Field Investigation of a Hanging Dam in the St. Lawrence River, Winter of 1981-82

Hung Tao Shen, William A. VanDeValk, Gordon B. Batson and Iury L. Maytin

Department of Civil and Environmental Engineering Clarkson College of Technology Potsdam, N.Y. 13676



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Final Report

Prepared for

U.S. Department of Transportation St. Lawrence Seaway Development Corporation Office of Plans and Policy Development Washington, D.C. 20591

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16. Abstract	
In this study, field surveys were carried	out for a large hangin
ice dam in the St. Lawrence River near Sparrow	hawk Point, during the
winter of 1981-82. Cross section profiles of	the dam and the channel
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the headloss characteristics, and velocity profiles underneath the dam were monitored during the season. Based on the field data, hydraulic resistance characteristics of the dam are analyzed. The effect of channel geometry and flow pattern on the formation and shape of the ic dam are discussed.

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inches (in)	centimeters (cm)	2.540
inches (in)	meters (m)	0.0254
feet (ft)	meters (m)	0.305
miles (miles)	kilometers (km)	1.61
square inches (sq in)	square centimeters (cm^2)	6.45
square feet (sq ft)	square meters (m ²)	0.093
cubic feet (cu ft)	cubic meters (m ³)	0.028
Fahrenheit temperature (°F)	Celsius temperature (°C)	5/9 after subtracting 32
Btu	Calori e s	252
Btu/hr. ft. °F	Watts/cm °C	0.0173

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Chapter I

INTRODUCTION

The existence of hanging dams in rivers with an ice cover has been known to hydraulic engineers for many years. The presence of hanging dams in a river constricts the flow cross section and, in many cases, could cause large head losses in addition to the loss caused by a normal sheet ice cover. A number of field studies have been reported in the literature (5,6,8,9,13,26) on the formation and characteristics of hanging dams. In this report, a field study on the hydraulic characteristics of a hanging dam located in the St. Lawrence River near Sparrowhawk Point during the winter of 1981-82 is presented. Formation of Hanging Dams

Although all hanging dams are massive accumulations of ice particles on the underside of an ice cover, the formation processes of these hanging dams may be significantly different. According to their formation processes, hanging dams may be classified into two categories. The first type of hanging dams, which will be referred to as fragment ice hanging dams, are accumulations of large ice plates or frazil ice pans. These dams are formed near the leading edge of an ice cover during its upstream progression. This process can occur either at the beginning of the winter during the formation of the new ice cover or during the spring break-up period when ice floes, which are released from an upstream reach of a river, reaches a stationary downstream obstacle or ice cover. Fragment ice dams are formed either by the submergence, transport, and arrest of ice floes

underneath an ice cover or through the internal collapse and subsequent thickening of floating ensembles of large ice particles caused by the action of external forces. The second type of hanging dams, which will be referred to as frazil ice hanging dams, are formed by the deposition of suspended frazil ice particles which are produced in open water areas upstream of a stable ice cover during periods of supercooling of the river The deposition of frazil suspensions could occur when surface. suspended ice particles reach the undersurface of the ice cover. Near the leading edge of the ice cover, when frazil ice particles are active, particles that reach the ice-water interface are deposited. Further downstream, when ice particles become inactive they will deposit only in low flow velocity regions. For a fragment ice dam, which is located near the leading edge of an ice cover, and is formed at the beginning of the ice covered season, a relatively soft outer layer of frazil slush could form on the surface of the fragment ice dam due to the accumulation of active frazil ice particles produced in the open water area during the winter. In the St. Lawrence River both the fragment ice hanging dams and the frazil ice hanging dams, and the combination of both could form.

The St. Lawrence River Ice Cover

The Upper St. Lawrence River is being utilized for hydropower production and serves as the only navigation passage for shipping between the Atlantic Ocean and the Great Lakes. The existence of hanging ice dams in the river has been an important consideration in the planning and management of the river flow

during the winter. The flow of the river is regulated by the Moses-Saunders Power Dam located near Massena. This dam is part of the St. Lawrence Project, which was completed in 1959 as a joint venture of the United States and Canada. The planning for the flow regulation is closely related to the ice conditions in the river which affect the length of the navigation season and the amount of power that can be generated.

With the present regulation plan, ice covers are initiated at artificial obstacles such as dams and booms at the beginning of each winter (18). Stable ice covers are formed between Lake Ontario and Cardinal with the installation of Ogdensburg-Prescott ice booms in the Galop Island area. Ice cover conditions in the International Rapids Section (see Appendix C) between Cardinal and Massena, however, are more complicated. Early in the winter, when the water temperature is supercooled, frazil ice forms This frazil ice develops into large ice floes and in the river. fragmented ice sheets which flow into Lake St. Lawrence. Beginning initially at the face of the Power Dam, these ice floes pack against each other to form a solid cover. The ice front formed by the edge of the ice floes then progresses upstream as more ice floes arrive. Due to the low flow velocity, the progression of the ice cover from the Power Dam to Morrisburg can be completed with no flow regulation. When the leading edge of the ice cover reaches Morrisburg, the progression of the ice cover still further upstream is accomplished by a short period of flow reduction which is created when favorable cold weather developes. With the flow reduction, the ice cover will progress

upstream towards the region of high velocity flow near Pinetree Point. At this stage the incoming ice floes which pass through Iroquois Dam will become entrained by the flow and are transported under the ice cover. When these floes come to rest under the ice cover they form a hanging dam. The size of this hanging dam will increase either due to the arrest of additional ice floes underneath the hanging dam or due to the thickening caused by the collapse of the accumulated floating ice mass. The collapse of the ice mass occurs when the hydrodynamic drag and streamwise gravity force acting on the ice mass surpasses the strength of the fragmented ice cover. The lowering of the gates at the Iroquois Control Dam can control the size of the hanging dam at Pinetree Point by preventing the inflow of additional ice floes. The lowering of these gates also initiates the progression of an ice sheet upstream from Iroquois Dam towards Sparrowhawk Point and Cardinal. The ice cover upstream of Iroquois Dam can only progress to the vicinity of Sparrowhawk Point because of high flow velocities in the Galop Island Region. Another hanging dam will then form near Sparrowhawk Point.

Due to the high flow velocity, open water reaches always remain downstream of the Galop Booms and Iroquois Dam, throughout the winter. These open water areas can produce large amounts of frazil ice. Because of the relatively short reaches in these open water areas, the frazil ice generated in these areas is entrained in the flow in the form of buoyant suspensions. The suspended frazil ice particles will gradually accumulate under the ice cover and lead to the growth of existing hanging dams

and form additional hanging dams further downstream (3,4,21).

Field measurements in the St. Lawrence River indicated that major fragment ice hanging dams are formed each winter in the vicinity of Sparrowhawk Point and Pinetree Point. Frazil ice hanging dams have also been observed in the Ogden Island region (3,4). During the winter of 1981-82 a major hanging dam located near the Sparrowhawk Point was monitored and studied for its hydraulic characteristics. The result of this field study will be presented in the following chapters.

