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A SPRING-SUMMER (APRIL-SEPTEMBER) CLIMATOLOGY OF THE LAKE ERIE BASIN

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Technical Bulletin
OHSU-TB-7

Report No. AS-S-115

The Ohio State University
Sea Grant Program
August 1984





The Ohio State University
Ohio Cooperative Extension Service
Ohio Department of Natural Resources

Sea Grant Technical Bulletins are published by the Ohio Sea Grant Program at the Ohio State University and are partially supported through a grant from the National Sea Grant College Program at the National Oceanic and Atmospheric Administration (NOAA), U.S. Department of Commerce. These bulletins are designed to transmit research results from Sea Grant sponsored and related investigations to users of coastal and offshore resources. The U.S. Government is authorized to produce and distribute reprints for governmental purposes notwithstanding any copyright notation that may appear hereon.

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ABSTRACT

The effects of Lake Erie upon the meteorology and climatology of the lake basin during the spring and summer months are investigated. The major driving force behind these effects is seen to be the temperature difference between lake and land, the lake surface being cooler than the mean land temperature for all months during the period except September. Specific climatic variables are examined as well as any possible lake influence upon them. The report, therefore, constitutes a climatology of the Lake Erie region from the months of April through September.

This research was supported by the National Oceanic and Atmospheric Administration, Ohio Sea Grant Program, under Grant No. NA81AA-D-0095, and by the National Science Foundation under Grant No. AMT-8003376, Amendment No. 1.

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CHAPTER 1

INTRODUCTION

While the effects of the Great Lakes on the meteorology of the region during the cold season are reasonably well known [Petterssen and Calabrese (1979), Danard and Rao (1972), among others], comparatively little is known about their effects during the spring and summer months. This report surveys the current body of knowledge concerning the meteorology and climatology of the Lake Erie basin, in particular, during the months of April through September; it, therefore, constitutes a spring-summer climatology of this region. Such a survey is important in order to better understand the weather in the region and as a basis for possible future research studies which might lead to improved understanding of Lake Erie-related weather during this time of year. Consideration of the period from October through March will be treated in a later report.

1.1. General Description

Anyone living near the shoreline of Lake Erie can attest to the fact that its water can exert a marked cooling effect on the local climate, especially during the spring months. Trees and shrubs bloom one to two weeks later than they do inland and the lake breeze is a common occurrence. On a larger scale, Lake Erie also significantly influences the weather of northern Ohio, northwestern Pennsylvania, western New York, extreme southern Ontario, and southeastern Michigan.

There are a number of reasons why the study of "Lake Erie meteorology" is of practical importance. It is useful in agriculture for classifying areas suitable for certain crops and in deciding planting schedules.

Vineyards along the south shore of Lake Erie are examples of the benefits of Lake Erie in extending the frost-free period over that inland. Secondly, the large cities of Toledo, Cleveland (Ohio), Erie (Pennsylvania), Buffalo (New York), and Detroit (Michigan) are located on or near the shore of Lake Erie. Thirdly, the shoreline area is a popular holiday area. Finally, knowledge of lake winds is important to boating and shipping.

Figure 1 depicts the Lake Erie basin, while Table 1 lists some of the dimensions of Lake Erie. Covering an area of 9930 square miles, Lake Erie is only the fourth largest of the Great Lakes. It is the most southerly of the Great Lakes and also the shallowest with a mean depth of 58 feet. The western end of the lake is the shallowest part, averaging 30-40 feet in depth. Lake Erie comprises 30% of its water basin.

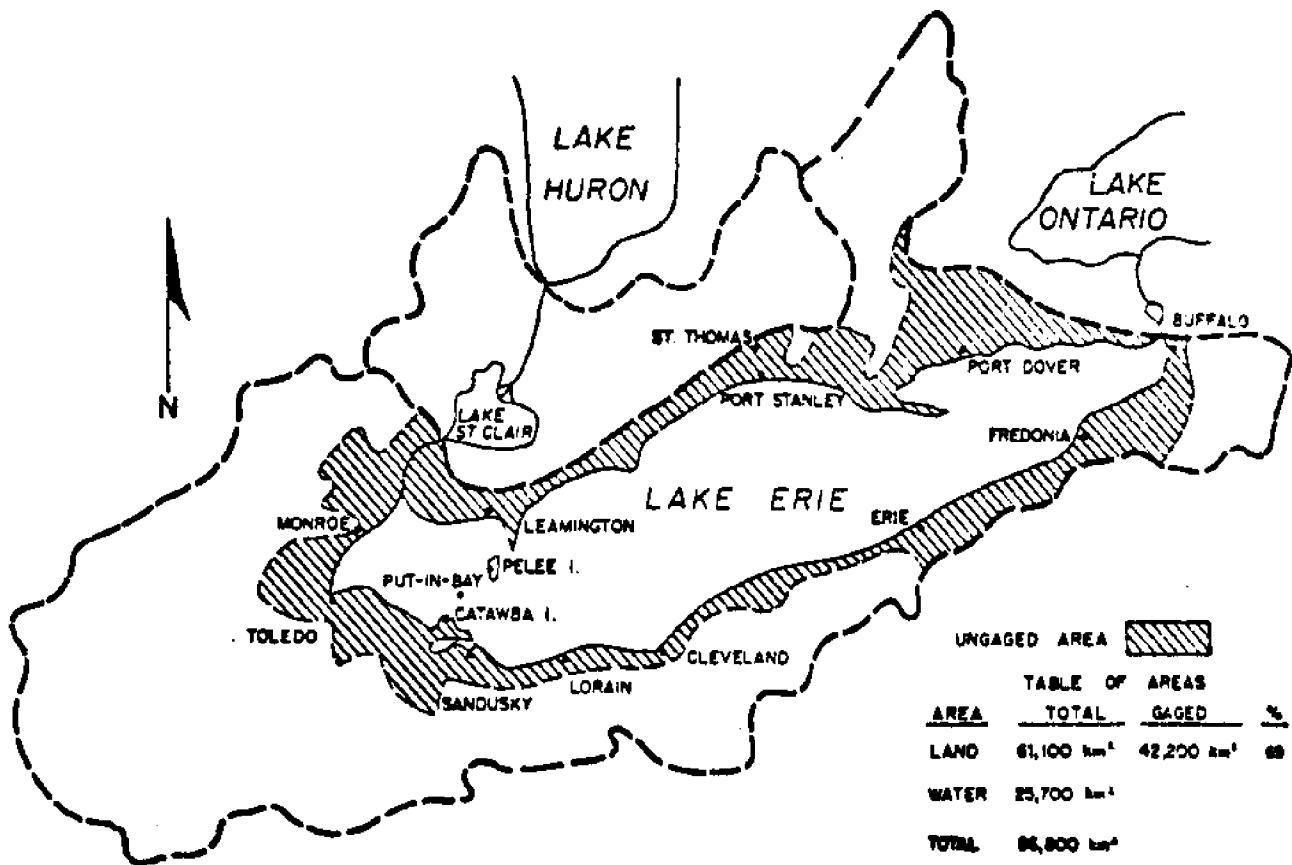


Figure 1. Lake Erie basin.
[Derecki (1964)]

TABLE 1. Dimensions of the Great Lakes.
[Jones and Meredith (1972)]

Lake	Length ^a (mi) ^c	Breadth ^a (mi) ^c	Area ^a		Average discharge 1946-1965 ^b (ft ³ /sec) ^e	Maximum depth ^a (ft) ^f	Mean depth ^a (ft) ^f
			Water Surface (mi ²) ^d	Drainage (ft ³ /sec) ^e			
Superior	350	160	31,820	80,000	78,000	1,333	487
Michigan	307	118	22,400	67,860		923	276
Huron	206	183	23,010	72,620	St. Clair 182,000	750	195
St. Clair	26	24	490	7430	Detroit 186,000	21	10
Erie	241	57	9930	32,490	Niagara 198,000	210	58
Ontario	193	53	7520	34,800	St. Lawrence 239,000	802	283

^aSource: U.S. Army Engineer Division, North Central (1965a).

^bSource: From data supplied by U.S. Department of Commerce, Lake Survey Center (personal communication).

^c1 mile (mi) = 1.61 kilometers (km).

^d1 square mile (mi²) = 2.59 square kilometers (km²).

^e1 cubic foot per second (ft³/sec) = 0.028 cubic meters per second (m³/sec).

^f1 foot (ft) = 0.305 meters (m).

1.2. Order of Investigation

Chapter 2 examines the physical basis behind the effects of Lake Erie on the meteorology during the spring and summer months. The thermal properties of water versus soil are looked at with a view to better understanding the lake-land temperature differences, which are the primary driving forces behind the lake effects.

Chapter 3 presents the general effects of Lake Erie on the meteorology and climate. The chapter first discusses the most local effects in horizontal scale, the lake and land breezes. Then, geostrophic flow across the lake is examined and, finally, the climatological effect of the lake on the large-scale pressure patterns is presented.

Chapter 4 looks at the various climatic variables and how Lake Erie affects them. Temperature, humidity, solar radiation, wind, evaporation, precipitation, and severe weather are each examined in turn.

Finally, Chapter 5 provides a brief summary of the report.

CHAPTER 2

PHYSICAL BASIS BEHIND THE METEOROLOGICAL INFLUENCE OF LAKE ERIE

Any discussion of the physical causes behind any effects of lake Erie on the region's meteorology must of necessity focus on the differences between the surface boundaries of lake versus land. In other words, one must understand the physical properties of both water and soil.

2.1. Surface Roughness

A fixed land surface provides a different lower boundary condition for the atmosphere than does a changing water surface. Roughness length z_0 is a rough measure of the friction exerted by the surface on the atmosphere and is mainly a function of the type of surface. For land, z_0 varies from 1 cm for short grass to 165 cm for a large city to 283 cm for a fir forest [Sellers (1965)]. For water, z_0 ranges from 1×10^{-4} cm to 50 cm [Munn (1966)]. Hence, the generally smoother lake surface means there is less friction to slow a given air flow. Accordingly, winds might be expected to be stronger over the lake. That this occurrence is not always true will be seen later in this report, where it is seen that the thermal structure over the lake plays a governing role as well.

2.2. Moisture Supply

A lake surface obviously provides a greater amount of water per surface area for evaporation than does land. In the winter this fact leads to greater evaporation over Lake Erie, resulting in lake-effect snow squalls on the lee side of the lake. During other seasons of the year, however, the greater availability of water on the lake surface

does not necessarily lead to greater evaporation over the lake. As will be seen later, the thermal stratification over the lake is again important and during the spring months the cooler lake suppresses evaporation and in April, may actually cause a net condensation.

2.3. Surface Temperature

2.3.1. Thermal Properties. Table 2 presents values of two physical properties for different types of soil and water. *Heat capacity C* is the amount of heat, in calories, needed to raise the temperature of one cubic centimeter of substance by 1°C. As can be seen in the table, the heat capacity of water is 1.00 whereas that of soil varies from 0.23 to 0.67 for the types listed. Thus land heats more rapidly than does water. This fact accounts for some of the observed temperature differences between the lake and land.

TABLE 2. Thermal properties of soil and water
[Adapted from Sellers (1965)].

<u>Soil</u>	Heat Capacity C (cal/°C cm ³)	Thermal Diffusivity κ (cm ² /s)
Dry quartz sand	0.23	.0030
Quartz sand, 40% moisture	0.65	.0090
Sandy clay, 40% moisture	0.67	.0066
Peat, 40% moisture	0.50	.0010
<u>Water</u>		
Still	1.00	.0014
Stirred (stable)	1.00	50.
Stirred (neutral)	1.00	300.

The other property is the *thermal diffusivity* κ , which is an inverse measure of the time it takes for surface temperature changes to penetrate to lower depths. For still water the κ values are of the same order as those of soil. However, the water in Lake Erie is not still and for such water the heat transfer occurs through turbulence and values for κ (now representing "eddy" diffusivities) are thousands of times that for land. Hence, heat flux downward within water occurs much faster than within soil. Accordingly, temperatures at the surface (the level of meteorological interaction) exhibit greater extremes in soil than in water, both diurnally and annually.

2.3.2. Lake-Land Temperature Differences. These two properties combine to make the average monthly Lake Erie surface temperature lower in summer and higher in winter than over land. Greater mixing within the lake also causes the surface temperature of Lake Erie to lag behind that of the land by about a month. These temperature differences are presented in Table 3, where Cleveland is chosen as the land station.

The meteorological and climatic controls exerted by Lake Erie on the region during the spring-summer months are due largely to such lake-land temperature differences. Since these differences are crucial, consider also Figure 2, which is a graph of the average maximum, mean, and minimum temperatures for Sandusky, Ohio, for the months of April through September. Also plotted is the average Lake Erie surface temperature for these months.

Note that the average daily high over land is greater than the lake temperature throughout the entire period, the greatest difference occurring in May (+19.3°F). The average mean temperature over land also exceeds that of the lake for all months except September; the greatest positive difference also occurs in May (+10.3°F). The average minimum

TABLE 3. Average monthly temperatures (°F) of Lake Erie and Cleveland, Ohio. [Adapted from Climates of the States (1974) and Richards & Irbe (1969)]

	<u>Cleveland</u>	<u>Lake Erie</u>
JAN	28.5	34
FEB	28.6	33
MAR	36.8	33
APR	47.3	38
MAY	59.1	49
JUN	69.4	63
JUL	73.7	70
AUG	71.9	72
SEP	65.5	67
OCT	54.4	59
NOV	41.7	49
DEC	30.9	39

temperature over land is lower than the lake temperature for all months except April and May.

