QC 807.5 .06 W6

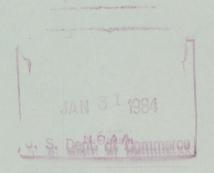
no.113

NOAA Technical Memorandum ERL WPL-113



BETA: A PROGRAM FOR CALCULATING AND ARCHIVING BACKSCATTERING PROFILES TAKEN WITH THE NOAA COHERENT LIDAR SYSTEM

Wave Propagation Laboratory Boulder, Colorado November 1983



NOAA Technical Memorandum ERL WPL-113

BETA: A PROGRAM FOR CALCULATING AND ARCHIVING BACKSCATTERING PROFILES TAKEN WITH THE NOAA COHERENT LIDAR SYSTEM

M. J. Post J. F. Morrow D. B. Jensen

Wave Propagation Laboratory Boulder, Colorado November 1983



UNITED STATES
DEPARTMENT OF COMMERCE
Malcolm Baldrige,
Secretary

NATIONAL OCEANIC AND ATMOSPHERIC ADMINISTRATION

John V. Byrne, Administrator **Environmental Research Laboratories** 

Vernon E. Derr, Director

1/1908

## NOTICE

Mention of a commercial company or product does not constitute an endorsement by NOAA Environmental Research Laboratories. Use for publicity or advertising purposes of information from this publication concerning proprietary products or the tests of such products is not authorized.

# CONTENTS

| ABSTRACT             | Page<br>1 |
|----------------------|-----------|
| INTRODUCTION         | 1         |
| TECHNICAL BACKGROUND | 2         |
| RUNNING BETA         | 7         |
| APPENDIX A           | 25        |
| APPENDIX B           | 34        |
| REFERENCES           | 44        |

BETA: A PROGRAM FOR CALCULATING AND ARCHIVING BACKSCATTERING
PROFILES TAKEN WITH THE NOAA COHERENT LIDAR SYSTEM

M. J. Post, J. F. Morrow, and D. B. Jensen

#### ABSTRACT

We document the interactive program BETA to provide to users information about its purpose and its inputs and outputs. More emphasis is placed on running the program and understanding the raw and processed data than is placed on programing particulars.

#### INTRODUCTION

NOAA has developed a pulsed, coherent, infrared Doppler laser-radar (lidar) system (Post et al., 1982) for measuring atmospheric winds to ranges of 10-20 km. Pulses of 3-µs duration are transmitted at 10.0-Hz pulse repetition frequency (PRF), and a very small amount of energy from the outgoing pulse is scattered by aerosol particles back towards the lidar. By measuring the frequency shift (Doppler shift) of this backscattered signal, we can deduce the velocity component of the aerosols which is in line with (radial to) the laser beam. Since the aerosols are so small that they move "exactly" with the winds, we say we are measuring the radial wind component. By scanning the lidar beam in a conical pattern, it is possible to deduce the average speed and direction of the winds in the pattern.

The whole process depends on having enough aerosols to backscatter energy from the transmitted pulse, so there is a genuine need to document the

backscattering properties of the aerosols and their statistical fluctuations. The backscattering coefficient  $\beta$  is a measure of the aerosol's backscattering ability. By definition it is the effective total cross-sectional area in  $m^2$  of all the particles in one cubic meter of air, divided by  $4\pi$ . The division by  $4\pi$  assumes the energy is scattered isotropically into a  $4\pi$  steradian solid angle (sphere), so  $\beta$  has units of  $m^2/m^3/sr$  or  $m^{-1}$  sr<sup>-1</sup>.

The most efficient way to measure  $\beta$  vs. altitude, or the  $\beta$ -profile, is to point the lidar beam vertically (angle from straight up = zenith angle = 0°), but any zenith angle is permissible.

We describe here the BETA program, which uses the recorded lidar signals and calculates the  $\beta$ -profile that had to exist in order to give rise to the signals. The program itself is listed in Appendix B.

### TECHNICAL BACKGROUND

The lidar's data-taking system digitizes an analog voltage originating on the heterodyne signal detector and records this signal on 9-track digital magnetic tape. The digitizer has 8-bit resolution, a minimum sampling interval of 100 ns (maximum rate of 10 MHz) and a sampling capacity of 2048 consecutive points. The data are complex; that is, for each data point, two numbers are generated (R and Q) which are 90° out of phase (hence they are in "quadrature" and represent the real and imaginary parts of a complex number). Signal power is computed by squaring both R and Q and taking the sum of the squares.

The gain of the receiving chain is adjusted so that when no signal is present, the mean background noise level uses only 1-3 bits of the dynamic

range of the digitizer, leaving 7-5 bits for the (additive) signal level. The maximum unsaturated signal plus noise is  $128^2$ , or 16,384.

We know the relationship between the volume backscatter coefficient  $\boldsymbol{\beta}$  and the system SNR. It is

SNR = 
$$\frac{\eta \ A_1 A_2 A_3 A_4 J \ a^2 \ \Pi \ c\tau\beta}{8 \ hv \ B \ R^2}$$

where

η = Detector quantum efficiency

 $A_1(R) = Gaseous absorption factor =$ 

exp 
$$-2 \int_{0}^{R} \sum_{i} \alpha_{i}(R^{i}) dR^{i}$$
 (i = species)

 $A_2(R) = Apodization factor$ 

 $A_3$  = Shot noise factor

 $A_4$  = Round trip optical loss factor

a = Beam radius at primary (m)

c = Speed of light  $(m s^{-1})$ 

τ = Pulse duration (s)

 $\beta$  = Backscatter coefficient (m<sup>-1</sup> sr<sup>-1</sup>)

h = Planck's constant (J s)

v = Laser frequency (s<sup>-1</sup>)

B = Bandwidth  $(s^{-1})$ 

R = Range (m)

J = Energy per pulse (J)

The A factors are discussed in more detail by Lawrence (1983). The combination of  $\eta$ ,  $A_3$ , and  $A_4$  is variable ETA in the BETA program, and is set to a value of 0.48, assuming the following baseline calibration using a small spinning disc:

CW SNR = 43 dB

Range = 159 m

CW power = 0.5 W

Shot noise = 3.7 dB.

Many of the terms in the SNR equation vary, and must be set properly when BETA is run. Among these, A<sub>1</sub> accounts for round-trip atmospheric absorption of the CO, radiation by H2O, CO2, and other gases. One of five AFCRL absorption models (McClatchey et al., 1972) may be chosen (1 = tropical, 2 = midlatitude summer, 3 = midlatitude winter, 4 = subarctic summer, 5 = subarctic winter), but only models 2 or 3 are appropriate for U.S. data. We investigated the use of radiosonde measurements to better account for H<sub>2</sub>O absorption effects, and found that the present models introduce no more than 2 dB error in  $\beta$  at extreme ranges (~20 km) and for extreme (unrealistic) mismatches between the models and the radiosonde-measured atmosphere. The mean errors between the midlatitude summer and winter models and the average radiosonde profiles for Denver increase monotonically from zero at 2 km altitude to 0.29 dB above 9 km altitude. A more complete description of continuum absorption by water vapor and its variability is given in Appendix A. A peculiarity we find is that the absorption coefficient is nearly constant with altitude above 10 km, despite decreasing density, owing to pressure and temperature effects on the CO, line shape and strength.

Another variable parameter is B, the system bandwidth. Noise power is proportional to B, but signal power is not a function of B if its frequency spectrum lies entirely within B. Typically, we set B to 1.5 MHz when the lidar is pointing vertically, since in this configuration we expect negli-

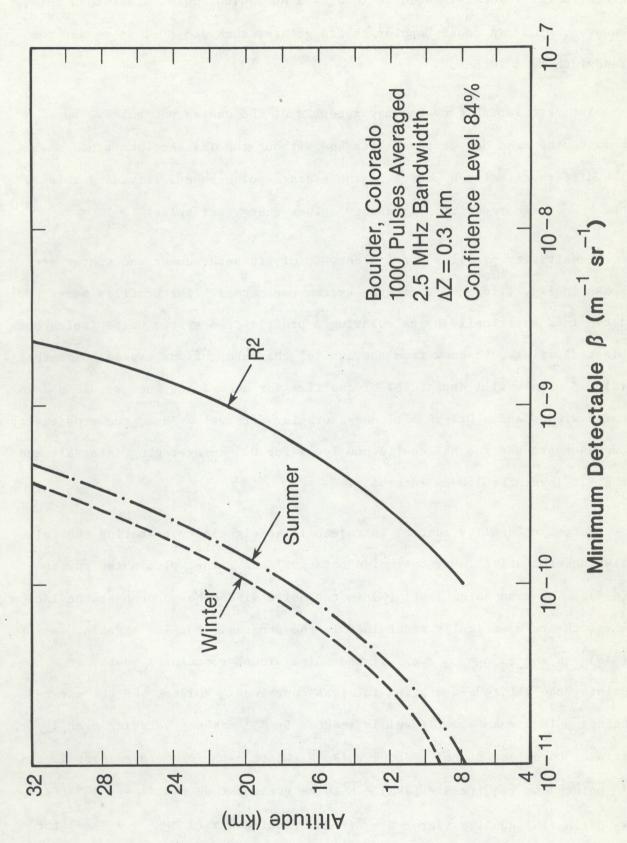
gible radial velocity. When we take data at zenith angles other than zero, however, wind can cause Doppler shifts greater than 1.5 MHz, so we set the bandwidth to  $10.0 \, \text{MHz}$ .

The last variable system parameter is J, the energy per pulse. We measure the mean output power with and without the TEA section turned on.

The difference between powers is the average pulse power. Dividing this by the PRF (pulse repetition frequency) gives energy per pulse.