Chapter II

FIELD MEASUREMENTS

During the winter of 1981-82, ice covers in the International Rapids Section of the St. Lawrence River were formed during the period between Jan. 11 and Jan. 18, 1982 (18). Preliminary surveys made by Ontario Hydro in late January showed that a large hanging dam existed near Sparrowhawk Point and a smaller one existed near Pinetree Point. Considering the time and resources available, the larger hanging dam near Sparrowhawk Point was selected for detailed study. A reconnaissance survey was then made on Jan. 28, 1982 to determine the exact location of the dam. This survey indicated that the hanging dam had a length of about 2,000 ft with a maximum thickness of more than 20 ft. Since a full three-dimensional mapping of the hanging dam is not feasible, a grid system, which includes a longitudinal control line and three transverse control lines, at the site of the dam was established with grid points marked on the ice surface to serve as a reference for locating measuring stations later. The location of the hanging dam and control lines for the grid system are presented in Figs. 1 and 2. Coordinates of grid points along control lines are summarized in Fig. 3. Based on the grid system defined, the hanging dam was monitored on a weekly basis during the whole period. The data measured include i) longitudinal and transverse profiles of the dam, the corresponding channel bottom profiles, ii) drop in the water surface level along the hanging dam, and iii) velocity distribution underneath the hanging dam. These measurements are discussed in detail in the following







Location of Control Lines at the Hanging Dam Site. Figure 2.



Figure 3. Coordinates of Grid Points on the Control Lines

Hanging Dam Profiles and the Channel Geometry

The field measurement procedures adopted in this study to obtain ice thickness and channel cross sectional geometries were the same as those reported in previous surveys conducted during the winters of 1977-78 and 1978-79 on flow and ice conditions in the Upper St. Lawrence River (3,4). When measurements were to be made, holes were first drilled through the ice sheet at a selected The thickness of the ice sheet was then measured grid point. with an L-shaped metal rod which was hooked onto the underside of The location of the undersurface of the hanging the ice sheet. dam and the depth to the channel bottom were measured using a Lowrance Fish Lo-K-Tor Sonic Meter. The measurement procedure consisted of first lowering the depth sensing portion of the meter through the ice mass with a scaled metal pipe. When the sensor was no longer inside the hanging dam, the meter would indicate the depth from the sensor to the river bottom. The thickness of the hanging dam and the flow depth could then be determined.

<u>Geometry of channel cross sections</u> - The elevation of the channel bottom along all four control lines was determined based on the average of flow depth readings obtained by the sonic meter during ice thickness measurements. Additional information obtained from the detailed hydrographic charts prepared by the U.S. Army Corps of Engineers in 1960 are also used. The channel bottom profiles given by the hydrographic charts compared very well with the sonic meter data. This shows that relatively little change in channel bottom profiles has occurred since 1960. All channel bottom elevation data were adjusted to measure from the 1955 IGLD.

Hanging dam profiles - Hanging dam thickness profiles along control lines were measured. Since the thickness of the hanging dam along both transverse lines 1-1 and 12-12 at the upstream and the downstream ends were small, only the longitudinal profile and the transverse profile along line 6-6 were monitored periodically. All measured profiles are presented in Appendix A. A summary of these results are presented in Figs. 4, 5, 6, and 7. These figures show that the hanging dam has a three-dimensional form. The shape of the dam is affected by the channel geometry, the amount of ice supply from the upstream, and the flow pattern in the vicinity and upstream of the site of the hanging dam. Both the length along the longitudinal control line and the width along the transverse control line 6-6 of the hanging dam decreased as winter progressed. The thickness of the dam along these two control lines, however, both increased and decreased during this The variation of thickness during the winter was closely period. related to the variation of the air temperature.

Water Levels

The resistance characteristics of a hanging dam is an important parameter in ice hydraulics. For this purpose, water levels at both the upstream and downstream ends of the hanging dam; i.e., at stations L-1 and L-12; were measured using a surveyor's level. The water surface slope along the hanging dam could then be determined. Table 1 summarizes measured water surface slopes.



Figure 4. Hanging Dam Profiles Along Longitudinal Line L-L





Elevation, Ft.

14

4000

Figure 6. Hanging Dam Profiles Along Transverse Line 6-6



(.JT)

Elevation



15

Date	Slope, s _f
Feb. 2, 1982	5.98x10 ⁻⁵
Feb. 4, 1982	4.98×10^{-5}
Feb. 9, 1982	5.48x10 ⁻⁵
Feb. 16, 1982	6.48×10^{-5}
Feb. 18, 1982	6.98x10 ⁻⁵
Feb. 23, 1982	8.72x10 ⁻⁵
Feb. 25, 1982	4.48x10 ⁻⁵
Mar. 2, 1982	8.47×10^{-5}
Mar. 12, 1982	3.49×10^{-5}

Velocity Profiles

Velocity profiles underneath the hanging dam were measured using a Marsh-McBirney electromagnetic current meter (3,4) and a gimbal-mounted, vane-type photoelectric current meter (12). Both meters were attached to a winch-cable assembly with a standard 75-1b USGS lead weight attached at the lower end of the cable to maintain the vertical position of the meters (Figs. 8 and 9). Even with the successive pounding of the 75-lb lead weight, it was often difficult to send the meter assembly through the inner portion of the hanging dam which consisted of hardened fragment ice pieces. Only a few measured velocity profiles were obtained along the longitudinal control line. These profiles are presented in Figs. 10-15. Most of these profiles have the shape of composite logarithmic profiles commonly observed in rivers with sheet ice cover (11). Velocity profiles at stations L3 and L5, however, are rather irregular. A velocity defect at the upper portion of the velocity profile at station L3 can be observed. This velocity defect is similar to that commonly observed in convergent pipe flows when the adverse pressure gradient exists (30).In the present case, although the flow area remains fairly uniform in the vicinity of station L3, an adverse pressure gradient may exist due to the change in pressure at the ice-water interface caused by the streamwise variation of ice thickness. At station L5, large fluctuations in velocity readings were experienced during the measurement. The mean velocity profile is extremely irregular. This is believed to be caused by the cross currents and large eddies which existed in this region where



Figure 8. The Assembly for Velocity Meters





Figure 9. Velocity Measurement in the Field













Figure 15. Velocity Profile at Station L-12; Feb. 18, 1982

the flow is distinctly three-dimensional.

Other Flow and Ice Condition Data

Since flow and ice cover conditions can provide useful information for interpreting the field survey results related to the hanging ice dam, additional data obtained by various agencies are presented in the Appendices. These data include: 1) water levels along the river and discharge at the Power Dam; 2) ice cover thickness measured near the Power Dam, Waddington, Pinetree Point, and Cardinal; and 3) ice cover condition maps.
Chapter III

ANALYSIS AND DISCUSSIONS

The importance of effects of hanging ice dams on winter flow conditions of a river have often been discussed in the literature on ice hydraulics. The present understanding on the formation process and hydraulic effect of hanging dams is, however, still rather qualitative. This is due partly to the complex nature of the evolution process, and partly to the difficulty involved in carrying out comprehensive field studies. In this chapter, the field data presented in Chapter II will be analyzed and interpreted. Formation and Geometry of the Hanging Dam