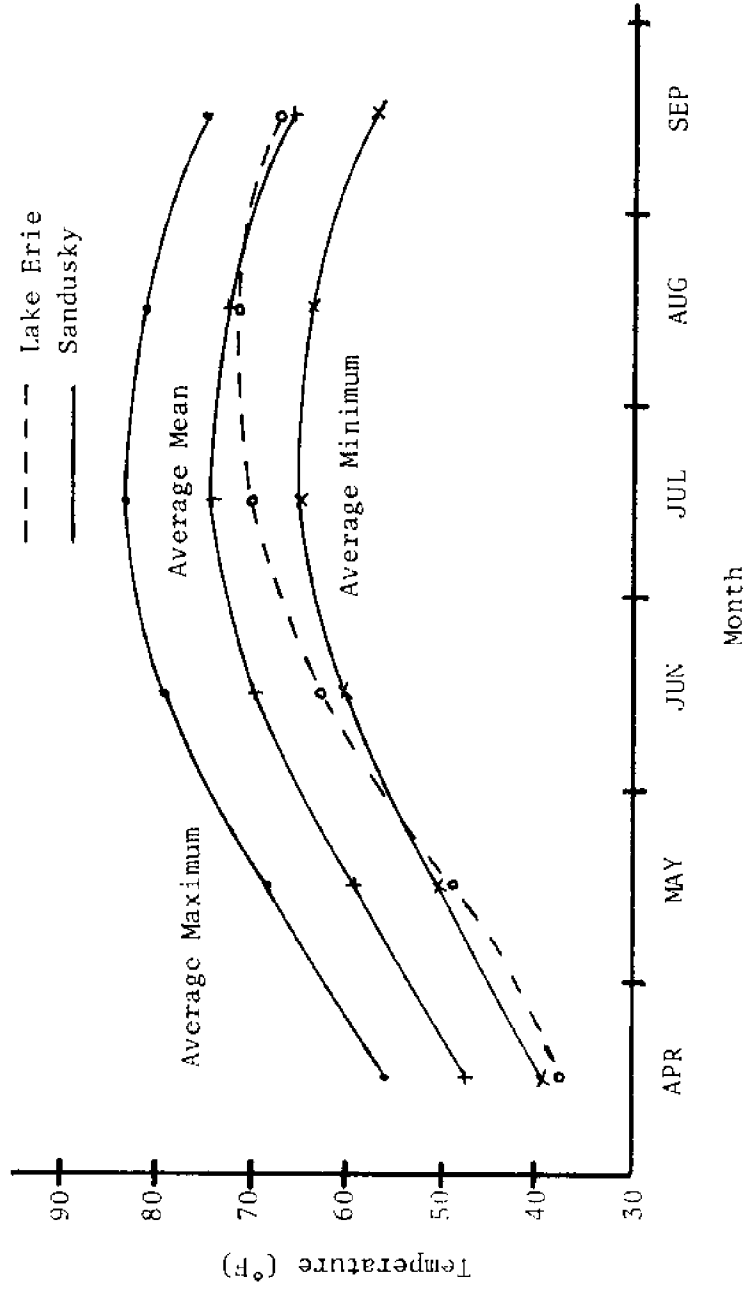


Figure 2. Lake-land temperatures for the spring and summer months.
 [Data from Climates of the States (1974) and Richards and Irbe (1969).]

CHAPTER 3

GENERAL METEOROLOGICAL EFFECTS OF LAKE ERIE UPON THE REGION

In this chapter some of the observed effects of Lake Erie upon the meteorology and climatology of the region are presented. Local winds are examined first. Then, effects of a greater horizontal scale are discussed.

3.1. The Lake and Land Breezes

During periods of light geostrophic winds, the lake-land temperature differences may lead to mesoscale circulation cells. Warmer temperatures over land cause the lake breeze phenomenon, with surface winds blowing from lake to land. Cooler land temperatures may induce a weaker land breeze, with surface winds blowing toward the warmer lake. From Figure 2, one can infer that, on the average, lake breezes are possible during the daytime of all the months studied, being strongest and most probable in the spring months. A land breeze may develop during the night hours in the summer.

An excellent observational study of the Chicago lake breeze was conducted by Lyons (1972). Analyzing mesoscale data from ten summer months (1966-68), Lyons found the lake breeze to occur on 36% of all days. Time of onset averaged from 0800-0900 LST and the lake breeze front was found to push inland anywhere from one block to over 40 km. The typical circulation pattern (see Figure 3) is seen to have an inflow layer depth of 500 m and a return layer aloft twice as deep. Surface inflow speeds average about 6 m s^{-1} with return speeds aloft half as large.

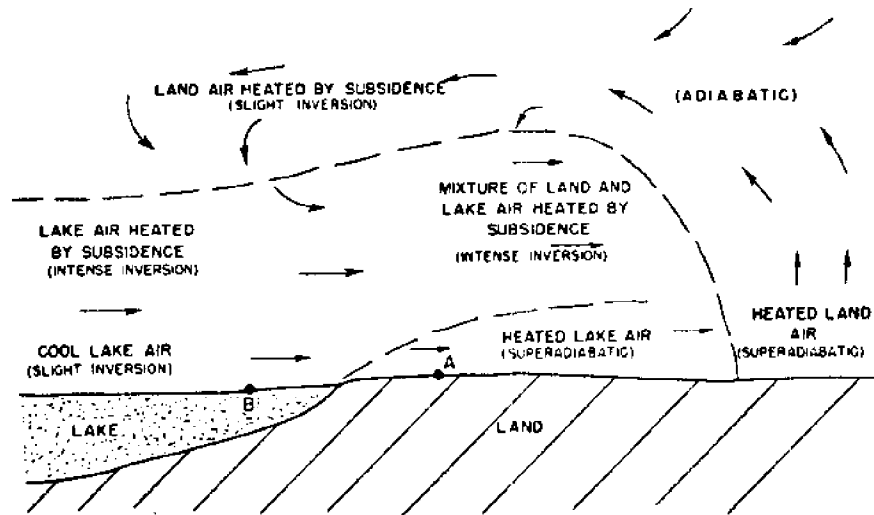


Figure 3. Schematic representation of a lake breeze with light geostrophic wind. [Munn and Richards (1964)]

The lake breeze front is a 1 - 2 km wide band of convergence with updrafts of more than 1 m s^{-1} , often accompanied by cloud bands due to the uplifting. Subsidence occurs over the lake, resulting in clear skies there and over the shoreline. A mesoscale high pressure cell develops over the lake. Lyons also found the lake breeze to veer with time in response to Coriolis accelerations, so that by evening the breeze is parallel to the shoreline.

Munn and Richards (1964), in a survey article on the lake breeze, indicated that passage of the lake breeze front is accompanied by a sharp fall in temperature and rise in relative humidity (due to the cooler temperature). They also observed "surges" of the lake breeze along the north shore of Lake Erie.

Numerical models have also been developed to study the lake breeze, among which are those of Estoque (1962), Moroz (1967), Mahrer and Pielke (1977), and Danard (1978).

3.2. Geostrophic Flow Over the Lake

During periods of substantive geostrophic winds, modification of the air being advected over the lake occurs. During the spring and early summer, such air is usually warmer than the underlying lake water.

Bellaire (1965) presents an observational study of a warm air mass moving directly northward over Lake Michigan in May of 1964. Wiresonde soundings from a ship were taken from the southern tip of Lake Michigan to a point 70 miles north. Figure 4 depicts the thermal cross-section observed. Air leaving the south shore is cooled, forming a temperature inversion whose top rises in height to an equilibrium value of about 350 feet at 50 miles downwind. Winds were observed to rapidly decrease over the lake; beyond 20 miles the water surface was rippled to smooth. Hence, warm air advection leads to a shallow dome of cold, stable air over the lake.

Lyons (1970), in another similar study, studied warm air advection over Lake Michigan in June of 1966. This time the air flow was southwesterly and wiresonde soundings were made along a line from Waukegan, Illinois, to Saugatuck, Michigan. Figure 5 presents the thermal cross-section. The inversion top in this case gradually increased to 150 m near the eastern shore in the afternoon. Vertical temperature gradients as large as $+9^{\circ}\text{C}/30\text{m}$ over the middle of the lake were observed between ship deck and the lake surface. Once ashore, the lake air undergoes remodification and an internal boundary layer characterized by convergence and uplift occurs.

Clouds moving over the lake were observed by Lyons to dissipate within 15 km of the windward lake shore and the air over the lake was cloud-free. Over the lake, surface wind speeds decreased due to the

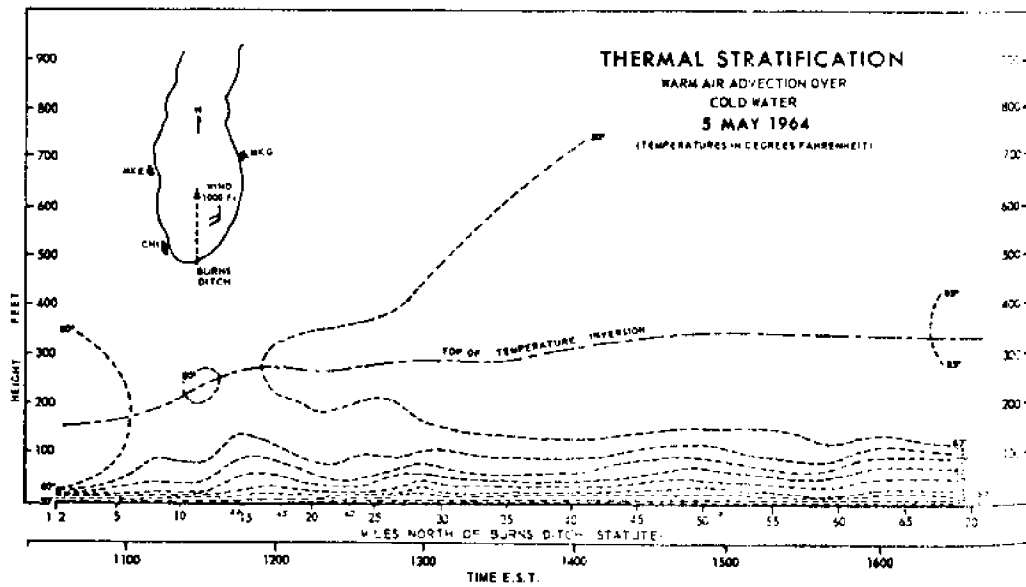


Figure 4. Temperature cross-section obtained from wiresonde soundings, May 5, 1964. [Bellaire (1965)]

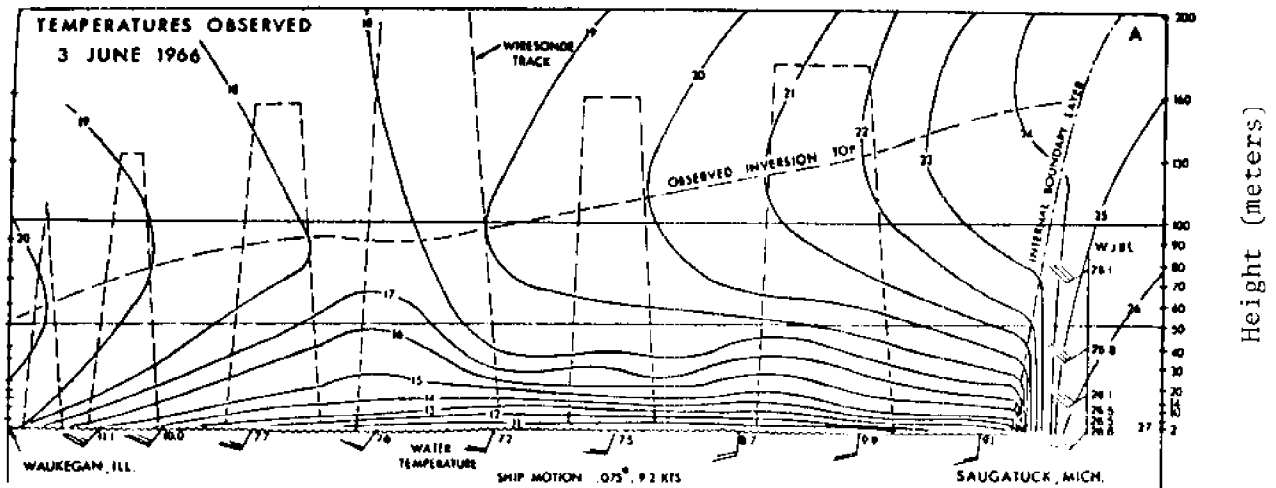


Figure 5. Temperature cross-section obtained from wiresonde soundings on June 3, 1966. One wind barb equals 5 knots; temperature is in °C. [Lyons (1970)]

greater stability of surface air. The wind also backed in direction over the lake due to the development of a lake mesohigh having a pressure excess of 2 - 4 millibars over the prevailing synoptic value.

The lake mesohigh is the subject of an article by Strong (1972). In an observational study over Lake Michigan, he found the warm offshore wind to experience a pronounced upward lifting as it approached the windward shoreline and then to undergo strong subsidence (as large as -1 m s^{-1}) as it moved over the lake. The most intense lake high observed during the study was one having a pressure excess of 2 mb. He also found the winds to back in direction as air moved over the lake.

3.3. Lake Effect on Climatological Pressure Patterns

From the discussion thus far it might be inferred that, on the average, the Great Lakes should be a region with higher pressure than surrounding areas during the spring and early summer months. Such is indeed the case. Charts from Klein's (1957) paper show local maxima over the Great Lakes during these months for the number of anticyclone centers observed in a 20-year period. Also, major tracks of anticyclones cross the region.

Carlson (1974) produced a quantitative argument for the existence of these high pressure zones during the months of April through June. Using surface heat budget models, he established the surface sensible heating fields over the Great Lakes region and from them calculated the thermally induced vertical motion fields at 900 mb. Figure 6 presents the results. In the absence of any synoptic systems, the surface heating fields cause subsidence over the cooler lake waters on the order of -0.1 to -0.4 cm s^{-1} , the same order as that observed within actual anticyclones. Such subsidence encourages anticyclones to develop and maintain themselves over the lakes.

