



# Evaluation of Unoccupied Aircraft System (UAS) Remote Sensing Reflectance Retrievals for Water Quality Monitoring in Coastal Waters

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Unoccupied aircraft systems (UAS, or drones) equipped with off-the-shelf multispectral sensors originally designed for terrestrial applications can also be used to derive water quality properties in coastal waters. The at-sensor total radiance a UAS measured constitutes the sum of water-leaving radiance ( $L_W$ ) and incident radiance reflected off the sea surface into the detector's field of view ( $L_{SR}$ ).  $L_W$  is radiance that emanates from the water and contains a spectral shape and magnitude governed by optically active water constituents interacting with downwelling irradiance while  $L_{SR}$  is independent of water constituents and is instead governed by a given sea-state surface reflecting light; a familiar example is sun glint. Failure to accurately account for  $L_{SR}$  can significantly influence Rrs, resulting in inaccurate water quality estimates once algorithms are applied. The objective of this paper is to evaluate the efficacy of methods that remove  $L_{SR}$  from total UAS radiance measurements in order to derive more accurate remotely sensed retrievals of scientifically valuable in-water constituents. UAS derived radiometric measurements are evaluated against *in situ* hyperspectral Rrs measurements to determine the best performing method of estimating and removing surface reflected light and derived water quality estimates. It is recommended to use a pixel-based approach that exploits the high absorption of water at NIR wavelengths to estimate and remove  $L_{SR}$ . Multiple linear regressions applied to UAS derived Rrs measurements and *in situ* chlorophyll *a* and total suspended solid concentrations resulted in 37 and 9% relative error, respectively, which is comparable to coastal water quality algorithms found in the literature. Future research could account for the high resolution and multi-angular aspect of  $L_{SR}$  by using a combination of photogrammetry and radiometry techniques. Management implications from this research include improved water quality monitoring of coastal and inland water bodies in order to effectively track trends, identify and mitigate pollution sources, and discern potential human health risks.

**Keywords:** multispectral, water quality, chlorophyll *a*, total suspended solids, unoccupied aircraft system, drones, coastal

**Abbreviations:** DLS, downwelling light sensor;  $E_d$ , downwelling irradiance;  $L$ , spectral radiance;  $L_{SR}$ , surface-reflected radiance;  $L_{sky}$ , sky radiance; LUT, lookup table;  $L_T$ , total radiance;  $L_W$ , water leaving radiance; NIR, near infrared;  $R_{rs}$ , remote sensing reflectance;  $R_{UAS}$ , UAS total reflectance; UAS, unoccupied aircraft system;  $\rho$ , effective sea-surface reflectance of wave facet

## INTRODUCTION

Unoccupied aircraft systems (UAS, drones) provide on-demand remote sensing capabilities at ultra-high resolution (<5 cm) without the challenges of cloud cover, land adjacency, and atmospheric effects associated with satellite and airborne remote sensing (Anderson and Gaston, 2013). UAS are becoming an integral tool in studying and managing coastal ecosystems (Johnston, 2019), with applications ranging from thermal remote sensing (Lee et al., 2016; Dugdale et al., 2019) and pollutant tracking (Arango and Nairn, 2019; Morgan et al., 2020), to habitat and population assessments (Gray et al., 2018; Windle et al., 2019). The increased spatial and temporal resolution provided by UAS remote sensing can enhance ecological and biogeochemical research of aquatic ecosystems. UAS have the potential to characterize the degree of eutrophication, identify the extent and movement of harmful algal blooms, and resolve fine-scale coupled biophysical processes in coastal and inland water bodies.

A number of recent studies have derived optical water quality parameters using UAS imagery and are listed in **Table 1**. In these studies, UAS are equipped with a variety of either multispectral or hyperspectral imagers that measure light at discrete wavebands in the visible and near-infrared (NIR) spectrum. Following the large body of research borne from earth observing satellites (Werdell and McClain, 2019), many of these studies use a combination of UAS optical payloads, calibrations, and numerical methods to determine remote sensing reflectance ( $R_{rs}$ ) defined as:

$$R_{rs}(\theta, \Phi, \lambda) = \frac{L_W(\theta, \Phi, \lambda)}{E_d(\lambda)} \quad (1)$$

where  $L_W$  ( $W m^{-2} nm^{-1} sr^{-1}$ ) is water-leaving radiance,  $E_d$  ( $W m^{-2} nm^{-1}$ ) is downwelling irradiance,  $\theta$  represents the sensor viewing angle between the sun and the vertical (zenith),  $\phi$  represents the angular direction relative to the sun (azimuth), and  $\lambda$  represents wavelength.

Like all above-water optical measurements, UAS do not measure  $R_{rs}$  directly as the at-sensor total radiance ( $L_T$ ,  $W m^{-2} nm^{-1} sr^{-1}$ ) constitutes the sum of  $L_W$  and incident radiance reflected off the sea surface into the detector's field of view, herein referred to as surface-reflected radiance ( $L_{SR}$ ).  $L_W$  is radiance that emanates from the water and contains a spectral shape and magnitude governed by optically active water constituents interacting with downwelling irradiance, while  $L_{SR}$  is independent of water constituents and is instead governed by a given sea-state surface reflecting light; a familiar example is sun glint. Here we define UAS total reflectance ( $R_{UAS}$ ) as:

$$R_{UAS}(\theta, \Phi, \lambda) = \frac{L_T(\theta, \Phi, \lambda)}{E_d(\lambda)} \quad (2)$$

where

$$L_T(\theta, \Phi, \lambda) = L_W(\theta, \Phi, \lambda) + L_{SR}(\theta, \Phi, \lambda) \quad (3)$$

As UAS measurements are typically performed close to the surface (e.g. United States Federal Aviation Administration's maximum allowable altitude of 122 m), atmospheric measurement effects are routinely assumed to be negligible and ignored (Zeng et al., 2017; Schneider-Zapp et al., 2019). Indeed, this is a key advantage of UAS imagery as atmospheric effects over coastal environments can introduce significant uncertainty in satellite-based measurements (Gordon and Clark, 1980). However, failure to accurately account for  $L_{SR}$  can significantly influence  $R_{rs}$ , resulting in inaccurate water quality estimates once algorithms are applied (Su, 2017; Zeng et al., 2017).

The objective of this paper is to evaluate the efficacy of methods that remove  $L_{SR}$  from total UAS radiance measurements in order to derive more accurate remotely sensed retrievals of scientifically valuable in-water constituents. UAS derived radiometric measurements are evaluated against *in situ* hyperspectral  $R_{rs}$  measurements to determine the best performing method of estimating and removing surface reflected light and derived water quality estimates.

**TABLE 1** | Summary of existing UAS aquatic remote sensing literature including the UAS sensor(s) used, radiometric quantity studied (where  $R_{rs}$  represents UAS derived remote sensing reflectance and  $R_{UAS}$  represents UAS derived total reflectance), whether the study accounted for surface reflected radiance ( $L_{SR}$ ), and the water quality parameter(s) derived.

Reference	UAS sensor(s)	Radiometric quantity	Removal of $L_{SR}$ ?	WQ parameter(s)
Zeng et al. (2017)	Ocean optics STS-VIS spectrometers (hyperspectral)	$R_{UAS}$	No	Chl <i>a</i> , CDOM, turbidity
Shang Z. et al. (2017)	AvaSpec-dual spectroradiometers (hyperspectral)	$R_{rs}$	Yes	Chl <i>a</i>
Su, (2017)	Canon powershot S110 RGB and NIR sensors	$R_{UAS}$	Yes	Chl <i>a</i> , secchi disk depth, turbidity
Choo et al. (2018)	MicaSense RedEdge and DLS (multispectral)	$R_{UAS}$	No	Chl <i>a</i>
Baek et al. (2019)	MicaSense RedEdge and DLS (multispectral)	$R_{rs}$	Yes	Chl <i>a</i>
Becker et al. (2019)	Ocean optics STS-VIS spectrometers (hyperspectral)	$R_{UAS}$	No	Cyanobacteria index, chl <i>a</i> TSS
Arango and Nairn. (2019)	MicaSense RedEdge and DLS (multispectral)	$R_{UAS}$	No	Secchi disk depth, chl <i>a</i> , TSS, TN, TP
Olivetti et al. (2020)	Parrot sequoia (multispectral)	$R_{UAS}$	No	TSS
McEliece et al. (2020)	Sentera multispectral sensor (4 visible bands)	$R_{UAS}$	No	Chl <i>a</i> , turbidity
Kim et al. (2020)	MicaSense RedEdge-M and DLS (multispectral)	$R_{rs}$	Yes	Chl <i>a</i> (but not focus of paper)
Castro et al. (2020)	MicaSense RedEdge and DLS (multispectral)	$R_{rs}$	No	Chl <i>a</i>
O'Shea et al. (2020)	Resonon Pika L spectrometer (hyperspectral) *deployed on a tower, not UAS	$R_{rs}$	Yes	Chl <i>a</i>

## Background/Theory

If a water surface was perfectly flat, incident light would reflect specularly and could be measured with known viewing geometries. This specular reflection of a level surface is known as the Fresnel reflection; however, most water bodies are not flat as winds and currents create tilting surface wave facets. Due to differing orientation of wave facets reflecting radiance from different parts of the sky,  $L_{SR}$  can vary widely within a single image. A common approach to model  $L_{SR}$  is to express it as the product of sky radiance ( $L_{sky}$ ,  $W\ m^{-2}\ nm^{-1}\ sr^{-1}$ ) and  $\rho$ , the effective sea-surface reflectance of the wave facet (Mobley, 1999; Lee et al., 2010):

$$L_{SR}(\theta, \Phi, \lambda) = \rho(\theta, \Phi, \lambda) * L_{sky}(\theta, \Phi, \lambda) \quad (4)$$

