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Technical Memorandum NWS NMC 63



DAY-NIGHT DIFFERENCES IN RADIOSONDE OBSERVATIONS OF THE STRATOSPHERE AND TROPOSPHERE

Washington, D.C. September 1979



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ABSTRACT. Day-night differences in temperature and height are used as the basis for a system designed to achieve compatibility between data measured by various radiosonde instruments. The resultant compatibility adjustments in reported data are a prerequisite to the analysis of stratospheric constant-pressure charts above the 100-mb level, and for some instrument types they are significant even at tropospheric levels.

1. INTRODUCTION

In the 1950's, when great numbers of radiosonde balloons were able to ascend for the first time above the 100-mb level and to provide stratospheric data on temperature and geopotential height, large differences became apparent between daytime and nighttime observations. Incompatibility was also observed between temperatures and heights measured during daylight hours by different radiosonde instruments used in adjacent countries. Teweles and Finger (1960) presented evidence that such irregularities at stratospheric levels were largely fictitious. A number of other studies indicated that the problem was most likely due to varying responses of different temperature sensors to solar radiation (Hayashi et al. 1956; Scrase 1956; Badgley 1957). Finally, Finger and McInturff (1968) showed decisively that the true diurnal temperature range of the ambient stratosphere at middle latitudes is much too small to account for the observed discrepancies. Values for this temperature range are approximately 1.0°C at 10 mb, and decrease with decreasing height to 0.5°C at 30 mb; these values are lower, often by an order of magnitude, than the day-night differences that were reported by stations throughout the Northern Hemisphere.

Since the mid-1950's, several national meteorological agencies have developed systems to correct measured temperatures and computed heights at individual stations. Such efforts, however, have generally failed to solve the problem of incompatibility between stratospheric observations. Correction systems based on laboratory tests invariably fail on account of the impossibility of simulating real atmospheric conditions.

McInturff and Finger (1968), following Hawson and Caton (1961), showed that compatibility between instrument types and between daytime and nighttime soundings could be improved through the application of mean day-night differences determined as functions of instrument type and of mean daytime solar elevation angle. A study involving 33 months of twice-daily radiosonde observations from 1964-66 provided the basis for sets of adjustment coefficients dependent on solar elevation angle. (These served to reduce sunlit observations to equivalent nighttime observations for levels from 100 to 10 mb.) This empirical adjustment scheme was used operationally in the Upper Air Branch (UAB) of the National Meteorological Center (NMC) until its recent replacement by the results reported here.

In 1977 it was decided that the problem of day-night differences in radiosonde observations would be reexamined. Reasons for this decision included: (1) developments in radiosonde instrumentation and reduction techniques which have rendered some of the old empirical adjustments obsolete; (2) changes in use of instruments by different national meteorological services, especially in parts of the world where high solar elevation angles might result in large empirical adjustments; (3) requirements within the UAB for extension of the stratospheric (100 to 10 mb) analysis system to the tropics and the Southern Hemisphere (areas of the globe from which little information was used in development of the original empirical adjustment scheme); (4) application of the correction scheme in tuning the regression system used in operational reduction of Vertical Temperature Profile Radiometer (VTPR) observations by the National Environmental Satellite Service (Werbowetzki 1975).

As in the earlier study (McInturff and Finger 1968), it is stressed that attainment of compatibility does not ensure accuracy. For further discussion of this point, as well as more background material on high-level meteorological observations, see Finger et al. (1978), which includes results obtained by Spackman and others in the United Kingdom.

2. TYPES AND NUMBERS OF RADIOSONDES UNDER STUDY

The present study is confined to the following radiosonde instruments: Finnish Vaisala, French Mesural, British Kew, West German Graw, U.S.S.R. A-22, U.S.S.R. RKZ, Japanese "codesending", U.S.A. NOAA, U.S.A. AN/AMT4, Chinese, Swiss, Sangamo (employed by Canada and Portugal), and Australian. Many of these instruments are used outside the country or countries in which they are manufactured, and they are often referred to by different names. For example, the Vaisala instrument is widely used throughout the Scandinavian countries, the Middle East, and South America; the Mesural is used throughout most of the French-speaking world; the Australian instrument is used also in New Zealand; and the U.S.A. AN/AMT4 is still widely used in the countries which at one time contained U.S. military bases. The instrument types listed above account for at least 95% of the soundings that attain levels above 100 mb.

Several instrument types included in the present study were not covered in the 1968 report by McInturff and Finger (example: Australia/New Zealand). In some cases, as in those of the French Mesural and the U.S.S.R. RKZ, the instrument was either still under development or else was not widely enough deployed during the years 1964-66 to provide a sufficiently large statistical sample. A significant gap in worldwide stratospheric radiosonde coverage has been filled with the acquisition of Chinese data. Unfortunately, the data for levels above 100 mb were not received until after the cut-off date for the sample used here.

Figure 1 and figure 2 show the distribution of upper air stations throughout the Northern Hemisphere and throughout the Southern Hemisphere, respectively. Table 1 provides a summary of current knowledge concerning deployment of instrument types. The remark made in 1968 concerning a similar figure and a similar table is applicable here as well: these representations cannot be definitive, in view of the changes that are continually taking place in the network.

A comparison of figures 3 and 4 shows what areas and on which dates various observation points have daylight and darkness. The sunset and sunrise lines are depicted for 0100 GMT (fig. 3), December 15 and June 15, and the sunset and sunrise lines for the sames dated at 1300 GMT (fig. 4). These times are those at which the radiosonde balloon with nominal 0000 GMT and 1200 GMT observation times will normally reach the 10-mb level. The migration of the sunset and sunrise lines for March 15 and September 15 would lie midway between the lines for June and December. It should be noted that some countries (e.g., the United Kingdom) are always in darkness at 0100 GMT and always in daylight at 1300 GMT. Other countries (e.g., the U.S.A.) have large areas where in summertime daylight occurs at both observation times, and in wintertime darkness occurs at both observation times. Such areas and periods of so-called "double daylight" and "double darkness" cannot be used for computing day-night differences. However, there are sufficient areas and adequately long periods wherein one observation is in daylight and the other, 12 hours later, is in darkness, to permit the calculation of vast numbers of day-night differences in temperature and geopotential height.

Table 2 is a summary of numbers of observed day-night height differences. It is important to keep these numbers in mind, not so much for the information they provide on the decline in quantity of data at levels above 850 mb, but for the interpretation of the scatter-diagrams and the applications of the tables of adjustments, which will be discussed further on. Sample size is crucial in determining confidence-levels, so it can be seen, for example, that one can place less reliance on any adjustment for the Mesural instrument used overseas (with only 356 day-night differences for a 19-month period at 100 mb) than on the Mesural instrument used in Metropolitan France (with 2,157 day-night differences at 100 mb for the same 19-month period).

The question of pretransmission corrections was dealt with in detail by McInturff and Finger (1968). In the perspective of the present work, this question appears to have little importance, since all instruments require post-transmission corrections for solar-induced day-night differences. Thus, even though pretransmission corrections in most cases result in dramatic reductions in day-night differences, adjustments based on studies such as the present one are still necessary. Anyone interested in pursuing the question of pretransmission corrections may refer to the paper by McInturff and Finger (1968). For information on changes made by particular countries, specialized publications by various meteorological agencies and radiosonde manufacturers may be consulted (see, for example, Suzuki and Asahi (1978)).

