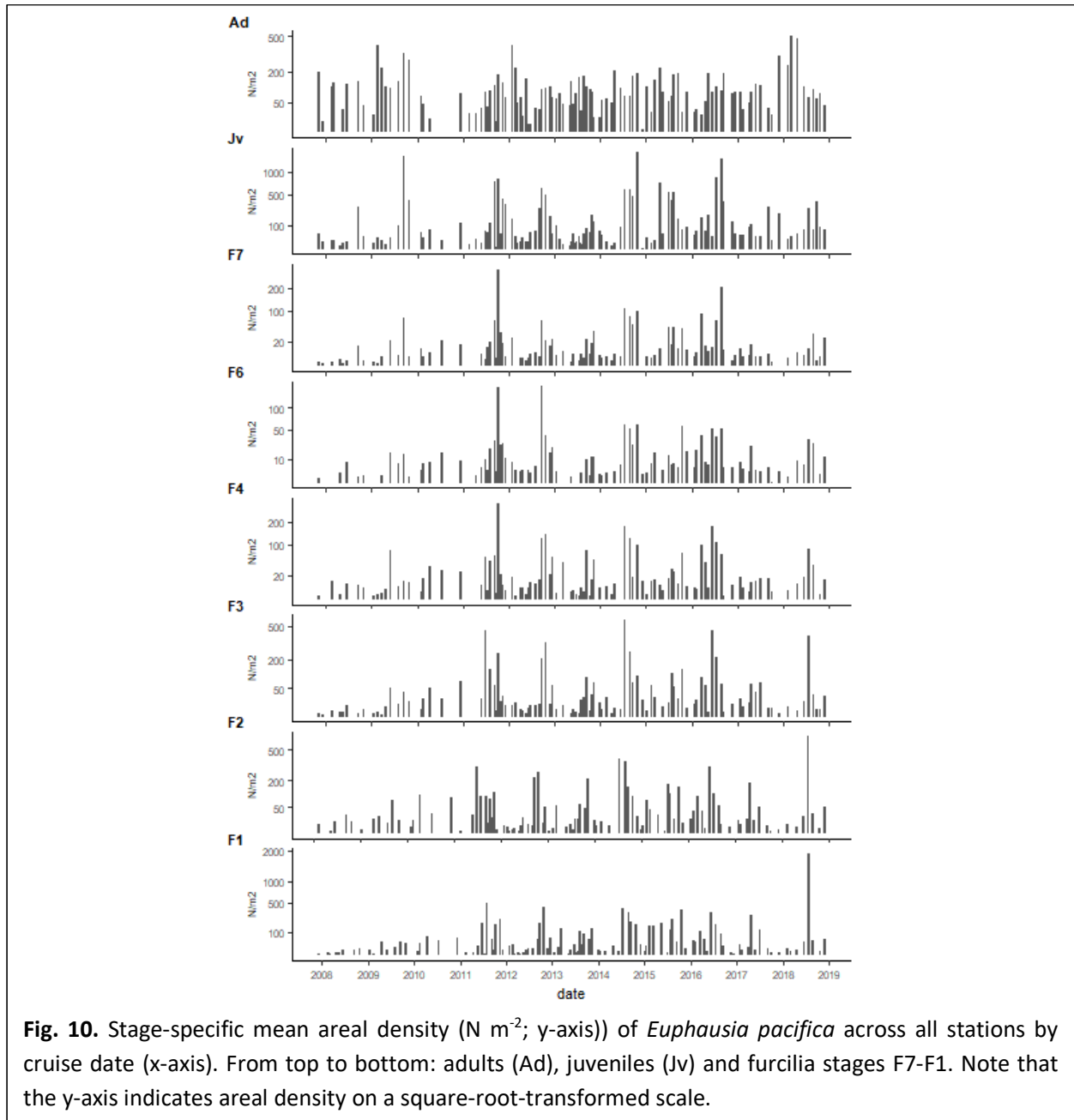
500

501 3.6 Abundance

502 Abundance of adult *E. pacifica* varies by over an order of magnitude over the course of the time series.

503 Adults are, in general, present year-round and a clear seasonal trend in abundance is not apparent.

504 Adults were absent from our collections on only rare occasions, most notably during a single cruise in
505 early 2016 that was strongly influenced by poleward transport associated with the 2015-16 El Niño.