Rearranging Eqs. 3 Eqs. 4,  $\rho$  can be derived by:

$$\rho(\theta, \Phi, \lambda) = \frac{L_T(\theta, \Phi, \lambda) - L_w(\theta, \Phi, \lambda)}{L_{sky}(\theta, \Phi, \lambda)} \quad (5)$$

Given measurements of  $L_{sky}$ , an accurate determination of  $\rho$  is critical to derive  $R_{rs}$  by:

$$R_{rs}(\theta, \Phi, \lambda) = R_{UAS}(\theta, \Phi, \lambda) - \frac{L_{sky}(\theta, \Phi, \lambda) * \rho(\theta, \Phi, \lambda)}{E_d(\lambda)} \quad (6)$$

Methods using the statistics of the sea surface, for a given wind vector, can be used to predict  $\rho$ ; tabulated values have been derived from numerical simulations with modelled sea surfaces, Cox and Munk wave states (wind), and viewing geometries (Cox and Munk, 1954; Mobley, 1999; Mobley, 2015). Mobley (1999) provides the recommendation of collecting radiance measurements at viewing directions of  $\theta = 40^\circ$  from nadir and  $\phi = 135^\circ$  from the sun to minimize the effects of sun glint and nonuniform sky radiance with a  $\rho$  value of 0.028. The suggested viewing geometries and  $\rho$  value from Mobley (1999) have been used to estimate and remove  $L_{SR}$  in UAS remote sensing studies (Ruddick et al., 2006; Shang S. et al., 2017; Baek et al., 2019; Kim et al., 2020). However, more recent studies have shown that  $\rho$  can be spectrally dependent due to the degree of sky polarization and the ratio of diffuse to direct light; assuming a spectally constant  $\rho$  in a roughened sea surface with multi-angular wave facets can lead to erroneous estimates of  $L_{SR}$  (Lee et al., 2010; Mobley, 2015). Lee et al. (2010) proposed a spectral optimization approach using spectral inherent optical properties to model  $R_{rs}$ , which has been applied in UAS remote sensing studies (O'Shea et al., 2020).

An alternative method to remove  $L_{SR}$  relies on the so-called dark pixel assumption that assumes  $L_w$  in the NIR is negligible due to strong absorption of water. Where this assumption holds, at-sensor radiance measured in the NIR is solely  $L_{SR}$  (Gordon and Wang, 1994; Siegel et al., 2000) and allows  $\rho$  to be calculated if  $L_{sky}$  is known. Studies have used this assumption to estimate and remove  $L_{SR}$ ; however, the assumption tends to fail in more turbid waters where high concentrations of particles enhance backscattering and  $L_w$  in the NIR (Siegel et al., 2000; Lavender et al., 2005). Other novel image-processing techniques such as nonlocal mean filtering and a matching pixel by pixel algorithm have been proposed to reduce the effects of  $L_{SR}$  variation from UAS imagery; however, these region-specific methods exhibit some limitations for applications in other water bodies (Su, 2017; Totsuka et al., 2019).

## METHODS

### Study Area

This study was conducted in the Choptank River, a major tributary of the Chesapeake Bay (United States). The 1,756 km<sup>2</sup> coastal plain watershed is dominated by agriculture and forest with a relatively low population density, transitions from non-tidal, freshwater reaches to a brackish, tidal mouth, and is eutrophic (Fisher et al., 2021). Data from eight stations located downriver and extending from the mouth of the Choptank River were collected on September 16, 2020 with a relatively clear sky, data from eight stations located up river were collected on October 1, 2020 with varying cloud cover, and data from thirteen stations located in between the downriver and upriver stations were collected on October 22, 2020 with a clear sky (Figure 1).

### In situ $R_{rs}$

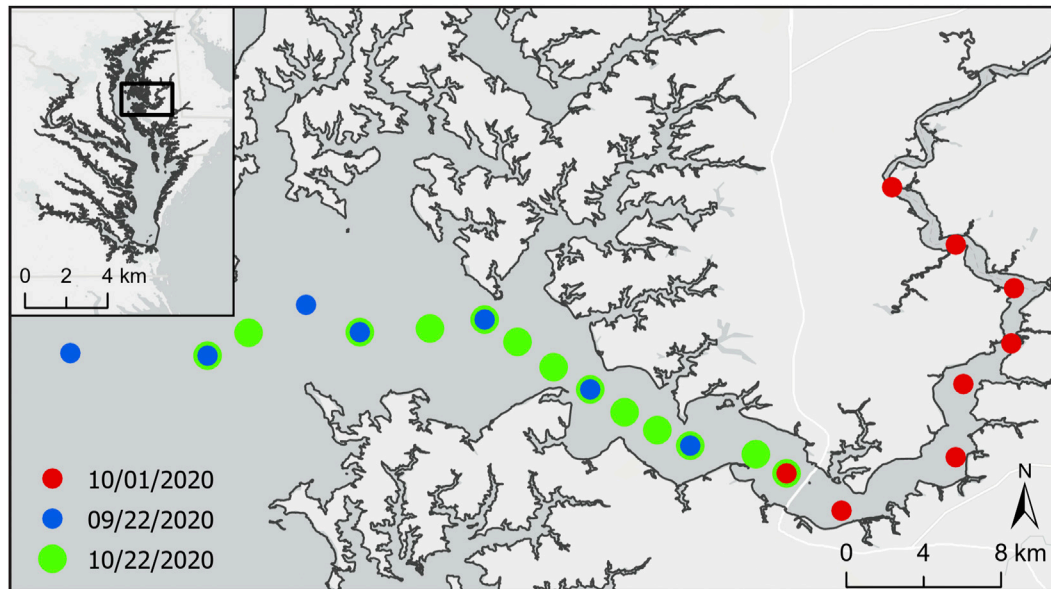
At each station, a set of hyperspectral radiometers (TriOS RAMSES; Rastede DE) deployed on a float provided *in situ*  $R_{rs}$  measurements at every station. The float held a downwelling irradiance sensor ( $E_d$ ) and an inverted radiance sensor ( $L_w$ ) positioned above the surface water and with a small black plastic cone that extended just below the surface to block  $L_{SR}$  (i.e. light blocking technique, Ahn et al., 1999; Lee et al., 2019). The TriOS radiometers collect 256 wavelength bands at 3.3 nm intervals within the 320–950 nm range. All *in situ* hyperspectral measurements were interpolated at 1 nm intervals. In order to compare these measurements to the UAS, spectral response functions (SRFs) of the five MicaSense wavebands were applied to the *in situ* hyperspectral  $R_{rs}$  data (Figure 2).

### In situ Water Quality Data

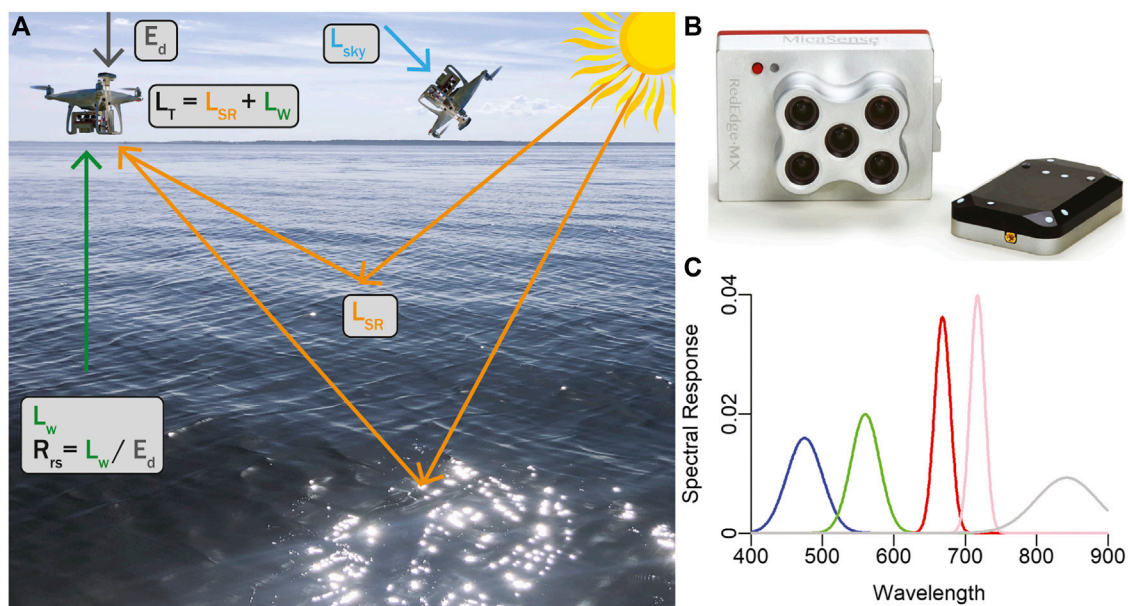
At each station, surface water grab samples were collected to measure chlorophyll *a* and total suspended solids (TSS) concentrations. Chlorophyll *a* concentration was measured in duplicate following EPA method 445.0 (Arar and Collins, 1997). Briefly, 100 ml of water from each station was filtered on 47 mm GF/F filters, immersed in 20 ml of 90% acetone and placed in a dark freezer for 24 h. Chlorophyll *a* fluorescence was measured on 5 ml of extract using a fluorometer (Turner 10 AU Fluorometer, San Jose CA) calibrated against chlorophyll *a* pigment standards (DHI, Horsholm DK) before and after acidification with 0.1 ml of 6 N hydrochloric acid. TSS was measured using a gravimetric analysis (American Public Health Association (APHA), 1995). 350 ml of water from each station was filtered using pre-weighed and combusted (450°C) GF/F filters and placed into a 105°C drying oven for at least 2 h. Filters were reweighed and the concentration was calculated by  $TSS\ (mg/L) = W_{post(g)} - W_{combust(g)} \times 1,000/V(L)$ .