3. PROCESSING AND ANALYSIS OF TEMPERATURE AND HEIGHT DATA

3.1 Data Processing

The data base for developing the empirical adjustment scheme presented here is the archive of twice-daily rawinsonde reports received operationally at the NMC, together with calculated mean monthly solar elevation angles at various atmospheric levels based upon assumed constant launch times and balloon ascent rates (see the appendix for details of solar elevation angle computations). The calculation scheme (fig. 5) was applied to 19 months of data from the period 1974-76.

The levels for which calculations were made are as follows: 1000, 850, 700, 500, 400, 300, 250, 200, 150, 100, 70, 50, 30, 20, and 10 mb. Monthly means of 12-hour day-night temperature changes (ΔT) were computed from every reporting upper-air station in the Northern Hemisphere, also for all such stations in Australia and New Zealand.

We write

$$\Delta T = T_d - T_n$$

where T_d is the temperature obtained in daylight and T_n the temperature obtained at night; the daylight observation occurs at either 00 GMT or at 12 GMT, depending on geographical location. The monthly-mean temperature difference for a particular station is given by

$$\overline{\Delta T} = \sum_{i=1}^{N} (\Delta T)_i, \text{ where } 1 \leq i < N$$

and N may vary from 1 to 60, depending upon the number of observations. This range of N results from an attempt to obtain the maximum number of observed 12-hour temperature differences, thus aiding in filtering out the effects of moving weather systems and other long-term trends. Accordingly, ΔT was determined for both the 12 hours preceding and the 12 hours following the time of T_d. The same procedure was followed in computing each mean height difference given by

$$\overline{\Delta H} = \sum_{i=1}^{N} (\Delta H)_i, \text{ where } 1 \leq i \leq N$$

and where $\Delta H = H_d - H_n$ (H_d is the height of a particular isobaric surface corresponding to T_d and H_n is the height corresponding to T_n).

To guard against the effects of gross errors within the data, any individual ΔT or ΔH which differed from ΔT or $\overline{\Delta H}$ by more than twice the standard deviation was rejected and a new monthly mean was calculated.

3.2 Plotting and Analysis Procedure

The output of the computer program is plotted automatically on graphs which show day-night differences in temperature and height (ordinates) versus mean daytime solar elevation angle (abscissas). The graphs (scatter diagrams) were prepared for all the pressure surfaces. All stations using a particular type of sonde were grouped together in a set of scatter diagrams (one for each of AT, AH at each pressure level). The analysis procedure was governed by the following principles: (1) In order for a monthly-mean value to qualify for representation on a scatter diagram, sufficient pairs of successive observations must be available at a particular station to give more than five temperature- or height-differences for the individual month. (2) Each monthly-mean value ΔT or of ΔH based on sample size n > 10 is marked by an asterisk in the scatter diagram, while each value based on a sample size n < 10 is marked by an oval. (3) The abscissa is divided into 10° intervals of solar elevation angle, whereas the ordinate ΔT is divided into 1°C-intervals and the ordinate ΔH into 10-meter intervals. (4) Averages for all the daynight differences are computed within each 10° interval of solar elevation angle, and are indicated in each interval by a plus-sign (+) or a minus-sign (-); the plus-sign is used for means calculated before data-rejection, and the minus-sign for means calculated after data-rejection. (If no data are rejected, only the plus-signs appear. If some data are rejected, normally both a plus- and a minus-sign appear in the column affected, except when they would be so close together as to interfere with each other.) (5) The standard deviations of all the day-night differences are computed within each 10°-interval of solar elevation angle, and an envelope with half-width equal to twice the standard deviation and whose upper and lower boundaries are marked by 'S' or 'T' is indicated; 'S' is used in case no data are rejected, 'T' is used in case some data are rejected in the column under consideration (in which case the value of the recomputed standard deviation is In case data are rejected for a particular column, both 'S' and 'T' used). will normally appear. Sometimes values of 'S' or 'T' will be so large as not to appear at all on the scatter diagrams, although usually they appear as marking the upper boundaries of the 20-envelopes. Most often the space left at the bottom of the figure is inadequate to accommodate the symbols 'S' and 'T.'

Examples of the computer-plotted output are shown in figures 6 to 12.

4. DISCUSSION OF SCATTER DIAGRAMS

Although all the empirically derived adjustments, for levels 700, 500, 300, 200, 100, 50, 30, 20, and 10 mb, will be presented as tables (in a form suitable for ready insertion into computer programs), it is instructive to examine a few of the scatter diagrams. It is not appropriate to present all of them since most of the information contained in these is presented in other forms throughout this paper. The diagrams shown here were chosen for no other particular reason than that each of these illustrates at least one important point. An understanding of these examples will facilitate the interpretation and application of the tables which follow.

Figure 6 is the diagram for the day-night differences in temperature at 200 mb (0000 GMT sunlight) for the U.S.A. NOAA (VIZ) instrument. The amount of data represented can be estimated from the knowledge that each asterisk and oval-shaped symbol stands for an entire monthly-mean day-night temperature difference for a single station; hence each symbol is derived from more than 5 and less than 60 observations, in accordance with the convention explained in the previous section. The decrease in the amount of data with height, and the increase in scatter for this particular instrument, can be estimated by comparing this figure with that for 10 mb (fig. 12). The amount of adjustment needed is also seen to be much less at 200 mb than at 10 mb. However, there is still a slight adjustment to be made even at 200 mb, where it is approximately 0.5°C for most daytime solar elevation angles.

The visible scatter in these diagrams is due to intermonth, interstation variability, since each symbol represents a monthly mean for one station of day-night temperature differences. Each symbol therefore corresponds to a stationmonth, and the amount of scatter of these symbols can provide a measure of intermonth, interstation variability. The intramonth, intrastation variability is of course determined by the second moment of daily difference values about their monthly mean; this variability is hidden from us in these diagrams. However, both intramonth, intrastation variability and intermonth, interstation variability are taken into account in computing standard deviations of day-night differences for various instrument types; these standard deviations will be discussed further on. For a more complete discussion of intramonth, intrastation variability and of intermonth, interstation variability, see McInturff and Finger (1968).

Figures 7 and 8, for metropolitan France and for overseas stations making use of the French instrument (the Mesural), respectively, both groups employing (presumably) the same type of instrument, illustrate how different the results can be in spite of supposed similarities in equipment. The scatter in figure 7 (as measured, for example by the distance between the mean and the S-symbols near the top) is much greater than in figure 8; also the sample used for figure 7 is much larger than that used for figure 8. Of course there is always the possibility that the overseas stations are using some of the old rawinsonde equipment, manufactured before the thermistor was changed (1970); in this case, figure 7 and figure 8 would reflect differences between two instrument types.

Figure 9 is an example of a scatter diagram in $\overline{\Delta T}$ for the Chinese instrument at 100 mb. It may be readily seen that the mean value of $\overline{\Delta T}$ hardly ever deviates significantly from zero. The sample size is large, but because of Earth-Sun geometry the data are restricted to low daytime solar angles.

Figures 10 and 11 are scatter diagrams of $\overline{\Delta T}$ for 30 mb generated by data from the West German and Japanese radiosonde instruments, respectively. The amounts of scatter are similar on the two diagrams. The one for the Federal Republic of Germany (fig. 10) shows the means of the $\overline{\Delta T}$'s (represented by the dash-symbol) constituting approximately a straight horizontal line; this is due to the fact that pretransmission corrections based on sets of historical day-night difference data have been applied. Ideally, the means of the ΔT 's would be zero; the fact that they are not suggests that the pretransmission adjustments are not quite so large as they should be.