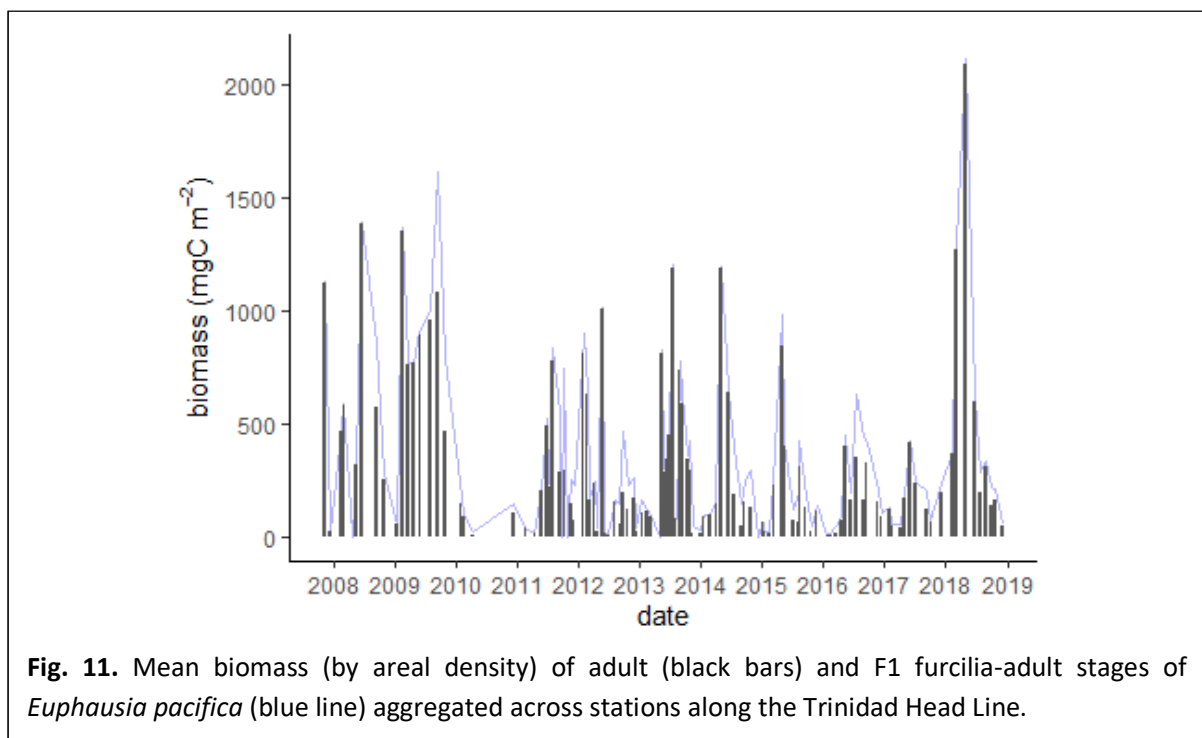


506
507 In contrast, immature life history stages exhibited much greater variability in abundance over time, (Fig.
508 10). All stages were encountered throughout the year across the time series, with strong seasonal
509 maxima in abundance typically observed during spring and summer. Abundance of furcilia, at least

510 during peak seasons, appears to have increased following the 2009-10 El Niño and to have declined
511 again during the 2015-16 El Niño. Following relatively high densities in summer 2016, furcilia abundance
512 was again relatively low throughout 2017. Except for a single cruise during summer, when densities of
513 F1-F3 furcilia were exceptionally high, furcilia abundance during 2018 remained relatively low.
514

515 3.7 Variability in individual and population biomass

516 The biomass (in $\mu\text{g C}$) of an adult *E. pacifica* of average size (2527 $\mu\text{g C}$) greatly exceeds that of an
517 average juvenile (332 $\mu\text{g C}$), intermediate furcilia (e.g., F4/5; 44 $\mu\text{g C}$), or early furcilia (F1; 8 $\mu\text{g C}$) (Table
518 1). Within the range of adult size, individual biomass can vary by a factor of 40 or more (e.g.,
519 approximately 225 $\mu\text{g C}$ at 6 mm to approximately 8700 $\mu\text{g C}$ at 19 mm). Accordingly, variability in
520 biomass of *E. pacifica* along the Trinidad Head Line is dominated by changes in adult abundance and
521 (especially) size (Fig. 11).