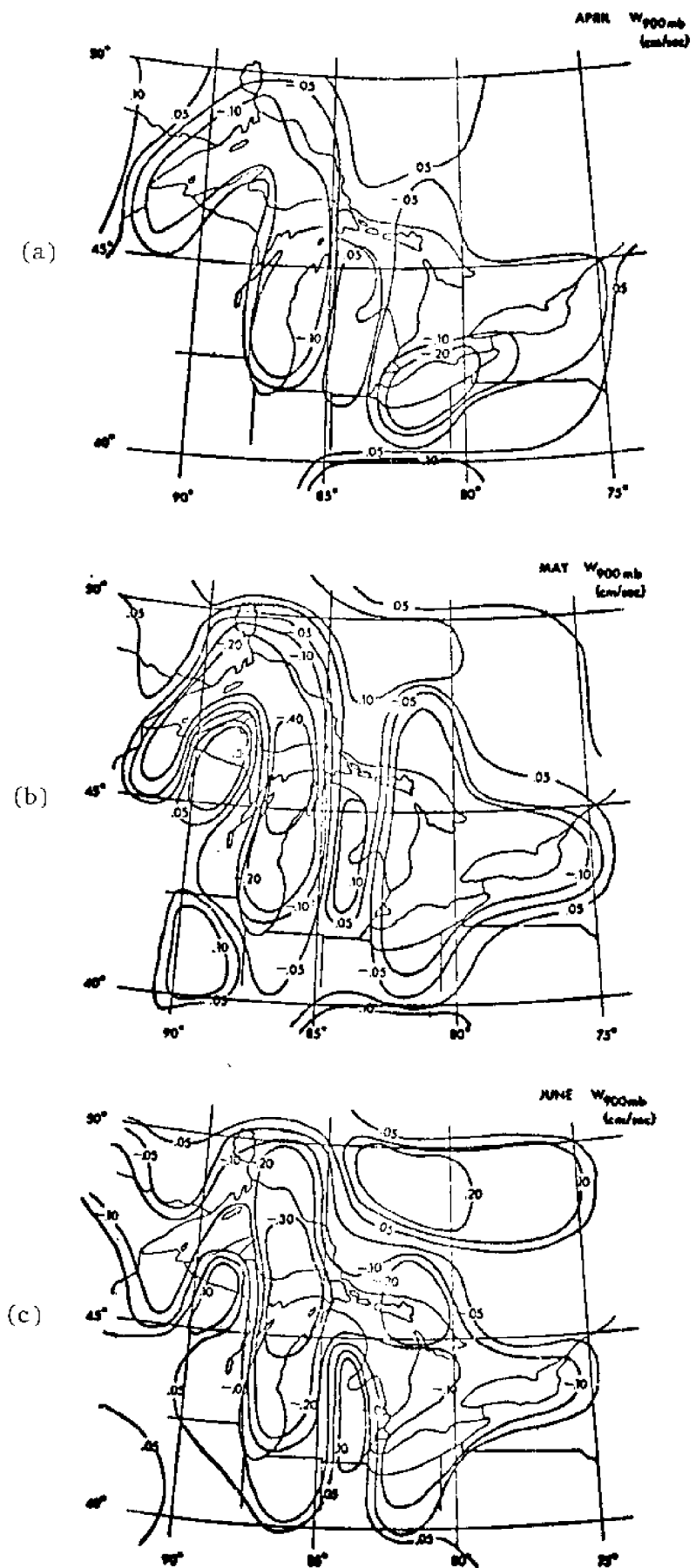


Figure 6. Mean induced vertical motions at 900 mb height over the Great Lakes area for April (a), May (b), and June (c). [Carlson (1974)]

CHAPTER 4

CLIMATIC VARIABLES IN THE LAKE ERIE BASIN

At the beginning of this section are presented the climatic data tables for Toledo, Sandusky, and Cleveland, Ohio, for Erie, Pennsylvania, for Buffalo, New York, and for Detroit, Michigan. These tables, Tables 4 - 9, respectively, are taken from Climates of the States (1974).

4.1. Temperature

Tables 4 - 9 give the average daily maximum, mean, and minimum temperatures for each month for each of the six cities mentioned above.

To see the effect of Lake Erie on temperatures near the shoreline, consider Figures 7 - 10, which are maps of the average daily maximum and minimum temperatures for April and July for the Great Lakes region [Phillips and McCulloch (1972)]. In April and July (see Figure 2) Lake Erie is cooler than the average maximum temperatures over land. Accordingly (Figures 7 and 9), the temperature contours reflect large gradients as Lake Erie is approached, with cooler shoreline temperatures during the daytime than further inland. At night in April (Figure 2), the nearly equal temperatures of lake and land result in negligible lake influence in April (Figure 8) on nighttime temperatures. In July, however, lake temperatures are warmer at night (Figure 2) and the temperature contours (Figure 10) reflect this fact, with shoreline night temperatures being warmer than further inland.

The influence of Lake Erie on temperatures can also be seen from Table 10, which lists the 1893-1921 monthly averages of temperature and precipitation for April through September for Cleveland and North Royalton,

TABLE 5. Climatic data for Sandusky, Ohio.
[Climates of the States (1974)]

SANDUSKY, OHIO
POST OFFICE BUILDING

LATITUDE 41° 27' N
LONGITUDE 82° 43' W
ELEVATION (ground) 803 Feet

Month	Temperature				Normal degree days (°F)	Precipitation						Relative humidity				Wind				Mean number of days											
	Normal		Extreme			Yearly	Maximum	Minimum	Monthly	Yearly	Maximum	Minimum	Yearly	7:00 a.m. - 1:31	7:00 p.m. - 1:31	Mean hourly	Prevailing	Speed	Direction	Fastest mile	Yearly	Clear	Partly cloudy	Cloudy	Precipitation 0.1 inch or more	Thunderstorms	Heavy fog	Max.	Min.		
	Daily maximum	Daily minimum	Record highest	Record lowest																										Maximum	Minimum
Jan.	37.6	35.2	70.1	-13.1	1036	6.27	0.63	1.74	1.74	1.74	0.63	1.74	81	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Feb.	35.3	32.8	68.3	-10.4	1039	6.27	0.63	1.74	1.74	0.63	1.74	81	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0		
Mar.	43.2	35.2	70.1	-13.1	1036	6.27	0.63	1.74	1.74	0.63	1.74	81	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0		
Apr.	49.3	43.2	76.8	-16.1	1079	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
May	59.3	49.3	80.7	-18.4	1122	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Jun.	63.5	51.5	80.7	-18.4	1122	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Jul.	73.0	50.7	80.7	-18.4	1122	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Aug.	79.0	50.7	80.7	-18.4	1122	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Sep.	80.7	50.7	80.7	-18.4	1122	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Oct.	80.7	50.7	80.7	-18.4	1122	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Nov.	80.7	50.7	80.7	-18.4	1122	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Dec.	80.7	50.7	80.7	-18.4	1122	5.83	0.69	1.904	5.83	0.69	1.904	5.83	71	71	75	10.8	SW	56	SW	1924	5	8	20	13	0	1	0	10	23	0	
Yearly	59.3	43.2	76.8	-16.1	1079	58.9	33.16	38.7	29.8	12.3	18.93	76.82	71	9.6	3.7	7.6	63	N	1879	71	4.7	11	14	6	10	0	7	0	5	0	0

TABLE 6. Climatic data for Cleveland, Ohio.
[Climates of the States (1974)]

LATITUDE 41° 24' N
LONGITUDE 81° 51' W
ELEVATION (ground) 787 feet

CLEVELAND, OHIO
CLEVELAND HOPKINS AIRPORT

Month	Temperature				Normal degree days	Precipitation				Relative humidity				Wind				Mean number of days																	
	Daily maximum	Daily minimum	Monthly	Extremes		Normal total	Maximum	Minimum	Monthly	Year	7:00 a.m. EST	1:00 p.m. EST	7:00 p.m. EST	Mean hourly	Prevailing direction	Speed	Direction	Fastest mile	Year	Clear	Partly cloudy	Cloudy	Precipitation	Thunderstorms	Heavy fog	90° and above	32° and below								
Jan	36.4	20.9	28.5	73 1950	1132	2.38	7.01	1950	0.79	1945	1.74	1952	18.18	18.18	8	16	16	16	16	16	16	16	16	16	16	16	16	16	16	16	16	16			
Feb	36.0	20.8	28.6	69 1957	1019	2.12	6.64	1950	0.73	1947	1.77	1950	80.82	73.70	12.5	7	39	SW 1950	27.8.1	3	5	23	16	4	2	0	12	27	2	0	0				
Mar	45.3	22.3	36.8	83 1943	874	2.89	6.07	1954	1.50	1947	2.78	1948	73.50	69.12	3.5	8	65	W 1936+	34.7.7	3	6	19	13	4	2	0	9	24	1	0	0				
Apr	57.1	37.5	47.3	69 1943	531	2.73	5.80	1950	1.15	1946	1.52	1950	77.78	64.83	12.0	4	41	W 1948	45.7.1	6	7	18	16	4	2	1	0	5	21	0	0	0			
May	69.9	46.2	59.1	82 1944	223	2.73	6.04	1947	1.04	1954	3.73	1955	82.78	55.40	10.5	3	62	SW 1957	50.8.9	5	17	15	1	4	2	0	0	8	0	0	0	0			
Jun	80.4	58.3	69.4	101 1944	46	3.05	6.07	1947	1.38	1948	2.79	1948	82.78	54.81	9.5	3	56	S 1953+	51.8.0	8	13	11	11	0	6	1	5	0	0	0	0	0	0		
Jul	84.7	62.9	73.7	103 1941	0	3.04	5.37	1943	1.23	1923	2.73	1950	81.79	52.158	8.9	8	65	W 1956	69.5.3	10	12	9	10	0	7	1	7	0	0	0	0	0	0	0	
Aug	82.7	61.1	71.9	103 1946	10	2.64	5.19	1947	1.61	1943	3.07	1947	83.82	53.183	8.5	8	61	W 1956	67.5.2	10	13	8	9	0	5	1	6	0	0	0	0	0	0	0	
Sep	75.6	55.1	65.5	101 1943	75	2.43	5.30	1943	1.57	1953	1.85	1945	81.83	53.87	9.3	8	45	S 1953+	63.2.3	10	10	10	9	0	3	2	0	0	0	0	0	0	0	0	0
Oct	64.3	44.5	54.4	90 1946	345	3.03	6.44	1943	0.9	1954	3.44	1954	78.82	55.70	10.2	3	43	W 1948	58.5.8	11	8	12	10	0	2	1	0	0	1	0	0	0	0	0	
Nov	49.3	34.2	41.7	82 1920	699	2.86	6.44	1950	0.9	1954	3.44	1950	77.80	64.73	12.8	5	59	W 1948	32.7.7	3	7	20	15	2	1	0	0	2	13	0	0	0	0	0	
Dec	37.5	24.3	30.9	69 1941	1057	2.29	5.80	1953	0.98	1953	1.26	1957	76.80	70.77	12.9	5	49	W 1948	28.9.0	3	6	22	16	5	1	0	10	25	1	0	0	0	0	0	
Year	59.9	41.3	50.6	103 1941	8006	32.08	9.50	1954	0.61	1953	3.73	1955	79.80	60.48	11.1	8	74	W 1948	49.6.6	75	106	184	166	20	36	14	20	38	119	4	0	0	0		

Means and extremes in the above table are from the existing or comparable location(s). Annual extremes have been exceeded at prior locations as follows:
 Lowest temperature -17 in January 1873; maximum monthly precipitation 9.77 in January 1902; minimum monthly precipitation 0.17 in August 1881; maximum precipitation in 24 hours 4.97 in September 1901; maximum monthly snowfall 30.5 in February 1908; maximum snowfall in 24 hours 17.4 in November 1913; fastest mile of wind 78 from Southwest in May 1940.

TABLE 8. Climatic data for Buffalo, New York.
[Climates of the States (1974)]

Station	BUFFALO, NEW YORK		GREATER BUFFALO INTL AIRPORT		Standard time used		EASTERN		Latitude		42° 46' N		Longitude		78° 44' W		Elevation (ground)		705 feet		
	Month	Yearly	Maximum	Minimum	Normal	Days (base 65°)	Normal	Maximum	Minimum	Yearly	Maximum	Minimum	Yearly	Maximum	Minimum	Yearly	Maximum	Minimum	Yearly	Maximum	Minimum
J	30.9	18.2	24.5	61	1987	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
F	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
M	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
A	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
M	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
J	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
J	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
A	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
S	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
O	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
N	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
D	30.0	17.2	24.1	50	1965	11	11	1968	11	11	11	11	11	11	11	11	11	11	11	11	11
YR	55.4	37.9	46.7	64	1966	120	120	1966	120	120	120	120	120	120	120	120	120	120	120	120	120

Means and extremes above are from existing and comparable exposures. Annual extremes have been exceeded at other sites in the locality as follows: Highest temperature 99 in August 1948; lowest temperature -21 in February 1934; maximum monthly precipitation 10.63 in August 1885; minimum monthly precipitation .05 in August 1876; maximum precipitation in 24 hours 4.28 in August 1883.

TABLE 9. Climatic data for Detroit, Michigan.
[Climates of the States (1974)]

Temperature		Extremes		Normal degree days		Precipitation						Relative humidity			Wind			Mean number of days																						
Month	Normal		Record highest	Record lowest	Year	Year	Year	Year	Year	Year	Year	Year	Snow, Sleet		Relative humidity		Wind			Sunrise to sunset	Cloudy	Precipitation	Snow Sleet	Thunderstorms	Heavy fog	32 and above	32 and below	Max.	Min.											
	Daily maximum	Daily minimum											Month	Year	Month	Year	Month	Year	Month											Year	Month	Year	Month	Year	Month	Year	Month	Year	Month	Year
J	33.9	19.1	26.5	62	1965	-17	1963	1194	1.93	3.63	1965	0.27	1961	1.40	1960	7.7	13.4	1959	5.3	1965	76.77	69.72	11.2	65W	5	7	19	12	3	0	17	30	4							
F	35.0	19.2	27.1	58	1966	6	1963	1061	1.95	2.54	1965	0.67	1963	1.23	1965	9.4	17.4	1962	10.3	1965	77.79	68.71	11.0	65W	3	7	17	12	3	0	13	27	2							
M	43.6	28.1	34.9	77	1963	1	1963	933	2.41	3.59	1965	0.92	1960	1.13	1965	6.6	16.1	1965	5.5	1965	76.80	63.67	11.4	65W	3	8	19	14	2	0	6	24	0							
A	57.7	36.6	47.2	85	1962+	17	1964	534	3.05	5.40	1961	1.80	1960	1.97	1965	1.6	7.4	1961	4.2	1961	77.91	67.62	11.1	65W	4	1	8	16	1	0	11	0	0							
M	69.7	47.1	58.4	92	1962	25	1966	239	3.54	4.09	1963	1.15	1965	1.72	1966	0	0	1963+	77.78	64.56	10.8	65W	6	1	4	1	10	10	0	0	0	0	0							
J	79.9	59.8	68.4	94	1965+	36	1966	57	3.31	4.60	1960	2.12	1959	2.62	1960	0.0	0.0	1963+	80.80	62.57	8.5	65W	5	0	1	4	0	0	0	0	0	0	0							
J	84.7	61.1	72.9	98	1966	61	1965	0	2.69	3.26	1966	1.11	1964	3.19	1966	0.0	0.0	1963	81.82	51.57	6.3	65W	5	1	5	0	0	0	0	0	0	0	0							
A	89.6	65.0	74.9	99	1966	81	1965	2.85	2.85	1.70	1965	1.44	1963	2.71	1964	0.0	0.0	1963	84.88	58.67	8.5	65W	3	0	3	0	0	0	0	0	0	0	0	0						
S	74.1	50.0	61.3	94	1960	37	1964	96	2.89	4.14	1965	0.35	1964	2.11	1965	0.0	0.0	1965+	84.88	57.68	8.5	65W	4	2	3	0	0	0	0	0	0	0	0	0						
O	64.8	43.1	54.0	91	1963	78	1965	353	2.57	4.14	1965	0.35	1964	2.11	1965	0.0	0.0	1965+	80.80	54.67	9.1	65W	2	3	3	0	0	0	0	0	0	0	0	0						
N	68.9	31.8	40.4	77	1964	9	1964	738	2.27	3.13	1966	0.50	1964	1.28	1961	3.4	11.8	1956	5.2	1965+	85.65	65.73	10.4	65W	1	3	5	0	1	3	3	0	1	18	0					
D	36.5	23.2	29.9	66	1966	9	1960	1088	1.92	4.00	1965	0.46	1960	3.71	1965	8.1	17.3	1962	5.7	1966	79.80	70.75	10.4	65W	3	0	3	0	14	27	2	0	0							
YR	56.6	39.7	49.7	98	JUL. 1946	-13	1963	6293	30.80	7.70	1964	0.27	1961	3.71	1965	37.0	17.4	1962	10.1	1965	80	62	60	66	9	8	5W	52	6.4	77	110	178	130	13	34	26	13	50	143	7

