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2	Geophysical Research Letters					
3	Supporting Information for					
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5	Multidecadal Intensification of Atlantic Tropical Instability Waves					
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30Text S1: Comparison of near-surface velocities from the drifter-wind-altimetry31synthesis and moored velocity observations

32 Near-surface velocities are derived from a synthesis of in-situ drogued drifter velocity 33 measurements, satellite winds, and altimetry-derived geostrophic velocity anomalies (Lumpkin & Garzoli, 2011; Perez et al., 2019) by simultaneously regressing the observed 34 35 velocities, low-passed at 1.5 times the local inertial period with a floor of 1 day and a 36 ceiling of 5 days, onto four spatially-varying, time invariant coefficients multiplying (a) 37 the wind speed from the ERA-interim product 38 (https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era-interim) the in 39 downwind and crosswind directions, and (b) the zonal and meridional geostrophic 40 velocities derived from altimetry as provided by the Copernicus Marine and Environment 41 Monitoring Service (CMEMS). These coefficients are then used to generate gridded, daily 42 near-surface velocities from the same satellite products. We note that Ekman dynamics are 43 not explicitly assumed, while CMEMS near-equatorial currents are derived from SLA 44 using an algorithm derived from Lagerloef et al. (1999). To address the reliability of near-45 surface velocities provided by this drifter-wind-altimetry synthesis, we compare this data set with velocity observations from current meters attached to the PIRATA moored surface 46 47 buoys at 4°N, 23°W and at 0°, 23°W, both installed at 10 m depth (Bourlès et al., 2019; 48 Foltz et al., 2019). In addition, we use current velocities from moored acoustic Doppler 49 current profiler (ADCP) measurements from a subsurface mooring at 0°, 23°W (Tuchen et 50 al., 2022b). Here, we are using ADCP data at the shallowest resolved depth at 20 m. Lastly, 51 surface velocities from another current analysis product, the Ocean Surface Current 52 Analyses Real-time (OSCAR) project, are used for comparison.

53 In Figure S1, horizontal current velocities are shown for one year in which all data 54 sets provided data (Fig. S1a-d). Overall, we find good agreement of zonal and meridional 55 velocities at 4°N, 23°W between the available data products (PIRATA, drifter-wind-56 altimetry synthesis, and OSCAR). OSCAR is provided at a 5-daily temporal resolution, 57 whereas PIRATA current meter data is generally provided with hourly resolution and is 58 then subsampled to 12-hour means (Fig. S1c-d). For this reason, some differences relative 59 OSCAR and the drifter-wind-altimetry synthesis can be expected. The comparison shows 60 that the occurrence of Tropical Instability Waves is reproduced by the products, but 61 underestimates the large meridional velocity amplitudes and eddy kinetic energy associated 62 with the waves (Fig. S1f). The correlation coefficients between PIRATA current velocities 63 at 10 m depth and near-surface velocities at an approximate depth of 15 m from the drifter-64 wind-altimetry synthesis (using all available data between 2005-2022) are 0.46 for zonal 65 and 0.60 for meridional velocities (Tab. S1). Correlations are slightly smaller between the 66 PIRATA velocities and OSCAR, 0.50 for zonal and 0.56 for meridional velocity.

67 The same comparison was carried out at 0°, 23°W (Fig. S1a-b,e). The agreement 68 between observed current velocities and those provided by the two data products is 69 considerably lower on the equator. The correlation coefficients between PIRATA current 70 velocities at 10 m depth and near-surface velocities at 15 m depth from the drifter-windaltimetry synthesis (using all available data between 2005-2022) are 0.21 for zonal and 71 72 0.23 for meridional velocities (Tab. S1). The correlation coefficients between moored 73 ADCP current velocities at 20 m depth and current velocities at 15 m depth from the drifter-74 wind-altimetry synthesis are 0.32 for zonal and 0.18 for meridional velocities. Although

the corresponding correlations are slightly larger between the moored and OSCAR
velocity estimates at 0°, 23°W, the correlations are still weaker than those found at 4°N,
23°W.

All correlation coefficients at 0°, 23°W and 4°N, 23°W shown in Tab. S1 are only slightly increased if a 5-day running mean is applied to the PIRATA, moored ADCP and drifter-wind-altimetry time series in order to adjust to the lower temporal resolution of

81 OSCAR.



84 Figure S1. Time series of (a-d) zonal (u) and meridional (v) velocity and (e-f) eddy kinetic energy (EKE; $({u'}^2 + {v'}^2)/2$) at 0°, 23°W (left panels) and at 4°N, 23°W (right panels). 85 Included are data from PIRATA current meter observations at 10 m (red), near-surface 86 87 velocities from the drifter-wind-altimetry synthesis at an approximate depth of 15 m (blue), 88 surface velocities from OSCAR (green), and at 0°, 23°W moored ADCP observations at 89 20 m (black). For both mooring sites, one year is chosen for which all products provided 90 complete records: 2013 at 0°, 23°W and 2019 at 4°N, 23°W.