### Multispectral Sensor: MicaSense RedEdge-MX Sensor

The MicaSense RedEdge-MX sensor (MicaSense, Seattle, Washington, United States) is a  $8.7 \times 5.9 \times 4.5$  cm multispectral camera capable of capturing five simultaneous bands on the electromagnetic spectrum in 12 bit radiometric



**FIGURE 1** | Location of Choptank River (38.63 N, -76.33 W) in Chesapeake Bay, Maryland, United States and locations of stations where data was collected.



**FIGURE 2** | (A) MicaSense RedEdge-MX multispectral sensor and downwelling light sensor (DLS) collects total radiance ( $L_T$ ) (sum of water-leaving radiance,  $L_w$  and surface reflected radiance,  $L_{SR}$ ) and downwelling irradiance ( $E_d$ ) measurements while in a low altitude flight. Sky radiance ( $L_{sky}$ ) is collected by positioning the sensor at  $40^\circ$  angle from zenith away from sun. (B) MicaSense RedEdge-MX multispectral sensor and DLS collects  $L_T$  and  $E_d$  in five wavebands: red, green, blue, red-edge, near-infrared. (C) Approximate spectral response functions of MicaSense RedEdge-MX sensor.

resolution: blue (475 nm center, 32 nm bandwidth), green (560 nm center, 27 nm bandwidth), red (668 nm center, 14 nm bandwidth), red edge (717 nm center, 12 nm bandwidth), and NIR (842 nm center, 57 nm bandwidth). The sensor has up to one capture per second, a  $47.2^\circ$  field of view with a spatial resolution of 8 cm per pixel at 120 m altitude. The sensor was mounted on a

DJI Phantom four Pro UAS using a  $10^\circ$  3D-printed mount which resulted in a direct nadir viewing angle while in flight (**Figure 2**). The sensor also includes a downwelling light sensor (DLS) which measures  $E_d$  in the same spectral wavebands during in-flight image captures. The DLS measures light incident on a diffuser, providing a downwelling hemispherical irradiance measurement

(Mamaghani and Salvaggio, 2019). The DLS was mounted above the UAS to eliminate shading and collected incident light at a 0° zenith angle. Images were collected at an average flight altitude of 70 m which resulted in 0.02 m pixel resolution (923 × 1219).

At each station, measurements of  $L_T$ ,  $L_{sky}$ , and  $E_d$  were collected by the MicaSense RedEdge-MX multispectral sensor and DLS. First, measurements of  $L_{sky}$  were collected by positioning the camera 40° from zenith with an approximate azimuthal viewing direction of 135° and taking several image captures (Figure 2).  $L_{sky}$  at each waveband was computed as the grand mean of all station-specific measurements. Next, the UAS was manually flown and images were automatically captured every 2 s. The MicaSense multispectral sensor has a built-in global positioning system (GPS) and inertial measurement unit (IMU) which recorded the positioning (latitude, longitude, altitude) and orientation (yaw, pitch, roll) of each image capture. Thirty image captures taken at the highest altitude from each station were used in subsequent analyses. This information along with solar elevation, image size, and coordinated universal time (UTC) time, were recorded in the metadata of each image capture.

## Pre-Processing

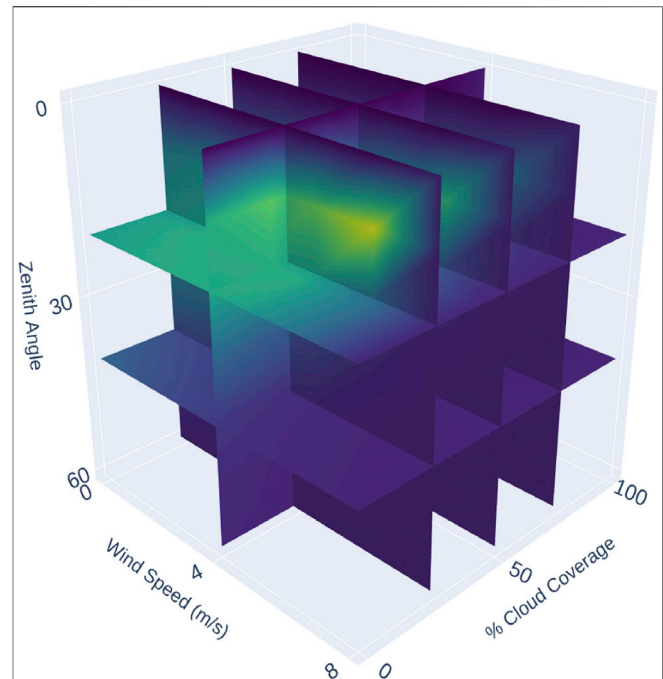
Data collected by the multispectral sensor were radiometrically calibrated using a Python workflow provided by MicaSense<sup>1</sup>. The process of converting raw pixel values (digital number, DN) into total spectral radiance ( $L_T$ ) values with units of W/m/sr/nm is given in Eq 7:

$$L_T = V(x, y) * \frac{a_1}{g} * \frac{DN - DN_{BL}}{t_e + a_2y - a_3t_e y} \quad (7)$$

The radiometric calibration compensates for dark pixel subtraction ( $DN_{BL}$ ), sensor gain ( $g$ ), exposure settings ( $t_e$ ), and lens vignette effects ( $V(x,y)$ ). Coefficients  $a_1$ ,  $a_2$ , and  $a_3$  are radiometric calibration coefficients and  $x,y$  is the pixel column and row number, respectively. Lens distortion effects, such as band-to-band image alignment, were removed from image captures by an unwarping procedure and wavebands were aligned to form a stacked TIFF for each image set.

A filtering procedure was applied to all images to remove specular sun glint and instances of the boat when it was present in the image. Studies have suggested filtering out high total radiance values, or instances of specular sun glint, before removing  $L_{SR}$  from  $L_T$  (Hooker et al., 2002). Specular sun glint arises when direct sunlight reflects off of a wave facet or surface at the viewing angle of the sensor. Sun glint pixels were masked using an empirical upper limit of  $R_{UAS}$  measurements in the NIR, and this mask was then propagated to pixels at all other wavebands. Specifically, *in situ* measurements and values from radiative transfer model simulations (see methods below) were used in Eq. 8 to solve for an upper limit of  $R_{UAS}(NIR)$  and pixels with values greater than this upper limit were masked in each waveband.

<sup>1</sup>MicaSense RedEdge Image Processing Tutorials. Retrieved online at <https://github.com/micasense/imageprocessing>.



**FIGURE 3** | Visualization of the three-dimensional  $\rho$  look-up table derived from HydroLight simulations with varying solar zenith angles, wind speed (m/s), and cloud cover (%) corresponding to a nadir viewing angle.

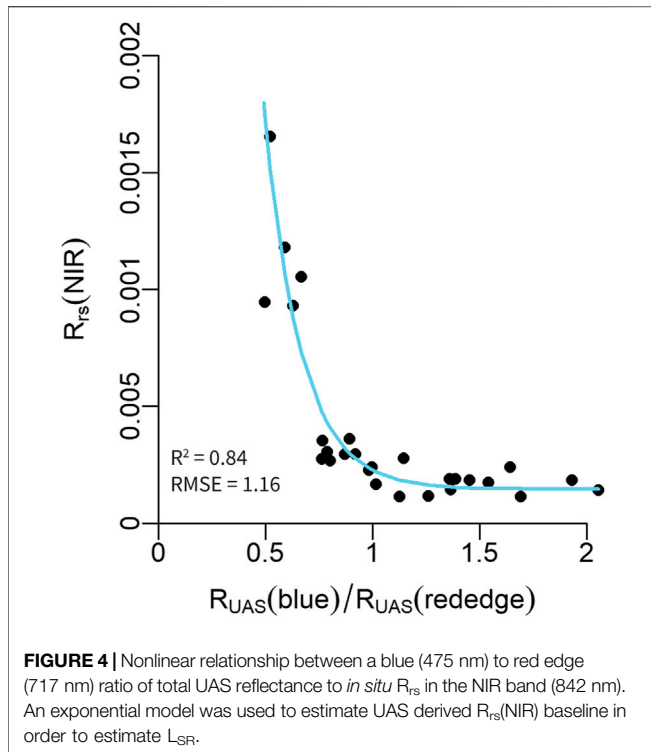
$$R_{UAS}(NIR) = R_{rs}(NIR) + \rho(NIR) * \frac{L_{sky}}{E_d} \quad (8)$$

Where  $R_{rs}(NIR) = 0.005$  is based on *in situ* measurements in this study and consistent with other turbid waters (Tzortziou et al., 2006),  $\rho(NIR)$  varied depending on station, and  $L_{sky}/E_d = 0.39$  is derived from radiative transfer model simulations (HydroLight v6.0, Numerical Optics Ltd., United Kingdom). A lower limit of  $R_{UAS}(green) = 0.007$  was used to mask out the dark canopy of the boat when present in images.

## Removal of Surface Reflected Light ( $L_{SR}$ )

At each station,  $L_{SR}$  was estimated using one of four different methods to ultimately derive  $R_{rs}$  following Eq. 6. These methods are provided below, and herein referred to as “ $\rho_{LUT}$ ,” “NIR = 0,” “NIR > 0” and “Deglinting.” Resultant UAS  $R_{rs}$  estimates were then compared to paired *in situ*  $R_{rs}$  data across stations and wavebands. Statistical evaluations to assess the relationship included root mean square error (RMSE), relative root mean squared error (RRMSE), coefficient of determination ( $R^2$ ), and  $p$ -values.

