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# A Numerical Scheme Using a Transmutation Approach for Underwater Sound Propagation 

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## 1. Introduction

In [8] we have described a procedure for solving the submersible identification problem in three steps. That is, for the first approximation we shall concentrate on an ocean which may be taken to be a slab of thickness $\mathrm{h}, R_{h}^{3}$, and has a constant index of refraction. This is reasonable as the most difficult aspects of our problem come from the way the ocean surface interferes with the propagation of sound off an arbitrary submersible.[6], [8] Because of the difficulties introduced by the ocean boundaries, analytical computation of the far field must be based on some sort of approximation scheme such as using a parabolic approximation, which tends to destroy vertical resolution; or using a trucated form of the modal expansion for the Green's function ( propagator):

$$
\begin{equation*}
G(r, \theta, z ; \rho, \phi, \zeta)=\frac{i \pi}{2} \sum_{n=0}^{\infty} \frac{\phi_{n}(z) \phi_{n}(\zeta)}{\left\|\phi_{n}\right\|^{2}} H_{0}\left(k a_{n}\left|r e^{i \theta}-\rho e^{i \phi}\right|\right) . \tag{1.1}
\end{equation*}
$$

Here the $a_{n}=\left[1-(n+1 / 2)^{2}(\pi / k h)^{2}\right]^{-1 / 2}$ are the modal eigenvalues and $\phi_{n}(z)=\sin \left(k\left(1-a_{n}^{2}\right)^{1 / 2} z\right)[1]$. By using the propagator $G(r, z ; \rho, \zeta)(1.1)$,

[^0]the Green's integral representation, and the asymptotic behavior of the Hankel functions, Gilbert and $\mathrm{Xu}[5]$, [6] have shown that the acoustic pressure has the asymptotic expansion
\[

$$
\begin{align*}
p(r, z, \theta) & =\frac{1}{r^{1 / 2}} \sum_{n=0}^{\infty} e^{i k a_{n} r} \sum_{m=0}^{\infty} \frac{f_{n m}(z, \theta)}{r^{m}} \\
& \simeq \frac{1}{r^{1 / 2}} \sum_{m=0}^{\infty} \frac{F_{m}(r, z, \theta)}{r^{m}}, \tag{1.2}
\end{align*}
$$
\]

where

$$
\begin{equation*}
F_{m}(r, z, \theta):=\sum_{n=0}^{N} e^{i k k_{n} \tau} F_{n m}(z, \theta) . \tag{1.3}
\end{equation*}
$$

The modes $\phi_{n}(z)$ for $n>N$ do not propagate, but rather decay exponentially and hence are not included in the sum (1.3). We refer to the term

$$
\begin{equation*}
F_{0}(r, z, \theta):=\sum_{n=0}^{N} e^{i k a_{n} r} F_{n 0}(z, \theta) \tag{1.4}
\end{equation*}
$$

as the finite-ocean far field pattern, which we also write in an array form as

$$
F(z, \theta):=\left[f_{00}(z, \theta), f_{10}(z, \theta), \ldots, f_{N 0}(z, \theta)\right] .
$$

Gilbert and $\mathrm{Xu}[5],[6]$ show that it has a modal representation of form

$$
\begin{equation*}
\nu(\mathbf{x}, z):=\sum_{n=0}^{N} \frac{\phi_{n}(\zeta)}{\left\|\phi_{n}\right\|^{2} \sqrt{a_{n}}} \int_{\partial Z_{1}} g(z, \phi) \phi_{n}(z) e^{i k_{n}\langle x, \xi\rangle} d \sigma_{\xi}, \tag{1.5}
\end{equation*}
$$

where $g(z, \phi)$ is called a propagating Herglotz kernel.
Having obtained an approximate far field pattern for various wave numbers $k$ and various $z$ dependencies in the incident "plane waves", in step two we may solve the inverse problem. Recall that the "plane waves" have modal components in the $z$ direction. If we are not considering axially
symmetric solutions then for each modal component and for each $k$ we use $2 n+1$ incoming waves with directions in the range coordinates

$$
\begin{equation*}
\alpha_{2}^{j}=[\cos (2 \pi j /(2 n+1)), \sin (2 \pi j /(2 n+1))], j=0,1,2, \ldots, 2 n \tag{1.6}
\end{equation*}
$$