On the other hand, the 30-mb scatter diagram of $\overline{\Delta T}$ for the Japanese instrument (fig. 11) shows unmistakable evidence of a pretransmission correction system which is adequate at daytime solar angles above 60° but not quite adequate at lower solar angles.

Figure 12, showing the ΔT -scatter diagram for the NOAA (VIZ) instrument at 10 mb, has already been compared with figure 6. 10 mb is the highest level for which we present results. The reason is not hard to find: data are already quite sparse at 10 mb (as evidenced by the preponderance of oval symbols), and become much too sparse at higher levels. The reader should compare figure 12 with figure 11 in McInturff and Finger (1968). (S)He will note the same general configuration of the curve of mean ΔT , indicating that this particular instrument is performing at 10 mb in much the same manner as it did 10 years earlier.

5. UNEXPLAINED VARIABILITY

As already indicated by McInturff and Finger (1968), at least five sources contribute to the standard deviations of ΔT about the mean of the day-night differences (for a standard height, a given daytime solar elevation angle, for any of the instrument types under study): (1) Differences between individual sondes of a given type and between items of ground equipment; (2) differences in station procedure, even for several stations using the same type of equipment; (3) day-to-day changes in the albedo of Earth and cloud, which cause variations in the amount of reflected sunlight reaching the radiosonde; (4) synoptic changes; and (5) the true diurnal temperature variation. Since it is impossible to separate out all these influences, or even to measure some of then precisely, it seems justifiable to combine them as factors in the unexplained variability of observations.

The standard deviations σ of $(\Delta T)_i$ about the average of the monthly means for all 10°-interval of daytime solar elevation angle have been calculated (for the 30-mb level). The results are shown in figure 13, where σ is the result of averaging the σ 's over all the intervals of daytime solar elevation angle. These results should be compared with those shown in figure 30 in the paper by McInturff and Finger (1968). In the latter, it should be emphasized, we dealt with average variances, for reasons given in the text. In the present study, owing to a higher degree of automation in dataprocessing, it was easy enough to calculate the standard deviations in a more straightforward fashion.

Figure 13 contains many interesting features, but perhaps the most significant is the behavior of the French instrument in the country of its manufacture in comparison to its behavior overseas; and the analogous behavior of the Vaisala instrument in the country of its manufacture (Finland) in comparison to its behavior in other countries. The utility of any posttransmission adjustment is inversely proportional to the σ -value associated with it.

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6. CONCLUDING REMARKS

Tables 3 to 19 summarize all the information from the scatter diagrams needed for adjustments in rawinsonde-reported temperatures and geopotential heights of constant-pressure surfaces. In this form, they are easily incorporated into programs for analyzing tropospheric and lower stratospheric data. The greater need for adjustments at stratospheric levels is clearly in evidence.

We recommend that the adjustment system for upper-air data presented here be applied only at analysis centers. Certainly it should not be applied as a pretransmission correction scheme in any one country; this would make it difficult to determine which stations require further adjustment and which do not.

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APPENDIX: COMPUTATION OF SOLAR ELEVATION ANGLE

The solar elevation angle α is calculated for the time at which the radiosonde balloon passes through each mandatory level at 100 mb and above. It is found from the formula:

 $\sin \alpha = \sin \phi \sin \delta + \cos \phi \cos h \cos \delta$

where ϕ is station latitude, δ is solar declination angle, and h is solar hour angle. The angle δ is obtained by 1

 $\sin \delta = \sin (23^{\circ}26'37.8'') \sin \sigma$,

where σ is in degrees and is given by

 $\sigma = 279.9348 + d + 1.914827 \sin d - 0.079525 \cos d + 0.019938 \sin 2d - 0.001620 \cos 2d:$

d is the number of the day in the year minus one, multipled by the constant 0.98565; e.g., for January 30, d = 29 x 0.98565. The solar hour angle h, the angular (longitudinal) distance of the sun from the observation point, can be expressed in terms of time of observation and longitude relative to Greenwich. The relation takes the form

h (deg.) = 15 (C + H - M) - L

where M, the time of meridian passage or true solar noon, is given by

M (hr.) = $12 + 0.123570 \sin \phi - 0.004289 \cos d$ + 0.153809 sin 2d + 0.060783 cos 2d.

with d as defined above; C (in hours) is a function of the difference between actual radiosonde release time and nominal observation time; and also of balloon ascent rate (table 25); H is nominal observation time, expressed as the number of hours after 0000 GMT; L is longitude of station is degrees and tenths, counted positive west of Greenwich.

Since station latitude and longitude are known, and solar declination angle and time of meridian passage can be determined with high precision, the only source of uncertainty in the calculation of solar elevation angle is the time of radiosonde arrival at each mandatory level. As no indication of this parameter is given in the coded rawinsonde message, an approximate release time and ascent rate must be assumed. From an inspection of individual station records, "normal" release times, to the nearest quarter hour, and typical rates of ascent have been determined for North American stations. However, since these records are not available for most of the remaining stations in the Northern Hemisphere, a release time of 20 minutes prior to nominal observation time is utilized.

 $^{^{1}}Relations$ involving δ and σ were derived from information provided by the U.S. Naval Observatory, Washington, D.C.

Table 1.--Types of radiosondes in use in the world. This representation cannot be considered current, in view of the changes that are continually taking place in the network.

Instrument type	Where employed	Pretransmission corrections applied?
Bendix-Friez duct- type (403 NHz)	Brazil Indonesia	No No
Chinese	People's Republic of China	Unknown
Diamond Hinman	Australia New Zealand	No No
Freiberg	German Democratic Republic (09486 only)	No
Graw M-60	Belgium Congo Federal Republic	No No Yes
	of Germany Mauritius Zaire	No Unknown
Indian	India	Yes
Japanese "code- sending" Type RSII56	Indonesia Japan	No Yes
Kew (Mark IIB)	Cyprus Gibraltar Ireland Malta Netherlands United Kingdom	Yes Yes Yes Yes Yes Yes

Table 1 (continued)

		Pretransmission
Instrument type	Where employed	corrections applied?
Mesural	Algeria	No
	Central African Empire	No
	Chad	No
	France	No
	Ivory Coast	No
	Mali	No
	Madagascar	Unknown
	Mauritania	No
	Morocco	No
	Niger	No
	Senegal	No
	United Republic of	No
	Camoroon	NO
	Vietnam	No
Pakistani	Pakistan (41594, 41675, 41756)	Unknown
Sangamo	Canada	N-
bailgamo	Bontucal	No
	Portugal	No
Swiss	Switzerland	Yes
U.S.A. AN/AMT-4	Austria	No
	Bahamas	No
	Bermuda	No
	Egypt	No
	Greece	No
	Greenland (04202 only)	No
	Iceland	No
	Italy	No
	Korea (Republic of) (47138)	No
	Netherlands Antilles	No
	Pakistan (41530 and	No
	41780 only)	NO
	Spain	No
	Taiwan	No
	Turkov	No
	Vietnem	NO
	Vicelant	NO
	TUgostavia	NO

Table 1 (continued)

