

1    Ratios of greenhouse gas emissions observed over  
2    the Yellow Sea and the East China Sea

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10

11    **Abstract**

12    During a cruise of the survey vessel Dongfanghong II on the Yellow Sea and the East China Sea  
13    in the spring of 2017 we performed accurate measurements of the mole fractions of carbon dioxide  
14    ( $\text{CO}_2$ ), methane ( $\text{CH}_4$ ), carbon monoxide ( $\text{CO}$ ) and nitrous oxide ( $\text{N}_2\text{O}$ ) using two types of Cavity  
15    Ring-Down Spectrometers (CRDS). The spatial variations of the mole fraction of the four trace  
16    gases were very similar. The emission sources of these gases were divided into several regions by  
17    using the NOAA HYSPLIT model. Then we analyzed the variations of the ratios of the mole  
18    fraction enhancements between every pair of trace gases downwind of these source areas. The  
19    ratios showed that the distributions of these trace gases over the Yellow Sea and the East  
20    China Sea in the spring were mainly caused by the emissions from Eastern China. The much

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21 higher enhancement ratio of  $\Delta\text{CO}/\Delta\text{CO}_2$  and the lower ratio of  $\Delta\text{CH}_4/\Delta\text{CO}$  observed in the air  
22 parcels from big cities like Beijing and Shanghai indicated high CO emission from the cities  
23 during our time of observation. Compared with the values of NOAA's Marine Boundary Layer  
24 (MBL), the ratios of the averages in the air coming from the Northern sector (Russia) were on  
25 average closer to the MBL, and the air that stayed over the Yellow Sea and the East China Sea was  
26 a mixture of emissions from wide regional areas. The highly variable  $\text{N}_2\text{O}$  data of the air from  
27 Qingdao and Shanghai showed much more fluctuation.

28 *Keywords:* mole fraction, calibration, source regions, enhancement ratios, shipboard air  
29 measurements, major greenhouse gases

30

31 **1. Introduction**

32

33 The problems of global warming and climate change are mainly caused by the greenhouse  
34 gas increases, and have become one of the most important challenges of the 21st century (Liu  
35 et al., 2009). Greenhouse gases are pervasively produced by human activities, such as coal  
36 combustion, vehicle traffic, heating/cooling and industrial processes. The Asian continent has  
37 become the largest source of anthropogenic pollutants. Emissions to the atmosphere have  
38 rapidly increased recently as a result of the economic growth in Asian countries, especially  
39 China (Sakata et al., 2013; Zhang et al., 2017). The industrialization and energy consumption  
40 in China are increasing rapidly, and the total  $\text{CO}_2$  emissions from fossil-fuels and cement  
41 production increased from 1.4 billion ( $10^9$ ) metric ton carbon (GtC) in 2004 to 2.8 GtC in  
42 2014 according to the data of Carbon Dioxide Information Analysis Center (Boden and

43 Andres, 2017). This is regarded as the main reason for a sharp increase in continental Asian  
44 air pollution (Junker et al., 2009).

45 The platforms for observations of the greenhouse gases and their temporal and spatial  
46 variations around the world are ground-based stations (Cheng et al., 2017; Gomez-Pelaez et  
47 al., 2013; Zellweger et al., 2016), aircraft (Chen et al., 2010; Daube et al., 2002; Machida et  
48 al., 2008), satellites (Butz et al., 2011; Guo et al., 2017; Wunch et al., 2017), and ships (Feely  
49 et al., 2006; Schuster et al., 2009). Observation data are essential for atmospheric inversion  
50 methods to estimate regional sources and sinks, regional fluxes and the transport of the  
51 greenhouse gases (Deng et al., 2014; Peylin et al., 2013; Sawa et al., 2012).

52 Akira Wada et al. ( 2007; 2011) analyzed seasonal and geographical variations of the  
53 enhancement ratios of trace gases and found that the Asian continental pollution was  
54 influenced by the seasonal variations of the emission ratios in the source regions, with  
55 increased fluxes of methane ( $\text{CH}_4$ ) in summer and carbon monoxide (CO) in winter. The  
56 Asian influence on CO in the North Pacific troposphere is at a maximum during spring and at  
57 a minimum during summer (Liang et al., 2004), and observed and simulated concentrations of  
58 CO<sub>2</sub> in the Asian outflow would imply increases in Chinese anthropogenic CO<sub>2</sub> emissions  
59 (Suntharalingam et al., 2004).