Now let $F_{n m}^{j} 0 \leq n \leq N,-n \leq m \leq n$ be the coefficients of the spherical harmonic approximation of far field pattern $F^{j}$ generated by the plane wave with the direction $\alpha_{j}$. By expanding the propagating Herglotz kernels, and the parametric representation of the submersible's surface $\rho=f(\theta, \phi)$ in terms of surface harmonics we are led to consider a minimization problem of the form

$$
\begin{gather*}
\mu(F)=\min _{\left(k^{2},,, p\right)}\left\{\sum_{j=1}^{2 n}\left|\int_{\theta Z_{1}} F\left(z, \phi ; k, \alpha_{2}^{(j)}\right) g(z, \phi) d z d \phi\right|^{2}\right. \\
\left.\quad+\int_{0}^{2 \pi} \int_{0}^{\pi}\left|T^{-1} V(f(\phi, \theta), \phi, \theta)\right|^{2} \sin ^{2} \theta d \theta d \phi\right\} \tag{1.7}
\end{gather*}
$$

where

$$
V(y):=\sqrt{\frac{2}{i k \pi}} \int_{\partial Z_{1}} g(z, \phi) G(\mathbf{x}, z, \xi, \zeta) d \sigma_{\xi}
$$

is the propagating (entire) Herglotz function, and $T^{-1}$ is the inverse transformation which relates the coefficients of the starting field in spherical coordinates to those in cylindrical coordinates. Other minimization problems might be considered instead, for example see $\mathrm{X}_{\mathbf{u}}$ [10].

In step three, we consider the case with an index of refraction which is depth dependent. We must make certain alternations in STEP ONE and STEP TWO. In this step we need to find a complete system of the solutions to replace the complete family (1.9) by another one, which must be a solution of the depth-dependent Helmholtz equation

$$
\begin{equation*}
\Delta u+k^{2} n^{2}(z) u=0 \tag{1.8}
\end{equation*}
$$

Such a family may be generated by means of the transmutation

$$
\Xi \Omega:=\Omega(r, z, \phi)+\int_{z=h}^{z} K(z, s) \Omega(r, s, \phi) d s
$$

where the kernel $K(z, s)$ satisfies the Gelfand-Levitan equation

$$
\begin{equation*}
\frac{\partial^{2} K}{\partial z^{2}}-\frac{\partial^{2} K}{\partial s^{2}}+k^{2}\left[n^{2}(z)-1\right] K=0 \tag{1.9}
\end{equation*}
$$

and the characteristic conditions [4]

$$
\begin{gather*}
2 \frac{\partial}{\partial z} K(z, z)+k^{2}\left[n^{2}(z)-1\right]=0  \tag{1.10}\\
2 \frac{\partial}{\partial z} K(z,-z+2 h)+k^{2}\left[n^{2}(z)-1\right]=0 \tag{1.11}
\end{gather*}
$$

that characterize a hard ocean boundary at $z=h$. The functions $\Omega(r, z)$ are solutions of (1.8) with $n(z)=1$ and boundary conditions. In order to obtain the propagating far-field patterns, we represent the propagating solution by

$$
\begin{equation*}
\sum_{n=0}^{N_{2}} \sum_{m=-M}^{M} \beta_{m n} \phi_{n}(z) H_{m}^{(1)}\left(k a_{n} r\right) e^{i m \theta} \tag{1.12}
\end{equation*}
$$

where the $\phi_{n}(z)$ are the modal solutions (eigenfunctions of the separated z-equation) for the variable index $n(z):[4],[6]$

$$
\begin{equation*}
q_{z z}+k^{2}\left(n^{2}(z)-a^{2}\right) q=0 \tag{1.13}
\end{equation*}
$$

and the boundary conditions

$$
\begin{equation*}
q_{z}(k a, h)=0 \tag{1.14}
\end{equation*}
$$

and

$$
\begin{equation*}
q(k a, 0)=0 \tag{1.15}
\end{equation*}
$$

It is clear that if we want to study the wave propagation and its farfield patterns, we must know the normal modes in the stratified ocean. Therefore, as an important component of step three, we need to construct $\phi_{n}(z), n=0,1,2, \ldots$ numerically for $n(z) \neq 1$.