		Pretransmiss	ion
Instrument type	Where employed	corrections ap	plied?
W. G. A. NOAA	Ango10	Unknown	
U.S.A. NOAA	Colombia	No	
	Costa Picc	NO NO	
	Cube (Cuentanamo Bay NAS)	No	
	Cuba (Guantanamo Bay NAS)	No	
	Dominican Republic	No	
	Egypt (62378 and 62414)	No	
	Guadeloupe (France)	NO	
	Guam	NO	
	Guatemala	NO	
	Honduras	NO	
	Israel	NO	
	Jamaica	NO	
	Korea (Republic of)	No	
	Mexico	No	
	Mozambique	Unknown	
	Netherlands	No	
	Panama	No	
	Portugal (08509 only)	No	
	Spain (08001 and 08302)	No	
	Trinidad	No	
	U.S.A.	No	
	Venezuela	No	
IL S. S. R. A-22	Afghanistan	No	
0101011111111	Bulgaria	No	
	Czechoslovakia	Yes	
	Hungary	Yes	
	Poland (12425)	No	
	Romania	No	
	U.S.S.R.	Yes	
U.S.S.R. RKZ	German Democratic	Unknown	
	Kepublic	Unknorm	
	Hungary Poland (12330 and 12374)	Unknown	
	U.S.S.R.	Yes	

Instrument type

Väisälä

	Pretransmission
Where employed	corrections applied?
Argonting	Voc
Brazil	Ies
Burma	Voc
Denmark	Vec
Ethiopia	Ves
Finland	Vec
Greenland (except 0/202)	Ves
Hong Kong	Vec
Indonesia	Voc
Trad	Ves
Iran	Vec
Jordan	Ves
Kenva	Ves
Lebanon	Vec
Libya	Voc
Malaysia	Ves
Nigeria	Unknown
Norway	Voc
Philippines	Vec
Saudi Arabia	Voc
South Africa	Vec
Sudan	Vec
Sweden	Ves
Svria	Ves
Tanzania	Ves
Thailand	Ves
Tunicia	Ves
Ilganda	Vec
Zambia	Voc
L'and La	165

Instrument type	-10-0	0-10	10-20	20-30	ar angle 30-40	40-50	<u>(s)</u>	60-70	70-80	80-90	Total	85(as
W. German Graw 850 mb 100 mb 10 mb	000	001	605 612 99	824 654 147	693 567 167	750 606 192	467 224 55	242 145 0	000	000	3581 2808 661	r .
U.K. Kew 850 mb 100 mb 10 mb	000	333 188 38	1592 1399 142	1192 1042 153	1376 1081 111	1521 1292 152	631 541 33	282 224 0	63 33 0	000	6990 5800 629	
French Mesural (metropolitan France) 850 mb 100 mb 10 mb	000	000	148 86 62	662 578 81	563 471 53	526 331 127	634 477 87	274 214 13	000	000	2807 2157 423	
French Mesural (overseas) 850 mb 100 mb 10 mb	00	00	00	22 19	189 48 INSUFFIC	131 81 IENT DAT ¹	43 84	116 77	94 47	00	356	
Finnish Väísälä 850 mb 100 mb 10 mb	1400 1216 38	3625 3368 108	4075 3582 119	4372 2715 147	3559 2436 118	2333 1762 45	1766 697 4	1251 436	541 53 2	101 0 0	23023 16265 586	
Japanese "code-sending" 850 mb 100 mb 10 mb	000	172 2 0	4190 494 0	2905 3150 19	4176 2375 35	943 3370 35	291 835 37	0 144 8	000	000	12677 10370 134	
U.S.S.R. A-22 (after- noon daylight) 850 mb 100 mb 10 mb	2633 1927 167	3227 2850 212	3916 2878 225	3182 1808 156	2445 1090 75	1617 317 42	712 158 0	133 0 0	000	000	17865 11028 877	

Table 2.---Numbers of observed height differences

					1 104 1000	יובדמיור מ	riterence	ss (conti	(panur			
Instrument type	-10-0	0-10	10-20	20-30	lar angle 30-40	e (degree	<u>50-60</u>	60-70	70-80	80-90	Total	850-mb as 100%
U.S.S.R. A-22 (morning day1ight) 850 mb 100 mb 10 mb	3636 2552 340	3987 2989 527	2666 2712 576	1925 1831 468	839 1151 282	487 444 105	64 257 26	0001	000	000	13604 11936 2343	1002 882 172
U.S.S.R. RKZ (afternoon daylight) 850 mb 100 mb 10 mb	1463 2378 144	3183 4207 241	5870 4682 287	6426 3299 239	4745 2430 160	3178 1205 22	1675 297 0	279 0 0	000	000	26819 18498 1093	100% 69%
U.S.S.R. RKZ (morning daylight) 850 mb 100 mb 10 mb	1772 1830 80	2225 2137 305	2931 2487 414	1491 1231 337	1222 1150 301	447 648 234	45 274 123	0 61	000	000	10133 9757 1855	1002 96% 18%
U.S.A. NOAA (afternoon daylight) 850 mb 100 mb 10 mb	9026 6725 458	13022 7830 343	14319 6296 224	10798 3248 67	8046 2045 118	4270 1554 236	1818 1517 229	1378 746 139	982 450 84	672 94 13	64331 30505 1911	100% 47% 3%
U.S.A. NOAA (morning daylight) 850 mb 100 mb 10 mb	4266 5522 471	1897 5732 810	764 2878 766	427 1388 617	241 586 259	251 156 79	279 111 7	188 66 0	27 0	000	8340 16439 3009	100% 197% 36%
J.S.A. AN/AMT 850 mb 100 mb 10 mb	523 137 16	1351 785 19	3023 1298 217	3217 2458 574	2567 2154 539	1972 1398 541	1878 1102 261	836 548 55	233 58 18	000	15600 9938 2240	100ž 64% 14%

Table 2. -- Numbers of observed height differences

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Instrument type	-10-0	0-10	10-20	<u>20-30</u>	lar angle 30-40	(degrees	50-60	60-70	70-80	80-90	Total	850-mb as 100%
Australian "Diamond Hinman" 850 mb 100 mb 10 mb	0	149	605 204	484 423	741 343 INSUFFICI	1059 512 ENT DATA	527 638	221 254	189 72	22 17	4058 2463	1002 612
Chinese 850 mb 100 mb 10 mb	8483 1724	15481 4400	6965 7159	2508 5894	0 3268 INSUFFICI	0 170 ENT DATA	00	00	00	00	3 3437 22615	1002
Sangamo (Canada and Portugal) 850 mb 100 mb 10 mb	4057 3263 80	2558 2580 86	1603 1488 46	885 471 74	847 301 65	253 132 14	223 94 20	111 108 21	37	000	10537 8474 415	1002 802 42

Table 2. -- Numbers of observed height differences (continued)

Table 3.--W. German "Graw" instrument. Values of mean $\overline{\Delta T}$ and mean $\overline{\Delta H}$ as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. $\overline{\Delta T}$'s are given on upper line, $\overline{\Delta H}$'s on lower line. Values in italics are best estimates based on small samples.