Sensitivity profiles for Boulder, Colo., in both summer and winter are shown in Fig. 1 for the indicated system parameters. The profiles were determined empirically by calculating  $\beta$  profile from system noise (telescope-blocked) files. We know from the "white" character of our system noise that false  $\beta$  values will appear 16% of the time for such files when we use a 1 km noise window and a 0.1-km altitude-averaging interval. Thus, the sensitivity curves represent the 84% confidence level for 0.1-km averaging intervals and the 92% level for 1.0-km intervals.

We use a "quality factor" to reject false signals by comparing the relative uncertainties in system noise power (N) and signal plus noise powers (S+N). The mean noise level  $(\overline{N})$  in the noise window has a standard deviation  $\sigma_N$ ; both are numerically determined in the program. For low signals, (S+N)  $\approx \overline{N}$  and  $\sigma_{S+N} \simeq \sigma_N = \sigma$ . If the noise window contains  $\ell_W$  data points, the altitude-averaging interval contains  $\ell_B$  points, and the uncertainty  $\sigma$  in a mean power level is reduced by  $\ell^{-1/2}$  when averaging over  $\ell$  points, to detect signal  $\overline{S}$  above noise  $\overline{N}$  (where each is averaged over  $\ell_B$  and  $\ell_W$  points) we require  $\overline{S} = (S+N) - \overline{N}$  to be greater than  $\sigma(\ell_B^{-1/2} + \ell_W^{-1/2})$ . We define the quality factor  $0 = \overline{S}/\sigma$  and require  $0 > (\ell_B^{-1/2} + \ell_W^{-1/2})$  for the observed data before they are accepted for inclusion in a  $\beta$  profile.



noise and two different atmospheric absorption models. Range effect (R2) is shown for reference only. Figure 1. -- System sensitivity curves for NOAA's coherent lidar, determined empirically using system

# RUNNING BETA

To run the BETA program on the NOAA Eclipse, you must be in the LIDAR directory.

Type DIR DPØ:LIDAR.

If hard copies of the graphics output are desired, use the Tektronix terminal with attached hard copy unit, pre-warmed. Make sure the data tape is mounted and on the load point, and the tape drive is on line.

Type BETA to start the program.

Figures 2-10 show sequential outputs from the terminal screen and user responses to illustrate the discussion that follows.

When the screen clears and asks for the tape drive unit number,

Enter Ø or 10 to select unit the tape is mounted on.

The screen will then clear again and ask for calibration data.

Enter the most pertinent calibration data, which may be found on the calibration log sheets.

After you enter these data, the computer calculates and prints out the factor CAL, which is used to correct ETA. Beta values are multiplied by this factor near the end of the program to account for variations in calibration.

Enter the bandwidth in MHz from the tape log sheet next.

This is typically 10.0 for Doppler mode data or 1.5 for  $\beta$  mode data.

TAPE UNIT NUMBER 0
CW SNR (DB) ? 48.5
RANGE (M) ? 1.8
CW POWER (W) ? 1.2
SHOT NOISE (DB) ? 2.5
CALIBRATION FACTOR IS .707139

BANDWIDTH (IN MHZ) 1.5

ALTITUDE OF LIDAR (KM ASL) 1.684

INITIAL RANGE (KM) 1.891 END RANGE (KM) 30.450 WHICH ATMOSPHERIC ABSORPTION MODEL TO USE 1=TROPICAL 2=MIDLATITUDE SUMMER 3=MIDLATITUDE WINTER 3

FILE NUMBER ? 1 START RECORD NUMBER ? 1 NUMBER OF RECORDS TO PROCESS ? 200 ZENITH ANGLE OF SCANNER (DEGREES) ? 0
PER PULSE ENERGY (J) WAS ? .09

CONTINUE (1) OR GO BACK AND START OVER (8)? 1

Figure 2.--Sample terminal printouts and user responses for initialization of program.

Enter the lidar altitude above sea level (ASL) in kilometers.

This permits properly initializing the absorption models with respect to sea level.

Enter beginning and ending ranges.

These values will set the vertical scaling of the plots to follow. Typically, the minimum useful beginning range is 1.5 km because the tail of the optical pulse saturates the detector. The maximum ending range is determined by the length of the record and the digitization rate. It is about 31.5 km for 2048 point seconds and a 10.0 MHz rate. Normally no signal is found in the stratosphere, or after 10 to 15 km range, but volcanic loading of the stratosphere has produced usable signals to 29 km range, and hard targets (mountains, towers, etc.) are visible at even greater ranges. You may choose a narrower range increment positioned anywhere as well, for example, to examine a cloud layer. For archiving ß data, we want to average over 1-km altitude increments, centered on even kilometers above sea level. This requires setting the proper beginning range, which can be found with the aid of a sketch showing the altitude of the lidar ASL and its beam angle. The beginning altitude (and corresponding beginning range) will occur at a half integer value. Typical settings for the various field locations are written on a card near the terminal.

Enter the number of the atmospheric absorption model to be used.

Enter file and start record number to be processed. It is permissible to spin tape forward or backward.

Enter number of records to process in the file (this number must be less than or equal to the number of records in the file).

To compute altitudes ASL and to account for slant path absorption,

Enter the lidar beam angle from the zenith (vertical) in degrees, as indicated on the data sheet.

The computer will calculate the correct absorption profile and store it for use later, when  $\beta$ 's are computed.

Enter the per-pulse energy from the data sheet.

Check for errors.

If no errors have been made in all your entries,

Enter 1 and the tape will be properly positioned.

If you made an error,

Enter  $\emptyset$ , and make all entries again.

Next, set the scaling for the graphic output by answering the appropriate questions on the terminal (see Fig. 3). Full (HI) scale of the digitizer is 0-20,000, but most interesting data are viewed on the 0-400 scale (LO). Also available is logarithmic scaling, 0-40 (HI) or 0-12 (LO) dB. Once set, scaling cannot be changed during a run.

Set switch  $\emptyset$  on the front panel of the computer. Turning the switch up selects plot averages, while turning it down selects single plots

Switch Ø may be changed at any time. Normally, start with it up, but occasionally flip it down during a run to see if the data are good, to find out what record in the file is currently being processed, or simply to avert

SET SWITCH (0): UP FOR AVERAGE PLOTS ONLY. DOWN FOR SINGLE RECORD PLOTS. SWITCH (0) MAY BE CHANGED DURING PROCESSING. NOTE:

THE FOLLOWING SCALING OPTIONS CANNOT BE CHANGED.

LINEAR (1) OR DB (0) SCALING ? 1

HI (1) OR LO (8) SCALING ? 8

PLOTTING FORMAT IS SET FOR:

AUERAGE PLOTS ONLY X-AXIS 0-400 LINEAR IF NOT, ENTER 0. IF THIS FORMAT IS OK, ENTER 1. Figure 3.--Sample terminal printouts and user responses for scaling of graphical outputs.

boredom. It should be flipped back up immediately again, however, as processing slows to the graphical frame rate while it is down.

Enter 1 if you are satisfied with the scaling format.

Enter Ø if you wish to reselect the format.

Next, the file header is read by the computer and its contents are displayed (Fig. 4).

Check the variables to make sure you are working on the right file.

If satisfied,

Strike any key and processing will commence.

The tape moves in spurts—once or twice per second. Nothing will appear on the screen unless switch Ø is depressed for single records (see Fig. 5) or until all records are processed. Then a time history of the total backscattered power (one point for each record's average power) is displayed (see Fig. 6). Laser power fluctuations, "losing lock," intrusion of clouds, etc., are easily seen on this plot. If you decide some bad records were included, deduce which ones (time in seconds and record number are in one—to—one correspondence). Note the bad records, if any, on the log sheet.

Make a hard copy of this plot to document data quality.

Strike the page key to clear the page.

To eliminate sections of bad records,

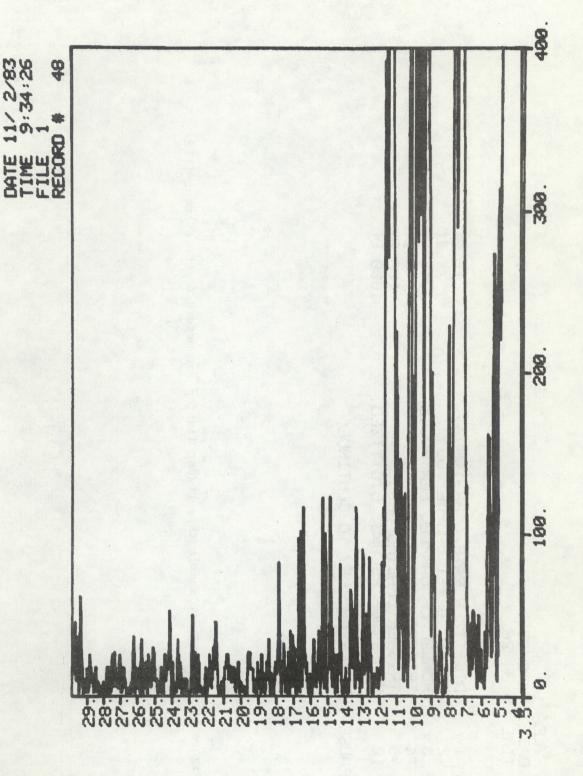
Enter Ø, which permits reprocessing the correct records.