60 In this paper, we study the data of carbon dioxide (CO<sub>2</sub>), methane (CH<sub>4</sub>), carbon monoxide  
61 (CO), and nitrous oxide (N<sub>2</sub>O) obtained on the ship over the Yellow Sea lying between  
62 northeastern China and the Korean Peninsula (Li et al., 2016) and the East China Sea (32.07°  
63 N, 125.10° E) in the spring (Han et al., 2015). We designed and built a ship-based continuous  
64 observation system for the main atmospheric greenhouse gases and CO in order to explore the

65 factors that affected the distribution of the gases. Firstly we defined the stability of the two  
66 analyzers during the cruise and then we analyzed the mole fraction distribution of the four gas  
67 species over the Yellow Sea and the East China Sea during the cruise. We used two methods  
68 to calculate the enhancement ratios for every pair of gases, and analyzed the results for  
69 understanding emission characteristics of the source regions.

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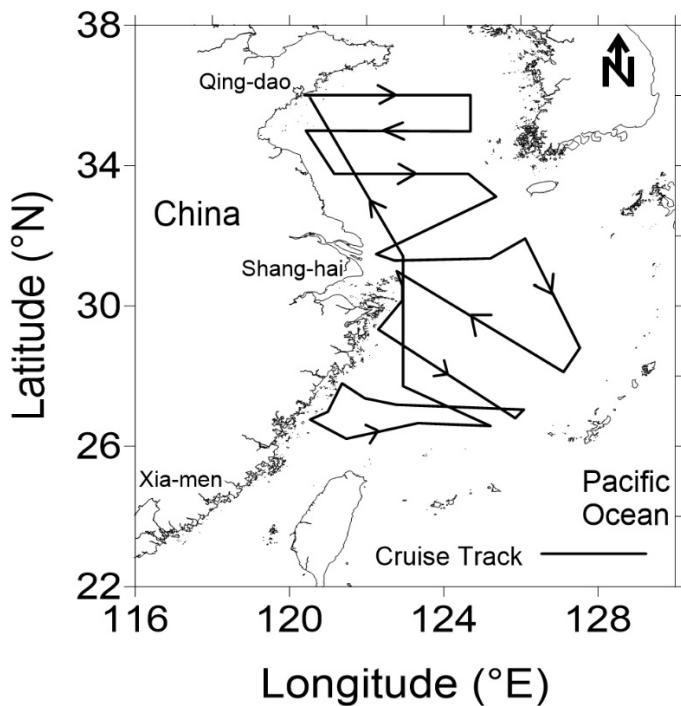
71 **2. Observation**

72

73 *2.1. Observation area*

74

75 We performed accurate measurements of CO<sub>2</sub>, CH<sub>4</sub>, CO, and N<sub>2</sub>O in air during an  
76 oceanographic cruise on board the survey vessel Dongfanghong II, organized by the Ocean  
77 University of China from 27 March 2017 to 15 April 2017. The ship track is shown in Figure  
78 1.



79

80 **Fig. 1.** Cruise track between the Pacific Ocean and the Asian continent. The survey vessel  
 81 sailed from Qingdao and finally returned to Qingdao.