Solar elevation					Press	ure le	vel (m	b)		
angle (degrees)		700	500	300	200	100	50	30	20	10
-5°	$\Delta T \over \Delta H$	1.2	-	_	-		-	-	-	1
50	$\frac{\Delta T}{\Delta H}$	5	1	-	_	-	-	0.9	1.0	-
15°	$\frac{\Delta T}{\Delta H}$	0.1	0 -1	0 -1	0.1	0.4	0.6	0.6	0.5	0.7
25 ⁰	$\frac{\Delta T}{\Delta H}$	-0.1 0	-0.1 1	0.1 -1	0.2	0.5	0.5	0.6	0.6	0.3
35 ⁰	$\frac{\Delta T}{\Delta H}$	0 2	-2	0 4	0.3	0.3	0.6	0.8	0.6	0.7
450	$\frac{\overline{\Delta T}}{\Delta H}$	0.1	-3	0 6	0.3	0.4	0.2	0.1 11	-0.1 13	-0.2 10
550	$\frac{\Delta T}{\Delta H}$	0 4	2	0 4	0.1	0.3	0.4	0 15	0.4	-0.3
65 ⁰	$\frac{\Delta T}{\Delta H}$	-0.2	0.2	0 5	0.2	0.6 17	0.5	-	Ξ	-
750	ΔT ΔH	-	-2	-	-	-	5	-	-	-

Table 4.--U.K. Kew instrument. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line. Values in italics are best estimates based on small samples

Solar elevation				<u>P</u>	ressure	level	<u>(mb)</u>			
(degrees)		700	500	300	200	<u>100</u>	50	<u>30</u>	20	<u>10</u>
-50	$\frac{\Delta T}{\Delta H}$	1	Ξ	- 1	-	- 1	-	-	-	-
5 ⁰	$\frac{\Delta T}{\Delta H}$	-0.3 -3	-0.2 -5	-0.2	-0.1	-0.2 -12	-1.0 -20	-0.5 -31	0.3 -39	3.2 10
150	$\frac{\Delta \mathbf{T}}{\Delta \mathbf{H}}$	-0.1 -1	0 -2	0 -1	-0.1 -1	0.1 -1	-0.3 -3	-0.1 -14	0.4 -18	3.3 36
250	$\frac{\Delta T}{\Delta H}$	-0.1 -1	0 -1	0 1	0 -2	0.3 1	0.1	0.8	1.6 27	4.3 113
350	$\frac{\Delta T}{\Delta H}$	0 -1	0 -4	0.2 -1	0 3	0.2	-0.1 -1	0.3	1.5 24	3.3
450	$\frac{\Delta T}{\Delta H}$	-0.1 -2	0 -3	0.1 -3	0 -3	0	-0.3 -5	0.1 -2	1.1 11	2.8 56
550	$\frac{\Delta T}{\Delta H}$	-0.1 -2	0 -2	0.1 -3	-0.2	-0.3 -1	0 -9	0.4 -10	0.8	1.4 34
65 ⁰	$\frac{\Delta T}{\Delta H}$	-0.1 -2	-0.1 -2	0 -5	0 -4	-0.1 4	0.3	0.1 1	-	-
75 ⁰	$\frac{\Delta T}{\Delta H}$	-2	-0.6	-0.3	-12	1		-	-	-

Table 5.--French Mesural instrument used in Metropolitan France. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line.

Solar elevation angle					Pressu	re lev	el (mb	<u>)</u>		
(degrees)		700	500	300	200	100	50	30	20	10
-50	ΔT	-	1 -	-	-	_	-	1>	-	-
	ΔH	-	-	-	-	-	-	-	-	-
50	ΔT		-	-	-	-	-	-	-	-
	ΔH	-	-	-	-	-	-	-	-	-
15 ⁰	$\frac{\Delta T}{\Delta H}$	0.3	0.4	0.3	-0.1 7	0.4	0.9 14	1.1	1.2	1.9 10
25 ⁰	$\frac{\Delta T}{\Delta H}$	0.2	0.2	0.2	0.5	0.6 19	0.8	1.2	1.3	2.2 12
350	$\frac{\Delta T}{\Delta H}$	0.3	0.3	0.4	0.5	0.6	0.8	1.0	1.8	2.6 10
450	$\Delta T \Delta H$	0.2	0.6	0.5	-0.1 21	0.3	0.8	1.3	1.2	3.2 10
550	ΔT ΔH	0.2	0.3	0.3	0.5	0.6	0.8	1.1	1.6	1.1 8
65 ⁰	$\frac{\Delta T}{\Delta H}$	-0.2	0.2	0.1	0.2	0.5	1.2 25	0.9	1.8	0.9 14
75 ⁰	$\frac{\Delta T}{\Delta H}$	-	÷	Ξ		-	-	-	-	-

Table 6.--French Mesural instrument, used outside France. Values of mean $\overline{\Delta T}$ and mean $\overline{\Delta H}$ as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. $\overline{\Delta T}$'s are given on upper line, $\overline{\Delta H}$'s on lower line. Values in italics are best estimates based on small samples.

Solar					Pressu	ire lev	vel (mb)		
elevation								1		
(degrees)		700	500	300	200	100	50	30	20	10
- 5 ⁰	ΔT	-	-	-	-	-	-	-	- ²² -	-
	∆H	-	-	-	-	-	-	-	-	-
5 ⁰	$\overline{\Delta T}$	-	-	-	-	-	-	-	_	-
	ΔH	-		-	-	-		-		-
15 ⁰	$\overline{\Delta T}$	-	-	-	-	-	-	_	· · · -	-
	ΔH	-	-	-	-	-	· • •	-	1. · · -	,
25 ⁰	ΔT	0.8	0.5	1.3	2.3	3.7	6.2	7.6	11.4	-
	∆H	10	13	27	40	123	288	279	510	-
35 [°]	ΔT	0.6	0.8	0.9	2.2	2.5	4.2	6.6	8.8	-
	∆H	8	15	27	41	105	16	229	333	
45 [°]	ΔT	0.8	0.9	1.4	5.5	2.4	4.3	5.1	8.5	-
	∆H	12	8	31	36	60	22	323	407	-
55	$\overline{\Delta T}$	0.5	0.7	1.1	4.9	2.7	0.4	5.1	9.3	-
	ΔH	9	19	32	46	81	96	220	314	
65 [°]	$\overline{\Delta T}$	0.2	0.5	0.8	7.6	2.3	3.7	5.3	5.0	-
	ΔH	8	12	26	31	54	86	102	212	-
75 [°]		0	0.4	0.6	3.8	1.2	2.4	-	-	-
	AH	1	3	8	8	30	38	-	-	-

Table 7.--Finnish Väisälä instrument used both in Finland and several foreign countries. Values of mean ΔT and mean $\overline{\Delta H}$ as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line. Values in italics represent best estimates based on small samples.

Solar elevation angle	Pressure level (mb)												
(degrees)		700	500	300	200	100	50	30	20	10			
-5 ⁰	$\frac{\overline{\Delta T}}{\Delta H}$	-0.2 -2	0.2	-0.1 -5	0 -3	0.1	0.1 -6	0.1	0.1	2.6 83			
50	$\frac{\overline{\Delta T}}{\Delta H}$	0.1 -1	0.1	0.1	0.1 2	0.3	0.8 11	2.0 28	1.9 42	3.1 83			
15°	$\frac{\overline{\Delta T}}{\Delta H}$	0.1 1	0 2	0.1 4	0.1	0.3	1.0 18	2.4	2.4	4.0 127			
25 ⁰	$\frac{\overline{\Delta T}}{\Delta H}$	0.3	0.1	0.2	0.1	0.4	1.0 26	2.2 43	2.0	1.8 91			
350	$\frac{\Delta T}{\Delta H}$	0.1	0.2	0.2 11	0.2 11	0.6	1.2 32	2.0	2.0 74	3.3 77			
450	$\frac{\Delta T}{\Delta H}$	0.2	0.2	0.3	0.3 15	0.5 15	1.2 30	1.5 43	1.6	1.4 53			
55 ⁰	$\overline{\Delta T}$ $\overline{\Delta H}$	0.3	0.3 13	0.3	0.5	0.8	1.0 11	0.4 29	0.4 38	2.5 73			
65 ⁰	$\Delta T \Delta H$	0.5	0.4	0.5	0.6	1.3 37	0.8	0.1 29	0.1 61	-			
75 ⁰	$\frac{\Delta T}{\Delta H}$	0.3 10	0.6	0.6	0.6	1.3	-	1	-	-			
85°	$\overline{\Delta T} \over \Delta H$	0.2	- 11	-	-	÷	-	-	-	-			

Table 8.--Väisälä instrument used in Finland. Values of mean ΔT and mean $\overline{\Delta H}$ as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line.