FILE 1 2/8: DATE 11/2/8: TIME 9:34:20

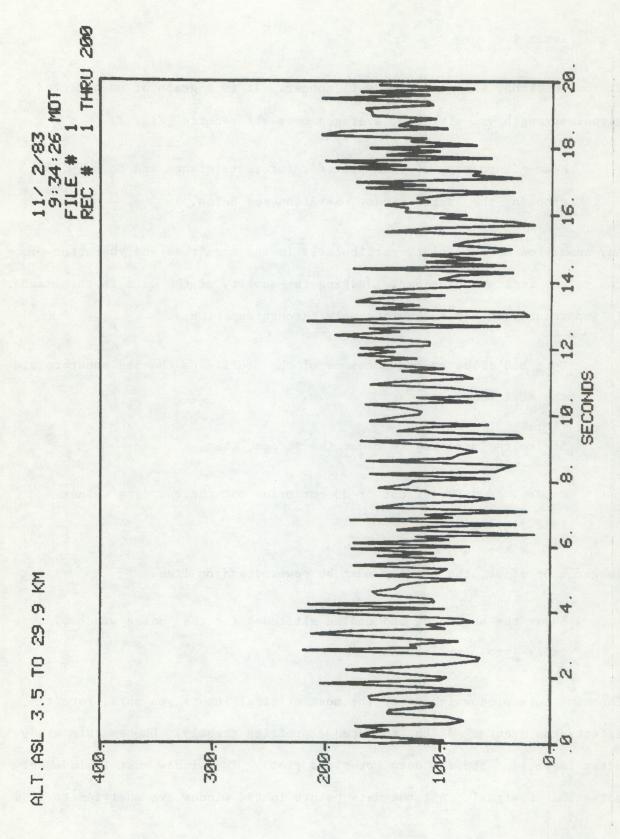
| 18.88                       | .15                      | 2048.00                     | 3.00         | 100.00               | 1999.99                      |
|-----------------------------|--------------------------|-----------------------------|--------------|----------------------|------------------------------|
| DIGITIZER RATE ( .08-10MHZ) | DELAY TO 1ST SAMPLE (US) | NUMBER OF SAMPLES PER PULSE | TRIGGER MODE | TRIGGER SPACING (MS) | NUMBER OF RECORDS DIGITIZED. |
| 1                           | N                        | כיו                         | 4            | S                    | 10                           |

PAUSE "STRIKE ANY KEY TO CONTINUE"

Figure 4.--Printout of sample tape header listing parameters of the current file.



 $\mathbb{R}^2+\mathbb{Q}^2$  on horizontal axis, altitude (km) on vertical axis. Figure 5.--Sample single pulse profile.



sents the average returned signal power for the pulse occurring at the indicated time. Display is useful for detecting signal variability throughout a file, whether due to atmospheric or system changes. Figure 6.--Sample plot of signal power vs. time (or power vs. pulse number in file). Each point repre-

To continue, (when all records appear satisfactory),

Enter 1.

If you continue, the next plot will appear. It is a graph of uncorrected signal strength vs. altitude, averaged over all records (Fig. 7).

Make a hard copy of this plot for future reference and to help in choosing the "noise window" as discussed below.

Any anomalies on this plot, particularly in the magnitude and character of the noise, indicate problems. Judging the quality of the data in this manner is important, and can be learned only through experience.

Note bad files or bad portions of the profile on the log sheet to aid the next user.

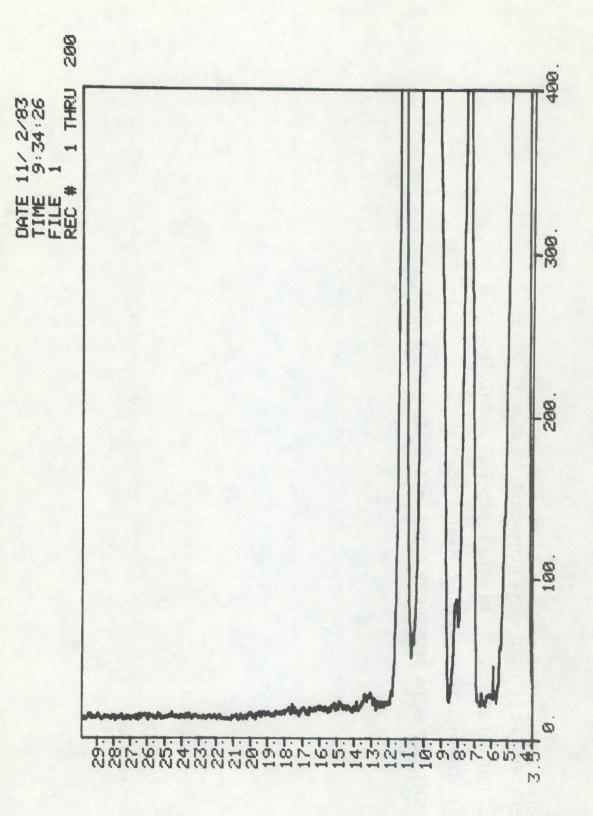
Strike the page key to clear the screen, then

Enter 1 or Ø (print out or do not print out the raw data values (see Fig. 8)

To get a printout, the printer must be powered and on line.

Enter the beginning and ending altitudes for the "noise window," typically 1 to 2 km apart.

Choosing this window is one of the most critical inputs you make, for it affects the accuracy of the computed  $\beta$  profiles greatly. Choose this window after inspecting the raw data (previous plot). The window must include only noise and no signal. All the data points in the window are analyzed to find



Incoherent Figure 7.--Sample plot of the average of 1000 single pulse profiles (as in Fig. 5). averaging reduces noise fluctuations and permits signals to appear.

ALTITUDE RESOLUTION IN KM (1.005 FOR ARCHIVING) ?.3 PRINT RAW DATA ? YES = 1 NO = 0 0 BEG, END ALTITUDES (KM) FOR NOISE WINDOW ? 27,29 0 8 = Q Q-FACTOR MUST BE GREATER THAN .35 YES = 1PRINT OUT BETA PROFILE ?

Figure 8. -- Sample terminal output and user responses for controlling processed data outputs, both tabular and graphical.

the mean noise level and the standard deviation of the noise. (A historical note: When the lidar system was first built, the mean noise versus time (or range) was not constant, and a stored, average noise profile had to be scaled and adjusted to match the noise level in the window. It was then subtracted from all the raw data to recover the signal power.) After the preamplifier was upgraded, the noise became nearly constant in time (range), permitting the mean noise level in the window to be subtracted from all data in the profile, a simpler operation. If you can find no window containing only noise, correct  $\beta$ 's cannot be computed.

Enter the altitude increment for averaging, or the height resolution you desire on the  $\beta$  plot.

Choose 0.3 km for matching to the lidar pulse length (and hence to the inherent system resolution). However, when archiving you must average over 1.005 km increments. This makes the reported  $\beta$  values occur at even kilometers altitude above sea level (after roundoff errors), assuming you entered the correct beginning range.

From the mean noise in the window, its standard deviation, and the altitude increment over which you are averaging the signal, the computer calculates and prints out the "quality factor" discussed earlier. The signal must exceed this criterion in order for us to be more than 84% confident that we are indeed sensing signal power and not simply statistical fluctuations in the noise. If you desire the  $\beta$  profile to be printed out at this point,

Enter 1, provided the printer is on line and has power.

Enter Ø to proceed without printout.

Next the  $\beta$  profile is plotted (Fig. 9). An experienced user will see anomalies immediately, e.g., spikes of increasing magnitude with height, which indicate bad data or an improper choice of the noise window. The signal must exceed the quality factor in two consecutive altitude increments in order for  $\beta$  to be accepted for either. That is, no isolated  $\beta$ 's will be accepted; this helps eliminate occasional noise spikes. If the  $\beta$  for an altitude increment is rejected either for this reason or for not passing the quality factor test, it is set to  $1.0 \times 10^{-15}$  m<sup>-1</sup> sr<sup>-1</sup>. If you are not happy with the  $\beta$  plot, you may try to improve it by choosing another noise window and/or altitude increment. To do this,

Strike the page key.

Choose new noise windows and/or altitude increments until you are satisfied that  $\beta$ 's have been calculated correctly.

Figure 10 shows the same  $\beta$  profile as Fig. 9, but averaged over 1.0 km in preparation for archiving. When viewing the  $\beta$ -profiles, bear in mind the system sensitivity curves of Fig. 1.  $\beta$ 's falling just below the curve or intermittant  $\beta$  points generally following it are suspect. Again, experience and comparison with the raw data profile are invaluable in judging the quality of the retrieved  $\beta$  profiles.

Once you are satisfied with the  $\beta$  profile,

Make a hard copy and store it with the other plots in the log sheet binder.

Strike the page key.

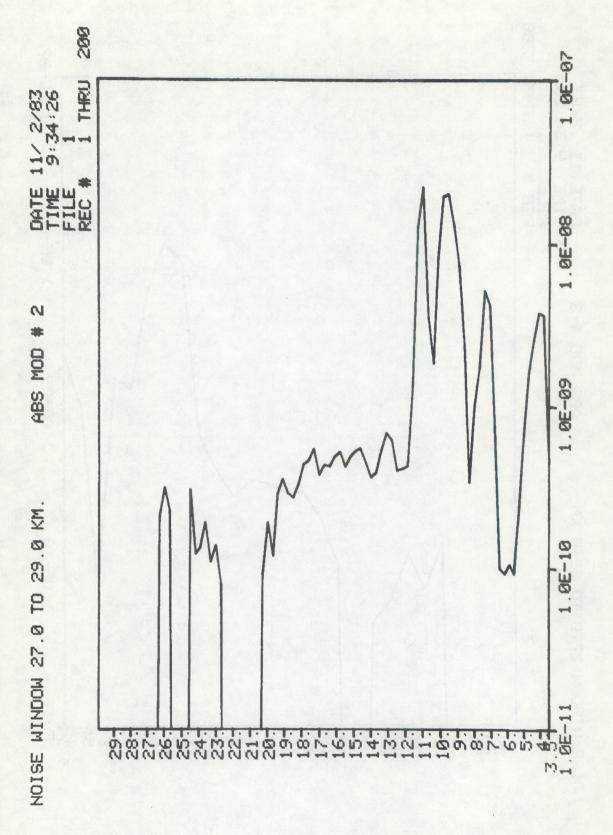


Figure 9.--Sample 8 profile output with 0.3-km height resolution.