## 2 Numerical Transmutation Method

As discussed in [4], a solution of (1.13) $\sim(1.15)$ is given by

$$
\begin{equation*}
q_{1}(k a, z)=p_{1}(k a, z)+\int_{h}^{z} K(z, s) p_{1}(k a, s) d s \tag{2.1}
\end{equation*}
$$

where $K(z, s)$ is a solution of $(1.9) \sim(1.11)$ and $p(k a, z)$ satisfies

$$
\begin{equation*}
p_{z z}+k^{2}\left(1-a^{2}\right) p=0, \tag{2.2}
\end{equation*}
$$

and the boundary condition

$$
\begin{equation*}
p_{z}(k a, h)=0 . \tag{2.3}
\end{equation*}
$$

Notice that $q_{1}(k a, z)$ will satisfy the boundary condition (1.14) at $\mathrm{z}=\mathrm{h}$ regardless of the value of $a$ because $p_{1}(k a, z)$ has the same property and our transmutation preserves boundary conditions at $z=h$. The boundary condition (1.15) at $\mathrm{z}=0$ is then used to determine the values of $a_{n}$ that are roots of

$$
\begin{equation*}
p_{1}(k a, 0)+\int_{h}^{0} K(0, s) p_{1}(k a, s) d s=0 . \tag{2.4}
\end{equation*}
$$

Since $n(z)$ may be $\geq 1$ or $\leq 1$ in general, we may have that $a_{n}^{2} \geq 1$ or $a_{n}^{2} \leq 1$. In fact, there is a documented example [13] where

$$
n^{2}(z)=1-0.1 \cos \left(\frac{2 \pi z}{h}\right)
$$

where $h=2000, k=8 \pi / 2000$, for which we have

$$
\begin{aligned}
& \left(k a_{0}\right)^{2}=1.6555156 \times 10^{-4} \\
& \left(k a_{1}\right)^{2}=1.5101910 \times 10^{-4}
\end{aligned}
$$

that is,

$$
\begin{aligned}
& a_{0}^{2}=1.0483675>1 \\
& a_{1}^{2}=0.9563396<1
\end{aligned}
$$

Therefore, $p_{1}(k a, z)$ may generally be given by either

$$
\begin{equation*}
p_{1}(k a, z):=\cos \left[k\left(1-a^{2}\right)^{1 / 2}(h-z)\right], \text { if } a<1, \tag{2.5}
\end{equation*}
$$

or

$$
\begin{equation*}
p_{1}(k a, z):=\cosh \left[k\left(a^{2}-1\right)^{1 / 2}(h-z)\right], \text { if } a>1 . \tag{2.6}
\end{equation*}
$$

For $a^{2}<1$, let $\lambda=k\left(1-a^{2}\right)^{1 / 2} h$, (2.4) becomes

$$
\begin{equation*}
\cos (\lambda)+\int_{h}^{0} K(0, s) \cos \left[\lambda\left(1-\frac{s}{h}\right)\right] d s=0 \tag{2.7}
\end{equation*}
$$

For $a^{2}>1$, let $\lambda=k\left(a^{2}-1\right)^{1 / 2} h$, (2.4) becomes

$$
\begin{equation*}
\cosh (\lambda)+\int_{h}^{0} K(0, s) \cosh \left[\lambda\left(1-\frac{s}{h}\right)\right] d s=0 . \tag{2.8}
\end{equation*}
$$

Obviously, in order to find all propagating eigenvalues we have to find all $\lambda \geq 0$ such that either (2.7) or (2.8) is satisfied. To avoid this inconvenience, we can rewrite the equation (1.13) as

$$
\begin{equation*}
q_{x x}+k^{2}\left[\left(n^{2}(z)-n_{0}^{2}\right)-\left(a^{2}-n_{0}^{2}\right)\right] q=0 \tag{2.9}
\end{equation*}
$$

where $n_{0}^{2}>0$ can be chosen so that $a^{2}<1+n_{0}^{2}$.