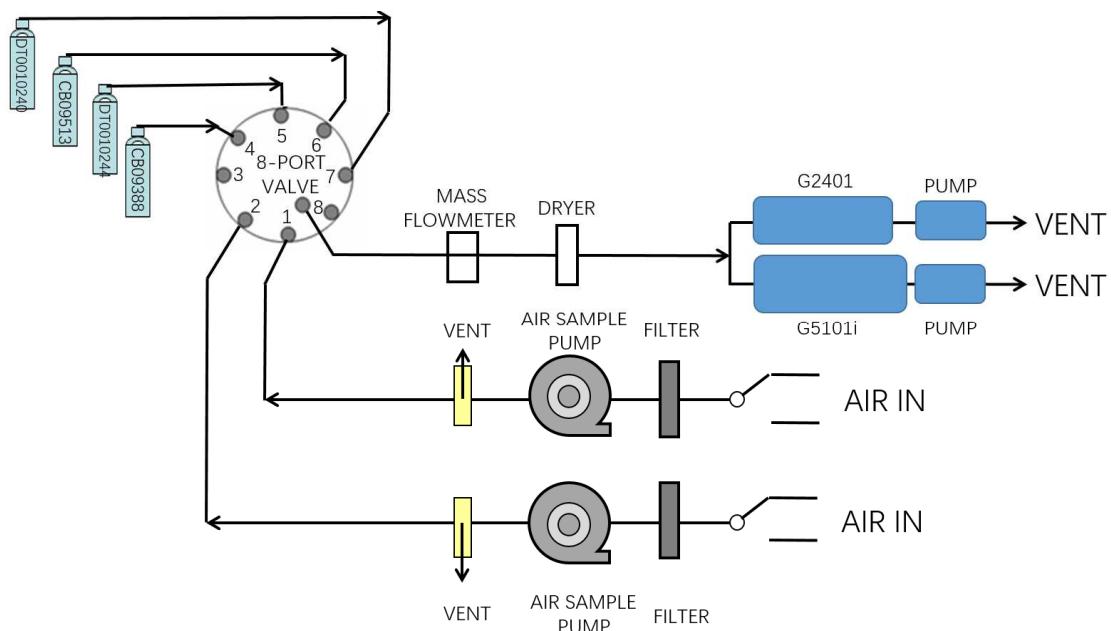
82

83 *2.2. Observation methods*

84

85 We took a Picarro G5101i analyzer for  $\text{N}_2\text{O}$ , and a Picarro G2401 analyzer for  $\text{CO}_2$ ,  $\text{CH}_4$ ,  
 86 and  $\text{CO}$ , as well as four high pressure calibration standard gas cylinders, each containing very  
 87 accurately known amounts of the four gases in air, from 394.24 to 434.12 micromol/mol (ppm)  
 88 for  $\text{CO}_2$ , from 1935.00 to 2288.89 nanomol/mol (ppb) for  $\text{CH}_4$ , from 317.20 to 878.51 ppb for  
 89  $\text{CO}$ , and from 309.75 to 609.24 ppb for  $\text{N}_2\text{O}$ . The calibration of these standard gases was  
 90 propagated by the Chinese Academy of Meteorological Sciences, China Meteorological  
 91 Administration (CAMS-CMA). The CAMS standard gases were in turn calibrated using the  
 92 primary standards of the World Meteorological Organization Global Atmospheric Watch  
 93 (WMO-GAW), maintained by NOAA/ESRL ([www.esrl.noaa.gov/gmd/ccl/](http://www.esrl.noaa.gov/gmd/ccl/); Fang et al., 2014;

94 Tans and Zellweger, 2014). Sample air was provided by using a sample inlet installed above  
 95 the top deck, about 10 m in front of the ship's engine exhaust stack. The schematic diagram of  
 96 the measurement set up was as follows. The 1.0  $\mu\text{m}$  membrane filter between the air inlet and  
 97 the pump was used to remove particles. One of the air inlets was used as a backup to avoid  
 98 special situations during the voyage. The sample or standard gas flowed through the eight port  
 99 multi-position valves into a mass flowmeter which controlled the flow rate at 350  $\text{mL}\cdot\text{min}^{-1}$ ,  
 100 then flowed into a dryer to remove water vapor, a tube filled with magnesium perchlorate  
 101 ( $\text{Mg}(\text{ClO}_4)_2$ ). Glass cotton on both sides of the dryer was used to avoid pumping the  
 102 magnesium perchlorate particles into the analyzers. The cavity pressure of G2401 was  
 103 maintained at a stable value of 140 torr by the vacuum pump, while the cavity operating  
 104 pressure of G5101i was at 100 torr (Crosson, 2008; Erler et al., 2015)



105  
 106 **Fig. 2.** Schematic diagram (not to scale) of the ship-based atmospheric  $\text{CO}_2$ ,  $\text{CO}$ ,  $\text{CH}_4$ , and  
 107  $\text{N}_2\text{O}$  observation system. The flow rate was controlled at 350  $\text{mL}\cdot\text{min}^{-1}$  by the mass  
 108 flowmeter.