$\rho_{LUT}$ . The  $\rho_{LUT}$  method follows from Mobley (1999) and involved developing a look-up table (LUT) of  $\rho$  values using HydroLight simulations. HydroLight is a numerical model that solves the radiative transfer equation to compute the radiance distribution within and at the surface of a water body. Inputs include absorbing and scattering properties of a water body, the nature of the wind-blown sea surface (Cox-Munk sea surface slopes), and the sun and sky radiance incident on the sea surface (Mobley, 1999). Outputs include the full radiance distribution,



including an effective  $\rho$  value for a range of geometries and wavelengths. Batch HydroLight runs incorporating varying solar zenith angles (0, 10, 20, 30, 40, 50, 60°), wind speeds (0, 2, 4, 6, 8 m/s), and fractional cloud cover (0, 0.2, 0.4, 0.6, 0.8, 1%) were used to develop a multidimensional look-up table of  $\rho$  values that correspond to a nadir viewing angle (Figure 3). HydroLight returned spectrally explicit  $\rho(\theta, \phi, \lambda)$  values and MicaSense SRFs were applied to  $\rho(\theta, \phi, \lambda)$  for each waveband.  $\rho(\theta, \phi, \lambda)$  values were obtained for each station according to the solar zenith angles obtained from MicaSense metadata, wind speed was collected from the Global Forecast System<sup>2</sup>, and cloud cover was determined by the  $L_{sky}$  images. Data in this study were collected at solar zenith angles ranging from 36.5 to 58.6°, wind speeds ranged from 1.1 to 3.1 m/s and cloud cover ranged from 0 to 60%.

$NIR = 0$ .  $L_{SR}$  was also estimated using a pixel based dark pixel assumption to derive  $\rho$ . Assuming  $R_{rs}$ (NIR) equals 0, Eq. 6 can be rearranged to solve for  $\rho$  (Eqs. 9 and 10). This  $\rho$  value is used to calculate  $R_{rs}$  across all wavebands (Eq. 11).

$$0 = L_T(NIR) - L_{sky}(NIR) * \rho \tag{9}$$

$$\rho = L_T(NIR) / L_{sky}(NIR) \tag{10}$$

$$R_{rs}(\theta, \Phi, \lambda) = R_{UAS}(\theta, \Phi, \lambda) - \frac{L_{sky}(\theta, \Phi, \lambda) * (L_T(NIR) / L_{sky}(NIR))}{E_d(\lambda)} \tag{11}$$

<sup>2</sup>Wind speed data was retrieved from <https://earth.nullschool.net/about.html>.

$NIR > 0$ . Since the dark pixel assumption is invalid in turbid waters including the Chesapeake Bay (Siegel et al., 2000) an alternative pixel based approach was developed to instead estimate a baseline  $R_{rs}$ (NIR). Specifically,  $R_{rs}$ (NIR) was estimated empirically by a nonlinear regression using *in situ*  $R_{rs}$ (NIR) and  $R_{UAS}$  data (Figure 4). The ratio of  $R_{UAS}$ (blue)/ $R_{UAS}$ (rededge) led to the most robust predictions of  $R_{rs}$ (NIR) using Eq 12:

$$R_{rs}(\theta, \Phi, \lambda) = 0.025e^{-5.469 * R_{UAS}(blue) / R_{UAS}(red\ edge)} + 0.00013 \tag{12}$$

Conceptually, this relationship makes sense because with increasing particle concentrations, water becomes less blue and  $R_{rs}$ (NIR) increases.

“Deglinting.”  $L_{SR}$  was also estimated following the “deglinting” methods of Hochberg et al. (2003) and Hedley et al. (2005). For each station, a minimum NIR value was determined by finding the lowest 10% of  $R_{UAS}$ (NIR) across all images. For each band, a linear regression was made between all  $R_{UAS}$ (NIR) and  $R_{UAS}$ (visible) values and the slope ( $b_i$ ) was determined. Each pixel was corrected by subtracting the product of  $b_i$  and the NIR brightness of the pixel (Hedley et al., 2005):

$$R_{rs}(i) = R_{UAS}(i) - b_i(R_{UAS}(NIR) - \min(R_{UAS}(NIR))) \tag{13}$$

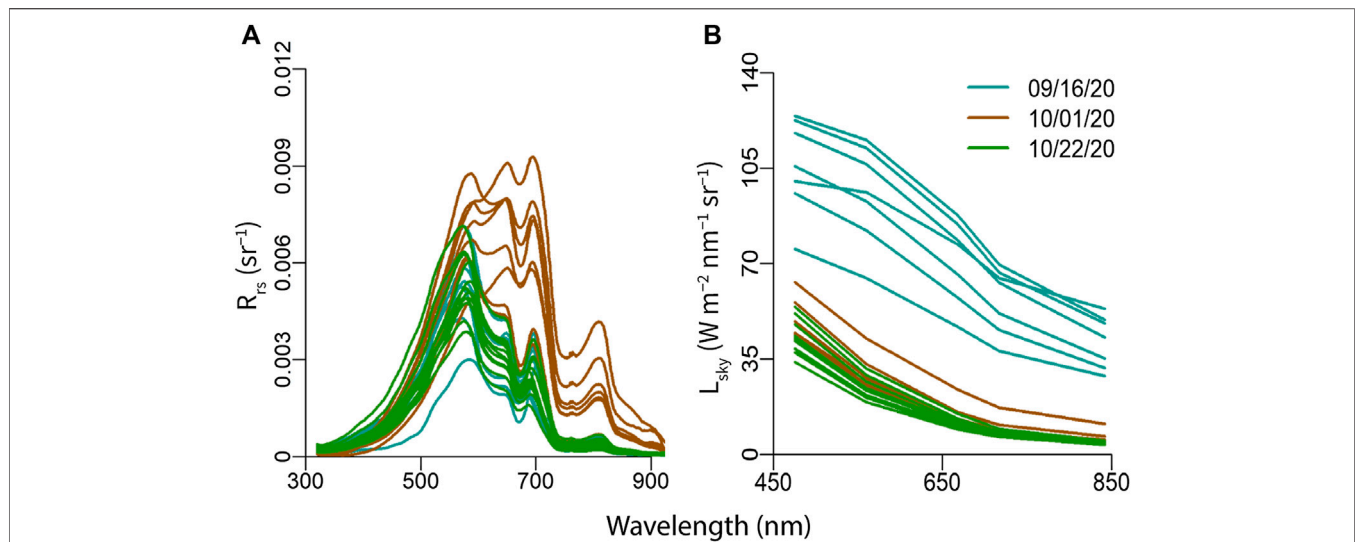
## Water Quality Retrievals

$R_{rs}$  values from each  $L_{SR}$  removal method were compared against *in situ* chlorophyll *a* and TSS data (n = 28) using multiple linear regressions. The best performing model (highest  $R^2$  and lowest RMSE, RRMSE, *p*-values) was used in a stepwise model selection by Akaike information criterion (AIC, “stepAIC” in R). The AIC stepwise regression iteratively added and removed wavelengths in order to determine the combination of data that resulted in the best performing model with low prediction error, while taking into account model simplicity.  $R_{rs}$  values were used as input into optical algorithms derived from the best performing multiple linear regressions and mean chlorophyll *a* and TSS concentration at each station was obtained by averaging values across all images. The resulting arrays were georeferenced using the Python libraries “CameraTransform” (Gerum et al., 2019) and “Rasterio” using archived metadata including latitude, longitude, image width, image height to position the images accurately in a known coordinate system (WGS84). Georeferenced arrays were exported as individual TIFFs and mapped using ArcGIS Pro (ESRI Inc. Redlands, CA, United States).

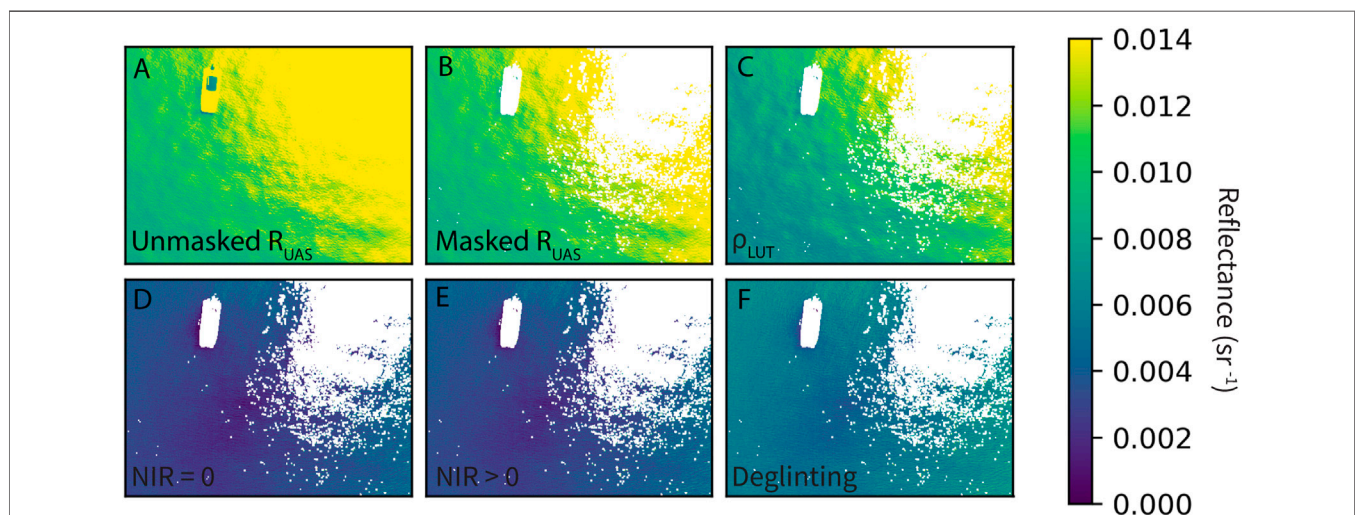
## RESULTS

### *In situ* Data

Chlorophyll *a* concentration ranged from 5.19 to 53.30 ug/L with an average of  $17.13 \pm 10.96$  ug/L. TSS concentration ranged from 19.94 mg/L to 39.69 mg/L with an average of  $28.25 \pm 5.21$  mg/L. In general, *in situ*  $R_{rs}$  spectra were representative of a eutrophic system (Figure 5, Gitelson et al., 2007; Spyarakos et al., 2018). Due to absorption of chromophoric dissolved organic matter (CDOM) and chlorophyll *a* in lower wavelengths,  $R_{rs}$  is low in the blue region with a distinct peak in the green region