Let $\hat{n}^{2}(z)=n^{2}-n_{0}^{2}, \hat{a}^{2}=a^{2}-n_{0}^{2}$, the solution of (2.9) can be written as

$$
\begin{equation*}
q(k \hat{a}, z)=p(k \hat{a}, z)+\int_{h}^{z} \hat{K}(z, s) p(k \hat{a}, s) d s \tag{2.10}
\end{equation*}
$$

where $\hat{K}(z, s)$ satisfies

$$
\begin{equation*}
\frac{\partial^{2} \hat{K}}{\partial z^{2}}-\frac{\partial^{2} \hat{K}}{\partial s^{2}}+k^{2}\left[n^{2}(z)-n_{0}^{2}-1\right] \hat{K}=0 \tag{2.11}
\end{equation*}
$$

and the characteristic conditions [4]

$$
\begin{gather*}
2 \frac{\partial}{\partial z} \hat{K}(z, z)+k^{2}\left[n^{2}(z)-n_{0}^{2}-1\right]=0  \tag{2.12}\\
2 \frac{\partial}{\partial z} \hat{K}(z,-z+2 h)+k^{2}\left[n^{2}(z)-n_{0}^{2}-1\right]=0 \tag{2.13}
\end{gather*}
$$

The eigenvalues are the roots of equation (2.7) with $\lambda=k\left(1-\hat{a}^{2}\right)^{1 / 2} h$. We can normalize the problem by letting $\hat{z}=z / h, \hat{s}=s / h, \xi=(\hat{z}+\hat{s}-$ 2) $/ 2$, and $\eta=(\hat{z}-\hat{s}) / 2, k=k_{0} h$; then (2.11) $\sim(2.13)$ can be reduced to

$$
\begin{gather*}
\frac{\partial^{2} M}{\partial \xi \partial \eta}+k^{2}\left[n^{2}(1-\xi-\eta)-n_{0}^{2}-1\right] M=0,0 \leq \xi+\eta \leq 1  \tag{2.14}\\
M(\xi, 0)=\frac{1}{2} k^{2} \int_{0}^{\ell}\left[n^{2}(1-t)-n_{0}^{2}-1\right] d t, 0 \leq \xi \leq 1  \tag{2.15}\\
M(0, \eta)=\frac{1}{2} k^{2} \int_{0}^{\eta}\left[n^{2}(1-t)-n_{0}^{2}-1\right] d t, 0 \leq \eta \leq 1 \tag{2.16}
\end{gather*}
$$

We use a finite difference method developed in [2] to compute $M(\xi, \eta)$ as well as $K(z, s)$. Following Aziz and Hubbard [2], we define for $0 \leq m+n \leq$ $L$, where $L$ is the number of interpolation points and $d z=1 / L$, we see that

$$
\begin{aligned}
& M\left(m d z-\frac{1}{2} d z, n d z\right)=\frac{1}{2}[M(m d z, n d z)+M(m d z-d z, n d z)] \\
& M\left(m d z, n d z-\frac{1}{2} d z\right)=\frac{1}{2}[M(m d z, n d z)+M(m d z, n d z-d z)]
\end{aligned}
$$

$$
\begin{gathered}
M\left(m d z-\frac{1}{2} d z, n d z-\frac{1}{2} d z\right)= \\
\frac{1}{2}\left[M\left(m d z-\frac{1}{2} d z, n d z\right)+M\left(m d z-\frac{1}{2} d z, n d z-d z\right)\right] \\
=\frac{1}{2}\left[M\left(m d z, n d z-\frac{1}{2} d z\right)+M\left(m d z-d z, n d z-\frac{1}{2} d z\right)\right] \\
=\frac{1}{4}[M(m d z, n d z)+M(m d z-d z, n d z) \\
+M(m d z, n d z-d z)+M(m d z-d z, n d z-d z)], \\
M_{\xi \eta}\left(m d z-\frac{1}{2} d z, n d z-\frac{1}{2} d z\right) \\
=(d z)^{-2}[M(m d z, n d z)-M(m d z-d z, n d z) \\
-M(m d z, n d z-d z)+M(m d z-d z, n d z-d z)],
\end{gathered}
$$

where

$$
c_{m n}=k^{2}\left[n^{2}(N d z-(m+n-1) d z)-n_{0}^{2}-1\right]
$$

Now (2.14) may be approximated by

$$
\begin{gathered}
(d z)^{-2}[M(m d z, n d z)-M(m d z-d z, n d z)-M(m d z, n d z-d z) \\
+M(m d z-d z, n d z-d z)]+c_{m n} \frac{1}{4}[M(m d z, n d z)+M(m d z-d z, n d z) \\
+M(m d z, n d z-d z)+M(m d z-d z, n d z-d z)]=0 .
\end{gathered}
$$