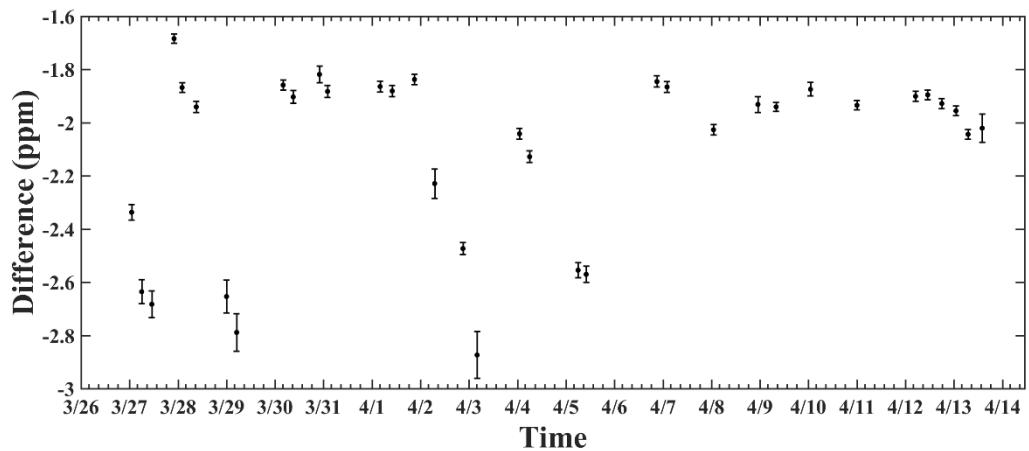
110 2.3. *Calibrations*

111

112 Using the calibration gases, the response functions of the analyzers were calculated 35  
113 times during the whole voyage. The Picarros recorded measurement values approximately  
114 every 5 seconds. Every standard gas ran for ten minutes each time, and to ensure complete  
115 flushing of the measurement cell only the last 5 minutes of data were used. Then the averages  
116 of the stable five minutes of calibration data was calculated for each standard gas. All four  
117 standards were used during each calibration to generate a response curve for each instrument  
118 over the ranges mentioned above. During sample air measurements, in between the  
119 calibrations, the averages of the bracketing calibrated response curves were used to correct the  
120 observation data.

121 Figures 3-6 showed the average difference and standard deviation during the 5 min.  
122 averaging periods between the value indicated by the instrument and the known value of the  
123 standard gas for all 35 calibrations, for example cylinder number CB09388. The other three  
124 standard gases behaved similarly to CB09388, with dips in a few cases at the same times. The  
125 uncertainty of the air measurements was larger between times in which successive calibration  
126 curves showed a larger change. Compared to the other three gases, CO<sub>2</sub> showed relatively  
127 larger changes.

128

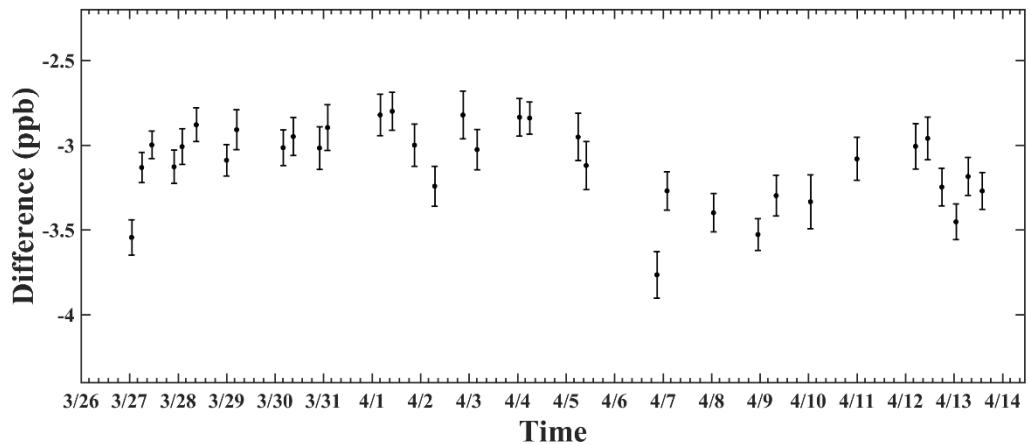


129

130 **Fig. 3.** Stability of successive CO<sub>2</sub> calibrations, the error bars represent the standard deviation

131 of 5-second averages, 35 in total.