0 Data from Detroit City Airport through April 1966.

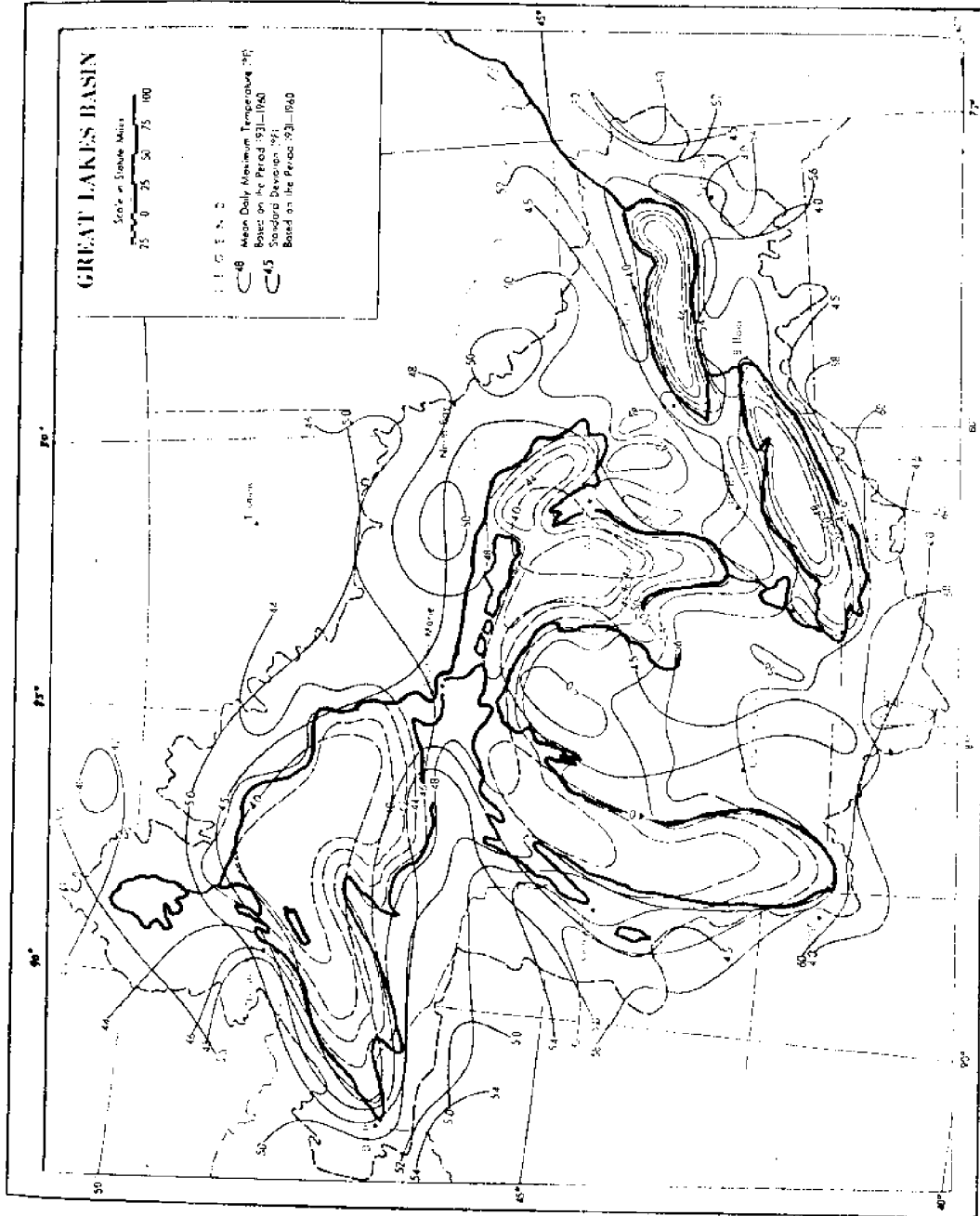


Figure 7. April mean daily maximum temperature and its standard deviation (°F).
[Phillips and McCulloch (1972)]

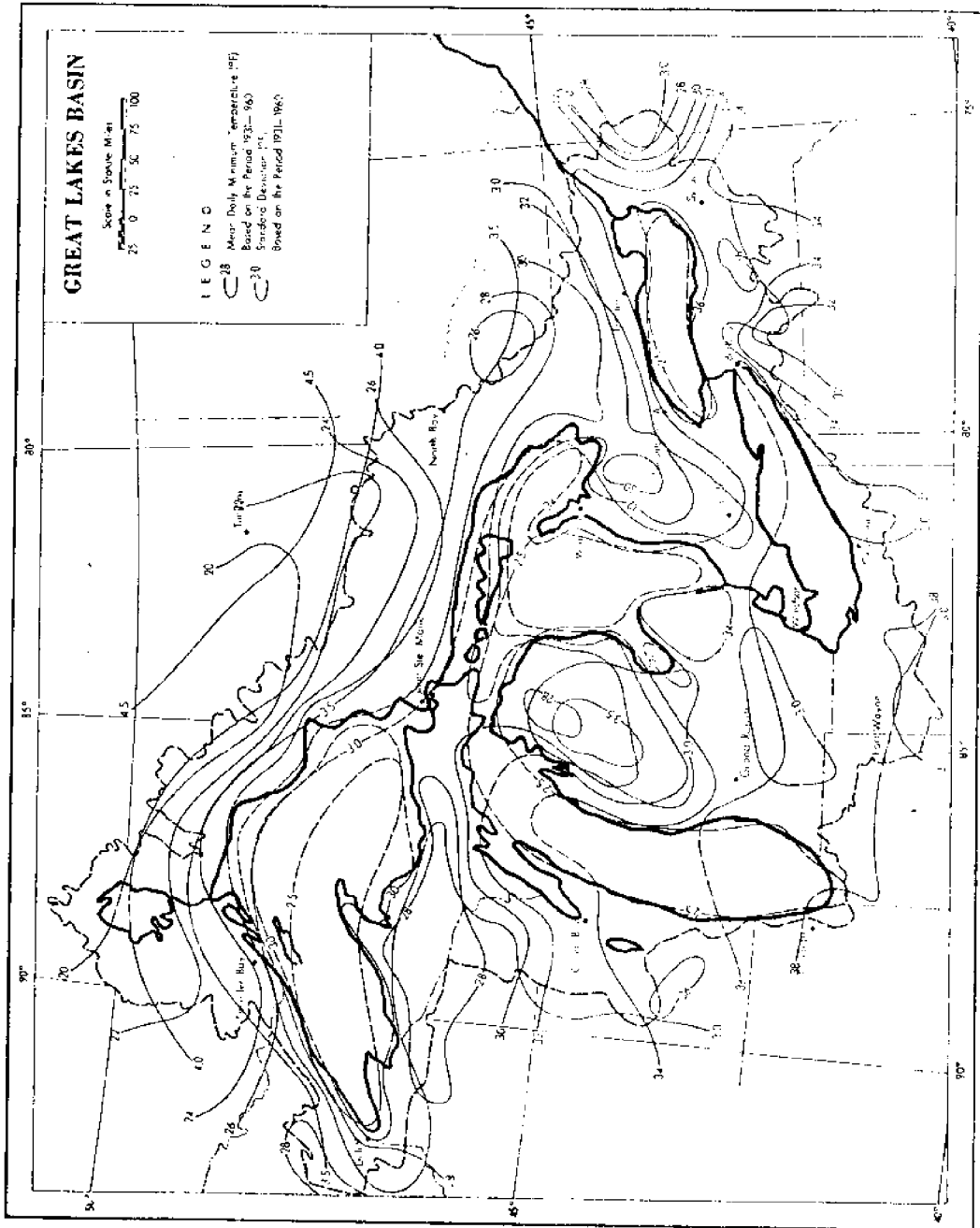


Figure 8. April mean daily minimum temperature and its standard deviation (°F).
[Phillips and McCulloch (1972)]

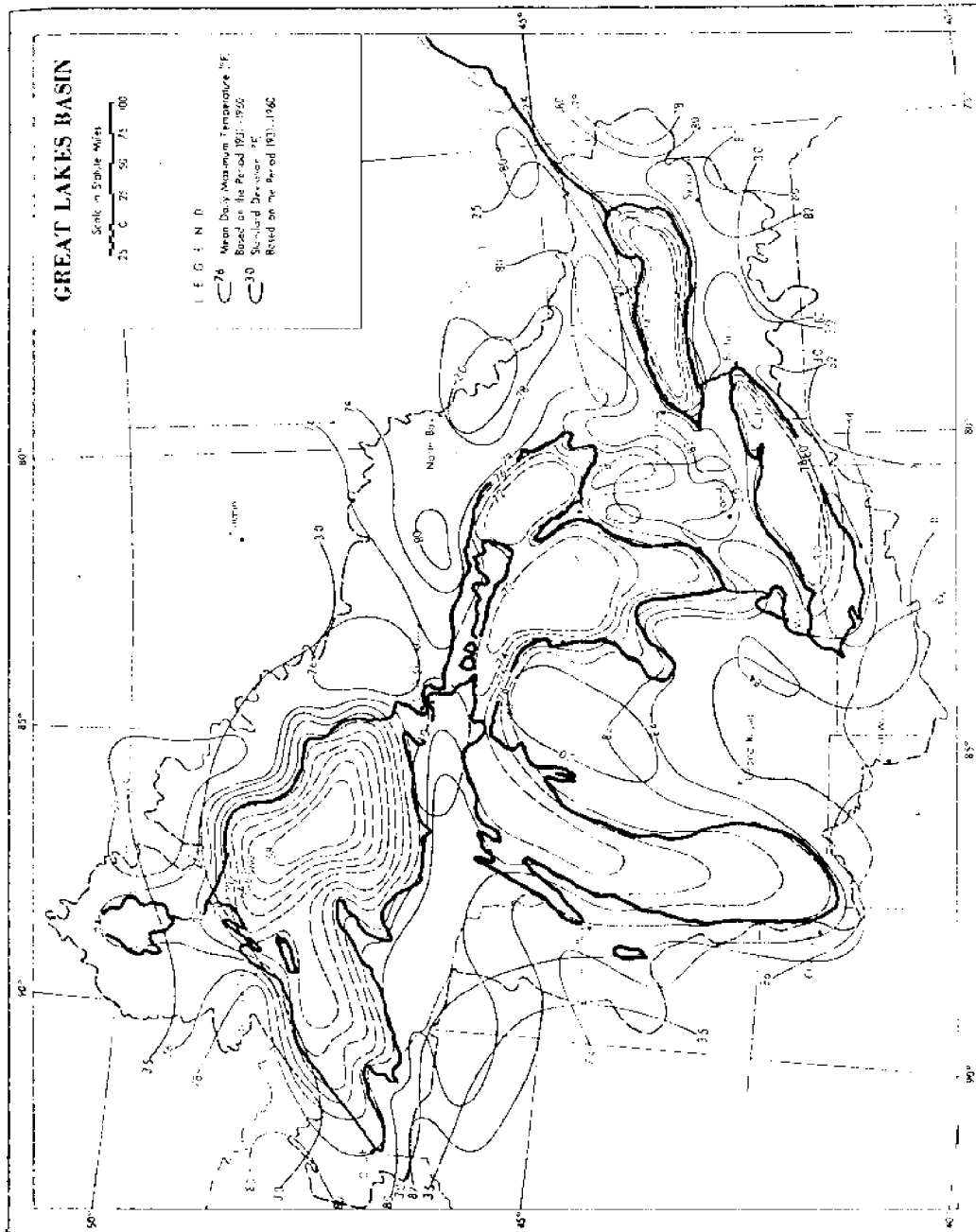


Figure 9. July mean daily maximum temperature and its standard deviation (°F).
[Phillips and McCulloch (1972)]