**FIGURE 5 | (A)** *In situ* remote sensing reflectance ( $R_{rs}$ ) spectra collected with hyperspectral TriOS radiometers and **(B)** Sky radiance ( $L_{sky}$ ) spectra collected with MicaSense RedEdge-MX multispectral camera at stations along the Choptank River on 9/16/20 (blue), 10/01/20 (brown), and 10/22/20 (green).



**FIGURE 6 |** Example of an individual UAS image (green band) with different radiometric values: **(A)**  $R_{UAS}$ , **(B)**  $R_{UAS}$  with initial sun glint masking and **(C–F)** remote sensing reflectance ( $R_{rs}$ ) using various methods to remove surface reflected light: **(C)**  $\rho$  look-up table (LUT) from HydroLight simulations, **(D)** Dark pixel assumption with NIR = 0, **(E)** Dark pixel assumption with NIR > 0, **(F)** Deglinting methods following Hochberg et al. (2003).

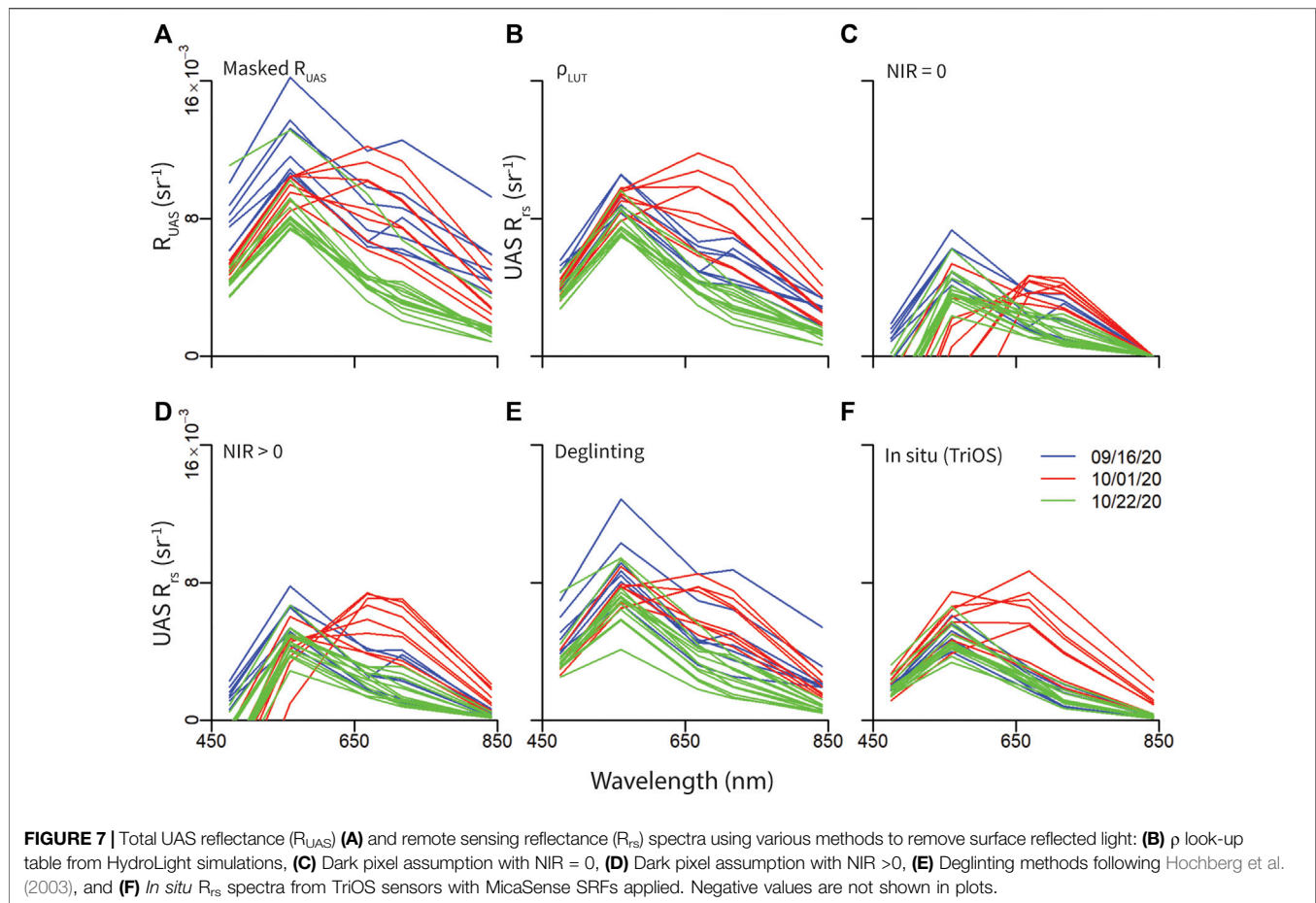
(~550 nm). A secondary peak in the red region (~700 nm) corresponds to chlorophyll fluorescence emission (Spyrakos et al., 2018). Overall, spectra collected from upriver stations on 10/01/20 were larger in magnitude and contained a tertiary peak in the NIR region (~800 nm).

### Removal of Surface Reflected Light ( $L_{SR}$ )

UAS measurements of  $L_{sky}$  were approximately proportional to the fourth power of the wavelength ( $L_{sky} \approx \lambda^{-4}$ , Figure 5) and were used with estimations of  $\rho$  to calculate and remove  $L_{SR}$  from  $R_{UAS}$  measurements at each station. Figure 6 illustrates the initial masking and methods to remove  $L_{SR}$  using an individual UAS

image taken in the green band. When  $L_{SR}$  was estimated using the  $\rho$  look-up table ( $\rho_{LUT}$ ),  $R_{rs}$  values did not change much from  $R_{UAS}$  values; however,  $R_{rs}$  values using the pixel-based dark pixel assumptions (NIR = 0, NIR > 0) and “deglinting” approach declined and became more homogenous across pixels (Figure 6).

UAS derived reflectance measurements from each  $L_{SR}$  removal method were plotted as spectra (Figure 7) and compared against *in situ*  $R_{rs}$  values (Figure 8). UAS reflectance spectra are similar in shape to *in situ*  $R_{rs}$  spectra and display a distinct peak in the green band (560 nm) and often a secondary peak in the red band (668 nm) corresponding to chlorophyll *a* reflectance and fluorescence,



respectively (Figure 7). Spectra collected on the second sampling day (upriver sites) shift to longer wavelengths likely due to scattering of inorganic particles (Figure 7).  $R_{UAS}$  spectra are higher in magnitude and are overestimated when compared to *in situ*  $R_{rs}$  due to the inclusion of  $L_{SR}$  ( $R^2 = 0.55$ ,  $RMSE = 0.004$ ,  $p < 0.05$ ). When  $L_{SR}$  is estimated using the  $\rho_{LUT}$ , spectra are similar in shape to *in situ* spectra, but are higher in magnitude and overestimated compared to *in situ* measurements ( $R^2 = 0.85$ ,  $RMSE = 0.003$ ,  $p < 0.05$ ). When  $L_{SR}$  is estimated with the dark pixel assumption (NIR = 0),  $R_{rs}$  spectra are lowest in magnitude, had the poorest fit to *in situ* data ( $R^2 = 0.18$ ,  $RMSE = 0.004$ ,  $p < 0.05$ ), and produced negative  $R_{rs}$  values in the lower wavelengths. This is not surprising given that the *in situ* measurements in the NIR (Figure 5A) are not negligible which leads to an overestimation in  $L_{SR}$ . When  $L_{SR}$  is estimated using a baseline NIR value (NIR >0),  $R_{rs}$  spectra are higher in magnitude with a lower error but still indicate negative values in the lower wavelengths, indicating  $L_{SR}$  was still slightly overestimated ( $R^2 = 0.50$ ,  $RMSE = 0.002$ ,  $p < 0.05$ ).  $R_{rs}$  spectra with the “Deglinting” approach are similar in shape to *in situ* spectra but are slightly higher in magnitude (Figure 7). This approach contained the second highest correlation and lowest RMSE ( $R^2 = 0.65$ ,  $RMSE = 0.002$ ,  $p < 0.05$ ), but still tended to overestimate  $R_{rs}$  values (Figure 8).