522
523 Biomass of adult *E. pacifica* tends to be relatively low during fall and winter months and high during
524 spring and early summer months, and responds sharply to rapid transitions in population size structure,
525 such as the sharp drop in biomass concurrent with the similarly dramatic decline in mean adult size
526 during the arrival of warm blob in late 2014. Except for a seasonal pulse related to strong upwelling in

527 2015, biomass remained relatively low throughout the 2014-16 MHW. A brief increase in biomass during
528 spring 2018 reflects unusually high abundance of adult *E. pacifica* (Fig. 10) coupled with an increase in
529 adult size following the MHW (Fig. 6).

530

531 **4. Discussion**

532

533 Through the development and analysis of detailed information on length, life-history stage, and
534 abundance, we document the strong effect of warm climate events on the size distributions of
535 *Euphausia pacifica* in coastal waters off northern California. This pattern was especially apparent during
536 the 2014 – 16 MHW, when we captured numerous mature individuals smaller than any previously
537 reported in the literature and larger size classes were conspicuously absent or rare. Further, by including
538 in our analysis assessment of earlier life history stages, we find that furcilia exhibit a contrasting
539 response to climatic variability, in which larger size classes of furcilia are associated with warmer
540 conditions.

541

542 Our analysis reveals that climate variability modulates seasonal trends in size distributions of *E.*
543 *pacifica* in coastal waters off northern California, and indicates that changes in both temperature and
544 chlorophyll concentration (or some unmeasured factor correlated with these parameters) drive these
545 dynamics. *We focus on the hypothesis that variability in temperature is the main driver of shifts in size*
546 *distributions across years. This is in part, a recognition that temperature is a powerful proxy for ocean*
547 *dynamics that influence ecosystem state, as well as a reflection of clear patterns in the data:*

548 *temperature varies dramatically across years, especially during non-upwelling seasons, whereas*
549 *chlorophyll concentration does not exhibit a consistent response to climate forcing (Fig. 5). Including*
550 *chlorophyll concentration as a covariate in models relating size to environmental conditions reflects the*
551 *common observation that, all else constant, more food typically supports greater growth. In contrast,*
552 *allowing the response to temperature to include random year-to-year (i.e., climate) effects is consistent*
553 *with the role of temperature in controlling organisms' metabolism, thereby determining the rate at*
554 *which individuals can grow when food is not limiting, and more generally, the amount of surplus energy*
555 *available from a given amount of food. Surplus energy, whether as an instantaneous measure or as a*
556 *predictor of future resources, is a powerful driver of life history decisions regarding size-at-maturation*
557 *and the allocation of energy towards growth or reproduction (Mangel and Ludwig, 1992, Siegel, 2000;*
558 *Poleck and Denys, 1982; Jager and Ravagnan, 2015; Constable and Kawaguchi, 2018). This effect of*

559 temperature on population structure appears to be fundamental to the patterns we observed in our
560 data.

561
562 Before proceeding, we note that various sources of potential sampling bias (e.g., net avoidance, vertical
563 migration) appear to have little substantive effect on our data and analysis. While mid- and inner-shelf
564 stations are sampled during daylight hours when net avoidance by larger euphausiids is more likely than
565 during night-sampling, adult *E. pacifica* are seldom found during either day or night in shelf waters (L.
566 Feinberg *personal communication* in Gómez-Gutiérrez et al., 2005; Bjorkstedt and Robertson, in prep.;
567 unpublished data from THL cruises when all stations were occupied during darkness). Moreover, shelf
568 samples containing adult *E. pacifica* and *T. spinifera* indicate our gear is capable of capturing adult
569 euphausiids during daylight hours when they are present in waters over the shelf (Bjorkstedt and
570 Robertson, in prep.). Because sampling at TH03 occurs just prior to darkness and the entire water
571 column is not sampled, underestimation of adult abundance may occur. However, consistent size
572 distributions between TH03 and offshore stations suggest samples from TH03 accurately represent
573 population size structure (Fig. 4).