Denote $u_{m n}=M(m d z, n d z)$, then

$$
\begin{equation*}
u_{m n}=\frac{4-c_{m n}(d z)^{2}}{4+c_{m n}(d z)^{2}} u_{m(n-1)}+\frac{4-c_{m n}(d z)^{2}}{4+c_{m n}(d z)^{2}} u_{(m-1) n}-u_{(m-1)(n-1)} \tag{2.17}
\end{equation*}
$$

with

$$
\begin{equation*}
u_{m 0}=f\left(m d z-\frac{1}{2} d z\right) \tag{2.18}
\end{equation*}
$$

and

$$
\begin{equation*}
u_{0 n}=g\left(n d z-\frac{1}{2} d z\right) \tag{2.19}
\end{equation*}
$$

where

$$
\begin{aligned}
& f(\xi)=\frac{k^{2}}{2} \int_{0}^{\xi}\left[n^{2}(1-t)-n_{0}^{2}-1\right] d t, 0 \leq \xi \leq 1 \\
& g(\eta)=\frac{k^{2}}{2} \int_{0}^{\eta}\left[n^{2}(1-t)-n_{0}^{2}-1\right] d t, 0 \leq \eta \leq 1
\end{aligned}
$$

As showed in [2], the global error of this scheme is $O\left((1 / L)^{2}\right)$. A FORTRAN program using this scheme is in Appendix A.

We can also use finite difference approximation to the equation (2.1) as well as (2.12)(2.13). Let $d z=d s=1 / l$,

$$
\begin{gathered}
K_{z z}(i d z, i d s)=\frac{1}{d z^{2}}\left(K_{i+1, j}-2 K_{i, j}+K_{i-1, j}\right) \\
K_{s s}(i d z, i d s)=\frac{1}{d z^{2}}\left(K_{i, j+1}-2 K_{i, j}+K_{i, j+1}\right) \\
K(i d z, i d s)=\frac{1}{2}\left(K_{i, j+1}+K_{i, j-1}\right)
\end{gathered}
$$

then (2.11) follows
$K_{i+1, j}+K_{i-1, j}-K_{i, j+1}-K_{i, j-1}+(d z)^{2}\left[n^{2}(i)-n 0-1\right]\left(K_{i, j+1}+K_{i, j-1}\right)=0$,

$$
\begin{equation*}
0 \leq i+j \leq 2 L, \quad j \geq i \geq 0 \tag{2.20}
\end{equation*}
$$

and

$$
\begin{gathered}
K_{i, i}=-\frac{k^{2}(d z)}{2}\left[\frac{\hat{n}^{2}(i)+\hat{n}^{2}(0)-2}{2}+\sum_{j=1}^{i-1}\left(\hat{n}^{2}(j)-1\right)\right] \\
K_{i, 2 L-i}=-\frac{k^{2}(d z)}{2}\left[\frac{\hat{n}^{2}(i)+\hat{n}^{2}(0)-2}{2}+\sum_{j=1}^{i-1}\left(\hat{n}^{2}(j)-1\right)\right], \\
\text { fori }=1,2, \ldots, 2 L .
\end{gathered}
$$

A PC-Matlab program using this scheme is shown in Appendix B.

## 3 Examples

In this section we present two examples and compare them with typical normal mode computation results.

Example 1: The results presented in this example are based on an idealized ocean model with a symmetric sound changel. The index of refraction is

$$
n^{2}(z)=1-0.1 \cos \left(\frac{2 \pi z}{h}\right)
$$

where the depth of ocean $h=2000$, and $k=8 \pi / 2000$.
Figure 1: The transmutation kernel $K(z, s)$ for example 1.
Figure 2: The eigenvalues shown bracketed by two os, which are the zeros of

$$
\cos (\lambda)+\int_{h}^{0} K(0, s) \cos \left[\lambda\left(1-\frac{s}{h}\right)\right] d s=0
$$

The dotted curve is $\cos (\lambda)$ and the solid curve is the integral term.
Table 1: Unperturbed eigenvalues $\left(\times 10^{-3}\right)$ for example 1. The index of refraction $n^{2}(z)=1$.