132

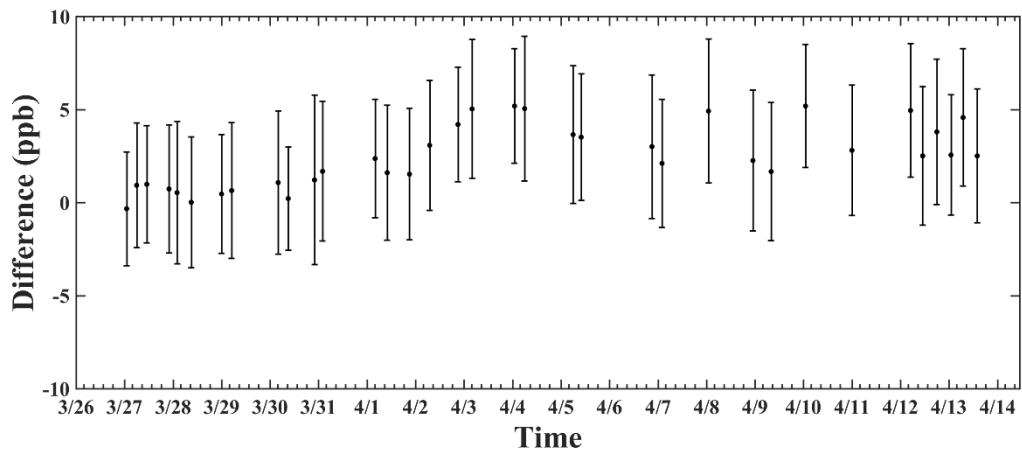


133

134 **Fig. 4.** Stability of successive CH<sub>4</sub> calibrations, the error bars represent the standard deviation

135 of 5-second averages, 35 in total.

136

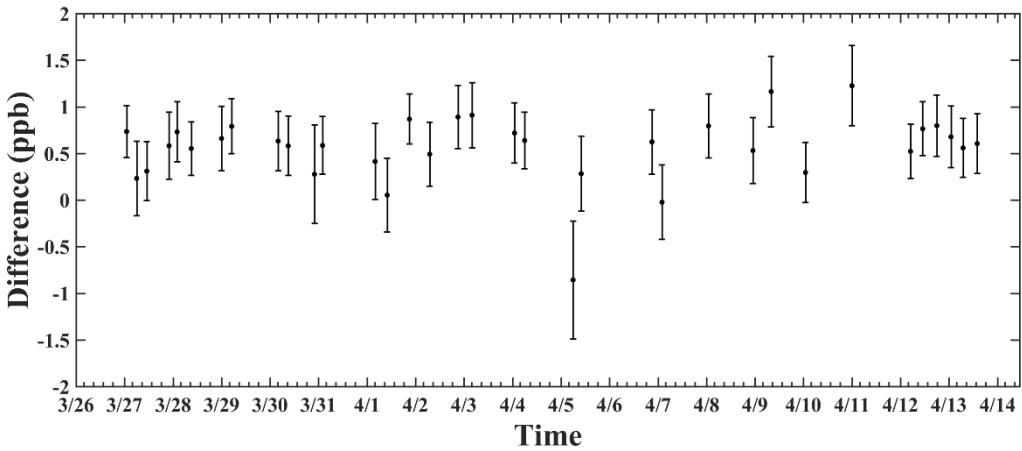


137

138 **Fig. 5.** Stability of successive CO calibrations, the error bars represent the standard deviation

139 of 5-second averages, 35 in total.

140



141

142 **Fig. 6.** Stability of successive N<sub>2</sub>O calibrations, the error bars represent the standard deviation

143 of 5-second averages, 35 in total.