574

575 **4.1 Seasonal, Interannual, and Ontogenetic Patterns**

576 Size distributions of adult *E. pacifica* reached seasonal maxima during spring and summer, coincident
577 with periods of higher productivity and conditions supportive of growth (Smiles and Pearcy, 1971;
578 Tanasichuk, 1998). Climate variability modulates this seasonal pattern, as demonstrated by sharp
579 declines in the size of adult *E. pacifica* – declines on the order of 30-40% of average size concurrent with
580 the arrival in late 2014 of unusually warm waters associated with the ‘warm blob’ (Bond et al., 2015).
581 The shift towards small size distributions of adult *E. pacifica* persisted into and through continued warm
582 conditions associated with the 2015-16 El Niño (Chao et al., 2017). Similar dynamics were reported from
583 collections along the Newport Hydrographic Line off Oregon (44.2°N), where in 2015 the average length
584 of adult *E. pacifica* was several millimeters smaller than any previous observations (Peterson et al.,
585 2017). Climate variability also drives more subtle variability in size distributions among years which
586 manifests as smaller sizes during years with delayed onset of upwelling (e.g., 2012; Wells et al., 2013) or
587 recovery from moderate El Niño conditions (e.g., 2010) and larger size during years marked by intense
588 upwelling (e.g., 2008, 2013). In general, these patterns are consistent with negative correlations
589 between individual adult *E. pacifica* weight and SST (Abraham and Sydeman, 2006), and similar, though

590 less clearly resolved responses in adult *Thysanoessa spinifera* and *Nematoscelis difficilis* from our
591 collections off northern California (unpublished data).

592

593 Our analysis revealed ontogenetic trends in correlation between *E. pacifica* body size and temperature:
594 adults and juveniles exhibited a clear shift to smaller size distributions under warmer conditions while
595 the opposite pattern was observed for furcilia. These patterns are consistent with stage-dependent
596 temperature-size relationships reported for diverse marine taxa, including crustaceans (Forster and
597 Hirst, 2012). These observations, in conjunction with the ability of adult *E. pacifica* to shrink in response
598 to food limitation or elevated temperature (Pinchuk and Hopcroft, 2007; Marinovic and Mangel, 1999)
599 suggest that *E. pacifica* has a highly plastic life history, in which the timing and size of stage transitions—
600 including, especially, maturation—are governed by a temperature-dependent balance between energy
601 intake and metabolism (Angilletta and Angilletta, 2009) in the context of size-dependent predation
602 (Peterson and Wroblewski, 1984) and reproductive output (Gómez-Gutiérrez et al., 2006).

603

604 The magnitude of change in biomass in response to temperature estimated in this study ($-24\% \text{ }^{\circ}\text{C}^{-1}$ for
605 adults) exceeds estimates of similar responses in marine crustaceans (maximum decline of 8%; $\sim 2\%$ for
606 *E. pacifica*; Forster et al., 2012). We suspect that part of this discrepancy is due to our use of average
607 water temperature over the entire water column (1:100 m max) rather than temperature specific to
608 where individuals reside. Consequently, estimates of percent change in relation to temperature
609 reported here should be treated with caution when compared to laboratory experiments.

610

611 **4.2 Mechanisms**

612 Several non-exclusive mechanisms offer plausible explanations for observed shifts in the size structure
613 of euphausiid populations in the upper water column off northern California, and may be broadly
614 categorized as those dominated by movement of water masses (i.e., advection or displacement) and
615 those arising from ecological interactions and physiological processes (Dorman et al., 2011). Here we
616 describe and evaluate these hypotheses, keeping in mind that the data available to us represent
617 snapshots of growth processes that integrate organisms' exposure to multiple factors (i.e., temperature
618 and food) over time in a highly dynamic environment.

619

620 Horizontal advection of warm water masses into the sampling region is, we believe, the dominant cause
621 of rapid shifts in size distributions of adult *E. pacifica* and in particular the presence of small individuals.