In order to better interpret the field data, a brief description of the formation process of a typical fragment hanging ice dam will be given. This description is based on the understanding accumulated from previously reported analytical, laboratory, and field studies. Most of these studies are limited to idealized uniform rectangular channels. As winter begins, large quantities of ice floes will form in the river. These floes are carried downstream until they reach an artificial obstacle or form an ice bridge. With any additional supply of ice floes from upstream, the front of the ice cover will progress upstream against the river flow. In river reaches with low velocity, a relatively thin initial ice sheet can be formed and progress smoothly in the upstream direction. This upstream progression will be impeded when the leading edge of the cover enters a fast flowing reach where the Froude number exceeds a critical value (1,9,10,16,17, 27,29). At this time additional incoming ice floes will be

carried under the ice cover to form a hanging dam. The local thickening of an ice cover can also be induced by internal collapse of the ice cover when the total force acting on the cover exceeds its strength (16,17,28). During the build up of the hanging dam the surface slope of the river will change. This change will alter the Froude number upstream of the leading edge, the strength of the cover, and magnitude of forces acting on the All of these changes could again allow the cover to cover. progress in the upstream direction. The thickness of the hanging dam is limited by the available supply of ice floes, and the stability of the ice floes within the accumulation. This stability criteria may be expressed in terms of a critical Froude number, a critical velocity, or a critical shear stress (2,7,9, 10,14,25,29).

In the reach of the St. Lawrence River between Cardinal and Iroqouis Dam, a hanging ice dam usually forms in the vicinity of Sparrowhawk Point. No effort has been made in the past to determine the location of the hanging dam. The only hanging dam location surveyed other than that reported in Chapter II is the hanging dam of the winter of 1978-79 (3). The location and typical cross sectional geometry of both the 1978-79 hanging dam and the 1981-82 hanging dam are presented in Fig. 16 along with major currents estimated from stream-tube analysis (22). The dam of 1978-79 was formed between Jan. 16, 1979 and Jan. 18, 1979 when the discharge was 219,000 cfs. The dam of 1981-82 was formed between Jan. 15, 1982 and Jan. 18, 1982 when the discharge was 210,000 cfs (18). The larger discharge in 1979 is believed to

have caused the hanging dam of that season to form approximately 3,000 ft downstream from the 1981-82 hanging dam. The leading edge of the ice covers of both winters are located approximately two thousand feet upstream from the mid-section of the hanging The cross sectional profiles of the hanging dam also dams. indicate that these hanging dams have irregular three-dimensional The shapes of these hanging dams are governed by the shapes. flow pattern upstream as well as in the region where the hanging dams are formed. A comparison of Fig. 16 with the channel bottom topography, shown in Fig. 17, indicates further that when a hanging dam is formed in a river reach which has nearly constant width and total river channel cross sectional area, the dam tends to form at a cross section where there exists a large variation in depth across the river's width. This phenomena can be explained by noticing that the transverse distribution of the depth-averaged velocity at a given cross section can be approximately described by Eq. 1.

$$\frac{V(z)}{U} = \left[\frac{d(z)}{D}\right]^{2/3}$$
(1)

in which, z = transverse distance from a bank; V(z) = depthaveraged flow velocity at z; d(z) = depth of flow at z; U = crosssectional average velocity, Q/A; and D = average flow depth ofthe flow cross section. Eq. 1, which can be obtained from theManning Equation by neglecting interfacial shear stresses (22),indicates that the local depth-averaged flow velocity increaseswith the local flow depth. Thus, at a cross section where a







Figure 17. Channel Bottom Topography

large transverse variation in depth exists, there will usually be areas of high velocity which may cause the underturning of ice floes. The underturning of ice floes will prevent further progression of the ice cover until the formation of a hanging dam changes the flow condition.

In order to obtain quantitative estimates of the critical Froude number for leading edge progression and the critical velocity of ice deposition, depth-averaged velocity distributions at all four cross sections shown in Fig. 16 are determined for both the free surface condition and the condition with hanging dams. Flow distributions for free surface conditions at times of ice cover initiation are presented in Appendix A. These figures indicate that the critical Froude number for the progression of the leading edge of the ice cover is approximately equal to 0.06 for the 1978-79 ice cover and is 0.052 for the 1981-82 ice cover. Flow distributions at channel cross sections underneath the hanging dam are presented in Appendix A. To determine the critical velocity of deposition from these figures the criteria used are: a) at each cross section, only the velocity at the location where the ice thickness is a maximum will be considered; b) when the hanging dam was decaying, the velocity distribution was not used to determine critical velocities; c) during cold periods, active frazil ice produced in the upstream open water area may remain active when the ice particles reach the upstream end of the hanging dam. The attachment of active frazil ice particles to the existing hanging dam is not then limited by the critical velocity of deposition. Based on these considerations, velocity

distributions presented in Appendix A are examined along with the air temperature data given in Fig. 18. The depth-averaged velocity distributions underneath the maximum thickness of the ice cover at each cross section are summarized in Table 2. Based on this information it is estimated that the critical velocity of deposition for inactive frazil ice particles is approximately equal to 3 fps, which is in agreement with the observation of Michel and Drouin (13) in the La Grande River.

Hydraulic Resistance of the Hanging Dam

The hydraulic resistance of a hanging ice dam can have an important effect on the flow conditions in an ice covered river. Data presented in Chapter II will be used to obtain some quantitative information on the hydraulic resistance of the hanging dam.

Local roughness coefficients - The local roughness of the undersurface of the ice cover and the channel bottom are calculated from information obtained from the measured velocity profiles. These roughnesses, expressed in terms of Manning's coefficient, are determined using the information developed by Larsen (11). As summarized in Table 3, ice surface roughness coefficients at the upstream side of the dam are lower than those at the downstream side. The average ice cover roughness coefficient is 0.03, which is slightly lower than the value reported by Beltaos and Dean (5) for a Smoky River hanging dam. The average value of bed roughness coefficients at the hanging dam site is 0.047, which is much higher than the gross bed roughness coefficient value of 0.026 for the reach between

Cross Section	Date	Velocity	Remarks
Line l	Feb. 11, 1982	3.7 fps	Accumulation of active frazil
Line 6	Feb. 3, 1982	3.1 fps*	Cold weather.
	Feb. 9, 1982	2.9 fps*	Hanging dam size increased from Feb. 3, 1982.
	Feb. 16, 1982	2.3 fps	Hanging dam size decreased slightly from Feb. 9, 1982.
	Feb. 23, 1982	2.3 fps	Hanging dam size decreased from Feb. 16, 1982.
	Mar. 2, 1982	<u>~</u> 3.0 fps*	Very cold weather. Dam size increased. Measured dam cross section was not complete.
	Mar. 12-19, 1982		Dam size decreased continuously.
Line 2	Feb. 11, 1982	3.0 fps*	Very cold weather.

Table 2. Depth-Averaged Velocity Underneath the Maximum Ice Thickness

*Estimated critical velocity for the deposition of inactive frazil ice particles.

Station No.	Date	ni	n _b
Ll	Feb. 24, 82	0.028	0.048
L3	Feb. 24, 82	0.026	0.055
L7	Feb. 25, 82	0.022	0.057
L9	Feb. 18, 82	0.037	0.029
L12	Feb. 18, 82	0.036	0.047

Table 3. Local Roughness Coefficients Calculated from Measured Velocity Profiles

Cardinal and the Iroquois Dam reported by Potok (19).