## Water Quality Retrieval Algorithms

$R_{rs}$  measurements with  $L_{SR}$  estimated using a baseline NIR value (NIR >0) performed best when compared to *in situ* chlorophyll *a* data ( $R^2 = 0.37$ ,  $RMSE = 5.89$ ,  $RRMSE = 37\%$ ,  $p < 0.05$ ). A stepwise model selection by AIC demonstrated that the green, red edge, and NIR bands were most important in estimating chlorophyll *a* concentration (Figure 9,  $R^2 = 0.42$ ,  $RMSE = 5.90$ ,  $RRMSE = 37\%$ ,  $p < 0.05$ ) and a remotely sensed chlorophyll *a* algorithm was determined as:

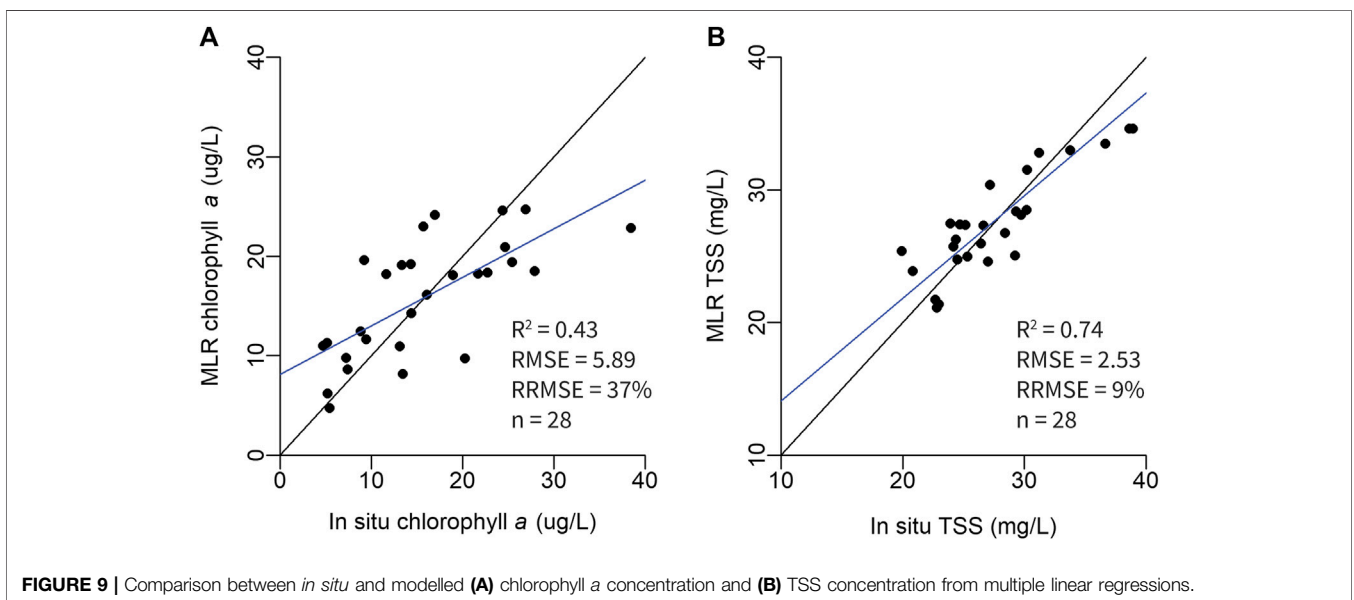
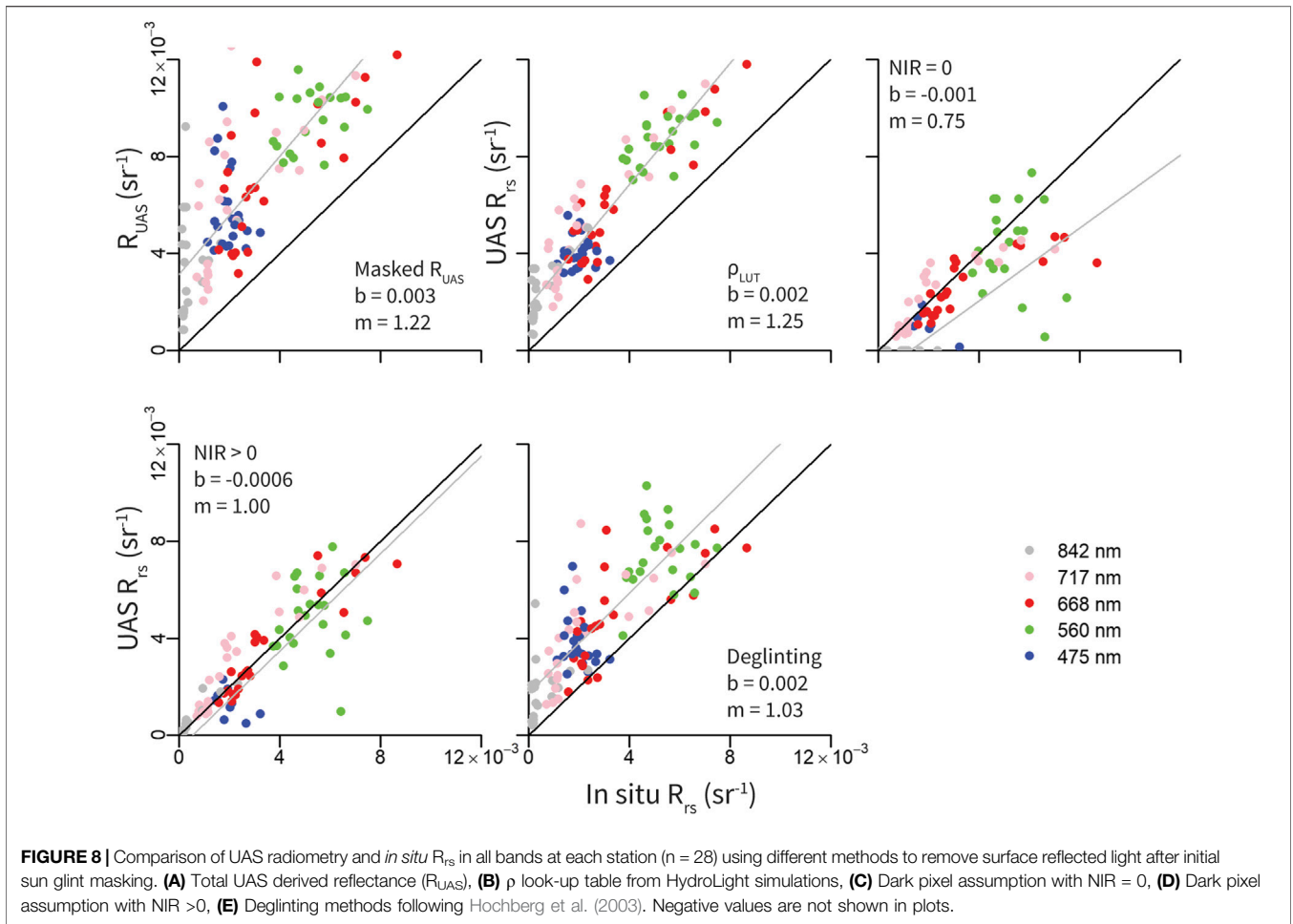
$$\text{Chlorophyll } a \text{ (ug/L)} = 24.02 + R_{rs}(560) * -4337.88 + R_{rs}(717) * 9639.75 + R_{rs}(842) * -2922.80$$

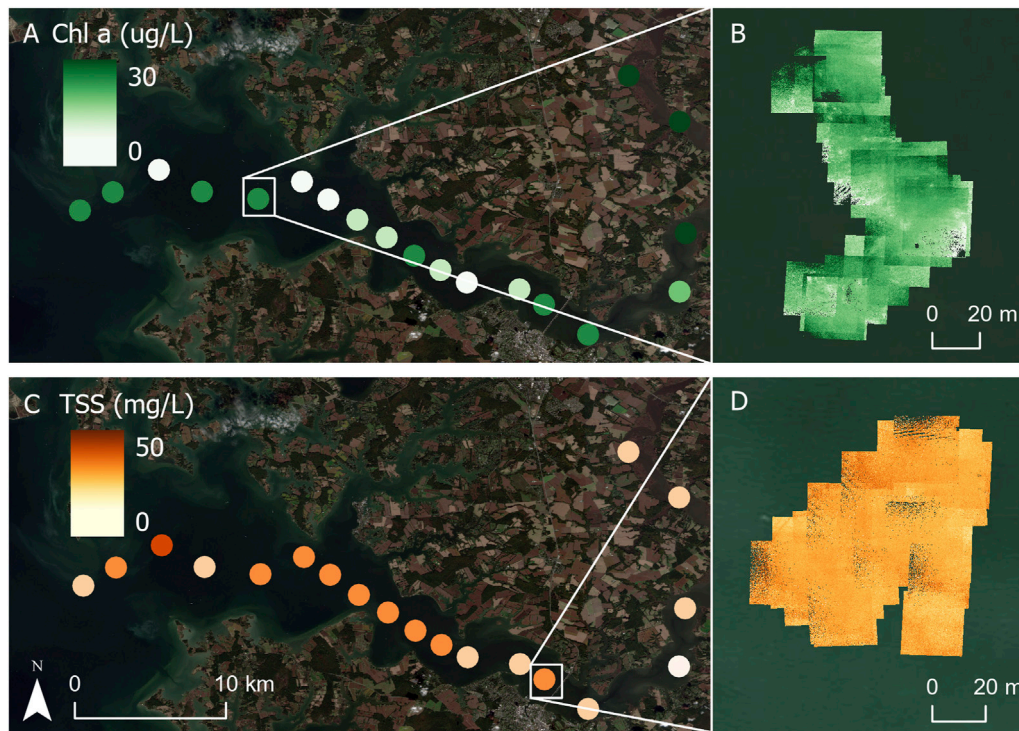
$R_{rs}$  measurements with the “deglinting” technique performed best when compared to *in situ* TSS data ( $R^2 = 0.72$ ,  $RMSE = 2.53$ ,  $RRMSE = 9\%$ ,  $p < 0.05$ ). A stepwise model selection by AIC demonstrated that the blue, red, red edge, and NIR bands were most important in estimating TSS concentration (Figure 9,  $R^2 = 0.73$ ,  $RMSE = 2.53$ ,  $RRMSE = 9\%$ ,  $p < 0.05$ ) and a remotely sensed TSS algorithm was determined as:

$$\text{TSS (mg/L)} = 30.57 + R_{rs}(475) * 1364.86 + R_{rs}(668) * -5255.88 + R_{rs}(717) * 2548.08 + R_{rs}(842) * 4579.36$$