Solar elevation				Pre	essure 1	Level (mb)	
angle (degrees)		700	500	300	200	100	50	10
-50	$\frac{\Delta T}{\Delta H}$	-0.1 - 5	-0.1 - 4	-0.3 - 3	-0.1 2	0.0	-0.4 - 2	0.3 12
5 ⁰	$\frac{\Delta T}{\Delta H}$	0.1	0.2	0.0	0.1	-0.1	-0.2 - 18	0.0 - 23
15°	$\Delta T \over \Delta H$	0.1 3	0.4	0.2	-0.2	0.0	0.0	-0.4 8
25 ⁰	$\Delta T \over \Delta H$	0.1 3	-0.2	0.1	-0.2 8	-0.1 7	0.1 14	0.5 7
35 ⁰	$\frac{\Delta T}{\Delta H}$	0.1	0.4	0.1 10	0.1 11	0.0	0.4 21	0.2 21
450	$\frac{\Delta T}{\Delta H}$	0.1	2	52	2 E	-	-	-

<u>Note</u>: Data were insufficient above 30 mb for derivation of meaningful numbers. Data were also insufficient for daytime solar elevation angles higher than 50°.

Table 9.--Japanese "code-sending" instrument. Values of mean ΔT and mean $\overline{\Delta H}$ as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line. Values in italics are best estimates based on small samples.

Solar elevation angle	Pressure level (mb)										
(degrees)		700	500	300	200	100	50	30	20	10	
-5 [°]	ΔT ΔH	Ξ	-	-	-	-	-	1	2	-	
5 ⁰	ΔT ΔH	0.3	- 3	-	Ξ.	-	2	-	-	-	
15 [°]	<u>∆</u> ∆H	0 2	0.2	0.6	0.4	0.7	1.3 43	2.6 82	-	-	
25	<u>∆</u> T ∆H	0.1	0.3	0.4	0.6	0.8	1.2 43	1.9 70	3.0 102	4.1 173	
35°.		0.2	0.3	0.3	0.6	0.7	1.0 44	1.6	2.8	3.5 142	
45 [°]		0.2	0.4	0.4	0.4	0.8	0.8	1.2 53	1.9 70	1.6 94	
55 ⁰		0.2	0.5	0.5	0.4	0.8	0.9 31	1.1 43	1.3	1.8	
65 [°]	ΔT ΔH	0.3	0.7	0.6	0.5	0.4	0.2	1.0	1.4 84	1.4 128	
75 [°]		-	-	Ξ	:	-	0.1	-0.5	-1.0 20	-	

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Table 10.--Afternoon daylight (12Z), U.S.S.R. A-22 instrument. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation angle of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line. Values in italics are best estimates based on small sample sizes.

Solar Elevation	Pressure level (mb)												
angle (degrees)		700	500	300	200	100	50	30	20	<u>10</u>			
-50	$\frac{\Delta T}{\Delta H}$	0.1	0.2	0.1	0.1 7	0.2	0.3 20	0.2	0.3 42	-0.2 18			
50.	$\frac{\Delta T}{\Delta H}$	0.3	0.2	0.3 11	0.3 14	0.5	0.6 30	0.3	0.5	0.2			
150	$\frac{\overline{\Delta T}}{\overline{\Delta H}}$	0.3	0.3	0.2	0.2	0.4	0.7	0.8	1.0 63	2.5 121			
250	$\frac{\overline{\Delta \mathbf{T}}}{\overline{\Delta \mathbf{H}}}$	0.5	0.3	0.1	0.3	0.5 20	0.7 32	0.9 45	0.1 54	0.3 84			
35°	$\frac{\Delta T}{\Delta H}$	0.2	0.2	0.4 13	0.1 15	0.5 23	0.9 31	0.9	1.2 62	1.8 80			
45 ⁰	$\frac{\Delta T}{\Delta H}$	0.2	0.2	0.4	0 17	0.3	0.6	0.5 18	1.0 28	0.2			
559	$\frac{\Delta \mathbf{T}}{\Delta \mathbf{H}}$	0.1	0.3	0.2	-0.2 12	0 5	0.4	0.8	0.2	-			
650	$\frac{\Delta T}{\Delta H}$	0.1	0 5	0.4		-	-	Ξ	2	-			
750	AT AH	-	-	-	2	-	2	-		-			

Table 11.--Morning daylight (00Z), U.S.S.R. A-22 instrument. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line.

Solar elevation angle			1							
(degrees)		700	500	300	200	100	50	30	20	10
-5 [°]	$\frac{\Delta T}{\Delta H}$	0 -4	-0.1 -5	0 -4	0 -5	0 -6	-0.1 -7	-0.2 -9	-0.1	0.1 -16
5 ⁰	$\frac{\Delta \mathbf{T}}{\Delta \mathbf{H}}$	0.1 -3	0 -3	0 -5	0.2	0.2 -1	0.3	0.4	0.6 18	0.9
15 ⁰	$\frac{\Delta T}{\Delta H}$	0 -2	0 -3	0 -4	0.1 -3	0.2	0.5 5	0.6	0.9	1.9 51
25 ⁰	$\frac{\Delta T}{\Delta H}$	-0.1 0	-0.1 -2	0.2 -2	0.2 -4	0.3	0.5	0.5	0.9	2.8 56
35 [°]	$\frac{\Delta T}{\Delta H}$	0 1	0 0	0.1	0.2	0.5	0.8 17	0.7	1.1 33	0.5 48
45 [°]	$\frac{\Delta T}{\Delta H}$	-0.1 -1	0 0	0.2 -3	-0.1 -5	0.1 -3	0.5	0.6	1.2 32	2.3 38
55 ⁰	$\Delta T \over \Delta H$	-0.2 -3	-0.2	0.1 -3	-0.2 -7	0.3 -8	0.3	0.4	0.9	2.4
65 ⁰	$\Delta T \Delta H$							0.3 -10	-11	-0.6
75 ⁰									-	

Table 12.--Afternoon daylight (12Z), U.S.S.R. RKZ instrument. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation and of pressure level. Units are degrees Velsius and meters. ΔT 's are given on upper line, ΔH 's on lower line.

Solar elevation					Pressure level (mb)						
(degrees)		700	500	300	200	100	50	30	20	10	
-5°	$\frac{\Delta T}{\Delta H}$	0.1 5	0 5	0.1	0.2	0.2	0.5	0.7	0.8	1.8 82	
50	$\Delta T \Delta H$	0.2	0.2	0.4	0.3	0.5	0.8	1.0 54	1.1 67	2.8 120	
15 ⁰	$\frac{\Delta T}{\Delta H}$	0.3	0.3 10	0.3	0.4	0.5	0.7	0.9	1.1 53	1.7 73	
25 ⁰	$\Delta T \Delta H$	0.3	0.3	0.3	0.2 18	0.2	0.5	0.7	0.6	0.5	
35 ⁰	$\frac{\Delta T}{\Delta H}$	0.3	0.3	0.3	0.3	0.1 20	0.2	0.4 28	0.6	0.5	
45 ⁰	$\frac{\Delta T}{\Delta H}$	0.3 8	0.3 10	0.4	0.2 18	0 18	0.3	0.2	0.1	0 8	
550	$\frac{\Delta T}{\Delta H}$	0.1	0.2	0.2	0 12	0 12	0.3	-0.2	-	-	
650	$\frac{\Delta T}{\Delta H}$	0.1	0.1	2	Ξ	-	-		-	-	
75 ⁰		-	-	-		-	-	-	-	-	

Table 13.--Morning daylight (00Z), U.S.S.R. RKZ instrument. Values of mean $\overline{\Delta T}$ and mean $\overline{\Delta H}$ as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. $\overline{\Delta T}$'s are given on upper line, $\overline{\Delta H}$'s on lower line.