<u>Roughness coefficient of the ice cover</u> - The gross Manning roughness coefficients of the ice cover in the reach between Cardinal and Iroquois Dam, n_i^{CI} , are calculated for the entire winter using the scheme developed by Shen and Ruggles (24). Water levels at Cardinal and Iroquois headwater, and the discharge at the Iroquois Dam are used in the calculation. Gross Manning's coefficients of the ice cover between upstream and downstream ends of the hanging dam, n_i^D , are also determined from the measured water surface slopes given in Table 1. The Manning Equation and Belokon-Sabaneev (15) formula are used in the calculation.

$$Q = \frac{1.49}{n} A R^{2/3} s_{f}^{1/2}$$
(2)

$$n_{i}^{D} = n_{b} \left[2 \left(\frac{n}{n_{b}} \right)^{3/2} - 1 \right]^{2/3}$$
(3)

Both n_i^{CI} and n_i^{D} are computed using $n_b = 0.026$ and the channel geometries given by Potok (19), as shown in Table 4. The calculated gross roughness coefficients of the ice cover are presented in Fig. 18 along with the discharge at the Iroquois Dam and the air temperature at Massena.

As shown in Figs. 4 and 6 the size of the hanging dam varies with time even though the general shape of the hanging dam changes very little before the break up period. It is, therefore, reasonable to assume that the form loss of the hanging dam remains nearly constant. The variation in the gross hydraulic resistance of the hanging dam will be mainly due to the variation in surface roughness of the dam. Fig. 18 shows that the resistance coefficient of the ice cover at the hanging dam section, n_{i}^{D} , remains relatively constant during a major portion of the winter. The range of the variation of the magnitude of n_{i}^{D} is about the same as that of n_{i}^{CI} . A comparison between n_{i}^{D} and the air temperature also indicates that the variation in n_{i}^{D} is affected by the air temperature. This appears to be reasonable since the air temperature governs the production of frazil in the open water area upstream and hence influence the surface roughness of the ice cover.

Table 4. Cross-Sectional Geometry at the Reference Water Level (19)

Parameter	Cardinal	Iroquois H.W.	
Reference Water Level	240.80 ft.	239.90 ft.	
Width	2620 ft.	2620 ft.	
Wetted Perimeter	2690.916 ft.	2683.1298 ft.	
Cross Sectional Area	92900 sq. ft.	82700 sq. ft.	

*Water surface levels are measured from the International Great Lakes Datum (1955).



Chapter IV

SUMMARY AND CONCLUSIONS

In this report, results are presented and analyzed of a field survey on a large hanging ice dam in the St. Lawrence River near Sparrowhawk Point during the winter of 1981-82. The major conclusions obtained are: a) the channel bottom topography may be used to provide a convenient guide for determining the location where a hanging dam will form beneath the ice cover in the study reach; b) the shape of a hanging dam is affected by the pattern of currents immediately upstream; c) the critical Froude number for the progression of the ice cover and the critical velocity of deposit of ice particles underneath the dam in the reach between Cardinal and Iroquois Dam are approximately 0.06 and 3 fps, respectively; d) the ice surface roughness at the upstream side of the hanging dam is smaller than that of the downstream side; e) the gross roughness coefficient of the hanging dam remains relatively constant during the winter even though it is slightly affected by the air temperature. These conclusions provide useful insights to the formation process and resistance characteristics of hanging dams in the Upper St. Lawrence River. The information obtained will lead to the further development of simulation models for the dynamics of ice cover formation and its hydraulic effects in the Upper St. Lawrence River. It would be desirable to conduct additional field surveys of hanging ice dams in subsequent winters in order to substantiate and extend these conclusions.

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APPENDIX A

Flow Distributions and Hanging Dam Profiles

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Λετοςτελ

(Ft./Sec.)



Velocity (Ft./Sec.)

Depth (Ft.)



Velocity and Froude Number Distributions at Cross Section 1-1; Feb. 11, 1982



Velocity, fps

ъерсћ, Ft.







Velocity and Froude Number Distributions at Cross Section 6-6; Feb. 23, 1982



Velocity and Froude Number Distributions at Cross Section 6-6; Mar. 2, 1982



Velocity and Froude Number Distributions at Cross Section 6-6; Mar. 12, 1982



Velocity and Froude Number Distributions at Cross Section 6-6; Mar. 16, 1982


















Figure A21. Longitudinal Hanging Dam Profile Along Line L-L; Mar. 12, 1982



Figure A22. Longitudinal Hanging Dam Profile Along Line L-L; Mar. 16, 1982



Longitudinal Hanging Dam Profile Along Line L-L; Mar. 19, 1982 Figure A23.