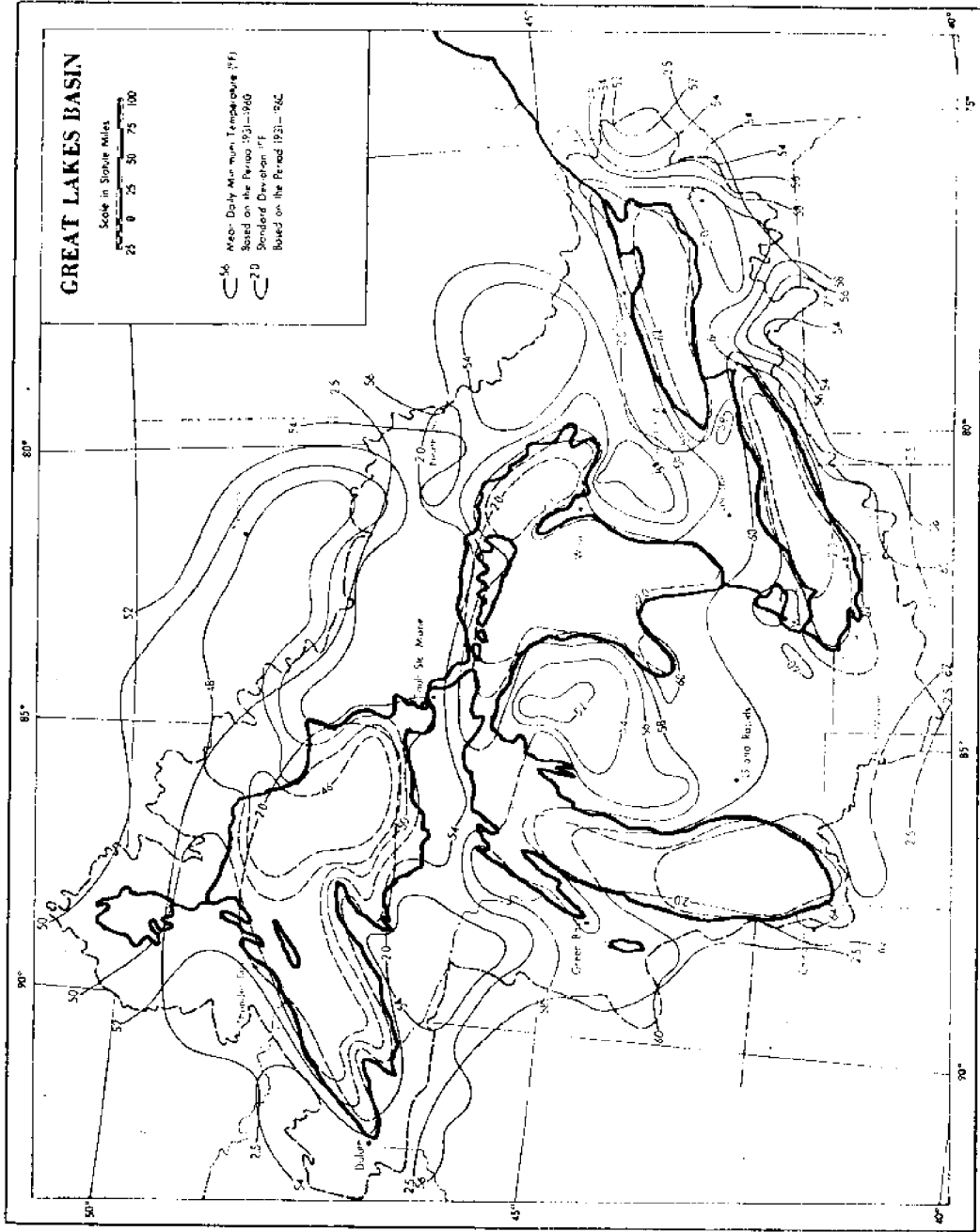


Figure 10. July mean daily minimum temperature and its standard deviation (°F).
[Phillips and McCulloch (1972)]

TABLE 10. Average monthly values of temperature and precipitation for Cleveland and North Royalton, Ohio (1893-1921 means). Means are calculated from yearly values given in Alexander (1923).

		<u>Temperature</u> (°F)					
		APR	MAY	JUN	JUL	AUG	SEP
Cleveland		46.6	58.0	66.9	72.0	70.0	64.6
N. Royalton		47.2	58.5	67.3	72.2	70.3	64.5

		<u>Precipitation</u> (inches)					
		APR	MAY	JUN	JUL	AUG	SEP
Cleveland		2.59	2.92	2.72	3.34	2.93	2.80
N. Royalton		3.10	3.87	3.44	4.29	3.68	3.33

Ohio. Both cities are located in Cuyahoga County, but Cleveland is on the lakeshore while North Royalton is in the southern part of the county, about 12 miles inland. Note Cleveland's mean monthly temperatures are cooler than those of North Royalton from April through August; in September, the slightly warmer lake (Table 3) gives Cleveland a warmer mean temperature.

To see the lake effect on an island station, consider Table 11, which lists the 1921-1950 mean monthly temperatures and precipitation for Gibraltar Island and Sandusky, Ohio. Sandusky is a shoreline station whereas Gibraltar is a small island on Lake Erie. From April through June the mean temperature on Gibraltar Island is lower, but from July through September it is higher, presumably due to the greater solar radiation received over the lake versus shoreline and inland stations.

TABLE 11. Average monthly values of temperature and precipitation for Gibraltar Island and Sandusky, Ohio (1921-1950 means). [Ref. (8)]

	<u>Temperature</u> (°F)					
	APR	MAY	JUN	JUL	AUG	SEP
Gibraltar I.	45.7	57.9	69.3	74.7	74.2	67.4
Sandusky	47.9	59.3	69.9	74.6	72.8	66.5

	<u>Precipitation</u> (inches)					
	APR	MAY	JUN	JUL	AUG	SEP
Gibraltar I.	2.98	3.19	3.11	2.46	2.88	2.66
Sandusky	2.96	3.32	3.73	3.45	2.81	3.26

4.2. Humidity

The amount of moisture in the air is important in determining the rate of evaporation and also for determining the convective stability or instability of air, an important factor in precipitation development.

Vapor pressure e is one measure of the amount of water vapor in the air. To see the lake effect on vapor pressure, consider Figure 11, which is a graph of the mean vapor pressure at Wiarton and Killaloe, Ontario. Wiarton is virtually surrounded by lakes (Georgian Bay and Lake Huron) while Killaloe is 150 miles inland. Note that vapor pressures are virtually the same at both locations from May through July, but higher beginning in August. Hence, in spring and early summer, the cooler lake waters have little effect on the relative amount of water vapor in the air, but beginning about mid-summer, evaporation over the lakes supplies more atmospheric moisture than does evaporation over land.

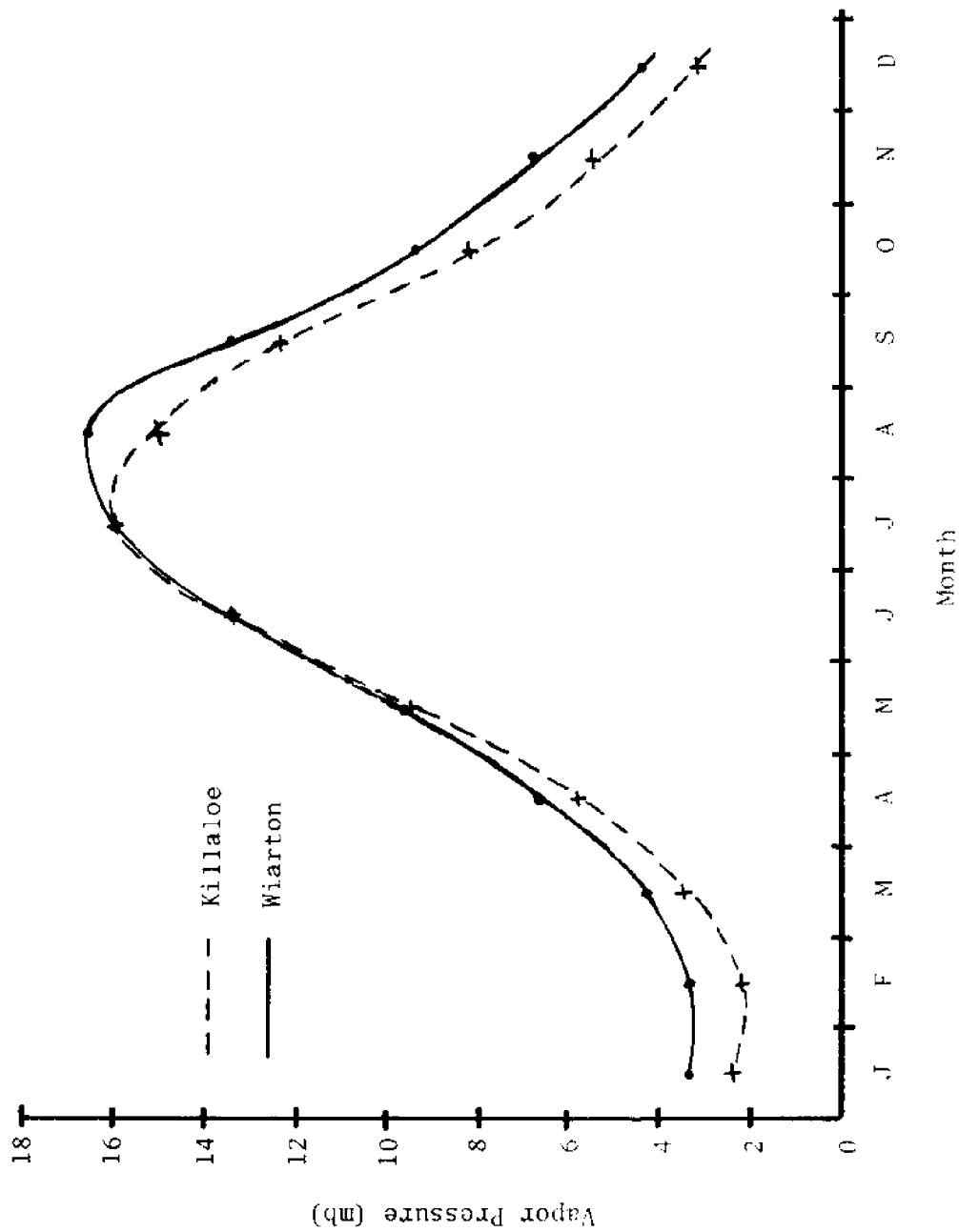


Figure 11. Mean Monthly vapor pressures at Wiarnton and Killaloe, Ontario.
 [Adapted from Phillips and McCulloch (1972)]

In contrast to vapor pressure, *relative humidity* r , defined as

$$r \equiv e/e_s ,$$

where e_s is the saturation vapor pressure, may in the spring be greater over Lake Erie than over land, since the overlying lake air, being cooler, has a lower e_s value than over land. That such is the case is confirmed by relative humidity increases after passage of a lake-breeze front [Munn and Richards (1964)]. Tables 4 - 9 give the average monthly relative humidity at four times of the day for the six major cities surrounding Lake Erie.

4.3. Solar Radiation

Due to frequent subsidence over Lake Erie in the spring and early summer (see Chapter 3), it might be inferred that areas over and near the lake should have more sunshine than areas further inland.

Figure 12 depicts the mean number of hours of sunshine for July in the Great Lakes region. Note the local maximum over Lake Erie. In addition, Tables 4 - 9 list the mean monthly percentage of possible sunshine for the six cities around the lake.

Figure 13 shows the mean monthly daily solar radiation at Sault Ste. Marie, Michigan, and Ottawa, Ontario. Ottawa is an inland station, only 1° of latitude further south. Note the greater daily amounts at Sault Ste. Marie during April through August, when the cooler Great Lakes which surround that city encourage subsidence and cloud dissipation.

4.4. Wind

The mean monthly wind speed and prevailing wind direction for the six major cities surrounding Lake Erie are given in Tables 4 - 9.

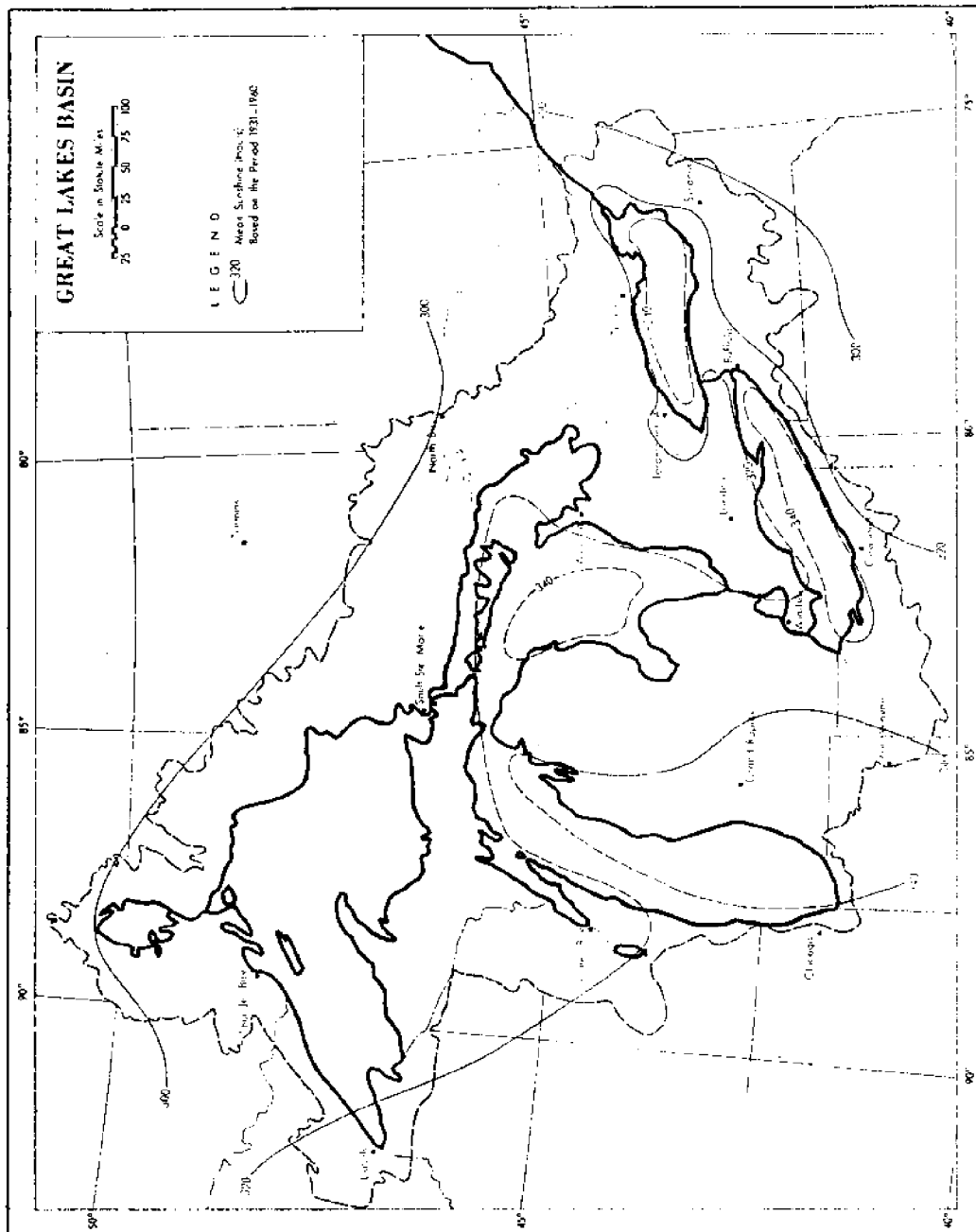


Figure 12. Mean number of hours of sunshine for July.
[Phillips and McCulloch (1972)]