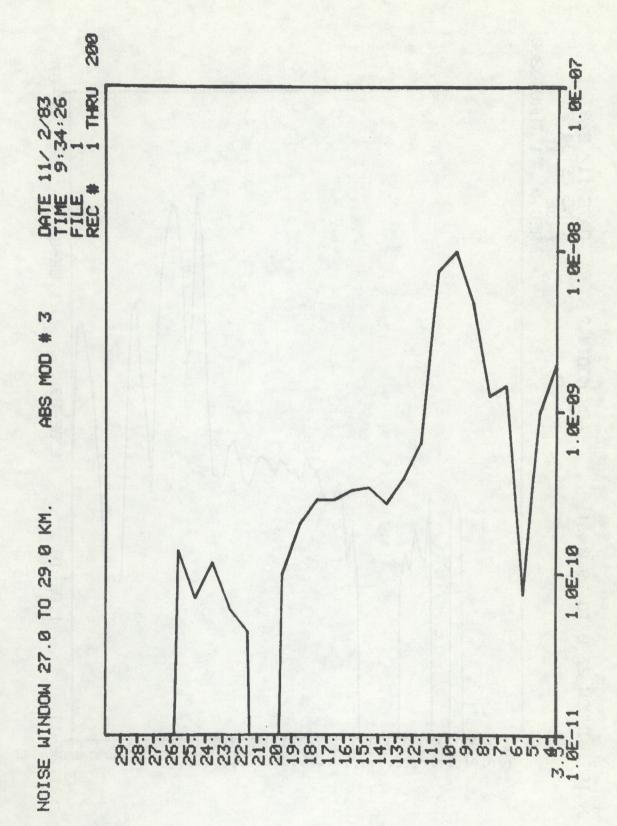


Figure 10.--Sample & profile output (as in Fig. 9), with 1.0-km height resolution.

Decide whether to archive the data, provided the  $\beta$  values are positioned on even kilometers above sea level and are spaced 1 km apart.

Enter 1 to archive, Ø to continue without archiving.

It is highly desirable to keep the archivings sequential. When archived, the raw SNR's, quality factors, and  $\beta$  values are recorded as a single record on a disc file labeled "ARCH", together with many "housekeeping" factors such as date, tape file number, noise window, and absorption model. If you did archive,

Record the archiving immediately on the master archive list to ensure up-to-date documentation of this important file.

The screen will display the pertinent parameters for the master list.

Another program, CPYARCH, copies the highest quality and statistically independent records from the ARCH file onto file SUPARCH for later processing.

After recording,

Strike the page key.

Enter 1 to process other sections of the tape (essentially starting the program over again).

Enter Ø to stop the program.

By stopping, you initiate an automatic tape rewind and return the terminal to the RDOS command level. In this case,

Take the tape drive unit off line (press the "on line" button)

Off-load the tape by pressing the "rewind" button.

Dismount and store the tape.

Finally, return the log sheet binder to its proper shelf space.

#### APPENDIX A

## · EFFECTS OF VARIABLE ATMOSPHERE WATER VAPOR ON β-PROFILES

In Boulder, Colo., we use either the midlatitude summer (2) or midlatitude winter (3) model of AFCRL (McClatchey et al., 1972) to correct for the effects of atmospheric absorption when computing  $\beta$ -profiles. The major gaseous constituent of the atmosphere most likely to depart from these models on a given day is water vapor. This Appendix quantifies the statistical variability between our models and the real atmospheric water vapor absorption, (as computed from measurements of temperature, pressure, and dew point by radiosondes released from Denver, Colo.). While making these calculations, we noted that at  $\lambda = 10.6~\mu m$ , water vapor continuum absorption was more than 100 times greater than water vapor line absorption, so we ignored the latter in studying these effects.

If the water vapor profiles in the AFCRL model atmospheres are integrated, they amount to the various precipitable water vapor (W) amounts listed in Table AI. For Boulder, Colo., data taken at an elevation of 1.67 km, we eliminate the lower 1.67 km of water vapor from these models and are left with 1.30 cm for the midlatitude summer model and 0.42 for the midlatitude winter model.

TABLE AI--Precipitable water vapor in AFCRL models

| (0 - | 10 | km | al | ti | t | ude) | ) |
|------|----|----|----|----|---|------|---|
|------|----|----|----|----|---|------|---|

Model Precipitable Water (cm)

Tropical 4.15

Midlatitude summer 3.00

Midlatitude winter 0.86

Subarctic summer 2.10

Subarctic winter 0.42

TABLE AII--Precipitable water by month in Denver, Colorado

| ( | 1. | 6 | - | 7 | .5 | km | al | ti | tude | ) |
|---|----|---|---|---|----|----|----|----|------|---|
|   |    |   |   |   |    |    |    |    |      |   |

|       | W                  | $\sigma_{ m w}$ |
|-------|--------------------|-----------------|
| Month | Precip. Water (cm) | Std. Dev. (cm)  |
| 1     | 0.502              | •220            |
| 2     | 0.476              | .183            |
| 3     | 0.527              | .195            |
| 4     | 0.696              | .237            |
| 5     | 1.040              | .344            |
| 6     | 1.444              | .384            |
| 7     | 1.897              | •485            |
| 8     | 1.813              | .498            |
| 9     | 1.302              | •422            |
| 10    | 0.886              | •301            |
| 11    | 0.649              | •223            |
| 12    | 0.499              | •205            |
|       |                    |                 |

The profiles of water vapor in the models lead to the two-way absorption curves (continuum only) shown in Figs. Al and A2.

A National Weather Service report prepared by Lott (1976) lists the monthly mean precipitable water vapor from the surface in Denver (1.611 km) to 400 mb (~7.5 km), along with its standard deviation. Essentially all water vapor is contained in this region. This listing is duplicated in Table AII. The summer average for June, July, and August is 1.72 cm, and the winter average for December, January and February is 0.49 cm; both are close to the model amounts.

From the radiosonde profiles of temperature, pressure, and dew point depression, we use Zuev (1972) to compute the water vapor partial pressure E at a given altitude Z:

$$E = E_o 10^{b\tau(c + \tau)}$$

where

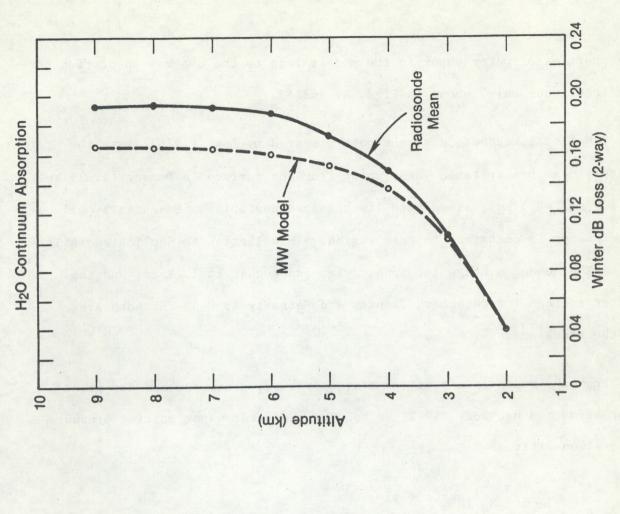
$$E_0 = 6.11 \text{ mb}$$

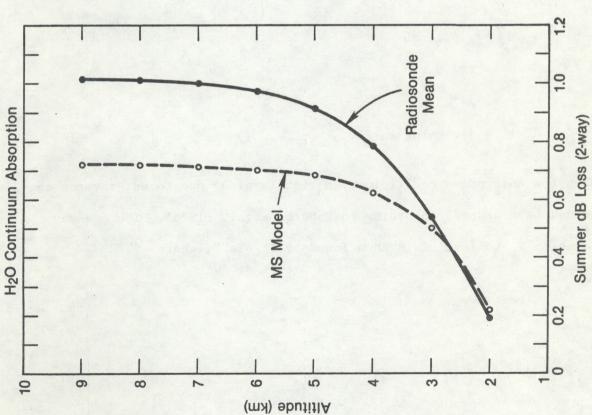
$$b = 7.5$$

$$c = 237.3$$
°K

$$\tau = \text{dew point in } ^{\circ}\text{C.}$$

Then the absorption coefficient,  $\alpha_n$  (at level n) due to water vapor continuum can be computed, according to Roberts et al. (1976). The one-way transmission  $T_i$  to level i is then found by  $T_i = e^{-x}$ , where





Figures Al and A2.--Two-way water vapor continuum absorption loss vs. altitude for the AFCRL summer (A1) and winter (A2) models for a vertically pointing lidar of  $10.6~\mu m$  wavelength at an altitude of 1.67~km.

$$x = \sum_{n=1}^{\infty} \alpha_n \Delta Z_n \sec \theta$$

and

 $\theta$  = zenith angle.