622 The underlying assumption that dissimilar water masses host populations with distinct size distributions
623 is consistent with generally strong biogeographic associations between water mass characteristics and
624 zooplankton assemblages or indicator species (Hooff and Peterson, 2006; Kiester et al., 2005). This
625 hypothesis draws further support from physical modeling studies (e.g., Chao et al., 2017) that indicate
626 anomalous onshore transport of warm water during the arrival of the warm 'blob' in 2014 and poleward
627 transport of warm water during the 2015-16 El Niño. Concurrent shifts in copepod assemblages, and the
628 sequential arrival of *Euphausia recurva* (in early 2015) and *Nyctiphanes simplex* (in late 2015), indicators
629 of onshore and poleward transport, respectively, in the northern CCE (Keister et al., 2005), further
630 corroborate the influence of anomalous transport on coastal waters (McClatchie et al., 2016; Wells et
631 al., 2017; Bjorkstedt and Robertson, in prep.). Warm water anomalies also brought a highly unusual
632 pelagic assemblage dominated by pyrosomes and other gelatinous taxa into coastal waters (Sakuma et
633 al., 2016).

634
635 We suspect that the disappearance of larger individuals from our samples is also, at least in part, driven
636 by vertical displacement associated with the arrival of warm water masses and depression of the
637 thermocline. If large adults are closely associated with cooler water, such displacements could push
638 large adults beyond the reach of our net tows. It is difficult to assess this possibility without additional,
639 deeper sampling, or possibly comparison to coincident acoustic observations. However, large adults
640 rapidly disappeared from our collections in late 2014, but made an equally rapid return and
641 disappearance during a brief window of strong upwelling in early 2015, suggesting that these individuals
642 had not been displaced substantial distances, nor had they suffered acute mortality.

643
644 Robust correlation in mean size across larval stages (Table 2) is also consistent with
645 advection/displacement hypotheses in that it suggests that individuals spanning several life history
646 stages have experienced similar conditions for periods extending from days to weeks, but is also likely to
647 reflect coherence in production and growth at seasonal scales (Fig. 5). Correlations tend to be strongest
648 between sequential larval stages, where time between stages is on the order of days (Ross, 1981;
649 Feinberg et al., 2006) and similar environmental exposure histories may be expected to yield parallel
650 growth patterns. The correlation between F1 furcilia and adults and breakdown of correlations between
651 adults and more advanced furcilia stages might reflect, at least in part, a divergence in environmental
652 conditions experienced over time, and is consistent with observations of declining spatial overlap
653 between adult and larval stage *E. pacifica* elsewhere in the California Current (Décima et al., 2010).

654

655 It is less clear that shrinking in response to warming temperatures or reduction in food availability
656 (Marinovic and Mangel, 1999; Pinchuk and Hopcroft, 2007; Shaw et al. 2010) or ecological interactions
657 such as changes in size-dependent mortality (starvation or predation), might substantively account for
658 our observations, particularly in the case of rapid shifts in size structure. On several occasions, however,
659 rates of decline in mean size of euphausiids in our samples exceeded rates of shrinking reported for
660 individual *E. pacifica* (Shaw et al., 2010; Marinovic and Mangel 1999; Pinchuk and Hopcroft, 2007). We
661 have no data with which to rigorously assess short-term variability in predation. Our record does not
662 span conditions similar to those observed in 2005, in which persistent suppression of upwelling
663 throughout much of the spring and early summer is hypothesized to have driven extensive starvation
664 mortality of krill in the CCE (Dorman et al., 2011).

665

666 Although rapid growth- or predation-dependent mechanisms seem less plausible as a cause of rapid
667 shifts in size-structure of *E. pacifica*, ecological and physiological processes must underpin persistent
668 differences in size-structure related to water mass characteristics, and thus establish conditions
669 necessary for transport processes to drive temporal variability at local scales as discussed above. At
670 present, we are unable to resolve whether the small juvenile and adult *E. pacifica* observed in our
671 samples had previously achieved a larger size and then shrank or had developed along a fixed
672 temperature-dependent trajectory (but see Shin et al., 2002). Such information would support a more
673 definitive assessment of underlying mechanisms. The capacity for krill to shrink during bouts of intense
674 reproduction suggests the potential for even more rapid changes in body mass than might be expected
675 through somatic metabolism (Feinberg et al., 2007).