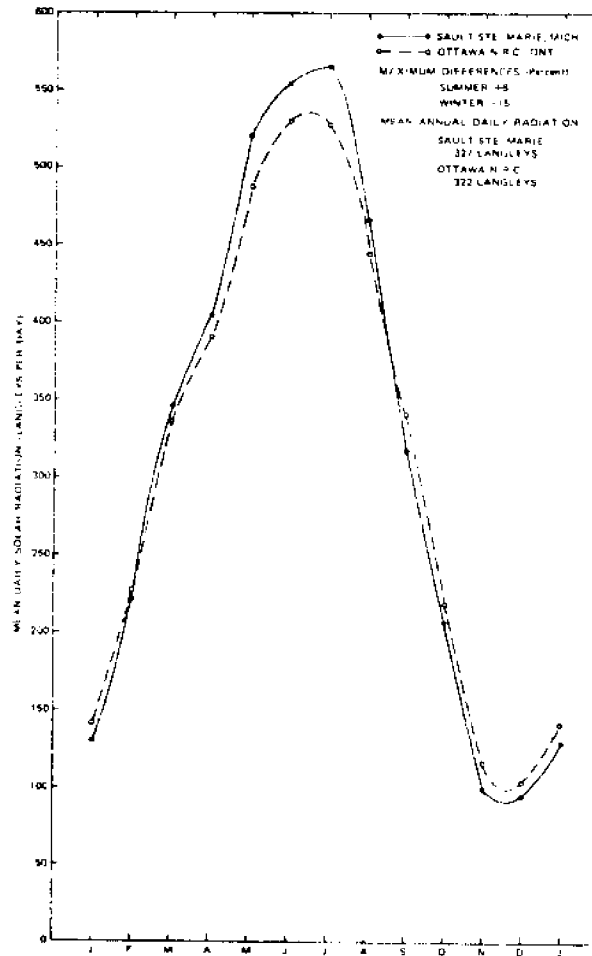


Figure 13. Mean daily solar radiation at Ottawa N.R.C., Ontario, and Sault Ste. Marie, Michigan. [Phillips and McCulloch (1972)]

April and July wind roses for the Great Lakes region are shown in Figures 14 and 15, respectively. The prevailing synoptic flow is out of the southwest during these months but lake breezes during these months lead to maxima in directions perpendicular to the shoreline as well (e.g., north wind at Cleveland, etc.).

Figure 16 depicts wind roses for April through September based on ship observations over western Lake Erie. Note the prevalence of southerly and southwesterly flows during May through August. The mean wind speed falls to a minimum during the summer months, as might be expected.

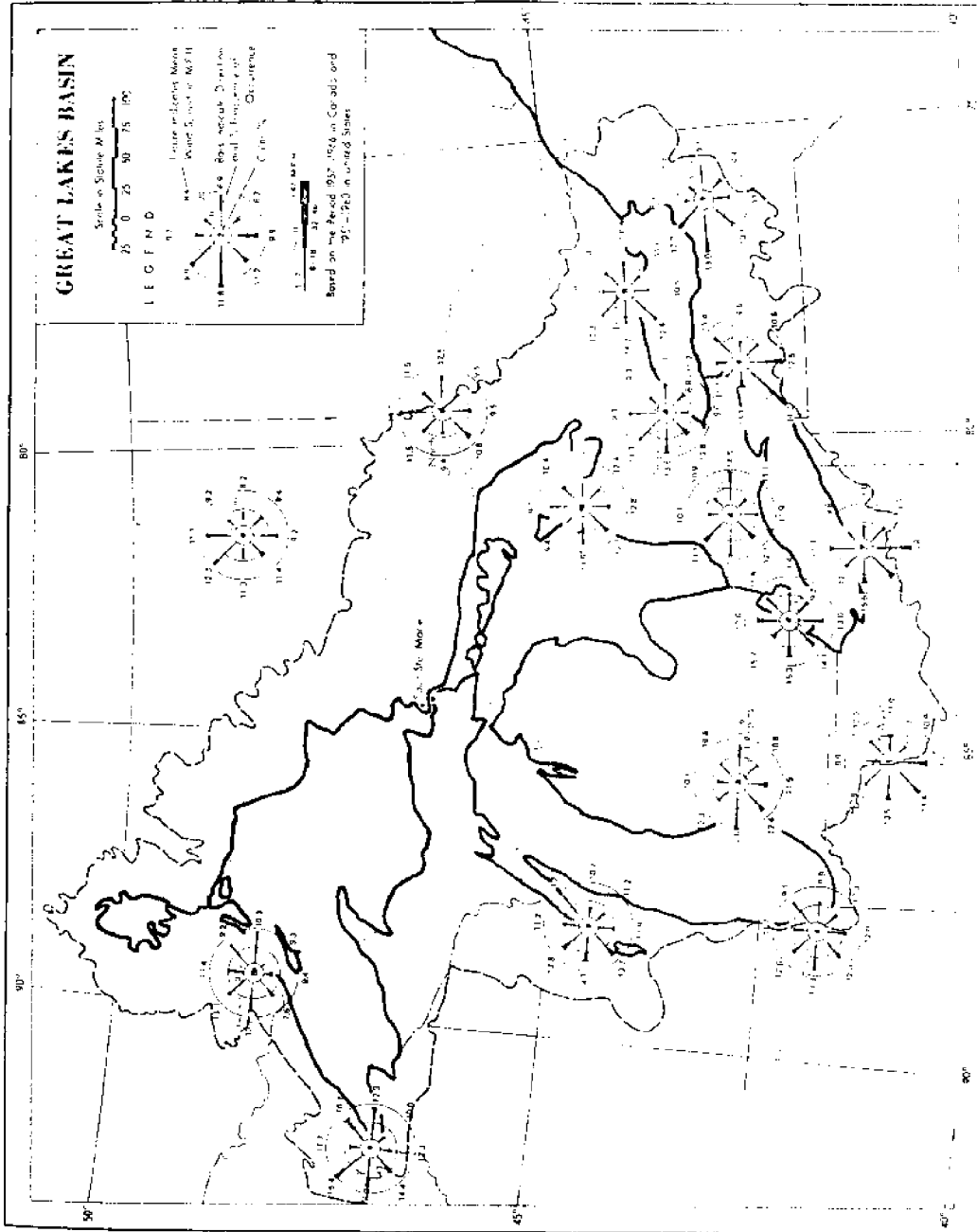


Figure 14. April wind roses for selected airport stations.
[Phillips and McCulloch (1972)]

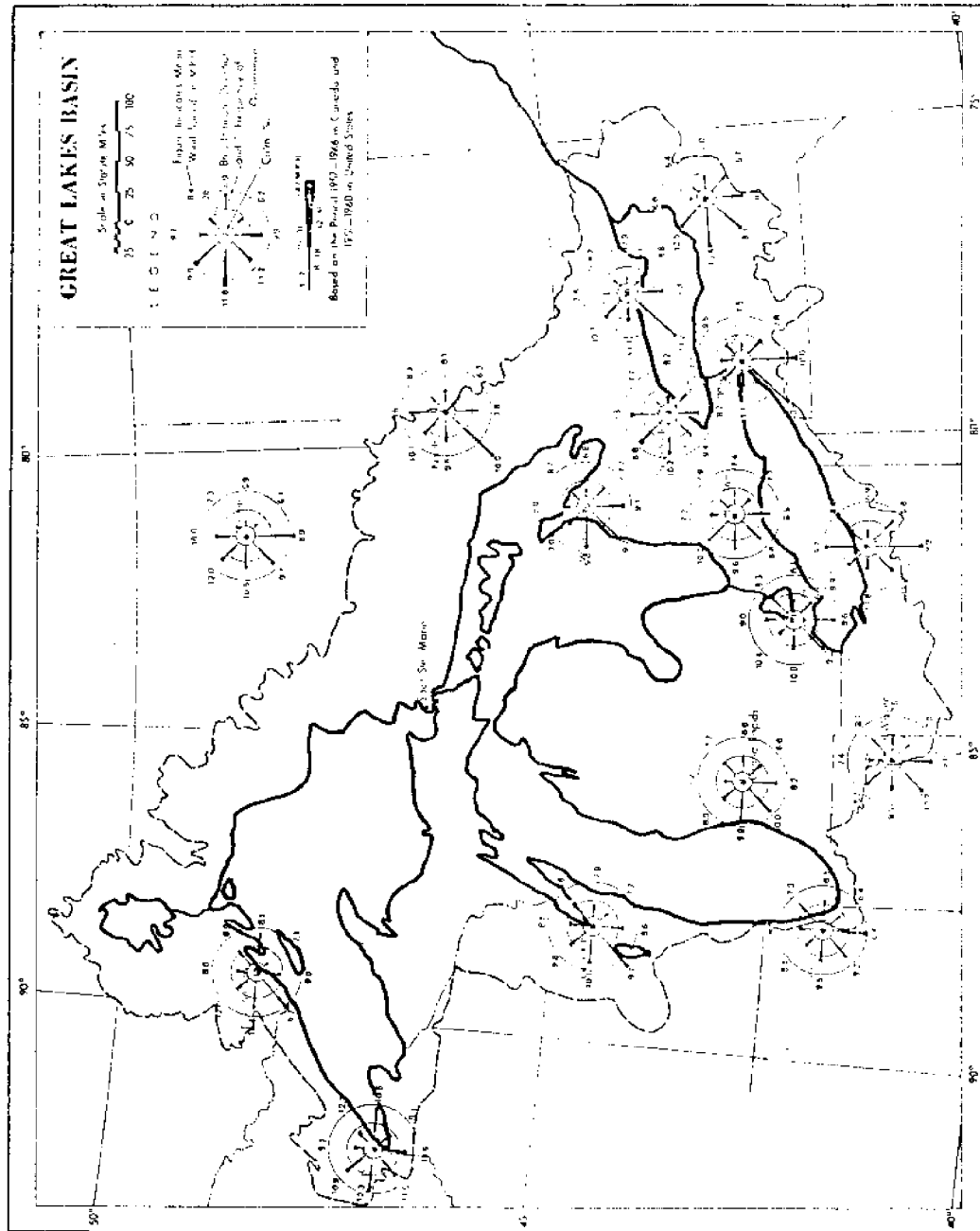


Figure 15. July wind roses for selected airport stations.
 [Phillips and McCulloch (1972)]

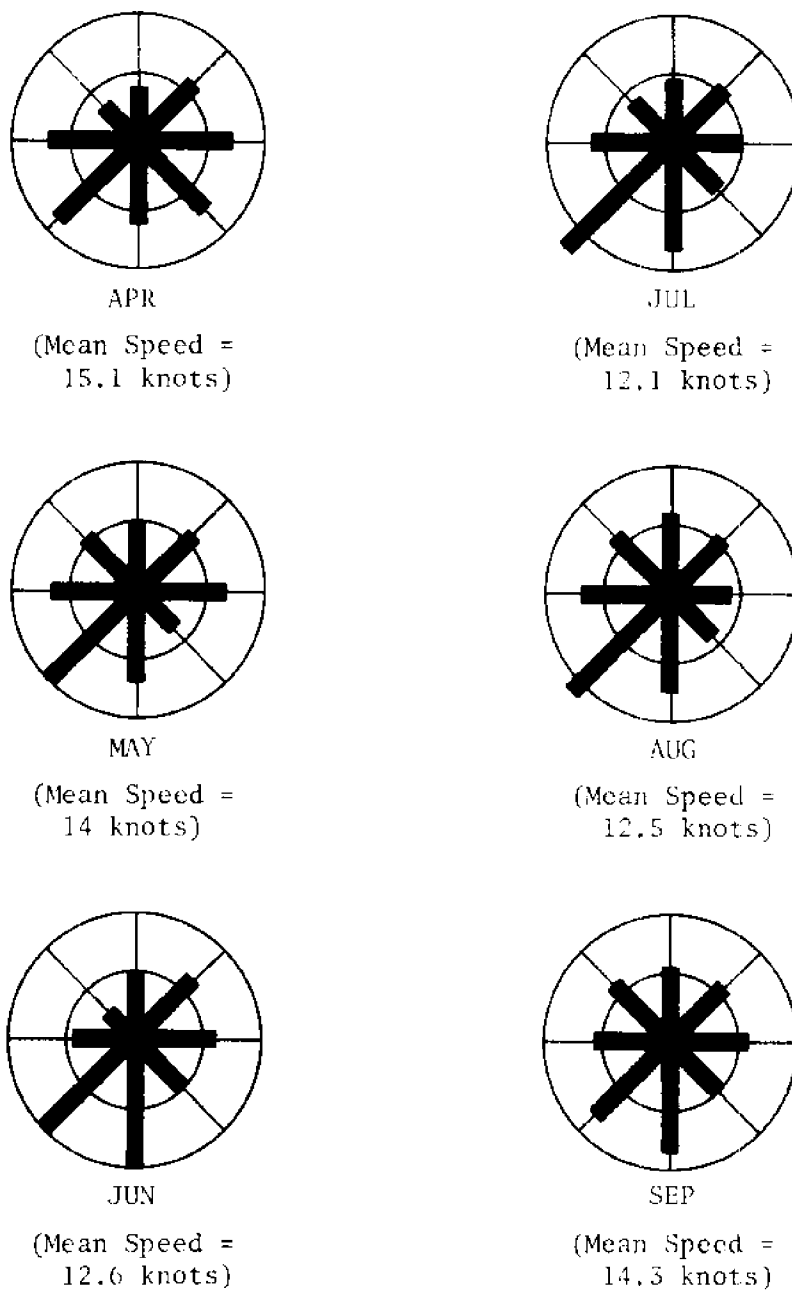


Figure 16. Wind roses for western Lake Erie. Data from 1963-1973 (June-September), 1964-1973 (May), and 1965-1973 (April). Contour interval is 10% frequency. [Adapted from Ref. (53)]

What is the lake effect on wind? Because the surface of a lake is much smoother than land and therefore exerts a lower stress on the atmosphere, one at first might expect speeds to increase as air flows over water. However, Lyons (1970) and Bellaire (1965), in their case studies, observed speeds to decrease because of the stable thermal stratification over the water. One might also expect winds to be more gradient in direction due to less friction (to veer over the water). Lyons (1972), as noted earlier, found the reverse because of wind adjustment to the lake mesohigh. However, with neutral or unstable lapse rates over the surface of the lake, one might indeed expect stronger winds over the lake due to reduced friction; winds should also be more nearly gradient in direction.