144

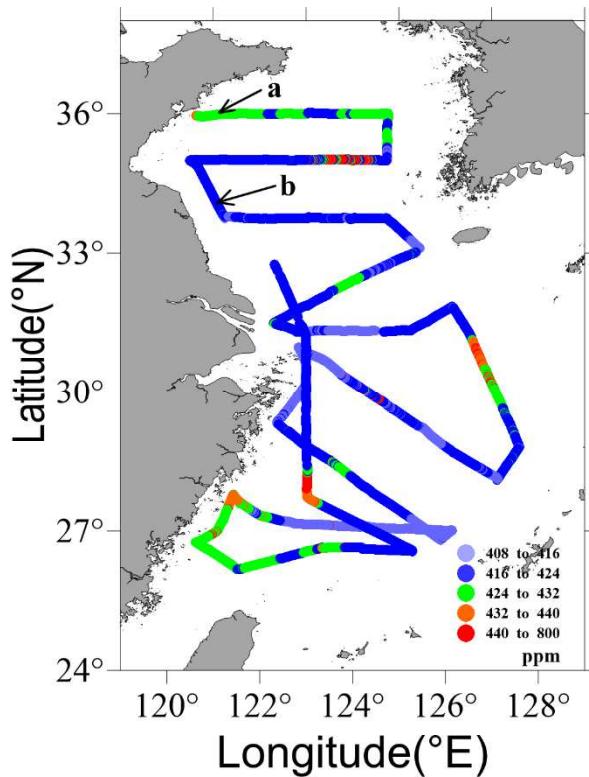
### 145 3. Results

146

#### 147 3.1. Observed mole fraction distribution

148

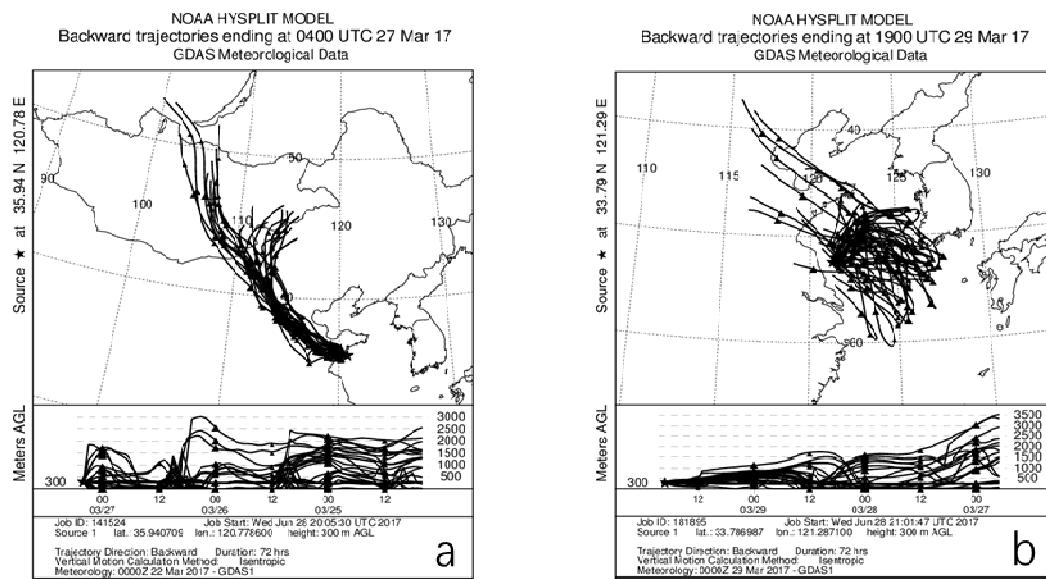
149 To avoid interference from the ship's engine exhaust, the data when the ship's speed was  
150 less than 3 knots was not used. Occasional contamination from the stack was characterized by  
151 very high frequency variability, and we did not use that data either. Fig.7 shows the mole  
152 fraction distribution of CO<sub>2</sub> during the voyage. The distributions of the other three gases had  
153 similar trends. The data were averaged over each minute and linked to the latitude and  
154 longitude. In order to smooth the mole fraction distributions, we calculated hourly average  
155 data. Then we used the NOAA HYSPLIT model to generate 72-hour backward air trajectories  
156 for each hour with its average longitude and latitude, with the starting point at 300 m altitude.  
157 Also, the CO<sub>2</sub>, CO, CH<sub>4</sub>, and N<sub>2</sub>O data were divided into two parts according to the hourly  
158 standard deviation, with the dividing lines respectively at 0.88 ppm, 6.35 ppb, 1.15 ppb and  
159 1.48 ppb. Above these values, the data were not very stable so we called them higher variability  
160 data while below those values we called them lower variability data. According to the calculated  
161 air trajectories, all the data was divided into ten distinct source regions that influenced the  
162 data at different times during the cruise. The locations of source regions and regional stations  
163 are showed in the Fig. 9. Fig. 8a is an example showing Beijing/Jinan and 8b showing Yellow  
164 Sea. The data of the Philippine Sea, at the eastern end of the cruise, were used as the base line  
165 for comparing all other data.