APPENDIX B

Water Levels and Discharge

WATER LEVELS AND DISCHARGE - ST. LAWRENCE RIVER

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22 44,30 0.0 38,60 37,32 229333 20 44,30 0.0 0.0 38,60 37,32 229333 21 44,10 0.0 41,41 40,16 0.0 38,60 37,32 229333 22 44,10 0.0 41,41 40,16 0.0 38,60 37,32 229933 23 44,51 0.0 41,41 40,16 0.0 0.0 37,32 229934 23 44,51 0.0 41,41 40,16 0.0 37,32 36,33 229934 23 44,51 0.0 41,41 40,16 0.0 37,32 36,33 22,9933 24 25 0.0 41,47 0.0 0.0 37,32 36,33 22,9934 25 44,57 0.0 41,47 0.0 0.0 37,32 36,33 22,9934 25 44,57 0.0 41,47 0.0 0.0 37,47 36,91 26,91 22,9933 25 44,27 0.0 41,43 39,40	- 0	2 F - 4 2		1 Z - E 7	0	42.18	41.03	0.0	0.0	0.0	39.06	37 . 87	215037
22 44.30 0.0 41.75 40.48 0.0 36.79 229938 22 44.10 0.0 41.41 40.16 0.0 37.94 26.79 22938 22 44.10 0.0 41.41 40.16 0.0 36.15 36.79 22938 22 44.51 0.0 41.41 40.17 0.0 37.94 36.79 229938 23 44.51 0.0 41.41 40.17 0.0 37.94 36.38 22993178 24 44.51 0.0 41.88 40.40 0.0 38.17 36.38 230178 25 44.51 0.0 41.88 40.40 0.0 38.17 36.38 230178 25 44.57 0.0 41.88 40.40 0.0 38.17 36.885 26.499 230178 25 44.57 0.0 41.88 40.12 0.0 0.0 38.17 36.885 230333 27 44.28 0.0 41.49 40.12 0.0 0.0 36.885 2309333	0	00 10 10		4 4 7	0	41 . 94	40.70	0.0	0.0	0.0	38,60	37,32	223333
22 44.09 0.0 41.41 40.16 0.0 37.94 36.49 230120 22 44.09 0.0 43.19 0.0 0.0 37.82 36.38 230170 22 44.61 0.0 41.41 40.17 0.0 0.0 37.82 36.38 230170 23 44.65 0.0 41.61 0.0 0.0 0.0 37.82 36.38 230170 24 44.51 0.0 41.68 40.47 0.0 0.0 37.82 36.38 230170 25 444.51 0.0 41.68 40.40 0.0 37.82 36.38 239990 26 44.51 0.0 41.68 40.40 0.0 38.22 36.38 229990 26 44.53 0.0 41.68 40.40 0.0 38.22 36.38 230170 26 44.53 0.0 41.43 40.12 0.0 0.0 37.85 36.38 230333 27 44.453 0.0 41.43 39.40 12 0.0 </td <td>h ⊂ + 0</td> <td></td> <td></td> <td>7 7 7 Z</td> <td>0.0</td> <td>41.76</td> <td>40.48</td> <td>0.0</td> <td>0.0</td> <td>0.0</td> <td>38.15</td> <td>36.79</td> <td>229938</td>	h ⊂ + 0			7 7 7 Z	0.0	41.76	40.48	0.0	0.0	0.0	38.15	36.79	229938
22 44.09 0.0 41.37 40.18 0.0 0.0 37.62 36.38 230120 23 44.51 0.0 41.41 40.17 0.0 0.0 37.65 36.38 230178 24 44.51 0.0 41.41 40.17 0.0 0.0 37.65 36.38 230178 25 44.51 0.0 41.68 40.40 0.0 0.0 37.65 36.38 230178 26 44.51 0.0 41.68 40.40 0.0 0.0 38.22 36.38 230178 26 44.55 0.0 41.68 40.40 0.0 38.22 36.38 230178 26 44.30 0.0 41.68 40.40 0.0 38.22 36.38 230178 27 44.35 0.0 41.43 399.93 0.0 0.0 37.85 36.78 230133 27 44.25 0.0 41.43 399.93 0.0 0.0 37.87 36.26 2402162 27 44.25 0.0 41.	2 -			01.54		41.41	40.16	0.0	0.0	0.0	46° 1E	04-90	229941
23 44.42 0.0 41.41 40.17 0.0 0.0 37.69 36.24 230178 25 44.51 0.0 43.37 0.0 0.0 0.0 38.22 36.83 229901 25 44.51 0.0 41.68 40.40 0.0 0.0 38.22 36.83 229901 26 44.55 0.0 41.68 40.40 0.0 0.0 38.22 36.83 229901 26 44.30 0.0 41.68 40.40 0.0 0.0 38.22 36.83 2299901 26 44.30 0.0 41.68 40.12 0.0 0.0 38.22 36.83 229930 27 44.25 0.0 41.63 40.12 0.0 0.0 37.87 36.48 230333 28 44.25 0.0 41.43 39.93 0.0 0.0 37.87 36.48 230333 28 44.25 0.0 41.43 39.78 0.0 37.87 36.48 240148 29 0.0 0.0	4 D 4 D	00.44			0.0	41.37	40,18	0.0	0*0	0.0	37,82	36,38	230120
24 44.51 0.0 41.69 40.64 0.0 0.0 38.22 36.83 229901 25 44.57 0.0 41.69 40.64 0.0 0.0 38.22 36.83 229901 26 44.30 0.0 41.65 40.12 0.0 0.0 38.17 36.91 229990 26 44.30 0.0 41.65 40.12 0.0 0.0 38.17 36.91 229990 26 44.30 0.0 41.65 40.12 0.0 0.0 38.17 36.91 229990 27 44.22 0.0 41.65 40.12 0.0 0.0 37.65 36.78 230.33 28 44.35 0.0 41.65 40.12 0.0 0.0 37.65 36.78 230.33 29 44.26 0.0 41.65 40.00 0.0 0.0 37.65 36.48 2400149 20 44.28 0.0 0.0 0.0 0.0 0.0 37.87 36.48 240076 20 44.04	4 C	44-42		с п С п С п С п	0.0	41.41	40.17	0.0	0.0	0.0	37.69	36,24	230178
25 44.25 0.0 41.68 40.40 0.0 0.0 38.17 36.91 229990 26 44.35 0.0 41.65 40.12 0.0 0.0 38.11 36.91 229990 26 44.35 0.0 41.65 40.12 0.0 0.0 38.11 36.85 230333 27 44.22 0.0 41.49 40.12 0.0 0.0 37.63 36.78 230333 28 44.35 0.0 41.43 39.93 0.0 0.0 37.63 36.78 230134 29 44.35 0.0 41.43 39.93 0.0 0.0 37.63 36.26 24026 29 44.28 0.0 41.03 39.56 0.0 0.0 37.87 26.11 239979 31 44.04 0.0 0.0 0.0 0.0 0.0 37.87 26.11 239979	10 4	44.51		43.81	0.0	41.89	40-64	0.0	0.0	•••	38.22	36.83	229901
27 44.30 0.0 41.65 40.12 0.0 0.0 38.11 36.85 230333 27 44.22 0.0 43.47 0.0 41.65 40.12 0.0 0.0 37.95 36.78 232162 28 44.22 0.0 41.49 40.12 0.0 0.0 37.95 36.78 232162 28 44.31 0.0 41.43 39.93 0.0 0.0 37.63 36.78 240208 29 44.31 0.0 41.43 39.93 0.0 0.0 37.63 36.26 240208 29 44.31 0.0 41.43 39.78 0.0 0.0 37.87 26.26 240076 31 44.04 0.0 0.0 0.0 0.0 37.87 26.11 239979	tu Jc			6.5		41.68	40.40	0.0	0.0	0.0	38.17	36*91	229990
27 44.22 0.0 41.49 40.12 0.0 0.0 37.95 36.78 232162 28 44.35 0.0 41.49 40.12 0.0 0.0 37.63 36.78 240208 28 44.31 0.0 41.49 40.12 0.0 0.0 37.63 36.26 240208 29 44.31 0.0 41.49 40.00 0.0 0.0 37.63 36.26 240208 20 44.28 0.0 0.0 0.0 0.0 37.87 36.48 240148 20 44.28 0.0 0.0 0.0 0.0 37.87 36.48 240076 31 44.04 0.0 0.0 0.0 0.0 0.0 37.87 36.11 235979) 4) 4	1 C 1 P 1		2 7 7 7 7 7 7		41.05	40.12	0.0	0.0	0.0	38 . 11	36.85	230333
28 41.55 0.0 0.1	2 N 4 C			5 F F	0.0	41.49	40.12	0 * 0	0.0	0.0	37 . 99	36.78	232162
29 44.31 0.0 41.49 40.00 0.0 0.0 37.80 36.48 240148 30 44.28 0.0 41.24 39.78 0.0 0.0 37.70 36.48 240076 31 44.04 0.0 0.0 0.0 0.0 0.0 37.87 36.11 235979	- a - a					41.43	10 0 M	0.0	0.0	0.0	37+63	36 . 26	240208
30 44.28 0.0 41.24 39.78 0.0 0.0 0.0 37.70 36.22 240076 31 44.04 0.0 0.0 0.0 010 010 010 235979						41.49	40.00	0.0	0.0	0.0	37.80	36.48	240148
	7 C 1 P					01.24	39.78	0.0	0.0	0.0	37.70	36 - 22	240076
) -) (10 56	0.0	0 0	0.0	37 . 87	36.11	239979
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WATER LEVELS AND DISCHARGE - ST. LAWRENCE RIVER