TABLE 1: (x10e-3)

| Mode | DSF <br> analytic | TRSM <br> fortran | TRSMf <br> error | TRSM <br> MC-Matlab | TRSM <br> error |
| :---: | :---: | :---: | :---: | :---: | :---: |
| $\mathbf{1}$ | 0.157297 | 0.157297 | 0.000000 | 0.157297 | 0.000000 |
| $\mathbf{2}$ | 0.152362 | 0.152362 | 0.000000 | 0.152362 | 0.0000000 |
| $\mathbf{3}$ | 0.142492 | 0.142492 | 0.000000 | 0.142492 | 0.000000 |
| $\mathbf{4}$ | 0.127688 | 0.127688 | 0.000000 | 0.127688 | 0.000000 |
| $\mathbf{5}$ | 0.107948 | 0.107948 | 0.000000 | 0.107948 | 0.000000 |
| 6 | 0.083275 | 0.083275 | 0.000000 | 0.083275 | 0.000000 |
| 7 | 0.053665 | 0.053665 | 0.000000 | 0.053665 | 0.000000 |
| $\mathbf{8}$ | 0.019122 | 0.019122 | 0.000000 | 0.019122 | 0.000000 |

Table 2: Eigenvalues $\left(\times 10^{-3}\right)$ for index of refractionn ${ }^{2}(z)=1-0.1 \cos (2 \pi z / h)$.
TABLE 2: (x10e-3)

| Mode | $\begin{gathered} \text { DSF } \\ \text { analytie } \end{gathered}$ | TRSM fortran | TRSMf <br> error | $\begin{gathered} \text { TRSM } \\ \text { MC-Matlab } \end{gathered}$ | TRSMm error |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 1 | 0.165552 | 0.165543 | 0.000009 | 0.165553 | 0.000001 |
| 2 | 0.151019 | 0.151031 | 0.000012 | 0.151062 | 0.000043 |
| 3 | 0.139981 | 0.139976 | 0.000005 | 0.140015 | 0.000034 |
| 4 | 0.126508 | 0.126466 | 0.000042 | 0.126436 | 0.000072 |
| 5 | 0.107284 | 0.107263 | 0.000021 | 0.107259 | 0.000025 |
| 6 | 0.082841 | 0.082835 | 0.000006 | 0.082869 | 0.000028 |
| 7 | 0.053359 | 0.053364 | 0.000005 | 0.053436 | 0.000077 |
| 8 | 0.018893 | 0.018910 | 0.000017 | 0.019030 | 0.000137 |

Figure 3 - 10: Propagating modes: solid curves computed by Fortran program and +++ curves by PC-MATLAB.

Figure 11: Transmision loss versus range r:

$$
T L=10 \log _{10}\left[\|p(z, r)\|^{2}\right]
$$

where

$$
p(r, z)=\frac{i}{4} \sum_{n=0}^{\infty} \phi_{n}(z) \phi_{n}\left(z_{0}\right) H_{0}\left(k a_{n} r\right)
$$

and source depth $z_{0}=500$, receiver depth $z=1000,200 \leq r \leq 20000$. The dotted curve is for a homogeneous ocean $n(z)=1$, and the solid curve is for the stratified ocean $n(z)=1-0.1 \cos (2 \pi z / h)$. We will consider an inverse problem to reconstruct the index of refraction by using this data.

Figure 12: Transmision loss versus range $r$ : where source depth $z_{0}=$ 500, receiver depth $z=1000,500 \leq r \leq 50000$.

Figure 13: Transmision loss versus range r : where source depth $z_{0}=$ 500 , receiver depth $z=1000, r_{0}=10,1000 \leq r \leq 100000$.

Figure 14: Finite difference scheme.