166

167 **Fig. 7.** CO<sub>2</sub> mole fraction distribution as a function of degrees north latitude and east  
 168 longitude. The point labeled “a” refers to Fig. 8a, and “b” to Fig. 8b, below.

169



170

171 **Fig. 8.** Three-day air back-trajectories of two locations, a (35.94° N, 120.78° E) and b  
 172 (33.77° N, 121.29° E).

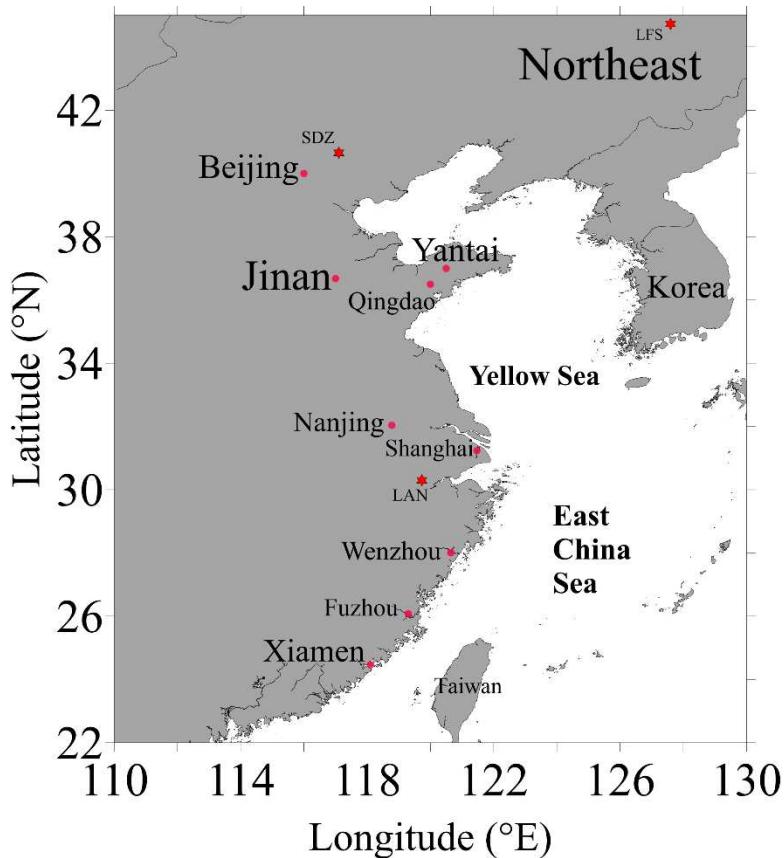


Fig. 9. The locations of the source reigns and regional stations.

176 *3.2. Ratios of mole fraction enhancements*

178 Figures 10 to 15 show the ratios of the observed mole fraction enhancements relative to the  
 179 Philippine Sea, pairwise for  $\Delta\text{CO}_2$ ,  $\Delta\text{CO}$ ,  $\Delta\text{CH}_4$ , and  $\Delta\text{N}_2\text{O}$ . The average values for the  
 180 Philippine Sea observed during the cruise are 410.7 ppm, 113 ppb, 1896 ppb, and 332.5 ppb,  
 181 respectively. NOAA's Marine Boundary Layer (MBL) values (Masarie, 1995), representative  
 182 of large ocean areas, at the same latitude zone and during the cruise, were 410.7 ppm, 128 ppb,  
 183 1911 ppb, and 330.2 ppb respectively. This indicates that our Philippine Sea data is a good  
 184 reference for this study. We used two different methods. In the first we plotted enhancements  
 185 for individual hourly averages of a pair of species directly against each other for each source