Algorithms were applied to respective UAS derived  $R_{rs}$  values and average chlorophyll *a* and TSS concentrations were mapped, along with mosaiced georeferenced TIFFs of







**FIGURE 10 |** (A) Average chlorophyll *a* concentration across image captures ( $n = 30$ ) at each station. (B) Example of mosaiced georeferenced TIFFs collected at one station with chlorophyll *a* algorithm applied. (C) Average TSS concentration across image captures ( $n = 30$ ) at each station. (D) Example of mosaiced georeferenced TIFFs ( $n = 30$ ) collected at one station with TSS algorithm applied.

individual image captures (**Figure 10**). Average chlorophyll *a* concentration tended to increase at upriver sites while average TSS concentrations were higher at downriver sites (**Figure 10**). This trend was also seen in the *in situ* data. Within the mosaic of individual TIFFs, slight variations in chlorophyll *a* and TSS concentration are visible (**Figure 10**).

## DISCUSSION

This study is one of the first to compare methods to remove surface reflected light from high resolution UAS multispectral measurements. An initial filtering procedure and four methods to remove  $L_{SR}$  were evaluated against *in situ*  $R_{rs}$  measurements. Water quality algorithm performance is compared to those found in the literature. Results have broad implications for improving UAS derived water quality measurements in coastal waters, but several aspects and caveats of this study merit additional discussion.

### Performance of Methods to Remove Surface Reflected Light ( $L_{SR}$ )

Removing surface reflected light from the total radiance measured by a UAS ensures only water-leaving information is used as input into water quality retrievals. Initial filtering

techniques successfully masked instances of specular sun glint and non-water objects (i.e. boat) in the UAS imagery using empirical upper and lower  $R_{rs}$  limits. This primary step ultimately improved final  $R_{rs}$  measurements and derived water quality products. While the magnitude of  $L_{sky}$  can impact  $L_{SR}$ , the spectral dependence ( $\sim \lambda^{-4}$ ) is constrained and can be measured; therefore, the effective sea-surface reflectance of a wave facet ( $\rho$ ) needs to be accounted for. Nonetheless, it is important to consider the impact of the varying absolute magnitude of  $L_{sky}$  and future work should analyze the effect of  $L_{sky}$  variability on UAS  $R_{rs}$  measurements.

The  $\rho_{LUT}$  method utilized a  $\rho$  lookup table approach developed from HydroLight simulations where  $\rho$  values were obtained depending on the solar zenith angle, wind speed, and cloud cover at the time of UAS data collection. Mobley (1999) HydroLight simulations led to recommendations of specific sensor viewing angles with a corresponding constant  $\rho$  value to reduce effects of  $L_{SR}$ ; however, most consumer grade UAS sensors and mounts do not have the ability to change viewing angles and are fixed at a nadir viewing angle. Therefore, the  $\rho$  values in the look-up table correspond to nadir viewing angles. This method resulted in UAS  $R_{rs}$  values that were generally greater than *in situ*  $R_{rs}$  values, indicating  $L_{SR}$  was underestimated. This can most likely be attributed to a constant  $\rho$  value that was used to estimate  $L_{SR}$  across all pixels in each image, which averaged out the multifaceted characteristic of a water surface.

The NIR = 0 and NIR >0 methods took advantage of the NIR waveband on the multispectral sensor and incorporated aspects of the dark pixel assumption which allowed  $L_{SR}$  to be removed from visible wavelengths by assuming that 1) a NIR signal is composed of  $L_{SR}$  and/or a spatially homogeneous NIR component in the water and 2) the amount of  $L_{SR}$  in the NIR band is linearly related to the amount of  $L_{SR}$  in the visible bands (Mobley, 1999; Siegel et al., 2000). Unlike the  $\rho_{LUT}$  method that applies a single value to an entire image, the dark pixel assumption approaches estimate  $L_{SR}$  on a per pixel basis which decreases resultant  $R_{rs}$  variability (Figures 6, 8). When NIR was assumed to be entirely negligible, calculated  $R_{rs}$  values were generally lower compared to *in situ*  $R_{rs}$  which can be attributed to an enhancement of NIR due to the presence of scattering inorganic particles in turbid waters. Thus, when a baseline NIR value was estimated, the method performed better (higher  $R^2$ , lower RMSE and RRMSE); however, negative  $R_{rs}$  values in the lower wavelengths indicated  $L_{SR}$  was still likely overestimated at some stations.

The “deglinting” method, modelled off a glint removal algorithm introduced by Hochberg et al. (2003), and later Hedley et al. (2005), calculated a constant ‘ambient’ NIR brightness level which was removed from all pixels in all wavelengths. This approach resulted in UAS  $R_{rs}$  measurements slightly overestimated compared to *in situ* values but produced the highest  $R^2$  and lowest RMSE of all methods. This method was applied by averaging pixel values across all thirty images collected at each station and is likely to perform better when applied on individual images.

The  $\rho_{LUT}$  method uses  $\rho$  values that are based on probability distributions of sea surface slopes which are related to the scale of satellite pixel resolution (100–1000 m) (Cox and Munk, 1954). For images with much higher pixel resolution (>1–10 m), statistical assumptions about the surface of water composed of many reflecting facets are less likely to hold (Kay et al., 2009). Therefore, to collect accurate water-leaving reflectance measurements from high resolution UAS imagery, it is recommended to use a pixel-based approach exploiting the high absorption of water at NIR wavelengths to estimate and remove  $L_{SR}$ . If a baseline NIR measurement can be retrieved, the dark pixel assumption should be used to remove  $L_{SR}$ . Otherwise, the ‘deglinting’ methods following Hochberg et al. (2003) and Hedley et al. (2005) are recommended. In open ocean waters without a strong influence of optically active properties, a pixel-based approach assuming NIR is negligible is expected to perform well.

It is important to note a potential caveat to the  $L_{SR}$  method evaluation. In this study, *in situ*  $R_{rs}$  measurements were provided by hyperspectral radiometers with a skylight blocking approach (Ahn et al., 1999; Lee et al., 2019). This approach consists of attaching an open-ended apparatus, or tube, to the front of a downward looking radiance radiometer and lowering it a few centimeters into the water, blocking surface-reflected light and allowing for a direct measurement of  $L_w$ . It is important to note that measurements from this technique are subject to instrument self-shading, which is a function of the water’s optical properties, sun elevation, and the size of the skylight-blocking cone (Zhang et al., 2017; Lee et al., 2019). Zhang et al. (2017) estimated that

self-shading accounts for approximately 1–20% error under most water properties and solar positions. Methods to correct for this self-shading have been derived (Zhang et al., 2017; Yu et al., 2021) and if applied, have the potential to improve relationships with UAS  $R_{rs}$  measurements.

## Performance of Water Quality Algorithms

Performance of the multiple linear regressions developed in this study were compared to existing chlorophyll *a* and TSS algorithms designed for coastal waters to determine if UAS measurements can produce accuracy within the range of other water quality algorithms. Performance of the UAS derived chlorophyll *a* multiple linear regression ( $R^2 = 0.43$ , RRMSE = 37%) is comparable to other chlorophyll *a* algorithms found in the literature (Ruddick et al., 2001; Gons et al., 2002; Gitelson et al., 2007). A three-band chlorophyll algorithm calibrated using a variety of coastal waters, including the Choptank River, resulted in a RRMSE of 51.9% (Gitelson et al., 2007), a two-band algorithm (red/NIR) with adaptive optimization in the second band calibrated with measurements from the North Sea and Lake IJless, Netherlands resulted in a RRMSE of 37% (Ruddick et al., 2001), and a two-band algorithm (red/NIR) designed for the Medium Resolution Imaging Spectrometer (MERIS) satellite sensor and calibrated using a variety of coastal and inland waters including Lake IJssel (Netherlands), the Chinese Lake Tai Hu, and the Hudson/Raritan Estuary (New York/New Jersey) resulted in a standard error of 9.2  $\mu\text{g/L}$  (Gons et al., 2002). Performance of the UAS derived TSS multiple linear regression ( $R^2 = 0.73$ , RMSE = 2.53, RRMSE = 9%) is also comparable to existing TSS algorithms found in the literature (Nechad et al., 2010; Novoa et al., 2017). Algorithms including a single-band (NIR) second-order polynomial and single-band (red/green) linear models calibrated with measurements from the Gironde Estuary, France resulted in RRMSE values ranging from of 9.11–16.41% (Novoa et al., 2017) and a non-linear regression calibrated with measurements from the Southern North Sea resulted in RRMSE values less than 30% (Nechad et al., 2010). Future work will include improving water quality algorithms.