Richards et al. (1966) performed an observational study on the influence of atmospheric stability on winds over Lakes Erie and Ontario. Using five years of ship data, they found wind speeds to decrease over water during stable stratification over the lakes and speeds to increase during unstable stratification. The greatest changes (in speeds over land to those over water) are experienced at lower wind speeds and occur over an over-water fetch of up to 25 miles, beyond which winds remain nearly constant.

Looking at Figure 2, one might, on the basis of these studies, make the following predictions for winds over Lake Erie during spring and summer. During the daytime with light synoptic winds, winds will decrease over Lake Erie. At nighttime with light synoptic winds, speeds will increase during the months of July through September.

4.5. Evaporation

Various investigators have provided estimates of the mean monthly

and annual evaporation from Lake Erie. The two most common techniques used are the water budget and mass transfer methods; these are well known and will not be discussed here.

Table 12 provides a summary of the evaporation estimates of the more comprehensive studies of Lake Erie. Monthly values for April through September and annual means are given. Note that the cooler lake suppresses evaporation in the spring months (in April, condensation is even indicated in three of the five studies); beginning in July, however, evaporation increases, reaching a maximum in early fall. Annual evaporation of around 35 inches occurs. These numbers should be regarded only as estimates, however, since both techniques involve the uncertain extrapolation of either precipitation or wind speeds over the lake.

Over the land surrounding Lake Erie, annual evaporation is much less, as might be expected. Phillips and McCulloch (1972) estimate a value of 22 inches. As inferred from Carlson (1974), average monthly land evaporation around Lake Erie for April through June is approximately 2.5 inches in April, 3.4 inches in May, and 3.9 inches in June. The comparisons in Table 12 clearly show that, in contrast to the annual results, more evaporation occurs over land than over Lake Erie in the spring months.

4.6. Precipitation

Tables 4 - 9 present the average monthly precipitation amounts for the six major cities surrounding Lake Erie. Figure 17 presents the monthly averages for Sandusky in histogram format. Amounts are rather evenly distributed with higher amounts in March through September.

TABLE 12. Estimates of evaporation for Lake Erie (inches).

Investigator(s)	Derecki (1964)	Richards & Irbe (1969)	Jones & Meredith (1972)	Perecki (1975)
Years of Data Used	1957-59 (23)	1950-68 (19)	1946-65 (20)	1937-68 (32)
Technique	Water Budget	Mass Transfer	Mass Transfer	Water Budget Mass Transfer
APRIL	0.00	-1.10	-1.08	0.31 -0.63
MAY	0.39	2.31	2.31	0.63 2.36
JUNE	1.18	2.91	2.86	1.14 2.24
JULY	3.54	2.65	2.50	3.78 2.20
AUGUST	5.51	4.45	4.46	5.35 4.21
SEPTEMBER	6.30	5.59	5.67	6.58 5.55
Annual	33.46	35.88	35.52	35.79 35.35

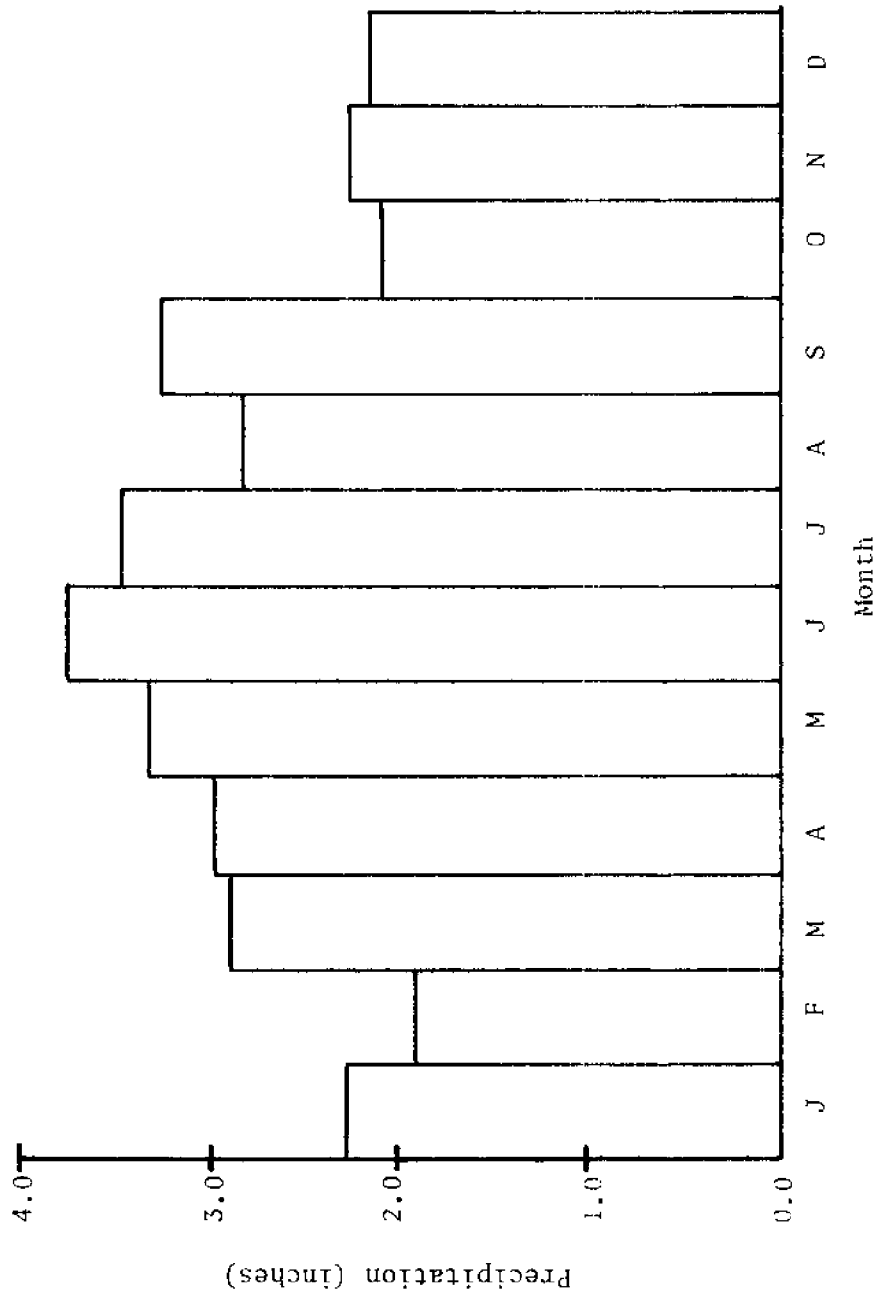


Figure 17. Histogram of average monthly precipitation amounts at Sandusky, Ohio.
 [Data from Climates of the States (1974)]

4.6.1. Long-Term Lake Effect. What is the effect of Lake Erie on precipitation during the spring and summer? On a long-term basis (average of many years) the cooler waters of Lake Erie during this period definitely reduce precipitation over the lake and shoreline areas as compared to further inland. Consider Table 10 again, this time with respect to precipitation. Monthly amounts at Cleveland are lower by 0.5 to 0.95 inches over those at North Royalton. Also, consider Table 11. During May through September, with the exception of August, monthly amounts are lower out over the lake (Gibraltar Island) than at the shore. Figure 18 depicts the mean July precipitation over the Great Lakes region. The lower amounts as Lake Erie is approached are clearly seen.

Blust and DeCooke (1960) made a comparison of precipitation over the islands of northeastern Lake Michigan with that over the lake perimeter. Their study covered the years 1952-1959. They found a decrease in precipitation of 6.2% over the islands as compared with over the lake perimeter during the summer months.

Using climatological data from land stations surrounding Lake Michigan, Changnon (1966) inferred that during the summer the lake reduces thunderstorms by more than 20% over and to the lee of the southern half of the lake. At night the warmer lake actually increases thunderstorm activity, but 30-50% reductions during the day cause a net summer reduction.

4.6.2. Short-Term Lake Effect. In any given year, however, it is unclear what effect a large cold lake will have on precipitation. For example, consider Table 13, which lists the observed monthly precipitation amounts from May through September 1916 for North Bass Island and Danbury, Ohio, a station on Sandusky Bay. North Bass is well out into Lake Erie and should reflect any lake effect. However, for this year, North Bass

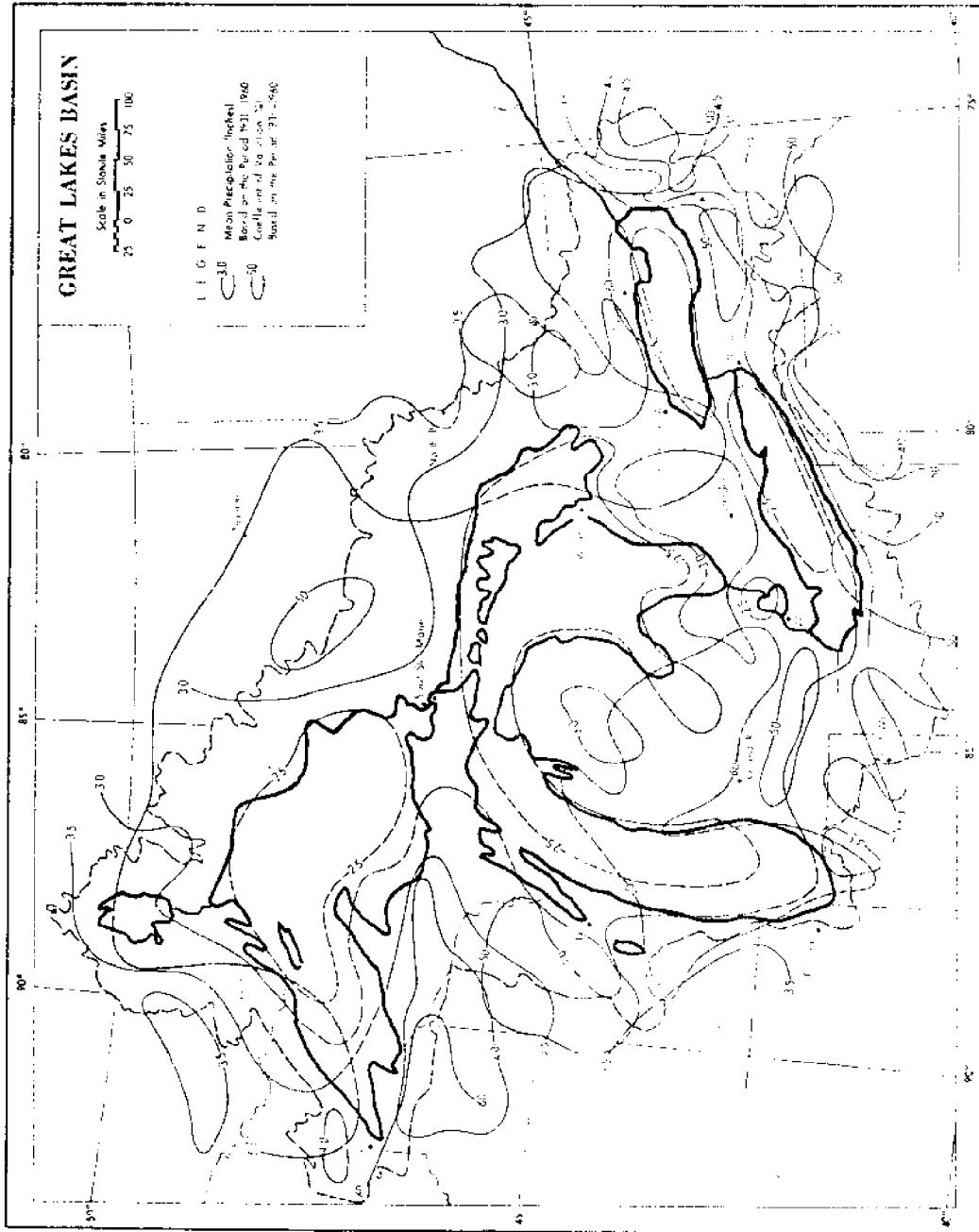


Figure 18. Mean July precipitation (inches) and its coefficient of variation (%).
 [Phillips and McGulloch (1972)]

shows greater precipitation than Danbury in all months except July. Also, consider Table 14, which gives the 1873-75 averages of monthly precipitation over Kelly's Island and Margaretta, Ohio, an inland station in Erie County. Note the island station has greater precipitation during June and July than does the inland station, the reverse of that expected.

TABLE 13. 1916 monthly precipitation amounts for North Bass Island and Danbury, Ohio (inches).
[Data from Alexander (1923)]

	MAY	JUN	JUL	AUG	SEP
North Bass	5.67	4.58	0.63	2.08	2.08
Danbury	4.49	3.89	0.73	1.91	1.62

TABLE 14. Three-year (1873-1875) monthly precipitation averages for Kelly's Island and Margaretta, Ohio (inches).
[Data from Alexander (1923)]

	APR	MAY	JUN	JUL	AUG	SEP
Kelly's Island	3.19	2.40	4.08	5.12	2.46	3.15
Margaretta	3.55	3.16	3.98	3.53	3.96	3.81

Why the discrepancy between what occurs in a given year or several year period and what occurs in long-term averages? One possibility is that over the long term the cooling effect of Lake Erie in spring and summer does encourage subsidence and inhibit precipitation, certainly the local summer convective type. However, in the short term the prevailing synoptic systems most likely override this "climatological" effect and may produce more rain over the lake. Lyons (1966) believes in some cases that the large lake may even intensify the precipitation, as the systems experience uplift over the cold air dome as they move

out over the lake. In other cases the lake may induce no effect at all, as the storms and advection move right over the cold dome, which serves to reduce friction. It may also be that the uplift associated with the lake breeze front encourages more precipitation over shoreline areas than over inland areas.