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				4 1 8 8 9		FEBRUAR	Y 1982					
DAY	K IN GS		D GDEN	CALOP GALOP	CARDI	IROHW	IRQTW	LEI SH	MADIN	MORIS	SAUND	DISCHARGE
 -			4 7. 3 0		41.15	39-21	0 0	0	0.0	37.25	35.52	240064
- n	10 10 10 10 10 10				0.04	39.03	0.0	0.0	0.0	36.67	35.30	240151
۳ נ			2 V		40-90	39.02	0.0	0.0	0.0	36.90	35.21	239864
n <			00.00		40.99	61 6E	0 0	0.0	0.0	37.11	35 • 48	239849
វជ					40 03	30.15	0.0	0.0	0.0	37.13	36°49	240386
n v					41.10	39.31	0.0	0.0	0.0	37.28	35 * 58	244987
0 1					41.21	39,36	0.0	0.0	0.0	37.35	35,61	244994
- 03					41.04	39.16	0.0	0.0	0.0	31+15	35.43	246968
0			4.1.4	0.0	40.70	38.69	0.0	0.0	0*0	36.47	34 - 67	250209
	44.42			0.0	40.77	38.75	0.0	0.0	0.0	36+42	34.61	250638
) - -) () () ()		40.73	38.73	0.0	0.0	•••	36.44	34,66	250054
- 0					40.69	38.68	0.0	0.0	•••	36.43	49.40	250018
4 - -	44.26		43.08	0	40.60	38. 50	0.0	0.0	0.0	36,20	34 . 29	254892
) 4 	2 2 2 2 2 2 2 2		43.03	0.0	40.56	38.47	0.0	0.0	0.0	36.14	34.19	255152
⊦ L° 	44.20	0.0	4 J. 0 D	0.0	40.54	38.43	0.0	0.0	0 . 0	36.07	34.11	255120
) (44.20	0-0	42.98	0	40.44	38.47	0*0	0*0	0.0	36,17	34.29	255056
	44.04		40.76		40.17	38.27	0.0	0.0	0.0	36.09	01* * 9	254974
- a	44.17		42.75		40.13	38.16	0.0	0.0	0.0	36.07	33.97	255171
20	44.25		42.94		40.31	38.37	0.0	0.0	0.0	36.16	34.16	255041
	A4.10		12.06	0.0	40.37	38,53	0.0	0 • 0	0.0	36.42	34 . 48	255064
2-					40.23	38.48	0.0	0.0	0.0	36.40	94 ***	255170
4 0 i €			40.04		40.29	10.00	0.0	0.0	••	36.46	34.54	254955
1 v 1 v			40.05		40.38	38.72	0.0	0.0	0:0	36.55	34.76	254978
10) <	10 10 10		4.2.78		40.21	38.63	0.0	0.0	0.0	36.60	94 - 79	255143
F UT J C	44.16	0.0		0 0	40.37	38.78	0 • 0	0.0	0.0	36.71	34.88	254728
) () (44.15		42.03	0.0	40.50	38,98	0.0	0.0	0.0	36.58	25,15	255488
) N 1 (47.08	0-0	40.56	39.08	0.0	0.0	0.0	37,15	36,35	255098
- 10 1 0	14 14 10		42.89	•	40.48	39.01	0.0	0.0	0.0	37.15	40°.04	254879