676

677 Our analysis identified a relationship between chlorophyll concentration and euphausiid size that is
678 consistent with the hypothesis that euphausiid growth in the CCE is driven by upwelling-fueled primary
679 production (Shaw et al., 2010), and might partly explain more subtle changes in size structure from year
680 to year (e.g., in response to the delayed onset of upwelling in 2012; Wells et al. 2013). However, high
681 chlorophyll concentrations during the warm anomaly were commonly associated with persistent,
682 recurring blooms of *Pseudo-nitzschia spp.* that generated high concentrations of domoic acid along the
683 THL during summer months from 2014 through 2016 (unpublished data), reaching record levels during
684 an extensive and persistent harmful algal bloom in 2015 (McCabe et al., 2016). Sustained exposure to
685 domoic acid have been shown to suppress feeding in euphausiids (Bargu et al., 2006), with obvious

686 consequences for growth rates. The potential for an interaction between elevated temperatures and
687 changes in food quality in driving the decline in adult size during the 2014-16 MHW cannot be ruled out,
688 leaving open the possibility that variability in size reflects a stronger and more variable influence of
689 trophic mechanisms at longer (interannual) time scales.

690

691 **4.3 Ecological Implications**

692 Regardless of specific mechanisms and drivers, changes in the size structure of critical forage species like
693 *E. pacifica* have the potential to trigger ecological responses in coastal waters of the CCE. Sustained
694 reduction in average size and biomass of *E. pacifica*, such as we observed during the 2014-16 MHW,
695 almost certainly reduced the availability of *E. pacifica* as forage for diverse predator species ranging
696 from generalists such as salmon or rockfishes to those that more specifically target euphausiids as prey
697 (e.g., some baleen whales and seabirds). Moreover, the general shift towards smaller size across the
698 population means that what biomass is available to predators comes in smaller packages, which
699 influences how efficiently predators, especially those that target individual prey, accumulate energy
700 (e.g., Lovvorn, 2010). Consequences for higher trophic levels can be profound. For example, a sustained
701 shift in the structure and distribution of euphausiid populations is suspected to have played a role in a
702 massive mortality event affecting Cassin's auklet (*Ptychoramphus aleuticus*) from northern California to
703 Alaska during the 2014-16 MHW (Jones et al., 2018). Moreover, seasonal trajectories of *E. pacifica* size
704 distributions appear to have been set early in the year, so that the size of adult *E. pacifica* during the
705 height of the upwelling season depends strongly on conditions during the preceding winter, whether
706 through persistent changes in ecosystem productivity and environmental conditions, as a consequence
707 of intrinsic demographic or physiological inertia, or arising from some other mechanism(s). From the
708 perspective of understanding impacts on higher trophic levels, our findings enrich support for
709 “preconditioning” hypotheses that emphasize the importance of environmental conditions in late
710 winter—prior to the onset of sustained seasonal upwelling—as a key determinant of productivity in the
711 California Current (Logerwell et al., 2003; Black et al., 2010; Schroeder et al., 2013).

712

713 Variability in the size structure of *E. pacifica* (and other euphausiids) also has the potential to impact
714 species upon which euphausiids prey, including copepods (Ohman, 1986; Stuart and Pillar, 1990) and
715 eggs and larvae of marine fishes (Bailey, et al. 1993, Theilacker, et al. 1993). Based on evidence that the
716 susceptibility of larval fish to predation depends on the size of the predator relative to its prey (Paradis

717 et al. 1996), changes in the average size of adult *E. pacifica* off northern California might contribute to
718 substantial variation in predation pressure on larval fishes in coastal waters.

719

720 **4.4 Connecting data to models**

721 Rigorous assessment of mechanisms that link population dynamics to environmental conditions will
722 require models that integrate the concurrent effects of multiple factors (i.e., temperature and food) on
723 the growth and development of *E. pacifica*. Coupling individual-based models and life history theory to
724 realistic models of ocean circulation and productivity offer a way to examine this dynamic under
725 different conditions and to parse out the contributions of advection versus production (Dorman et al.,
726 2015b). In more complex frameworks, modeling exposure to currents and predators will require
727 accounting for size-dependence in vertical distributions and migration (Duarte, 2007; O'Connor et al.,
728 2007). The validity and utility of such approaches, however, will depend on accurately representing
729 plasticity in the growth and maturation responses of *E. pacifica* to environmental conditions. Successful
730 development and evaluation of such models depends on high-quality empirical data, whether derived
731 from laboratory experiments (e.g., Ross 1979; Feinberg et al., 2006) or from careful analysis of samples
732 collected in the field to complement existing observations on spatial distributions (Gómez-Gutiérrez et
733 al., 2005; Lamb and Peterson, 2005; Lindsey and Batchelder, 2011)