Example 2 This example is an idealized ocean model for an ocean sound channel with its axis at one-fifth the depth of the ocean. The index of refraction is

$$
n^{2}(z)=1+0.1\left[\sin \left(\frac{5 \pi}{4 h} z\right)-\frac{2 \sqrt{2}}{5 \pi}\right]
$$

where the depth of ocean $h=5000$, and $k=8 \pi / 5000$.
Figure 15: The transmutation kernel $K(z, s)$ for example 2.
Figure 16: The eigenvalues shown bracketed by two os, which are the zeros of

$$
\cos (\lambda)+\int_{h}^{0} K(0, s) \cos \left[\lambda\left(1-\frac{s}{h}\right)\right] d s=0 .
$$

Table 3: Unperturbed eigenvalues ( $\times 10^{-4}$ ) for example 1. The index of refraction $n^{2}(z)=1$.

TABLE 3: (x10e-4)

| Mode | DSF <br> analytic | TRSM <br> fortran | TRSMf <br> error | TRSM <br> MC-Matlab | TRSMm <br> error |
| :---: | :---: | :---: | :---: | :---: | :---: |
| $\mathbf{2}$ | 0.251675 | 0.251675 | 0.000000 | 0.251675 | 0.000000 |
| 2 | 0.243779 | 0.243779 | 0.000000 | 0.243779 | 0.000000 |
| 3 | 0.227988 | 0.227988 | 0.000000 | 0.227988 | 0.000000 |
| 4 | 0.204301 | 0.204301 | 0.000000 | 0.204301 | 0.000000 |
| 5 | 0.272718 | 0.172718 | 0.000000 | 0.172718 | 0.0000000 |
| 6 | 0.333240 | 0.133240 | 0.000000 | 0.133340 | 0.0000000 |
| 7 | 0.085866 | 0.085656 | 0.000000 | 0.085656 | 0.0000000 |
| 8 | 0.030596 | 0.030596 | 0.000000 | 0.030596 | 0.000000 |

Table 4: Eigenvalues $\left(\times 10^{-4}\right)$ for index of refractionn $n^{2}(z)=1+0.1 \sin (5 \pi z / 4 h+$ $\pi / 4)-0.2 \sqrt{2} / 5 \pi$.

TABLE 4: (x10e-4)

| ModeDSF <br> analytic | TRSM <br> fortran | TRSMf <br> error | TRSM <br> MC-Matiab | TRSMm <br> error |  |
| :--- | :---: | :---: | :---: | :---: | :---: |
| $\mathbf{2}$ | 0.260430 | 0.260608 | 0.000178 | 0.260870 | 0.000440 |
| 2 | 0.237756 | 0.238109 | 0.000353 | 0.238656 | 0.0009900 |
| 3 | 0.221614 | 0.221957 | 0.000343 | 0.222472 | 0.000858 |
| 4 | 0.202254 | 0.202403 | 0.000149 | 0.202586 | 0.000332 |
| $\mathbf{5}$ | 0.171646 | 0.171741 | 0.000095 | 0.171872 | 0.000226 |
| $\mathbf{7}$ | 0.132547 | 0.132620 | 0.000073 | 0.132760 | 0.000213 |
| 8 | 0.085379 | 0.085445 | 0.000066 | 0.085621 | 0.000242 |
|  | 0.030234 | 0.030303 | 0.000069 | 0.030540 | 0.000306 |

Figure 17-24: Propagating modes: solid curves computed by Fortran program and +++ curves by PC-MATLAB.

Figure 25: Transmision loss versus range $r$ : where source depth $z_{0}=$ 500 , receiver depth $\mathrm{z}=1000, \mathrm{r}_{0}=10,200 \leq r \leq 20000$.

Figure 26: Transmision loss versus range r : where source depth $z_{0}=$ 500 , receiver depth $z=1000, r_{0}=10,500 \leq r \leq 50000$.

Figure 27: Transmision loss versus range $r$ : where source depth $z_{0}=$ 500 , receiver depth $\mathrm{z}=1000, \mathrm{r}_{0}=10,1000 \leq \mathrm{r} \leq 100000$.

Figure 28: Finite difference scheme.

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Figure 1


Figure 2


Figure 3




Figure 6





Figure 10


Figure 11




Figure 15






Figure 20



Figure 22







Figure 29


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