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WATER LEVELS AND DISCHARGE - ST. LAWRENCE RIVER

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Day KINGS VINCE GGEN GALCP CARDI INQUAL CESH ADIN NGRIS SAUND DISCHARGE 1 44.00 0.0 42.99 0.0 0.0 37.22 35.46 255912 2 444.03 0.0 42.99 0.0 40.53 39.01 0.0 37.22 35.46 255912 5 444.03 0.0 40.53 39.01 0.0 0.0 37.22 35.41 255912 5 444.03 0.0 42.97 0.0 0.0 37.22 35.41 25.51103 6 444.03 0.0 42.97 0.0 0.0 37.22 35.41 25.51103 6 444.03 0.0 42.87 0.0 0.0 37.22 35.41 25.51103 6 444.03 0.0 42.99 0.0 0.0 37.22 35.73 25.51103 11 444.03 0.0 42.94 0.0 0.0	 			1 1 1 1 1 1	} } ! !		MARCH	1982					
1 44.00 0.0 40.41 38.93 0.0 0.0 37.20 35.94 0.0 37.20 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 37.21 35.94 0.0 0.0 37.22 35.94 0.0 37.21 35.94 0.0 37.25 35.94 0.0 0.0 37.21 35.94 0.0 37.25 35.94 0.0 37.25 35.94 0.0 37.25 35.94 0.0 37.25 35.94 0.0 37.25 35.94 0.0 37.25 35.94 0.0 37.25 35.94 0.0 37.25 35.94 0.0 37.25 35.94 35.94 35.94 35.94 35.94 35.94 35.94 35.94 35.94 35.94 35.94 35.94 35.94 35.94	 DAY	KINGS		OGDEN	GALCP	CARDI	IRQHW	IRQTW	LEISH	N ADI N	MORIS	SAUND	D I SCHARGE
25:103 25:103	! ! ≁			42.84	0.0	40 • 41	38.93	0.0	0.0	0.0	36°6	35.18	255012
31 31 31 32 31 32 35 <td< td=""><td>• ^</td><td></td><td></td><td></td><td>0</td><td>40.57</td><td>41.6E</td><td>0.0</td><td>0.0</td><td>•••</td><td>37.20</td><td>35.46</td><td>255103</td></td<>	• ^				0	40.57	41.6E	0.0	0.0	•••	37.20	35.46	255103
7 444.05 0.0 40.53 38.94 0.0 37.21 35.31 0 7 444.12 0.0 42.95 0.0 40.53 39.12 0.0 37.21 35.53 0 9 444.10 0.0 42.95 0.0 40.55 39.28 0 0.0 37.22 35.55 0 0 0 37.23 35.55 0	J 🗠	44 03	0	42.94	0.0	40.45	39.06	0.0	0.0	0.0	37.22	35.40	255116
1 1	4	0 0 0 0	0	42.73	0.0	0E • 04	38.94	0.0	0.0	0.0	37.11	10.00	0
7 444.03 0.0 42.55 39.12 0.0 37.23 35.55 9 444.04 0.0 42.87 0.0 40.37.23 35.55 10 42.87 0.0 40.46 39.08 0.0 37.23 35.55 11 444.07 0.0 42.87 0.0 40.37.23 35.55 11 444.07 0.0 42.87 0.0 37.23 35.55 11 444.07 0.0 40.46 39.08 0.0 0.0 37.23 35.57 35.18 11 444.07 0.0 42.87 0.0 40.65 39.18 0.0 0.0 37.23 35.57 35.113 35.57 35.113 35.57 35.113 35.57 35.113 35.57 35.113 35.57 35.51057 35.51057 35.51057 35.51057 35.51057 35.51057 35.51057 35.51057 35.51057 35.51057 35.51072 35.51072 35.51072 35.51072 35.51072 35.51072 35.51072 35.51072 35.51072 35.51072 35.51072 35.51072<	n.	44 22	0.0	42.97	0.0	40.53	39.14	0.0	0.0	0.0	37.24	35.47	•
7 74412 0.0 77.28 35.55 9 944410 0.0 77.28 35.55 11 74.03 0.0 77.28 35.55 111 74.03 0.0 77.23 35.55 112 74.03 0.0 77.23 35.55 113 74.17 0.0 77.23 35.55 113 74.17 0.0 77.23 35.55 114 77.23 35.55 35.55 115 74.17 0.0 77.23 35.55 115 74.17 0.0 77.23 35.55 25.105 115 74.17 0.0 77.23 35.55 25.105 115 74.17 0.0 77.23 35.55 25.107 115 74.17 0.0 77.23 35.55 25.107 116 74.17 0.0 77.23 35.55 25.107 119 74.17 0.0 77.28 35.55 25.107 119 74.17 0.0 77.28 35.55 25.10	φ	44.03	0.0	40.04	0.0	40.47	39,03	0.0	0*0	0.0	37,32	35,59	•
9 444.04 0.0 40.45 39.24 0.0 37.23 35.73 111 44.07 0.0 42.881 0.0 0.0 37.23 355.73 355.73 112 44.07 0.0 42.881 0.0 0.0 37.23 355.73 355.73 111 44.07 0.0 42.881 0.0 0.0 37.23 355.73 355.73 113 44.17 0.0 42.887 0.0 0.0 37.53 355.73 </td <td>• •</td> <td>44.12</td> <td>0.0</td> <td>42.87</td> <td>0.0</td> <td>40.50</td> <td>39.12</td> <td>0.0</td> <td>0.0</td> <td>0.0</td> <td>37.28</td> <td>35 • 56</td> <td>•</td>	• •	44.12	0.0	42.87	0.0	40.50	39.12	0.0	0.0	0.0	37.28	35 • 56	•
0 42.79 0.0 42.79 0.0 47.23 35.52 112 44.03 0.0 42.88 0.0 47.23 35.55 25.123 35.55 113 44.17 0.0 42.88 0.0 40.45 39.08 0.0 37.23 35.55 25.1065 115 44.18 0.0 42.88 0.0 40.55 39.19 0.0 37.28 35.55 25.1065 115 44.18 0.0 40.58 399.19 0.0 0.0 37.58 35.56 25.1073 116 44.18 0.0 40.69 39.45 0.0 0.0 37.58 35.56 25.1065 117 44.16 0.0 40.68 39.53 0.0 0.0 37.58 35.56 25.1073 118 44.27 0.0 40.78 39.57 0.0 37.58 35.56 25.1073 118 44.27 0.0 41.77 399.53 0.0 0.0 37.56 25.1073 118 44.27 0.0 41.77 0.0	- 20	44.10	0.0	42.95	0.0	40.59	39.24	0.0	0.0	0.0	37.42	35.73	•
$ \begin{bmatrix} 1 & 44 & 07 & 0.0 & 42.81 & 0.0 & 40.46 & 39.08 & 0.0 & 0.0 & 37.29 & 35.55 \\ 1 & 44 & 07 & 0.0 & 42.81 & 0.0 & 40.51 & 39.18 & 0.0 & 0.0 & 37.28 & 35.55 \\ 1 & 44 & 25 & 0.0 & 42.81 & 0.0 & 40.58 & 39.19 & 0.0 & 0.0 & 37.28 & 35.55 \\ 1 & 44 & 25 & 0.0 & 42.87 & 0.0 & 40.58 & 39.19 & 0.0 & 0.0 & 37.28 & 35.55 \\ 1 & 44 & 25 & 0.0 & 42.87 & 0.0 & 0.0 & 0.0 & 37.58 & 35.58 & 251072 \\ 1 & 44 & 25 & 0.0 & 40.78 & 39.45 & 0.0 & 0.0 & 37.58 & 35.56 & 251072 \\ 1 & 44 & 27 & 0.0 & 42.87 & 0.0 & 0.0 & 0.0 & 37.56 & 35.56 & 251072 \\ 1 & 44 & 27 & 0.0 & 42.87 & 0.0 & 0.0 & 0.0 & 37.56 & 35.75 & 251072 \\ 1 & 44 & 27 & 0.0 & 42.87 & 0.0 & 0.0 & 0.0 & 37.56 & 35.75 & 251072 \\ 2 & 44 & 47 & 0.0 & 43.17 & 0.0 & 0.0 & 0.0 & 0.0 & 34.53 & 35.14 & 251072 \\ 2 & 44 & 47 & 0.0 & 43.53 & 0.0 & 0.0 & 0.0 & 0.0 & 34.53 & 35.14 & 251072 \\ 2 & 44 & 47 & 0.0 & 43.53 & 0.0 & 0.0 & 0.0 & 0.0 & 34.26 & 37.51 & 35.75 & 2550137 \\ 2 & 44 & 47 & 0.0 & 43.75 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 34.63 & 35.14 & 2511357 \\ 2 & 44 & 47 & 0.0 & 43.75 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 34.753 & 35.14 & 2551379 \\ 2 & 44 & 47 & 0.0 & 43.75 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 34.53 & 35.14 & 2551379 \\ 2 & 44 & 47 & 0.0 & 43.75 & 0.0 &$	م	44 . 04	0.0	42.79	0.0	40.36	39.04	0.0	0.0	0.0	37.23	35.52	•
$ \begin{bmatrix} 44, 07 & 0.0 & 2.81 & 0.0 & 40.45 & 39.08 & 0.0 & 0.0 & 37.23 & 35.57 & 251135 \\ 16 & 44, 17 & 0.0 & 42.86 & 0.0 & 40.51 & 39.19 & 0.0 & 0.0 & 37.28 & 35.51 & 2551055 \\ 16 & 44, 17 & 0.0 & 42.87 & 0.0 & 40.51 & 39.45 & 0.0 & 0.0 & 37.55 & 35.58 & 2551072 \\ 16 & 44, 17 & 0.0 & 42.87 & 0.0 & 40.58 & 39.45 & 0.0 & 0.0 & 37.55 & 35.58 & 2551072 \\ 17 & 44, 27 & 0.0 & 42.87 & 0.0 & 40.57 & 39.25 & 0.0 & 0.0 & 37.55 & 35.58 & 2551072 \\ 18 & 44, 17 & 0.0 & 43.21 & 0.0 & 40.57 & 39.25 & 0.0 & 0.0 & 37.55 & 35.58 & 2551072 \\ 19 & 44, 27 & 0.0 & 43.21 & 0.0 & 40.57 & 39.25 & 0.0 & 0.0 & 37.56 & 35.58 & 2551073 \\ 19 & 44, 27 & 0.0 & 43.21 & 0.0 & 0.0 & 0.0 & 37.56 & 35.75 & 2551073 \\ 22 & 44, 47 & 0.0 & 43.57 & 0.0 & 0.0 & 0.0 & 0.0 & 37.56 & 35.75 & 2551073 \\ 22 & 44, 47 & 0.0 & 43.57 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 37.56 & 37.56 & 2551073 \\ 22 & 44, 47 & 0.0 & 43.57 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 37.56 & 37.56 & 2551073 \\ 22 & 44, 47 & 0.0 & 43.57 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 37.56 & 2551073 \\ 22 & 44, 47 & 0.0 & 43.57 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 0.0 & 37.56 & 2551073 \\ 22 & 44, 45 & 0.0 & $	0	44.03	0.0	42,80	0.0	40.40	39.08	0.0	0.0	0.0	37,29	35.56	0
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$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	12	44.08	0.0	42,86	0.0	40.51	39.14	0 • 0	0*0	0.0	37.33	35.58	251065
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	Ē	44.17	0.0	42,88	0.0	40.58	39,19	•••	0.0	0.0	37+28	35.51	251062
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	14	44.25	0.0	43.11	0.0	40.89	39.45	010	0.0	0.0	37+53	35 ,84	251072
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17 44,20 0.0 42.92 0.0 40.74 39.55 0.0 37.26 35.43 251137 19 44,27 0.0 43.11 0.0 39.55 0.0 37.26 35.43 255103 20 44.33 0.0 43.19 0.0 0.0 37.26 35.47 2551048 20 44.33 0.0 41.17 39.90 0.0 0.0 38.23 35.47 2551048 21 44.49 0.0 41.27 49.50 0.0 0.0 38.64 2555918 2555913 22 444.45 0.0 41.56 0.0 0.0 38.65 2555913 2555913 22 444.45 0.0 41.51 0.0 0.0 38.65 2555913 2555913 22 444.45 0.0 41.66 0.0 0.0 38.65 2555913 2555913 22 444.45 0.0 41.65 0.0 0.0 38.65 2555913 2555913 23 444.75 0.0 41.65 0.0	16	44.05	0.0	42.87	0.0	40 - 68	39,36	0.0	0.0	0	37,50	35.75	251313
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22 44.49 0.0 43.53 0.0 41.70 40.87 0.0 38.93 37.21 256177 23 44.45 0.0 41.66 40.96 0.0 0.0 39.28 37.21 256003 24 44.45 0.0 41.66 41.06 0.0 0.0 39.28 37.21 256003 28 44.45 0.0 41.66 41.06 0.0 0.0 39.28 37.98 256003 28 44.45 0.0 41.66 41.06 0.0 0.0 39.28 37.98 256003 28 44.75 0.0 41.61 0.0 0.0 39.28 27.66 41.77 26.03 28 44.75 0.0 41.61 0.0 0.0 0.0 40.61 39.39 2561846 28 44.75 0.0 41.77 0.0 0.0 40.66 39.30 2661946 29 44.64 0.0 41.77 0.0 0.0 40.66 39.30 2661946 29 44.64 0.0	21	44 - 42	0.0	43.25	0.0	41,34	40.50	•••	•••	0.0	38*46	36 • 69	255918
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29 44.68 0.0 43.92 0.0 42.26 41.73 0.0 0.0 40.64 35.34 261973 30 44.64 0.0 43.74 0.0 42.05 41.52 0.0 0.0 40.46 39.17 262039 31 44.81 0.0 43.90 0.0 42.24 41.72 0.0 0.0 40.78 39.55 262023	8	44.70	0.0	43,92	0.0	42+27	41.70	0.0	0.0	0.0	40.61	39+30	262057
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		44 81	0.0	00	0.0	42.24	41.72	0.0	0.0	0.0	40.78	39 55	262023