734

735 This study clearly illustrates the potential return on such effort. Our data set is one of few that provide
736 detailed information on size distributions across a broad range of developmental stages of *E. pacifica*
737 (Boden, 1950; Brinton and Wyllie, 1976; Ross 1979). Moreover, detailed morphological analysis revealed
738 shifts in stage-specific size distributions of adults and juveniles that would have been misinterpreted as
739 changes in abundance for each life history stage had stage-designations been based solely on size
740 thresholds. Over the course of the time series, about 3% of adult *E. pacifica* were below minimum size
741 thresholds reported in the literature and often applied to delineate adults from juveniles in analysis of
742 plankton samples (e.g., 10-11 mm TL or about 8 mm BL; Brinton, 1976; Brinton and Wyllie, 1976;
743 Bollens, 1992; Ju et al., 2009). Conversely, about 26% of juveniles exceeded size thresholds and would
744 have been designated as adults. These proportions vary as size distributions shift over time. For
745 example, for 2013, 1% of adults and 34% of juveniles would have been misclassified, whereas 11% and
746 12% of adults and juveniles, respectively, would have been misclassified during the 2014-16 MHW.
747 Actual discrepancies in population structure will depend on the (true) ratio of adults and juveniles in the
748 population. Errors in stage designation will not affect estimates of biomass that do not discriminate

749 among stages or that are based on individuals that satisfy length criteria. Morphological stage-
750 designations, however, appear to be essential if data are to provide unbiased information to support
751 demographic models for euphausiid population dynamics or studies focused on understanding how life
752 history strategies and evolution are affected by environmental and climate variability.

753

754 **5. Conclusion**

755

756 Warm water events are expected to increase in frequency and duration with ongoing climate change
757 (Frölicher et al., 2018; Oliver et al., 2018). We show that such an event can drive strong shifts in size-
758 distributions of a dominant euphausiid species in coastal waters off northern California. Moreover, our
759 data confirm that warming events can disrupt or suppress typical seasonal patterns (i.e., the tendency
760 for mean size of adult euphausiids to increase through the spring and summer upwelling season),
761 leading to persistent transitions towards adult and juvenile length distributions dominated by smaller
762 size classes. Given the central role of *E. pacifica* in trophic dynamics of the CCE, and the sensitivity of
763 several higher trophic levels to the timing and magnitude of prey production (Abraham and Sydeman,
764 2004; Wolf et al., 2010, Wells et al., 2012), such dynamics have profound implications for ecosystem-
765 level responses to warming events and climate change in general. Accordingly, time series of adult *E.*
766 *pacifica* size have been incorporated into NOAA's California Current Integrated Ecosystem Assessment
767 as an indicator for lower trophic level productivity and ocean-ecosystem health (Harvey et al., 2019).
768 More generally, our results complement other work (Peterson et al., 2017; McCabe et al., 2016; Jones et
769 al., 2018; Cavole et al., 2016) in fulfilling the need for information on a multitude of processes, ranging
770 from individual to ecosystem, to understand marine ecosystem responses to climate and environmental
771 variability (Beaugrand and Kirby, 2018).

772

773 These insights would not be possible without detailed analysis of plankton samples, which highlights the
774 unique capability of long-term net sampling to reveal important biological signals not yet resolved by
775 other zooplankton sampling methods. Data from careful taxonomic and morphological analysis of in-
776 hand samples contain information essential for corroborating and enriching more extensive sampling
777 methods (e.g., acoustics and optics) in terms of community composition, demographic and size
778 structure, and for advancing the development of population and ecosystem models. For this reason,
779 high frequency, long-term sampling as part of a comprehensive observing program presently remains

780 essential for resolving the response of euphausiids and the broader zooplankton community to climate
781 variability and understanding their roles in dynamic ecosystems (Bjorkstedt and Peterson, 2015).

782

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794

795 **Appendix A. Supplemental Material**

796

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