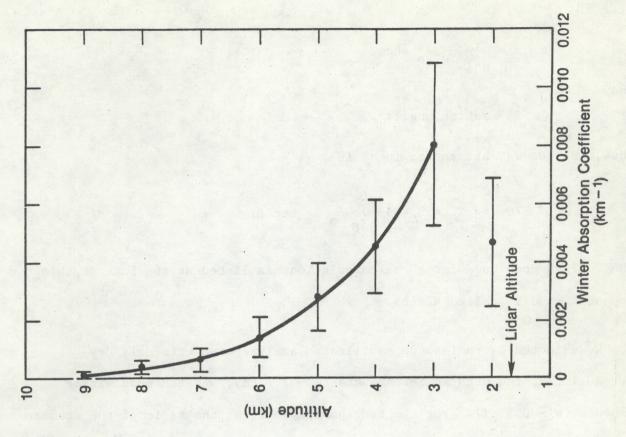
Thus, the two-way absorption in dB is

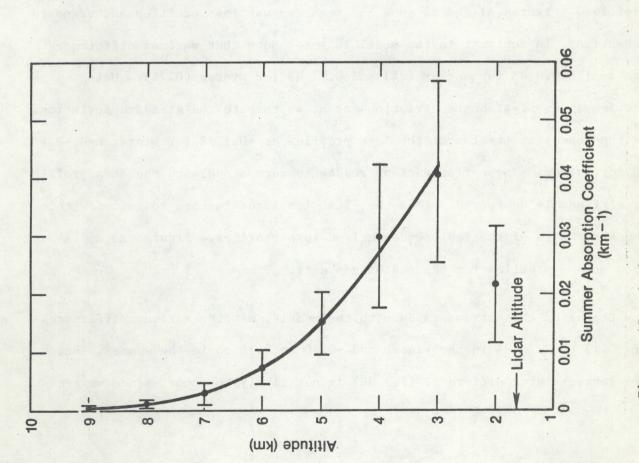
10 log 
$$T_i^2 = 8.686 \sum_{n=1}^{i} \alpha_n Z_n \sec \theta$$
.

The FORTRAN code used for these calculations is listed at the back of this appendix, and is called H2OABS.

We selected 24 radiosonde profiles to analyze--two extremely dry (W ~ 0.1 cm), ten mean summer amounts (W ~ 1.7 cm), and ten mean winter amounts (W ~ 0.5 cm). For the ten summer profiles, the ratio of the standard deviation of W to mean W was 0.013; the same ratio was 0.04 for the winter profiles. Figures Al and A2 show the mean two-way loss profiles and standard deviations, in contrast to the model losses. Note that maximum differences are small--0.3 dB for summer (7%) and 0.03 dB for winter (0.7%). Using Lott's values of standard deviations of W, we compute the standard deviations of the losses at the top of the loss profiles as ~0.1 dB for winter and ~0.4 dB for summer. These fluctuations reduce to zero along with the loss profile as altitude is decreased. Thus the AFCRL loss profiles are contained within the error bars of the radiosonde-derived loss profiles. Figures A3 and A4 show mean  $\alpha$  profiles for the radiosonde sets.

Comparing the extreme cases with the models, we find maximum differences of -0.15 to 2.70 dB in the winter and -0.70 to 2.15 dB in the summer. But the larger winter difference (2.7 dB) is not likely to occur, since we are





Figures A3 and A4.--Calculated absorption coefficients vs. altitude for ten summer (A3) and ten winter (A4) sample radiosonde profiles, each profiles having nearly the Denver mean precipitable water vapor for that season. Error bars show standard deviations of the coefficient at each altitude.

contrasting a summer radiosonde-derived loss profile with a winter model, and such a high W case will not occur in the winter. Reasonable bounds to put on all differences are thus -1.0 to +2.0 dB for the Denver-Boulder area and zenith pointing.

Errors introduced at other locations by using AFGL models instead of radiosonde profiles are likely to be greater, since the Denver area is relatively dry. Errors at any location will increase dramatically with increasing zenith angles, but such errors can be derived for any location and geometry using the techniques described here.

Listing of H2OABS follows on next page.

```
SUBROUTINE H2DABS (ABSRP)
  THIS SUBROUTINE FINDS THE H20 ABSORPTION FROM RAWIN DATA.
  THE DATA IS TYPED INTO A FILE OF ARBITRARY NAME CONTAINED
   IN ARRAY NENAM. THE FORMAT
  OF THE DATA IS PRESSURE, ALTITUDE IN METERS, TEMPERATURE IN
  DEGREES C. DEW POINT DEPRESSION.
   THE PROGRAM THEN FINDS THE PARTIAL PRESSURE OF WATER VAPOR
   (E) AND THE ABSORPTION DUE TO IT (SIGMA).
C
  SINCE THE RAWIN DATA IS AT
   RANDOM HEIGHTS AND WE WANT SPECIFIC INTERVALS, THE DO 20 LOOP
   INTERPOLATES SIGMA LINEARLY. WE CHECKED WHETHER IT WOULD BE
   BETTER TO INTERPOLATE BEFORE DOING CALCULATIONS
C
   AND THE RESULTS AGREED TO WITHIN ONE PERCENT.
C
C
       THE FORM OF THE EQUATIONS TO CALCULATE THE EXTINCTION
C
        COEFFICIENT SIGMA IS FROM ROBERTS ET AL: I.R. CONTINUUM
        ABSORPTION ETC., APPL. OPT. 15, P2085. EQUATIONS TO COMPUTE
C
        WATER VAPOR DENSITY FROM RAWINSONDE OUTPUTS IS FROM ZUEV:
        PROP. VIS. &I.R. WAVES IN ATMOS., P5-6. DIMENSIONS OF VAR-
C
       TABLES FOLLOWS.
C
C
       SIGMA IN KM-1 E IN MB
                                        TINK
                                                         GNU IN CM-1
                                     EO IN MB
C
                       TAU IN C
                                                        CINK
       P IN MB
       B IS NO DIM. CO IN MOL-1 M2 ATM-1
C
                                                        CON1 IN MOL G-1
C
        CON2 IN M KM-1 SIGMA IN KM-1 A IN G M-3 ATM
C
        DIMENSION P(40), T(40), DPDEP(40), ALT(40), ABSRP(20), SIGMA(40)
        DIMENSION C02(20,5), NFNAM(12)
        TYPE " DISC FILE NAME ?"
        READ(11, 3000) NFNAM(1)
3000
        FORMAT(S12)
        OPEN 1, NFNAM, ATT ="I", ERR=100
        ACCEPT "NUMBER OF LINES IN FILE ?", N1
        DO 4 I=1, N1
        READ FREE(1) P(1), ALT(I), T(I), DPDEP(I)
        ALT(I)=ALT(I)/1000.
        T(I)=T(I)+273.16
        CONTINUE
        CALL CLOSE (1, JER)
        TYPE "WHICH ATMMOSPHERIC ABSORPTION MODEL TO USE ?"
        TYPE "1=TROP 2=MS, 3=MW, 4=SS, 5=SW "
        ACCEPT MODE
        CON1=3. 327E22
        CON2=1. E3
        E0=6.11
        B=7. 5
        C = 237.3
        GAMMA=0. 002
        GNU=944. 194
        DO 10 NI=1, N1
        IF(DPDEP(I). LT. 100. )GO TO 8
        SIGMA(I)=0.0
        GO TO 10
8
        TAU=T(I)-DPDEP(I)-273.16
        E=E0*10. **(B*)AU/(C+TAU))
        CO=(1, 25E-26+2, 34E-23*EXP(-, 0083*GNU))*EXP(1800, *(1, /T(I)-1, /276, ))
        A=217.*E*(E+GAMMA*(P(I)-E))/(T(I)*1013.)
        SIGMA(I)=CON1*CON2*CO*A
10
        CONTINUE
        INTEGRATE ABSORPTION, INTERPOLATING BETWEEN RAWINSONDE
C
C
        LEVELS AND ACCOUNT FOR PARTIALS (SIG AND SIGLO) AT
        BEGINNING AND END OF EVEN KM INTERVALS.
```

```
C
        BE SURE RAWINSONDE DATA GOES AT LEAST ONE LEVEL ABOVE 10 KM.
C
        N=1
        NL=1
        ABSRP(1)=0.0
        SIGLO=0. 0
        ABTOT=0. 0
        WRITE(12,3001) NFNAM(1)
3001
        FORMAT(" OUTPUT FOR RAWINSONDE FILE ", S12, /)
        WRITE (12, 1001)
        DO 20 I=2, 10
        ABSRP(I)=SIGLO
25
        IF(ALT(N), GT. I) GO TO 27
        N=N+1
        GO TO 25
27
        NH=N-1
        DO 28 K=NL, NH
        ABSRP(I) = ABSRP(I) + (SIGMA(K) + SIGMA(K+1)) * (ALT(K+1) - ALT(K))/2.
28
        CONTINUE
        SIG=SIGMA(N)-(I-ALT(N))*(SIGMA(N)-SIGMA(N+1))/(ALT(N+1)-ALT(N))
        SIGLO=(SIGMA(N+1)+SIG)*(ALT(N+1)-I)/2.
        ABSRP(I) = ABSRP(I) + (I - ALT(N)) + (SIGMA(N) + SIG)/2.
        ABTOT=ABTOT+ABSRP(I)
        DBLOSS=8. 6859*ABTOT
        NL=N+1
1001
        FORMAT(" ALT.
                           ALPHA 2-WAY LOSS", /, " (KM) (KM-1)
     + (DB)",/)
        WRITE(12, 1000) I, ABSRP(I), DBLOSS
1000
        FORMAT(15, F10. 5, F7. 2)
20
        CONTINUE
        STOP
100
        TYPE "ERROR IN OPENING RAWINSONDE FILE"
        STOP
        END
```