4.6.3. Over-Lake Precipitation. Using ratio techniques (based on long-term precipitation averages over island versus shoreline stations) to extrapolate precipitation amounts from shoreline stations to areas over Lake Erie, various investigators have estimated precipitation over Lake Erie. Some of these estimates appear in Table 15 for April through September and also for the annual amount. Such techniques are highly suspect, however, because of the large geographical area involved in the extrapolation, and the numbers should be regarded merely as estimates.

TABLE 15. Estimates of precipitation over Lake Erie (inches).

Investigator(s)	Derecki (1964)	Jones & Meredith (1972)	Derecki (1975)
Years of Data Used	1937-59 (23)	1946-65 (20)	1937-68 (52)
APRIL	3.46	3.40	3.38
MAY	3.58	3.04	3.11
JUNE	3.11	2.89	3.03
JULY	3.19	2.89	2.91
AUGUST	3.19	3.14	3.25
SEPTEMBER	2.76	2.45	2.56
Annual	34.64	32.79	32.83

4.6.4. Intensive Precipitation. The Climatological Data--National Summary annual series provides a data base for examining intensive rainfall in the Lake Erie region during the spring and summer months.

Tables 16-21 summarize the findings for the six major cities surrounding the lake. In each table are listed the number of years (of the period studied) when precipitation greater than or equal to the given hourly intensity was observed at least once within that month. The results show that the peak months for intensive rainfall are June through August. The months with the greatest intensity (≥ 1 in h^{-1}) are July and August; these months, although past the severe weather peak, are months with much humidity and, therefore, water vapor supply. Also, remnants of tropical hurricanes occasionally influence the area.

Table 22, inferred from Hershfield (1961), presents the annual probabilities of intense rainfall (of the given amount) for the Lake Erie region for April through September. As can be seen, the greatest probabilities occur in June through August, which agrees with the results deduced from Tables 16-21.

4.7. Severe Weather

Tables 4-9 give the average monthly number of thunderstorm days for the six cities around Lake Erie. At most stations the number peaks during June and July with approximately 5 to 7 such days per month. During the six-month spring-summer period, the average number of thunderstorm days for the six stations is 30.3.

Information on Ohio tornadoes was obtained from the National Weather Service at Columbus. Table 23 presents the monthly tornado frequency. The peak month is May, with April having a nearly equal number of occurrences. Figures 19 and 20 depict the number of tornadoes reported in each county (1900-1975) and number of deaths (1916-1975), respectively. The Lake Erie region is by no means immune; in fact, some of Ohio's worst tornadoes (those of 1924, 1953, and 1965) have tracked through counties bordering on the lake.

TABLE 16. TOLEDO: Number of years with at least one hourly precipitation occurrence per month greater than or equal to given intensity. Twenty-five years of data were used (1950-1972, 1976, 1978). [Ref. (7)]

	Number of years with precipitation greater than or equal to				
	<u>.25 in/h</u>	<u>.5 in/h</u>	<u>.75 in/h</u>	<u>1.0 in/h</u>	<u>1.5 in/h</u>
APR	2	2	1		
MAY	10	7	4		
JUN	18	16	9	6	3
JUL	17	14	10	5	1
AUG	18	13	8	8	4
SEP	10	8	7	3	1

TABLE 17. SANDUSKY: Number of years with at least one hourly precipitation occurrence per month greater than or equal to given intensity. Fourteen years of data were used (1950-1963). [Ref. (7)]

	Number of years with precipitation greater than or equal to				
	<u>.25 in/h</u>	<u>.5 in/h</u>	<u>.75 in/h</u>	<u>1.0 in/h</u>	<u>1.5 in/h</u>
APR	3	3	1		
MAY	7	6	4	2	
JUN	8	6	4	2	
JUL	7	6	3	2	1
AUG	8	7	3	3	1
SEP	6	3	2	2	1

TABLE 18. CLEVELAND: Number of years with at least one hourly precipitation occurrence per month greater than or equal to given intensity. Twenty-five years of data were used (1950-1972, 1976, 1978). [Ref. (7)]

	Number of years with precipitation greater than or equal to				
	<u>.25 in/h</u>	<u>.5 in/h</u>	<u>.75 in/h</u>	<u>1.0 in/h</u>	<u>1.5 in/h</u>
APR	6	5	1		
MAY	13	11	8	2	1
JUN	18	15	6	2	
JUL	19	17	15	10	1
AUG	16	13	10	7	1
SEP	11	9	5	2	1

TABLE 19. ERIE: Number of years with at least one hourly precipitation occurrence per month greater than or equal to given intensity. Twenty years of data were used (1950-1952, 1958-1972, 1976, 1978). [Ref. (7)]

	Number of years with precipitation greater than or equal to				
	<u>.25 in/h</u>	<u>.5 in/h</u>	<u>.75 in/h</u>	<u>1.0 in/h</u>	<u>1.5 in/h</u>
APR	3	1			
MAY	6	6	5	2	
JUN	17	17	6	2	
JUL	15	13	8	2	
AUG	14	10	9	5	1
SEP	8	6	3	1	

TABLE 20. BUFFALO: Number of years with at least one hourly precipitation occurrence per month greater than or equal to given intensity. Twenty-five years of data were used (1950-1972, 1976, 1978). [Ref. (7)]

	Number of years with precipitation greater than or equal to				
	<u>.25 in/h</u>	<u>.5 in/h</u>	<u>.75 in/h</u>	<u>1.0 in/h</u>	<u>1.5 in/h</u>
APR	3	2	1		
MAY	6	5	1		
JUN	12	9	4	1	
JUL	15	12	7	5	1
AUG	19	15	8	6	2
SEP	11	9	7	4	1

TABLE 21. DETROIT: Number of years with at least one hourly precipitation occurrence per month greater than or equal to given intensity. Twenty-five years of data were used (1950-1972, 1976, 1978). [Ref. (7)]

	Number of years with precipitation greater than or equal to				
	<u>.25 in/h</u>	<u>.5 in/h</u>	<u>.75 in/h</u>	<u>1.0 in/h</u>	<u>1.5 in/h</u>
APR	3	2	2	2	
MAY	12	8	5		
JUN	15	12	7	5	1
JUL	18	13	10	3	2
AUG	21	19	13	10	3
SEP	11	8	3	2	2

TABLE 22. Annual probabilities for rainfall of specified intensity over the Lake Erie region. [Inferred from Hershfield (1961)]

Intense 1-hour rainfall (\geq 1 in.)					
APR	MAY	JUN	JUL	AUG	SEP
5%	10%	22%	26%	22%	13%

Intense 6-hour rainfall (\geq 1.5 in.)					
APR	MAY	JUN	JUL	AUG	SEP
5%	10%	17%	20%	22%	17%

Intense 24-hour rainfall (\geq 2.1 in.)					
APR	MAY	JUN	JUL	AUG	SEP
7%	12%	16%	16%	17%	15%

TABLE 23. Tornado frequency by month and year for Ohio.
[Ref. (25)]

<u>YEAR</u>	<u>JAN</u>	<u>FEB</u>	<u>MAR</u>	<u>APR</u>	<u>MAY</u>	<u>JUN</u>	<u>JUL</u>	<u>AUG</u>	<u>SEP</u>	<u>OCT</u>	<u>NOV</u>	<u>DEC</u>	<u>TOTALS</u>
1950	1	0	0	0	0	0	2	0	0	0	0	0	3
1951	0	0	0	1	0	2	0	0	0	0	0	0	3
1952	0	0	0	1	1	0	0	0	0	0	0	0	2
1953	0	0	0	1	0	2	0	0	0	0	0	0	3
1954	0	0	1	1	5	2	0	2	1	1	0	0	13
1955	0	0	5	0	0	0	0	0	0	0	1	0	6
1956	0	3	1	1	0	1	2	1	0	0	0	0	9
1957	0	0	0	0	1	2	1	0	0	0	0	0	4
1958	0	0	0	1	1	3	6	1	0	0	0	0	12
1959	0	1	0	2	4	0	1	1	0	1	0	0	10
1960	0	0	0	1	2	3	0	1	0	0	0	0	7
1961	0	1	1	4	4	2	4	3	0	0	0	0	19
1962	0	0	0	0	2	0	1	1	0	0	0	0	4
1963	0	0	0	10	2	2	2	0	3	0	0	0	19
1964	0	0	1	0	1	4	0	3	0	0	0	0	9
1965	0	0	1	15	4	0	2	2	0	0	3	0	32
1966	0	0	0	0	0	0	1	0	1	1	0	0	3
1967	0	0	1	0	0	3	2	0	0	0	0	0	6
1968	0	0	0	6	3	4	0	1	0	0	0	0	14
1969	0	0	1	1	8	2	4	1	0	0	0	0	17
1970	0	0	0	4	6	3	2	1	7	0	0	0	23
1971	0	3	0	0	4	2	3	1	1	1	0	0	15
1972	0	0	0	0	0	1	3	6	0	0	0	0	10
1973	0	0	0	0	13	15	10	3	0	0	0	2	43
1974	1	0	0	15	5	0	0	2	0	0	0	0	23
1975	0	0	0	1	0	2	1	4	3	0	1	0	12
1953-75	1	8	12	63	65	53	45	34	16	4	10	2	313
MEAN	+	0.4	0.5	2.8	3.0	2.3	2.0	1.4	0.6	0.2	0.4	0.1	13.7
MEDIAN	0	0	0	1	2	2	2	1	0	0	0	0	12
MOST	1	3	5	15	13	15	10	6	7	1	8	2	43

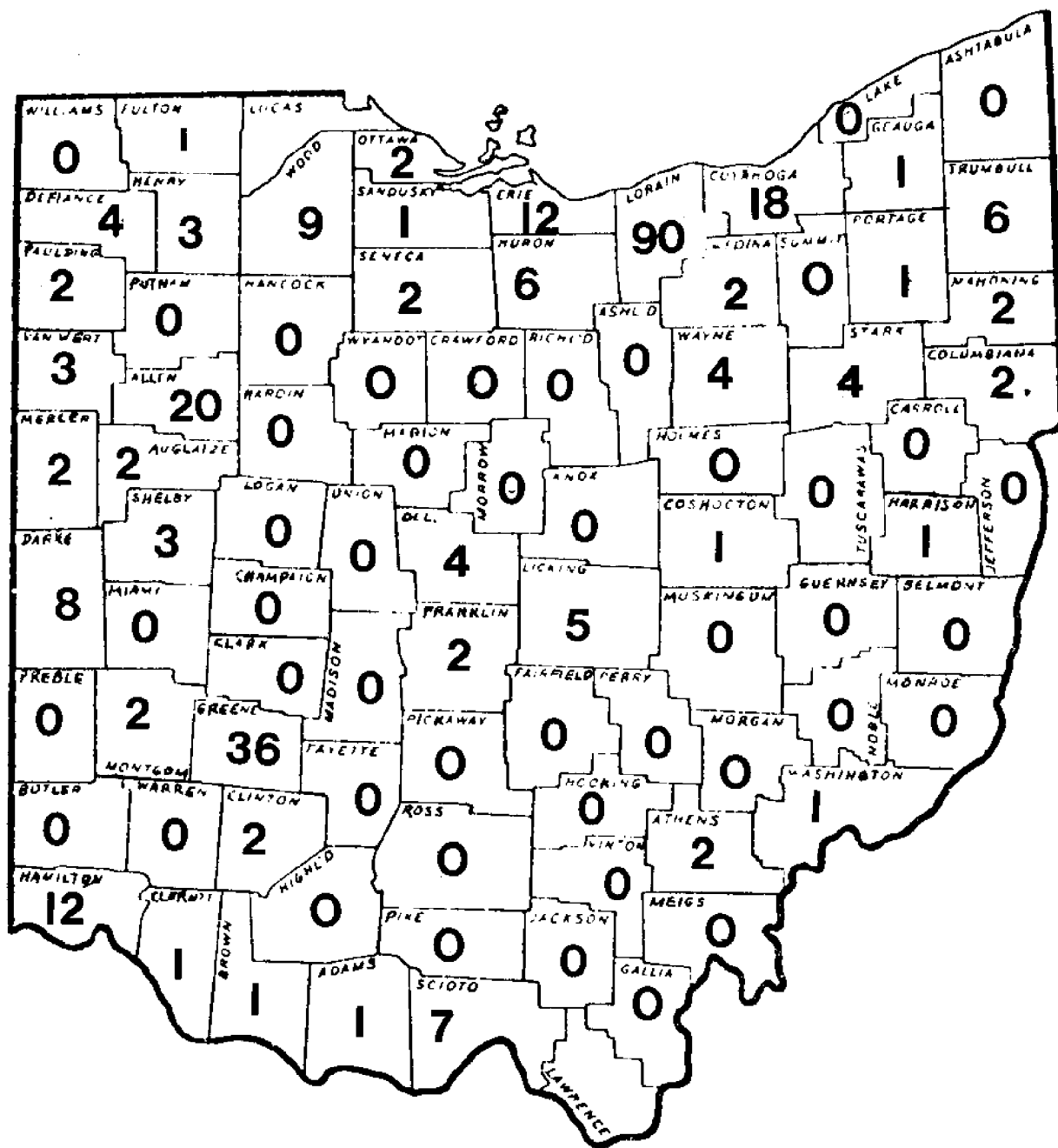


Figure 20. Number of deaths from tornadoes in each county of Ohio, 1916-1975. [Ref. (25)]

CHAPTER 5

SUMMARY

This report summarizes the climate of the Lake Erie region during the months from April through September and, as such, will provide the basis for future meteorological research in the lake basin. The major physical factors influencing the climate are identified in Chapter 2 and include the surface roughness, moisture supply, and temperatures of the lake versus the land surface. Chapter 3 examines some meteorological effects of Lake Erie: the lake and land breezes, the influence exerted by the lake on geostrophic flow, and its effect on pressure patterns, including the lake mesohigh and the climatological high. Lastly, Chapter 4 discusses the specific climatic variables and any possible lake influence upon them. They include temperature, humidity, solar radiation, wind, evaporation, precipitation, and severe weather. Results were obtained from a variety of sources and are frequently presented in both graphical and tabular form.

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