Correlation (u/v)	PIRATA	Moored ADCP	Drifter- altimetry	OSCAR
PIRATA	X	0.83 / 0.83	0.21 / 0.23	0.37 / 0.20
Moored ADCP		X	0.32 / 0.18	0.40 / 0.22
Drifter- altimetry	0.46 / 0.60		X	0.60 / 0.16
OSCAR	0.50 / 0.56		0.72 / 0.77	X

91 Table S1. Correlation coefficients for zonal (u) and meridional (v) velocity time series at

92 0°, 23°W (green shaded values) and at 4°N, 23°W (yellow shaded values) using all

93 available data between 2005-2022 at both mooring sites. Note that long-term moored

94 ADCP observations are only available at 0° , 23°W.



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96 Figure S2. Standard deviation (σ) of bandpass filtered anomalies of sea surface salinity 97 (SSS) for (a) the long-term mean seasonal σ in May to July, (b) spatial averages of monthly σ in the box and for the time period indicated in (a). (c) Decadal linear trend of 98 99 intraseasonal variability of SSS during May to July. Significant trends are indicated by grey dots and one background TIW variability contour line from (a) is shown (thin black 100 101 line). (d) Monthly SSS TIW variability trends spatially averaged over the box indicated in 102 (c). (e) Time series of the yearly composite value for the spatial and temporal averages 103 indicated in (c) and (d). Solid thin lines in (e) indicate the 95% confidence band. Primes 104 indicate intraseasonal (20-50 day) and zonal (4-20°) band-pass filtered variables. Overbars 105 indicate averages of the seasonal peak months highlighted in turquoise. Uncertainty 106 estimates are based on a linear regression analysis (95% confidence interval).



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109 Figure S3. Subsurface intraseasonal meridional velocity from moored observations at 0°, 23°W. (a) Time series of filtered (20-50 day) meridional velocity (v'), 30-day running 110 mean $(\overline{\nu'})$, and monthly standard deviation σ (red diamonds) as well as seasonal standard 111 deviation averages for May to July (green diamonds) and July to September (blue 112 113 diamonds). (b) Climatology of intraseasonal meridional velocity standard deviation. The 114 black box indicates the depth range used for (a) and the annual maximum standard 115 deviation from July to September used for the JAS composite in (a). Shallow measurements 116 at 10 m from current meter data are only included in the climatology. (c) Decadal linear 117 trend of intraseasonal meridional velocity standard deviation from 2008 to 2018. The black box indicates the decadal linear trend peak in May to July used for the MJJ composite in 118 (a). Significant trend values (95% confidence interval) are indicated by black crosses. 119

120Text S2: Covariance of zonal and meridional velocities from moored velocity121observations at 0°, 23°W

122 The analysis of moored subsurface velocities at 0°, 23°W reveals a pronounced 123 seasonal cycle in the covariance of zonal and meridional velocity fluctuations with 124 maximum values from July to August (u'v'; Fig. S4a) that is not reproduced by the drifter-125 wind-altimetry synthesis product on the equator (Fig. 3). This suggests that positive 126 barotropic energy conversion, meaning a generation site for TIWs, also occurs on the 127 equator given the positive meridional gradient of zonal velocity close to the equator. Such 128 a secondary band of barotropic energy conversion has been shown in the model study by 129 von Schuckmann et al. (2008).

Moored velocity observations further show that between 2008 to 2018 u'v' has increased in July and August during this approximately decadal time period (Fig. S4b). This indicates that barotropic energy conversion has likely strengthened in the nSEC/EUC region near the equator, and this increase was not captured by the drifter-wind-altimetrysynthesis product (Fig. 3).



137 **Figure S4.** Covariance of horizontal velocity fluctuations $(\boldsymbol{u}'\boldsymbol{v}')$ from moored velocity 138 observations at 0°, 23°W. (a) Mean $\boldsymbol{u}'\boldsymbol{v}'$ between 2001 to 2021. (b) Decadal linear trend 139 of $\boldsymbol{u}'\boldsymbol{v}'$ between 2008 to 2018. Significant trend values (95% confidence interval) are 140 indicated by black crosses.





150 (Hersbach et al., 2020) from 1980 to 2021 and have been subsampled to daily fields.



152 Figure S6. Zonal (left panels) and meridional (right panels) eddy temperature advection (ETA) in the tropical Atlantic due to Tropical Instability Waves at approximately 15 m 153 depth. Primes indicate intraseasonal (20-50 day) and zonal (4°-20°) band-pass filtered 154 155 variables. (a-b) Mean ETA from near-surface velocities and satellite SST averaged for June to July. Positive values indicate sea surface cooling, negative values indicate sea surface 156 157 warming. Also shown in (a) and (b) is the mean ETA at the 23°W mooring sites on the 158 equator and at 4°N using 10 m velocities and satellite SST. (c-d) Decadal linear trend of 159 ETA for June to July. (e) Long-term monthly trends of zonal ETA for the box indicated in (c). (f) Long-term monthly trends of meridional ETA in a north equatorial (NE) box and 160 161 an equatorial (EQ) box as indicated in (d). (g) Time series of the yearly composite value of 162 zonal ETA for the spatial and temporal averages indicated in (c) and (e). (h) Time series of the yearly composite value of meridional ETA for the NE box during June to July. Solid 163 164 thin lines in (g) and (h) indicate the 95% confidence band.

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