## APPENDIX B

LISTING OF PROGRAM BETA

```
COMMON/DD/ARAY(2048)
 COMMON/SIGNL/ PWR(1000), IDAT(2048), AVGPWR(2048)
 DIMENSION PARM(25), IHEAD(200), XABSRP(30,3)
 DIMENSION BEAM(30), ABSRP(30, 3), PWRAV(1000), AVEP(100)
 EQUIVALENCE (PARM(1), IHEAD(51)), (AVEP(1), IDAT(1))
LOGICAL ISW
DATA BEAM/1.081, 1.366, .921, .688, .551, .450,
 . 405, . 356, . 331, . 313, . 300, . 295, . 280, . 275,
. 265, . 260, . 256, . 25, . 246, . 241, . 24, . 239, . 239, . 238,
. 237, . 237, . 236, . 236, . 235, . 235/
DATA XABSRP/ .3669, .2310, .1386, .0819, .0592, .0475,; TROP
.0382, .0305, .0248, .0200, .0156, .0123, .0076, .0072,
.0054, .0040, .0032, .0034, .0041, .0049, .0059, .0070,
.0079, .0086, .0093, .0121, .0121, .0121, .0121, .0121,
. 2459, . 1519, . 0979, . 0685, . 0526, . 0424,
                                               ; MIDLAT SUM
.0351, .0287, .0232, .0187, .0148, .0116, .0090, .0080,
.0080, .0080, .0080, .0080, .0081, .0084, .0087, .0091,
.0096, .0102, .0106, .0118, .0118, .0118, .0118, .0118,
. 0689, . 0568, . 0473, . 0385, . 0309, . 0251,
                                               ; MIDLAT WIN
.0204, .0164, .0132, .0117, .0104, .0090, .0088, .0086,
.0085, .0083, .0081, .0101, .0102, .0077, .0077, .0077,
```

THIS PROGRAM LOOKS AT RAW SIGNALS AND COMPUTES BACK-SCATTERING FOR DATA COLLECTED WITH THE NOAA COHERENT DOPPLER LIDAR. MAXIMUMS: 1000 RECORDS/FILE, 2048 POINTS/RECORD, 30 KM RANGE. PROGRAMMED BY M. J. POST AND RON RICHTER.

ARRAY BEAM CONTAINS BEAM APODIZATION FACTORS FROM TRL (F = 99 KM, E-2 RADIUS = 0.15 M, MATCHED TX AND LO. GOOD TO 11 KM. INCREMENTS ARE 1 KM IN LIDAR RANGE (NOT ASL). CHANGED VALUES FOR 2-30 KM RANGE 3/15/83.

ARRAY ABSRP CONTAINS P20 ABSORPTION COEFFICIENTS FROM MJP, 8/5/82. INCREMENTS ARE 1 KM ASL. BECOMES THE 2-WAY PATH-INTEGRATED ABSORPTION ARRAY LATER IN THE PROGRAM.

.0077, .0077, .0081, .0081, .0081, .0081, .0081, .0081/

CALL VMEM(II, IER)

IF(II. LT. 12) TYPE "NOT ENOUGH MEMORY, SEE D. DAVIS"

IF(II. LT. 12) STOP

CALL MAPDF(12, ARAY, 4, 2, IER)

CALL INITT(30)

CALL ANMODE

ICURF=0

ICURF=0

ICHAN=22

CAL=1. O

ACCEPT "TAPE UNIT NUMBER ", ITU

CALL MTINIT(ITU, ICHAN, INTER, IOPER)

SNR EQUATION CONSTANTS AND FACTORS FOLLOW.

CALL CALIB(CAL)
TYPE
ACCEPT "BANDWIDTH (IN MHZ) ", BW
BW=BW\*1. E6
HNU=1. 88E-20
D=0. 3
RLAM=1. 059E-5

CC

C

1

C ACCOUNT FOR SULPHUR DEPOLARIZATION. C ETA=0. 048 PI=3. 14159 C=2. 998E8 P1=8. \*HNU\*BW P2=((PI\*D\*\*2)/(4.0\*RLAM))\*\*2 IPLOT = 0 ACCEPT "ALTITUDE OF LIDAR (KM ASL) ", ALTLID ACCEPT "INITIAL RANGE (KM) ", RANGE1 ACCEPT "END RANGE (KM) ", RANGE2 TYPE "WHICH ATMOSPHERIC ABSORPTION MODEL TO USE ?" TYPE "1=TROPICAL" TYPE "2=MIDLATITUDE SUMMER" ACCEPT "3=MIDLATITUDE WINTER ", MOD C C POSITION TAPE AT CORRECT FILE AND RECORD TYPE ACCEPT "FILE NUMBER ? ", IFN ACCEPT "START RECORD NUMBER ? ", IREC ACCEPT "NUMBER OF RECORDS TO PROCESS ? ", NREC TYPE ACCEPT"ZENITH ANGLE OF SCANNER (DEGREES) ? ", THETA ACCEPT "PER PULSE ENERGY (J) WAS ? ", EJ ACCEPT "CONTINUE (1) OR GO BACK AND START OVER (0)? ", IANS CALL NEWPAG IF (IANS, EQ. O) GO TO 1 CALL TAPOS (ICHAN, IFN, IREC, ICURF, ICURR, IHEAD, IDAT) IF(IHEAD(7), NE. IFN) TYPE"FILE # READ DOES NOT MATCH FILE REQUESTED" C INITIALIZE THE ABSORPTION COEFFICIENT ARRAY FOR C THE GIVEN LIDAR ALTITUDE (ALTLID) ABOVE SEA LEVEL C C (ASL) AND SCANNER ZENITH ANGLE (THETA). PRF=1. E6/PARM(5) P3=PI\*ETA\*EJ\*C\*D\*\*2 ZZZ=THETA\*0. 017453293 7ZZ=COS(ZZZ) IALT=ALTLID IALT1=IALT+1 DIFALT=IALT1-ALTLID IF (IALT. LE. 0) GO TO 15 DO 10 I=1, IALT ABSRP(I, MOD)=0.0 10 IF (IALT1. LE. 0) GO TO 20 15 ABSRP(IALT1, MOD) = XABSRP(IALT1, MOD) \*DIFALT IALT2 = IALT1+1 20 DO 25 KL=IALT2, 30 ABSRP(KL, MOD) = XABSRP(KL, MOD)CONTINUE 25 CUMK=0. 0 C INTEGRATE THE ABSORPTION COEFFICIENTS FOR ROUND-TRIP C

ETA CHANGED FROM 0.024 TO 0.048 ON 2/3/83 BY MJP TO

LOSS TO ALL RANGES.

CC

```
DO 30 I=1,30
                 CUMK=CUMK+ABSRP(I, MOD)/ZZZ
                 ABSRP(I, MOD) = EXP(-2, *CUMK)
30
        CONTINUE
        BFLAG=0.
C
C
                 INITIALIZE POWER ARRAYS
C
        DO 35 I=1,2048
                 AVGPWR(1) = 0.0
35
        CONTINUE
        DO 40 I=1, 1000
                 PWR(I) = 0.0
                 PWRAV(1)=0.0
40
        CONTINUE
C
C
                 SET SCALING OF PLOTS.
C
45
        TYPE"SET SWITCH (0): UP FOR AVERAGE PLOTS ONLY. "
        TYPE"
                               DOWN FOR SINGLE RECORD PLOTS. "
        TYPE
        TYPE "NOTE: SWITCH (0) MAY BE CHANGED DURING PROCESSING."
        TYPE
        TYPE
        TYPE "THE FOLLOWING SCALING OPTIONS CANNOT BE CHANGED."
        TYPE
50
        ACCEPT" LINEAR (1) OR DB (0) SCALING ? ", LINDB
        IF (LINDB, EQ. O. OR, LINDB, EQ. 1) GO TO 55
        GO TO 50
55
        TYPE
        ACCEPT" HI (1) OR LO (0) SCALING ? ", LOHI
        IF (LOHI, EQ. O. DR. LOHI, EQ. 1) GO TO 60
        GO TO 55
60
        TYPE
        TYPE
        TYPE"PLOTTING FORMAT IS SET FOR: "
        TYPE
        IF(ISW(O)) TYPE"
                              AVERAGE PLOTS ONLY"
        IF(.NOT. ISW(0)) TYPE" SINGLE RECORDS, THEN AVERAGE PLOTS"
        IF(LOHI. EQ. 1. AND. LINDB. EQ. 1) TYPE" X-AXIS 0-20000 LINEAR"
        IF(LOHI EQ. O. AND. LINDB. EQ. 1) TYPE" X-AXIS Q-400 LT
        IF (LOHI, EQ. O. AND, LINDB, EQ. O) TYPE" X-AXIS 0-400 LINEAR"

TYPE

X-AXIS 0-12 DB"
        TYPE
        TYPE
        ACCEPT "IF THIS FORMAT IS OK, ENTER 1. IF NOT, ENTER O. ", LSW
        IF (LSW. EQ. Q) CALL NEWPAG
        IF (LSW. EQ. 0) GD TD 45
C
C
                 PRINT TAPE HEADER
C
        CALL NEWPAG
        IF(PARM(1), GT. 10. O. OR. PARM(1), LT. O. OB) ACCEPT "DIGITIZER RATE ?"
        , PARM(1)
        WRITE(10, 1010) IFN, (IHEAD(I), I=1, 6)
        WRITE(10, 1011) PARM(1)
        WRITE(10, 1012) PARM(2)
        WRITE(10, 1013) PARM(3)
        WRITE(10, 1014) PARM(4)
        WRITE(10, 1015) PARM(5)/1000.
        WRITE(10, 1016) PARM(10)
        PAUSE "STRIKE ANY KEY TO CONTINUE"
```

```
/, " DATE ", I2, "/", I2, "/", I2, /" TIME ", I2, ": ", I2, ": ",
       12.//)
        FORMAT(1X, " 1", 5X, "DIGITIZER RATE (. 08-10MHZ) ", T45, F8. 2)
1011
        FORMAT(1X, " 2", 5X, "DELAY TO 1ST SAMPLE (US) "T45, F8. 2)
1012
        FORMAT(1X, " 3", 5X, "NUMBER OF SAMPLES PER PULSE ", T45, F8. 2)
1013
        FORMAT(1X, " 4", 5X, "TRIGGER MODE ", T45, F8. 2)
1014
        FORMAT(1X, " 5", 5X, "TRIGGER SPACING (MS) ", T45, F8. 2)
1015
        FORMAT(1X, "10", 5X, "NUMBER OF RECORDS DIGITIZED. ", T45, F8. 2)
1016
                 CALCULATE START, STOP INDICES
C
                FROM TAPE HEADER INFORMATION
C
C
        DELAY = PARM(2)
        DELT = 1.0/PARM(1)
        DELR = . 149896*DELT
        DELZ = DELR*ZZZ
        RR1=(DELAY-2.0)*. 149896
        STTIM = RANGE1 /. 149896 - DELAY + 2.0
        ISTART = INT( STTIM/DELT + 0.5)
        SPTIM = RANGE2/. 149896 - DELAY + 2.0
        ISTOP = INT( SPTIM/DELT + 0.5)
        IF(ISTOP. GT. 2048) ISTOP = 2048
        IF(ISTOP-ISTART+1 . GT. 2048) ISTART = ISTOP-2048+1
        IF(ISTART. LT. 1) ISTART = 1
        NPTS = ISTOP - ISTART + 1
        IF(ISTART. EQ. 1) RANGE1=RR1
        IF(ISTOP. EQ. 2048) RANGE2 = RR1 + NPTS/PARM(1) * 0.149896
        R1 = RANGE1*ZZZ + ALTLID
        R2 = RANGE2*ZZZ + ALTLID
C
                LOOP THRU REQUIRED NUMBER OF RECORDS (NREC) TO
C
C
                READ IN ALL REQUIRED DATA.
C
        ICMD = 2048
        DO 110 K=1, NRE.C
                KREC=K+IREC-1
                 CALL MTDID (ICHAN, ICMD, IDAT, ISTAT, IWRD, IER)
                 IF (IER. NE. 1) QO TO 215
C
                 DECODE R AND Q FROM IDAT ARRAY (USING R2Q2 FUNCTION )
                 CALCULATE POWERS AND SUM. REMOVE SPIKES.
C
                KM1=K-1
                 TKM1=KM1/10.
                 JJ = 1
                 IF (K. GT. 4) GO TO 75
                 DO 70 I=ISTART, ISTOP
                         RG=R2G2(I)
                         PWR(K)=PWR(K)+R@
                         ARAY(JJ)=RG
                        AVGPWR (JJ) =AVGPWR (JJ)+RQ
                         JJ=JJ+1
                 CONTINUE
70
                 GO TO 85
                 DO 80 I=ISTART, ISTOP
75
                         RQ=R2Q2(I)
                         THRESH=AVGPWR (JJ)/TKM1
                         IF (RQ. GE. THRESH) RQ=AVQPWR (JJ) /KM1
                         PWR(K)=PWR(K)+RG
                         ARAY(JJ)=RQ
                         DS+(UJ) SWGDVA=(UJ) +RQ
                 JJ = JJ + 1
```