ATER LEVELS AND DISCHARGE - ST. LAWRENCE RIVER

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APPENDIX C

Ice Cover Thickness at Selected Cross Sections*

*Data provided by Ontario Hydro. Cross section locations are given in Ref. 23.

CROSS SECTION A(Power Dam), Feb. 22, 1982

Location	Ice Thickness	and	Remarks
0+00	Barnhart Island		
1+50	16" Solid Ice		
3+00	17" Solid Ice		
4+50	19" Solid Ice		
6+00	17" Solid Ice		
7+50	15" Solid Ice		
9+00	18" Solid Ice		
10+50	18掉" Solid Ice		
12+00	25" Solid Ice		
13+50	21" Solid Ice		
15+00	21" Solid Ice		
16+50	18" Solid Ice		
18+00	21" Solid Ice		
19+50	16" Solid Ice		
21+00	21" Solid Ice		
22+50	26" Solid Ice		
24+00	20" Solid Ice		
26+50	18" Solid Ice		
28+00	21" Solid Ice		
29+50	Canadian Shore		

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CROSS SECTION Bl(Waddington), Feb. 22, 1982

Location	Ice Thickness	and	Remarks
0+00	U.S. Mainland		
1+50	21" Consolidated Broken Pack	No	Slush
3+00	16" Consolidated Broken Pack	No	Slush
4+50	19" Consolidated Broken Pack	No	Slush
6+00	18" Consolidated Broken Pack	No	Slush
7+50	18" Consolidated Broken Pack	No	Slush
9+00	18" Consolidated Broken Pack	No	Slush
10+50	18" Consolidated Broken Pack	No	Slush
12+00	20" Consolidated Broken Pack	No	Slush
13+50	21" Consolidated Broken Pack	No	Slush
15+00	Ogden Island		

CROSS SECTION C(Pinetree Pt.), Feb. 23, 1982

Location	Ice Thickness	and	Remarks
1+50	23" Solid Ice	No Slus	sh
3+00	17" Solid Ice	No Slus	sh
4+50	20" Consolidated Broken Pack	No Slus	sh .
6+00	18" Consolidated Broken Pack	Waferli evident	ke ice crystals in hole
7+50	18" Consolidated Broken Pack	Waferli evident	ke ice crystals in hole
9+00	22" Consolidated Broken Pack	No Slus	h
10+50	17" Consolidated Broken Pack	No Slus	h
12+00	22" Consolidated Broken Pack Underside of cove	No Slus r rough	h
13+50	30" Broken pack a	nd slush	- silt evident
15+00	19" Consolidated Broken Pack	No Slus	h
16+50	18" Consolidated Broken Pack	No Slus	h
18+00	17" Consolidated Broken Pack	No Slus	h
19+50	26" Broken pack a	nd slush	- silt evident
21+00	21" Consolidated Broken Pack	Waferli evident	ke ice crystals in hole
22+50	23" Solid Ice	No Slus	h
23+25	United States Mai	nland	

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CROSS SECTION D(Cardinal), Feb. 24, 1982

Location	Ice Thickness	and Remarks
1+50	17월" Solid Ice	No Slush
3+00	13" Consolidated Broken Pack	No Slush
4+50	15" Consolidated Broken Pack	No Slush
6+00	l6" Consolidated Broken Pack	No Slush
7+50	17" Consolidated Broken Pack	NO Slush
9+00	20" Consolidated Broken Pack	No Slush
10+50	20攴" Consolidated Broken Pack	No Slush
12+00	10戈" C onsol idated Broken Pac k	21' Heavy Slush
13+50	19攴" Consolidated Broken Pack	22'4" Heavy Slush
15+00	20" Consolidated Broken Pack	18'4" Heavy Slush
16+50	21" Consolidated Broken Pack	11'3" Heavy Slush
18+00	21" Consolidated Broken Pack	6'4" Medium Slush
19+50	5'8" Broken Pack	Light Slush
21+00	5'1" Broken Pack	Light Slush
22+50	18" Solid Ice	18" Light Slush
24+00	15" Solid Ice	No Slush
25+50	16" Solid Ice	No Slush
27+00	20" Solid Ice	12" Light Dirty Slush
27+50	American Shore	

APPENDIX D

Ice Cover Condition Maps*

*Based on aerial photographs by the St. Lawrence Seaway Development Corporation.

1982

Jan. 12,

DATE

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March 29, 1982



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