## UAS Sensor Considerations

The innovative use of UAS technology for environmental research is a relatively new field and researchers are only beginning to understand and alleviate the various methodological and sensor performance challenges. Sensors degrade over time from use and environmental conditions which can impact the accuracy of the data being collected. The most recent MicaSense model, the MicaSense RedEdge-MX, is integrated with low-cost, image-frame complementary metal-oxide semiconductor (CMOS) sensors, which compared to typical charged-coupled device (CCD) sensors, tend to generate more noise and have lower sensitivity levels (Mamaghani and Savaggio, 2019). In the present study, raw values were radiometrically calibrated using a workflow<sup>1</sup> provided by Micasense which implements default metadata parameters that remain the same unless a new factory calibration is performed. Due to sensor degradation, these values are likely to gradually

decline over time, lessening the accuracy of the radiometric calibration processing (Mamaghani and Savaggio, 2019). Studies have improved sensor performance by performing vicarious radiometric calibration using ground targets and panels with known radiometric accuracy, calibrating sensors using National Institute of Standards (NIST)-traceable equipment in a laboratory, and developing look-up tables for correction factors to update calibration parameters (Del Pozo et al., 2014; Mamaghani and Salvaggio, 2019; Cao et al., 2020). Baek et al. (2020) conducted an assessment on radiometric accuracy for the MicaSense RedEdge-MX sensor by comparing data to hyperspectral sensors with NIST-traceable calibration (TriOS RAMSES) and showed that MicaSense RedEdge-MX radiance is approximately 5–16% lower, and irradiance is approximately 1–20% lower, depending on wavelength (Baek et al., 2020). The radiometric accuracy of a new or recently/vicariously calibrated UAS sensor should meet the required radiometric accuracy of 5% that is expected with ocean color satellites (McClain et al., 1992).

Precise registration of multispectral bands within a UAS image capture is also important to derive accurate spectral radiometric values across pixels. Sources of misregistration include a difference in the lens location for each band, image acquisition times, and exposure times which can all influence  $R_{rs}$  and resulting water quality variables (Kim et al., 2020). In the present study, images collected in each band were registered using MicaSense's default alignment function<sup>1</sup>. This three step process unwraps images using built-in lens calibration, determines a transformation to align each band to a common band, and crops pixels which do not overlap in all bands. Although this method seemed to perform well in this study; it is acknowledged that this type of band registration can perform poorly with images of locational errors, such as moving water, and can produce noise even after removing surface reflected light (Kim et al., 2020). Kim et al. (2020) developed a novel morphological band registration technique designed for high resolution water quality analysis which effectively removes misregistration noise and improves the accuracy of  $R_{rs}$ . Future aquatic UAS remote sensing work should consider adapting this technique to improve UAS remotely sensed retrievals over water.

## Caveats and Considerations

Many UAS aquatic remote sensing studies use Structure-from-Motion (SfM) photogrammetric techniques to stitch individual UAS images into ortho- and georectified mosaics (Arango and Nairn, 2019; Castro et al., 2020; McEliece et al., 2020; Olivetti et al., 2020). This approach applies matching key points from overlapping UAS imagery in camera pose estimation algorithms to resolve 3D camera location and scene geometry (Westoby et al., 2012; Arango et al., 2020). Commonly used software (e.g. Pix4D) provide workflows that radiometrically calibrate, georeference, and stitch individual UAS images using a weighted average approach to create at-sensor reflectance 2D orthomosaics (Olivetti et al., 2020).  $L_{SR}$  removal methods and water quality algorithms can be directly applied to reflectance orthomosaics to effectively derive water quality products of an entire water body. However, current photogrammetry techniques

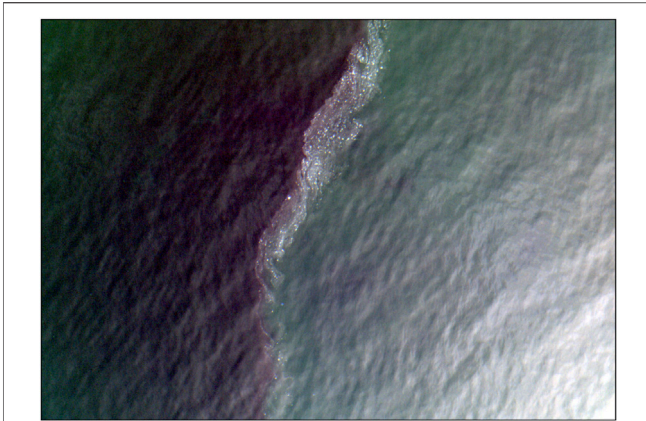
are not capable of stitching UAS images captured over large bodies of water due to a lack of key points in images of homogenous water surfaces (Arango et al., 2020). Orthomosaics of smaller water bodies or rivers can be created if UAS images contain enough surrounding land features containing keypoints that the photogrammetry software can use to successfully stitch the images containing water. This can be accomplished by increasing flight altitude, with the trade-off of lower spatial resolution. Alternative methods include a statistical interpolation method; however interpolated reflectance values can be imprecise when compared to true reflectance values (Arango et al., 2020).

## Management Implications and Future Research

Water quality monitoring is important for tracking water quality trends, identifying and mitigating pollution sources, and discerning potential human health risks. Traditional *in situ* based methods of sampling at discrete stations can be expensive due to high costs for boat time and analysis and can also potentially omit important water quality phenomena. Traditional satellite remote sensing can capture variability throughout time and space; however, limitations including the presence of clouds, atmospheric effects, land adjacency effects, and spatial resolution can hinder periodic monitoring (Shi and Wang, 2009; Becker et al., 2019). UAS fill an operational gap between *in situ* and satellite remote sensing methods. While the current available commercial multispectral UAS sensor technology is geared toward terrestrial applications, mostly precision agriculture, the spectral bands have been useful in retrieving water quality parameters in aquatic water bodies (Choo et al., 2018; Arango and Nairn, 2019; Baek et al., 2019; Castro et al., 2020; Olivetti et al., 2020). A UAS sensor package designed for aquatic environments would undoubtedly improve remotely sensed retrievals and water quality measurements.

Since UAS are deployed at a low altitude, atmospheric corrections and remedies to the land adjacency effect are eliminated. UAS are rapidly deployable and can provide the spatial and temporal variability required for useful water quality monitoring in a dynamic and rapidly evolving environment. UAS can enhance fine-scale physical oceanography research by resolving small-scale phenomena and physical processes such as patchy algal blooms, frontal structures, and turbulence characteristics (Figure 11, Shang Z. et al., 2017; Osadchiev et al., 2020). UAS remotely sensed water quality retrievals will also likely improve with the development of lightweight, off-the-shelf hyperspectral sensors, allowing for higher spatial and spectral resolution to better distinguish optical properties of the water (Shang S. et al., 2017; O'Shea et al., 2020).

Future research in turbid coastal waters may benefit from sensor packages that use longer wavelengths (e.g. shortwave infrared, SWIR). SWIR wavelengths have shown to be more effective in satellite atmospheric correction techniques in turbid waters due to the stronger water absorption relative to NIR (Shi and Wang, 2009). This will be advantageous for removing  $L_{SR}$  and will improve  $R_{rs}$  retrievals in coastal, turbid waters. Future research could also consider combining UAS radiometry with photogrammetry



**FIGURE 11** | Frontal process observed in UAS RGB imagery collected in the Choptank River, Maryland, United States with MicaSense RedEdge-MX multispectral camera.

computer vision to estimate  $L_{SR}$ . Schneider-Zapp et al. (2019) developed a method to estimate a hemispherical-directional reflectance factor (HDRF) from multi-angular UAS measurements. From a combination of photogrammetry and radiometry, a precise estimation of the downwelling light sensor position and orientation can be used to derive a multi-angular reflectance factor, which has the potential to significantly improve estimates of the multifaceted aspect of  $L_{SR}$ . An alternative approach to reduce  $L_{SR}$  is installing additional hardware to a UAS sensor. O'Shea et al. (2020) tested a hardware-based vertical polarizer on a hyperspectral spectrometer which effectively blocked horizontally polarized reflected skylight. This relatively simple approach could be an attractive solution to researchers and managers who are interested in applying a single algorithm to  $R_{rs}$  values to derive water quality parameters and should be further investigated.

## CONCLUSION

UAS-based applications of multispectral or hyperspectral remote sensing in aquatic remote sensing have the potential to effectively fill current observation gaps in aquatic remote sensing and provide critical information needed for water quality forecasting, ecosystem monitoring, and ultimately climate change research. While atmospheric effects can usually be ignored in low altitude UAS flights, the effect of sun glint and surface reflected light should be accounted for in order to obtain the highest accuracy of water quality data. The inclusion of surface reflected light can lead to an overestimation of  $R_{rs}$  and

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remotely sensed water quality retrievals. This study presents a comparison of four approaches to remove sun glint and surface reflected light that can be applied to UAS remote sensing to derive water quality parameters such as chlorophyll *a* and TSS concentration. Overall, the performance of the MicaSense RedEdge-MX multispectral sensor appears sufficient for providing high resolution water quality estimates of coastal water bodies when surface reflected light is removed. Of the four approaches examined, a pixel-based deglinting procedure utilizing the brightness of the NIR band performed best when compared to *in situ*  $R_{rs}$  measurements. This method also led to the best estimates of TSS while a pixel-based approach utilizing an ambient NIR signal to estimate and remove surface reflected light led to the best estimates of chlorophyll *a* concentration. Future work will include improving algorithms for water quality parameters. Future research should also consider the effects of sensor calibration and the residual misregistration between bands of a UAS multispectral camera.

## DATA AVAILABILITY STATEMENT

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

## AUTHOR CONTRIBUTIONS

Both AW and GS contributed to the development, planning, and data collection of the study as well as the data analysis, interpretation, and writing of the manuscript.

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**Conflict of Interest:** The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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