FORMAT(/, " FILE ", I2,

1010

| 80  | CONTINUE  |
|-----|---|
| C   |   |
| C   | SET SCALING AND PLOT SINGLE RECORDS IF REQUESTED  |
| 85  | IF(ISW(0)) GO TO 105<br>IPLOT = 1   |
|     | IF (LINDB. EQ. 1) QO TO 95 DO 90 I=1, NPTS  |
|     | PPWR=ARAY(I)  |
| C   | CALL VFETCH(PPWR, I)  |
| C   | IF(PPWR.LT.1.0) PPWR=1.0  |
| C   | IF(PPWR.EQ.1.)ARAY(I)=1. IF(PPWR.EQ.1.0) CALL VSTASH(PPWR,I)  |
|     | PPWR = 10.0 * ALDG10(PPWR)  |
|     | ARAY(I)=PPWR  |
| 90  | CONTINUE  |
|     | IF(LDHI. EQ. 1) YMAX = 40.0   |
|     | IF(LDHI.EQ.O) YMAX = 12.0   |
| 95  | IF(LOHI. EQ. 1) YMAX = 20000. 0   |
|     | IF(LOHI. EQ. 0) YMAX = 400.0  |
|     | CALL GRAFR (ARAY, NPTS, IFN, IHEAD, R1, R2, YMAX,   |
|     | + IPLOT, KREC, 1, DELZ, IREC, 0)  |
| 100 | GO TO 105 CALL GRAFR (ARAY, NPTS, IFN, IHEAD, R1, R2, YMAX,   |
| 100 | + IPLOT, KREC, 1, DELZ, IREC, 0)  |
| 105 | PWR(K) =PWR(K)/NPTS   |
| 110 | CONTINUE  |
|     | IF(LINDB. EQ. 1) GO TO 120  |
|     | DD 115 K=1, NREC<br>PWR(K)=10, 0*ALDG10(PWR(K))   |
| 115 | CONTINUE  |
| 120 | KREC=KREC-IREC+1  |
| C   |   |
| C   | CALCULATE AVERAGES AND PLOT   |
| C   | DO 125 I=1, NPTS  |
|     | AVGPWR(I)=AVGPWR(I)/NREC  |
| 125 | CONTINUE  |
|     | IF(LINDB.EQ. 1) GO TO 135 DO 130 I=1, NPTS  |
|     | AVPWDB = 10.0*ALDG10(AVGPWR(I))   |
|     | ARAY(I)=AVPWDB  |
| 130 | CONTINUE  |
| 135 | IF(LINDB. EQ. O. AND. LOHI. EQ. 1) YMAX = 40.0  |
|     | IF(LINDB. EQ. O. AND. LOHI. EQ. O) YMAX = 12. O<br>IF(LINDB. EQ. 1. AND. LOHI. EQ. 1) YMAX = 20000. O   |
|     | IF (LINDB, EQ. 1, AND, LOHI, EQ. 0) YMAX = 400.0  |
|     | XMAX = NREC/PRF   |
| C   |   |
| C   | PLOT TIME HISTORY OF INTEGRATED POWER PER RECORD.   |
| C   | CALL GRAFT (PWR, NREC, IFN, IHEAD, XMAX, YMAX, IREC, R1, R2)  |
|     | ICURR=IREC+NREC   |
|     | ACCEPT "CONTINUE (1) OR REPOSITION TAPE TO RE-RUN (0)? ", IANS  |
|     | CALL NEWPAG   |
|     | IF(IANS.EQ.O)GO TO 1  |
|     | IPLOT=O IF(LINDB, EQ. O)GO TO 140   |
| C   | are a rain a rain at a tour to the rain of the rain of the rain and the rain and the rain at the rain |
| C   | PLOT RAW POWER DATA AVERAGED OVER ALL PROCESSED RECORDS.  |
|     | CALL GRAFR (AVGPWR, NPTS, IFN, IHEAD, R1, R2, YMAX,   |

```
GD TO 145
140
        CONTINUE
        CALL GRAFR (ARAY (1), NPTS, IFN, IHEAD, R1, R2, YMAX,
     + IPLOT, KREC, 1, DELZ, IREC, 0)
C
C
                 OPTION TO PRINT DATA
C
        ACCEPT "PRINT RAW DATA ? YES = 1 NO = 0 ", IANS
145
        IREND=IREC+NREC-1
        IF (IANS. NE. 1) GO TO 150
        PRINT 1017, IFN, IREC, IREND, (IHEAD(I), I=1,6)
        FORMAT(1H1, "FILE ", 12, /, " RECS ", 13, " THRU ", 14,
1017
        /, " DATE ", I2, "/", I2, "/", I2,
        /, " TIME ", I2, ": ", I2, ": ", I2, //)
        PRINT 1018
        FORMAT(" AVERAGE POWER VS. RANGE", //)
1018
        PRINT 1019, (I, AVGPWR(I), I=1, NPTS)
1019
        FORMAT(5(8X, I3, 1X, F10, 2))
C
C
                 CHOOSE A NOISE WINDOW TO FIND MEAN NOISE LEVEL (AVNOIS)
C
150
        ACCEPT "BEG, END ALTITUDES (KM) FOR NOISE WINDOW ? ", AN1, AN2
        RNOS1=(AN1-ALTLID)/ZZZ
        RNOS2=(AN2-ALTLID)/ZZZ
        TNOS1=RNOS1/. 149896-DELAY+2. 0
        NOS1=INT(TNOS1/DELT+0.5)
        TNOS2=RNOS2/. 149896-DELAY+2. 0
        NOS2=INT(TNOS2/DELT+0.5)
        IF(NOS1. LT. 1)NOS1=1
        IF (NOS2, QT, 2048) NOS2=2048
        NNOIS=NOS2-NOS1+1
        IF (NNOIS . GT. 2048) NOS1 = NOS2 - 2048 + 1
        IF (NNOIS . GT. 2048) NNOIS = 2048
        AVNOIS=O.
        NOS1 = NOS1 - ISTART
        NOS2 = NOS2 - ISTART
        IF( NOS1 .LT. 1 ) NOS1 = 1
        DO 155 I=NOS1, NOS2
                AVNOIS=AVNOIS+AVGPWR(I)
155
        CONTINUE
        AVNOIS=AVNOIS/NNOIS
C
C
                 BEGIN CALCULATING BETA SNR'S FOR GIVEN ALTITUDE
C
                AVERAGING INTERVALS.
C
        TYPE
        ACCEPT "ALTITUDE RESOLUTION IN KM (1.005 FOR ARCHIVING) ?", DALT
        NINC=DALT/DELZ
        ZINC=NINC*DELZ
        NN=67*PARM(1)/10.
        QUAL=1. /SQRT(NN)+1. /SQRT(NINC)
        WRITE(10, 1020) QUAL
        FORMAT ("G-FACTOR MUST BE GREATER THAN", F5. 2)
1020
        ICOUNT=0
        DO 165 K=1, MPTS, NINC
                 ICOUNT=ICOUNT+1
                AVSIG=Q.
                DO 160 I=1, NINC
                         AVSIG=AVSIG+AVGPWR(K-1+I)
160
                CONTINUE
                AVSIG=AVSIG/NINC
```