Solar elevation	Pressure level (mb)												
angle (degrees)		700	500	300	200	100	50	30	20	10			
-5°	$\frac{\Delta T}{\Delta H}$	-0.2 -7	0 -8	0 -8	-0.1	0 -9	0.1 -10	0.1 -14	0.2 -20	0.5			
50	$\Delta T \Delta H$	0.1 -3	0 -5	0 -4	0.1 -5	0.3 -1	0.4	0.7 14	0.9	1.5 28			
15°	$\frac{\Delta T}{\Delta H}$	0 -3	0 -3	0.2 -3	0.1 -2	0.3	0.5	0.6 14	0.9 23	1.7 36			
25°	$\frac{\Delta T}{\Delta H}$	-0.1 -2	-0.1 -4	0.1 -3	0.4 -1	0.2 1	0.2	0.3	0.6 18	1.8 26			
35°	$\frac{\Delta T}{\Delta H}$	-0.3	-0.1 -5	0 -6	0.1 -3	0.1 1	0.1 1	0.1 -1	0.3	1.6 27			
45 ⁰	ΔT ΔH	-0.3	-0.1 -4	0.1 -4	0.1 -6	0 -5	0.1 -6	0 -1	0.1	0.5			
550	$\Delta T \Delta H$	-0.3	-0.3	-0.1 -4	-0.1 -5	0.2	0.1 -11	-0.1 -15	0.3 -17	0.1 -25			
65 ⁰	$\Delta T \\ \Delta H$	0 -	2	- 1	1	- 1	1	-0.2	0.2	0.9			
75 ⁰		2	-	-	-	2	-	-	-	-			

Table 14.--Afternoon daylight (00Z), U.S. NOAA instrument. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line. Values in italics are based on small sample sizes, but nevertheless are reasonable estimates.

Solar` elevation angle					Pres	sure 1	evel (mb)		
(degrees)		700	500	300	200	100	50	30	20	10
-50	$\frac{\overline{\Delta \mathbf{T}}}{\Delta \mathbf{H}}$	0.2 1	0.2 3	0.2 7	0.1 9	0.2 15	0.5 27	0.9 46	1.2 70	1.9 116
50	$\frac{\Delta T}{\Delta H}$	0.3 1	0.3	0.4	0.5	0.7 30	1.1 49	1.6 74	1.9 100	2.4
15 ⁰	$\frac{\Delta T}{\Delta H}$	0.5 3	0.4	0.6	0.6	0.9 38	1.3 60	1.8 88	2.3 116	2.9 178
25 ⁰	$\frac{\Delta T}{\Delta H}$	0.6 7	0.5 12	0.6	0.7	1.0 42	1.4	2.1 98	2.6 132	3.2 175
35 ⁰	$\frac{\Delta T}{\Delta H}$	0.6	0.5 13	0.6	0.7	1.0 40	1.4 65	1.9 94	2.4 129	2.6 172
450	$\frac{\Delta T}{\Delta H}$	0.3	0.4	0.5	0.6	1.0 35	1.2 56	1.4 88	1.8 107	2.8 158
55°	$\frac{\Delta T}{\Delta H}$	0.4	0.4 7	0.5	0.7	1.0 34	1.1 49	1.4 71	1.5	2.3
65 ⁰	$\frac{\Delta T}{\Delta H}$	0.5 3	0.5	0.5 15	0.7	1.0 36	1.3 49	1.0 65	1.3 95	2.5 134
750	$\frac{\Delta T}{\Delta H}$	0.5	0.5	0.5	0.7	1.2 34	1.0 49	1.6 72	1.3	1.1 151
85 ⁰	ΔT ΔH	0.4	0.6	0.6	0.8	1.0	0.9	1.2 80	2.2	3.2 -13

Table 15.--Morning daylight (12Z), U.S. NOAA instrument. Values of mean ΔT (upper line) and of mean ΔH (lower line) as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. Values in italics are besed on small sample sizes, but nevertheless are reasonable estimates.

Solar elevation				P	ressur	e leve	1 (mb)			
angle (degrees)		700	500	300	200	100	50	30	20	10
-50	$\frac{\Delta T}{\Delta H}$	-0.2 -3	-0.2	-0.2 -7	0 -8	0.1	0 -9	-0.1 -14	-0.2 -20	0 -39
5 ⁰	$\frac{\Delta T}{\Delta H}$	-0.1 -2	-0.1	.0 -6	0.2	0.5	0.5	0.6	0.8	1.3 -7
15 ⁰	$\frac{\Delta T}{\Delta H}$	0.2	0.2	0.2	0.5	0.7	0.7 18	1.0 25	1.1 29	1.3 30
25 ⁰	$\frac{\Delta T}{\Delta H}$	0.4	0.4	0.2	0.4	0.7	0.9 31	1.1 34	1.1 39	1.1 44
35 ⁰	$\frac{\Delta T}{\Delta H}$	0.5 10	0.4 13	0.4 17	0.8 18	0.8	1.1 36	1.4 44	1.5 64	1.3 71
45 ⁰	$\frac{\Delta T}{\Delta H}$	0.4	0.5	0.7 24	1.5 23	1.2 29	1.4 50	1.3 59	1.2 59	1.9 102
55 ⁰	$\Delta T \\ \Delta H$	0.8	1.0	1.2	1.0	1.6	1.0 181	2.2	1.3 54	3.5 119
65 ⁰	$\frac{\Delta T}{\Delta H}$	0.9	0.9	1.1	0.9	1.0	1.5	-0.4 47		

Table 16.--AN/AMT 4 instrument, used in various parts of the world. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line. Values in italics are best estimates based on small samples.

Solar elevation	Pressure level (mb)										
(degrees)		700	500	300	200	100	50	30	20	10	
-5°	ΔT ΔH	-0.1 -1	-0.5 -1	-0.2 -6	0 -9	-0.1 -11	0.2 -18	0.8 36	-	1.5	
5 ⁰	ΔT ΔH	0 -1	-0.1	0.1 -1	0.5	0.4	0.8 17	1.2 48	1.7 78	1.4 69	
15 ⁰	ΔT ΔH	0 3	0 2	0.4	0.6	0.8	0.8 31	1.7 50	1.8	2.9	
25 [°]	ΔT ΔH	0.2	0.2	0.4	0.5	0.8	1.1 43	1.4	1.8 76	1.7 109	
35 [°]	ΔT ΔH	0.3	0.4 7	0.5	0.5	0.9 28	1.1 42	1.8 68	1.6 91	2.0 119	
45 ⁰	ΔT ΔH	0.3	0.3	0.4	0.4	0.9 36	1.1 49	1.5 67	1.7 90	2.5 136	
55	ΔT ΔH	0.3 5	0.4	0.5	0.6	1.0 33	1.4 52	1.7 77	1.7 97	2.7 144	
65 ⁰	$\Delta T \over \Delta H$	0.2	0.4 9	0.6	0.7	0.8	1.5 55	1.8 43	1.6 72	2.3	
75 [°]	$\frac{\Delta T}{\Delta H}$	0.1	0.3	0.5	0.6	0.9	-	-	1.2	1.9	

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Table 17.--Australian ("Diamond Hinman") instrument, used also in New Zealand. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line. Values in italics are best estimates based on small samples.