IPLOT, KREC, 1, DELZ, IREC, 0)

|      | AVEP(ICOUNT)=AVSIG ALT=R1+(ICOUNT-1)*ZINC+ZINC/2. SIGNOS=AVNOIS/SGRT(NREC) GFAC=(AVSIG-AVNOIS)/SIGNOS CALL VSTASH(GFAC, 4096+ICOUNT) SNR=(AVSIG-AVNOIS)/AVNOIS IF(SNR.LT.O.O) SNR = O.O PWR(ICOUNT)=SNR RNG = (ALT-ALTLID)/ZZZ R=RNG*1000. |
|------|--|
| C    | ADD LIDAR ALTITUDE ASL FOR   |
| C    | INDEXING THE ABSORPTION ARRAY ONLY.  |
| C    | LINEARLY INTERPOLATE THE BEAM ARRAY  |
| C    | IN RANGE AND THE ABSORPTION ARRAY IN ALTITUDE  |
| C    | IRNG=RNG   |
|      | DIFRAN=RNG-IRNG  |
|      | IRNG1=IRNG+1   |
|      | IRA-ALT  |
|      | IRA1=IRA+1   |
|      | DRANG=ALT-IRA  |
|      | RKAP=(BEAM(IRNG1)-BEAM(IRNG))*DIFRAN+BEAM(IRNG) ABSP=(ABSRP(IRA1,MOD)-ABSRP(IRA,MOD))*DRANG+ABSRP(IRA,MOD) IF(IRNG.GE.30) RKAP=BEAM(30)  |
|      | IF(IRA. GE. 30) ABSP=ABSRP(30, MOD)  |
|      | P4=P1*(R**2+P2)/(P3*RKAP*ABSP)   |
| C    |  |
| C    | USE AVPWDB ARRAY AS THE BETA PROFILE ARRAY   |
|      | BETA=P4*SNR*CAL  |
|      | IF (QFAC. LT. QUAL) BETA=1. E-15   |
|      | CALL VSTASH(BETA, 2048+ICOUNT)   |
| 165  | CONTINUE   |
|      | ICM1=ICOUNT-1  |
| C    | The second s   |
| C    | ELIMINATE ISOLATED "VALID" BETA VALUES   |
| С    | DO 170 II-3 ICM1   |
|      | DO 170 II=2, ICM1<br>CALL VFETCH(BEF, II+2047)   |
|      | CALL VFETCH(BEF, II+2047)  |
|      | IF (BEF. EQ. 1. E-15. AND. AFT. EQ. 1. E-15) CALL VSTASH (BEF, II+2048)  |
| 170  | CONTINUE   |
|      | TYPE   |
| C    |  |
| C    | PRINTOUT OPTION  |
| C    |  |
|      | ACCEPT "PRINT OUT BETA PROFILE ? YES = 1 NO = 0 ", IANS IF(IANS. NE. 1) GO TO 180 PRINT 1017, IFN, IREC, IREND, (IHEAD(I), I=1,6)  |
|      | PRINT 1021   |
| 1021 | FORMAT (" ALT (KM) BETA(M-1 SR-1) Q-FACTOR SNR",/) DO 175 J=1, ICOUNT ALT=(J-1)*ZINC+ZINC/2.+R1  |
|      | CALL VFETCH(AVPWDB, J+2048)  |
|      | CALL VFETCH(PWRDB, J+4096)   |
|      | PRINT 1022, ALT, AVPWDB, PWRDB, PWR(J)   |
| 1022 | FORMAT (F7. 1, 7X, 1PE9. 2, 6X, 0PF6. 2, F12. 3)   |
| 175  | CONTINUE   |
| C    |  |
| C    | PLOT BETA PROFILE  |
| C    |  |
| 180  | DO 185 I=1, ICOUNT   |

```
AVPWDB=10. *ALOG10(AVPWDB)+110.
                 CALL VSTASH(AVPWDB, I+2048)
185
        CONTINUE
        YMAX=40.
        BFLAG=1.
        DO 190 III=1, ICOUNT
                 CALL VFETCH(ARAY(III), 2048+III)
190
        CONTINUE
        CALL GRAFR (ARAY, ICOUNT, IFN, IHEAD, R1, R2, YMAX,
     + IPLOT, KREC, 1, ZINC, IREC, BFLAG, AN1, AN2, MOD)
        TYPE
        ACCEPT "ANOTHER NOISE WINDOW (1) OR CONTINUE (0) ? ", IANS
        IF (IANS. EQ. 1) QO TO 150
C
C
                 ARCHIVE BETA DATA IF DESIRED, CHECKING FOR PROPER
C
                 POSITIONING AND INCREMENTING.
C
        TYPE
        ACCEPT "ARCHIVE THE DATA ? YES=1 NO=0 ", IANS
        IF (IANS. NE. 1) GO TO 210
        IALT=ALT
        DIFALT=ALT-IALT
        IF (DALT, EQ. 1, 005, AND, DIFALT, LT. . 2) 90 TO 195
        TYPE"ERROR IN ALTITUDE INCREMENT OR INDEXING"
        TYPE "DALT SHOULD BE 1.005, INDEXING ON EVEN KM (ASL)."
        TYPE
        GO TO 150
        DO 200 I=1, ICOUNT
195
                 CALL VFETCH(AVPWDB, 2048+I)
                 AVPWDB=10. **((AVPWDB-110.)/10.)
                 CALL VSTASH(AVPWDB, 2048+I)
200
        CONTINUE
        CALL APPEND(3, "ARCH", O, IER)
        IF (IER. NE. 1) STOP - ARCH OPEN ERROR
        WRITE BINARY(3) IFN, IREC, IREND, (IHEAD(I), I=1,6), ICOUNT,
        MOD, R1, DALT, ALTLID
        DO 205 I=1, ICOUNT
                 ALT=(I-1)*ZINC+ZINC/2. +R1
                 IALT=ALT
                 IALT1= IALT+1
                 DIFALT=ALT-IALT
                 IF (DIFALT. GE. O. 5) IALT=IALT1
                 ALT=IALT
                 CALL VFETCH(AVPWDB, 2048+I)
                 CALL VFETCH(PWRDB, 4096+I)
                 WRITE BINARY (3) ALT, AVPWDB, PWRDB, PWR(I), AVEP(I)
205
        CONTINUE
        CLOSE 3
        CALL NEWPAG
        TYPE"BE SURE TO RECORD THIS ARCHIVING ON THE MASTER LIST"
        WRITE(10, 1025) (IHEAD(I), I=1, 6), IFN, NREC, BW/1. E6, EJ
        FORMAT(/, "DATE", T25, I2, "/", I2, "/", I2, //"TIME", T25, I2, ": ",
1025
        12, ": ", 12, //"FILE NUMBER", T25, 18, //"NUMBER OF RECORDS", T25,
        18, // "BANDWIDTH (MHZ)", T25, F8. 1, // "LASER POWER", T25, F8. 3///)
210
        ICURR=IREC+MREC
        ACCEPT "PROCESS ANOTHER FILE (1) OR STOP (0)? ", IANS
        IF (IANS. EQ. 1) CALL NEWPAG
        IF (IANS. EQ. 1) GO TO 1
        CALL MTRWN (ICHAN)
        STOP
```

42

CALL VFETCH(AVPWDB, I+2048)
IF(AVPWDB, LE. O. ) AVPWDB=1. E-20

| C    | TAPE ERROR ROUTINE  |
|------|---|
| 215  | CALL NEWPAG   |
|      | WRITE(10, 1023) IER, ISTAT, IWRD                                      |
|      | WRITE(10, 1024) KREC  |
| 1023 | FORMAT(1X, "TAPE ERROR", /, "IER=", I4, "ISTAT= ", B16, " IWRD ", I5) |
| 1024 | FORMAT(1X, "TAPE ERROR OCCURED ON RECORD NUMBER ", 15)                |
| 550  | CALL MTRWN (ICHAN)  |
|      | STOP  |
|      | END   |

## REFERENCES

- Lawrence, T. R., 1983. Comparison of the observed antenna gain properties of the NOAA lidar with theory. To be published in Applied Optics.
- Lott, G. A., 1976. Precipitable water over the United States, Vol. 1, Monthly means. NOAA Tech. Rep. NWS-20, 174 pp.
- McClatchey, R. A., R. W. Fenn, J. E. A. Selby, F. E. Volz, and J. S. Garing, 1972. Optical properties of the atmosphere (Revised). AFCRL-72-0497, 108 pp.
- Post, M. J., R. A. Richter, R. M. Hardesty, T. R. Lawrence, and F. F. Hall, Jr., 1982. NOAA's pulsed, coherent, IR Doppler lidar--characteristics and data. Proc. Soc. Photo-Opt. Instrum. Eng. 300:60.
- Roberts, R. E., J. E. A. Selby, and L. M. Biberman, 1976. Infrared continuum absorption by atmospheric water vapor in the 8-12 µm window. Appl. Opt. 15:2085-2089.
- Zuev, V. Ye., 1972. Propagation of Visible and Infrared Waves in the Atmosphere. NASA Technical Translation NASA TT F-707, pp. 5-6.