Solar					Pressu	ire lev	vel (mb)	
elevation angle (degrees)		700	500	300	200	100	50	<u>30</u>	20
-5°	$\frac{\Delta T}{\Delta H}$	-	-		1	1	1	1	-
50	$\frac{\Delta T}{\Delta H}$	0.3	0 -2	0 -5	0.2 -3	- 2		-	-
15 ⁰	$\frac{\Delta T}{\Delta H}$	0	0.2	0.2	0.4	0.6 17	0.9 36	0.8 55	1.8
25 ⁰	$\frac{\Delta T}{\Delta H}$	0.3	0.1	0.2	0.4	0.8	1.0 44	1.7 52	1.7
35 ⁰	$\frac{\Delta T}{\Delta H}$	0.2	0.2	0.2	0.5	0.7	1.1 31	1.7 52	1.6 168
450	$\frac{\Delta T}{\Delta H}$	0.1	0.2	0.3	0.6	0.9	1.3 41	1.7 59	2.8 100
550	$\frac{\Delta T}{\Delta H}$	0.3	0.4	0.2	0.5	1.0 25	1.1 44	1.8 61	2.2 78
65 ⁰	$\frac{\overline{\Delta T}}{\Delta H}$	0.7	0.5	0.2 12	0.6	1.1 32	1.4 48	1.4 67	1.7 100
75 ⁰	$\frac{\Delta T}{\Delta H}$	0.3 1	0.4	0.1 5	0.5 14	1.0 35	1.4 10	2.4 48	1.4
85 ⁰	$\frac{\Delta T}{\Delta H}$	0.4	-0.3	-	0.9	0.8	-	-	

Table 18.--Canadian "Sangamo" instrument. Values of mean ΔT and mean ΔH as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. ΔT 's are given on upper line, ΔH 's on lower line. Values in italics represent best estimates based on small samples.

Solar elevation	Pressure level (mb)													
(degrees)	× .	700	500	300	200	100	50	30	20	<u>10</u>				
-5°	$\frac{\Delta T}{\Delta H}$	0 -1	0 -1	0 -1	0	0.1 1	0.1 3	0.4	0.5	1.5 68				
50	$\frac{\Delta T}{\Delta H}$	0.1 1	02	0.2	0.4	0.6	0.8	1.1 37	1.4 53	1.9 91				
150	$\frac{\Delta T}{\Delta H}$	0.3	0.4 7	0.4 17	0.4 17	0.7 28	1.1 39	1.3 49	1.4	1.3 65				
250	$\frac{\Delta T}{\Delta H}$	0.4	0.5 13	0.5 18	0.5	0.7 26	0.9 37	1.4 48	1.4 66	1.2 74				
35 ⁰	$\frac{\Delta \mathbf{T}}{\Delta \mathbf{H}}$	0.3	0.5 14	0.7. 19	0.6	0.9	1.2 44	1.2 49	1.3 62	2.2 90				
450	$\frac{\Delta T}{\Delta H}$	0.2	0.1	0.5	0.2 18	0.8	1.1 50	1.0	1.0 83	0.4 71				
55 ⁰	$\frac{\Delta T}{\Delta H}$	-0.1 -2	0.3 -1	0.5	0.6	1.3 31	0.7 19	0.8	1.4 95	1.7 98				
65 ⁰	$\frac{\Delta T}{\Delta H}$	-0.3 -2	0.1	0.6	0.9 11	1.0 22	0.9 36	1.3 54	1.4 71	1.0 102				
75°	$\frac{\Delta T}{\Delta H}$	12	2	-0.1	1.0 18	0.8	1.7	1.6	1.2	1.3				

Table 19.--Chinese instrument. Values of mean $\overline{\Delta T}$ and mean $\overline{\Delta H}$ as functions of mean daytime solar elevation angle and of pressure level. Units are degrees Celsius and meters. $\overline{\Delta T}$'s are given on upper line, $\overline{\Delta H}$'s on lower line. Values in italics are best estimates based on small samples.

Solar elevation	Pressure level (mb)								
(degrees)		700	500	300	200	100	Higher	levels missing	
-5 ⁰	$\frac{\overline{\Delta T}}{\Delta H}$	-0.2 -1	0.1 -3	0 -2	0.1 -5	0 1			
50	$\frac{\Delta T}{\Delta H}$	-0.2 -1	-0.1 -2	-0.4 -7	-0.4 -12	-0.2 -16			
15°	$\frac{\Delta T}{\Delta H}$	-0.1 0	-0.1 -1	-0.5 -5	-0.3 -11	0 -15			
25 ⁰	$\frac{\Delta T}{\Delta H}$	-0.1 1	0	-0.6 -2	-0.2 -9	0.1 -13			
350	$\frac{\Delta T}{\Delta H}$	-	-0.4	-0.3	-0.4 -6	0.4 -4			
45 ⁰	$\frac{\overline{\Delta T}}{\Delta H}$	- 2	-	-	-	0.6			
550	$\frac{\Delta T}{\Delta H}$	1	- 2	-	1	1			
650	$\frac{\Delta T}{\Delta H}$	-	2	2	-	Ľ.			
75 ⁰	$\frac{\Delta T}{\Delta H}$	-	2	1	-	-			



Figure 1.--Distribution of radiosonde instrument types throughout the Northern Hemisphere, as of July 1978. (This representation cannot be considered current because changes are continually taking place in the network.)



Figure 2.--Distribution of radiosonde instrument types over the Southern Hemisphere, as of July 1978. (Same <u>caveat</u> as for figure 1.)



Figure 3.--Yearly migration of 0100 GMT sunrise line (for nominal 0000 GMT observations) at 10 mb.



Figure 4.--Yearly migration of 1300 GMT sunset line (for nominal 1200 GMT observations) at 10 mb.

DATA PROCESSING--DAY/NIGHT DIFFERENCES



Figure 5.--Flow diagram for scheme of calculating day-night differences of temperatures and of geopotential heights.

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Figure 6.--Scatter diagram of $\overline{\Delta T}$ for the U.S. NOAA instrument (also known as the manufacturer's name VIZ) for 200 mb.

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Figure 7.--Scatter diagram of $\overline{\Delta T}$ for the Mesural instrument used in metropolitan France, 150-mb level.

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Figure 8.--Scatter diagram of $\overline{\Delta T}$ for the Mesural instrument used outside France, 150-mb level.

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Figure 9.--Scatter diagram of $\overline{\Delta T}$ for the instrument used by the People's Republic of China, 100-mb level.

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Figure 11.--Scatter diagram of $\overline{\Delta T}$ for the Japanese "code sending" radiosonde, 30-mb level.



Figure 12.--Scatter diagram of ΔT for the U.S. NOAA (VIZ) instrument, 10-mb level.



*U.S. GOVERNMENT PRINTING OFFICE: 1979 281-067/268 1-3

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(Continued from inside front cover)

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- NWS NMC 60 The LFM Model 1976: A Documentation. Joseph P. Gerrity, Jr., December 1977, 68 pp. (PB-279-419)
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- NWS NMC 62 Addition of Orography to the Semi-Implicit Version of the Shuman-Hovermale Model. Kenneth A. Campana, April 1978, 17 pp. (PB-286-009)

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