186 region, and we determined the slope, without paying attention to the average offsets (blue  
 187 bars), with the aim to emphasize more recent emissions that had undergone less mixing and  
 188 may have had larger hour-to-hour variations. For the purpose of calculating uncertainties from  
 189 the statistics, the hourly averages have the advantage that successive hours are independent of  
 190 each other whereas successive one minute averages are often very close together, and can not  
 191 be considered to be independent measurements. The blue and orange columns show lower  
 192 variability data and higher variability data respectively, again to try to separate emissions that  
 193 were perhaps closer to the measurement point from those further away. The heights of the  
 194 bars are the slopes of the relations between every two gas species. The slope is obtained by  
 195 first fitting  $\Delta Y = a \Delta X + b$  assuming errors only on the y-axis, and by then fitting  $\Delta X =$   
 196  $\frac{\Delta Y}{a'} - \frac{b'}{a'}$  assuming errors only on the x-axis. This gives different values for the slopes  $a$  and  
 197  $a'$ . We take the geometric mean of  $a$  and  $a'$ , keeping track of whether the slope is positive or  
 198 negative, as below.

$$199 \quad c = \text{sign}(a) * \sqrt{aa'} \quad (1)$$

200 The difference between the two slopes  $a$  and  $a'$  provides a measure of the uncertainty of  $c$   
 201 (denoted as “ $u(c)$ ”), namely  $\pm (a-a')/2$ , which is plotted on each bar in Figs 10-15.

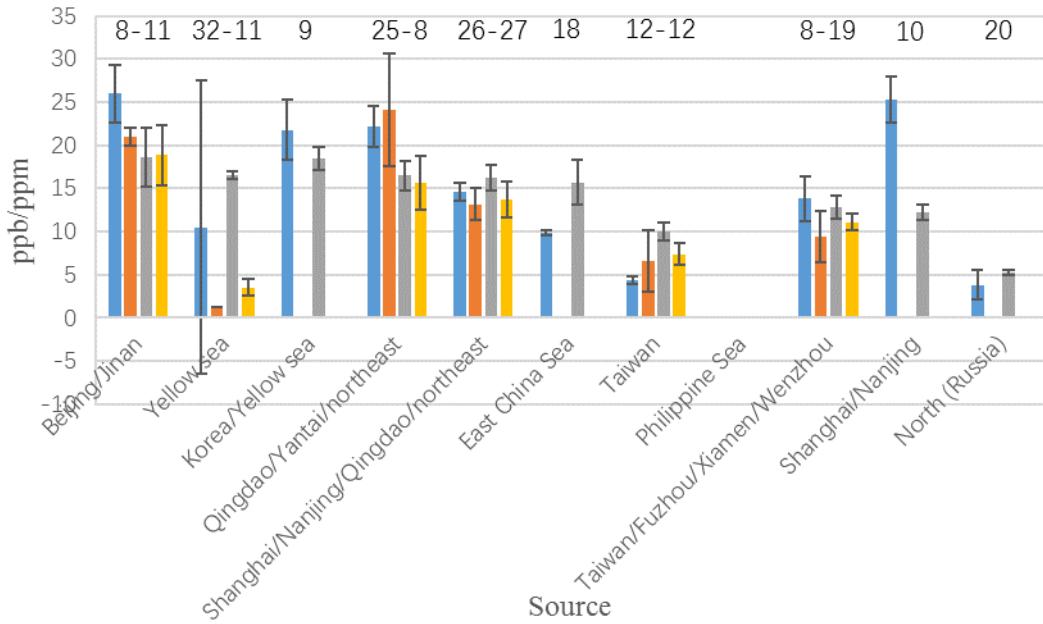
202 The gray and yellow column values show the ratios ( $Y$ ) of the averages (denoted by the  
 203 overbar  $\bar{Y}$ ) of the separate enhancements of two species  $X_1$  and  $X_2$ , namely  $\bar{Y} = \frac{\bar{X}_1}{\bar{X}_2}$ , for the  
 204 lower variability data (gray) and the higher variability data (yellow), respectively. The  
 205 uncertainty is estimated as the relative uncertainty of a product or quotient, as in the GUM  
 206 1995, as follows:

$$207 \quad \frac{u(Y)}{Y} = \sqrt{\left[ \frac{u(\bar{X}_1)}{\bar{X}_1} \right]^2 + \left[ \frac{u(\bar{X}_2)}{\bar{X}_2} \right]^2} \quad (2)$$

208 in which the standard error of the mean for both species is

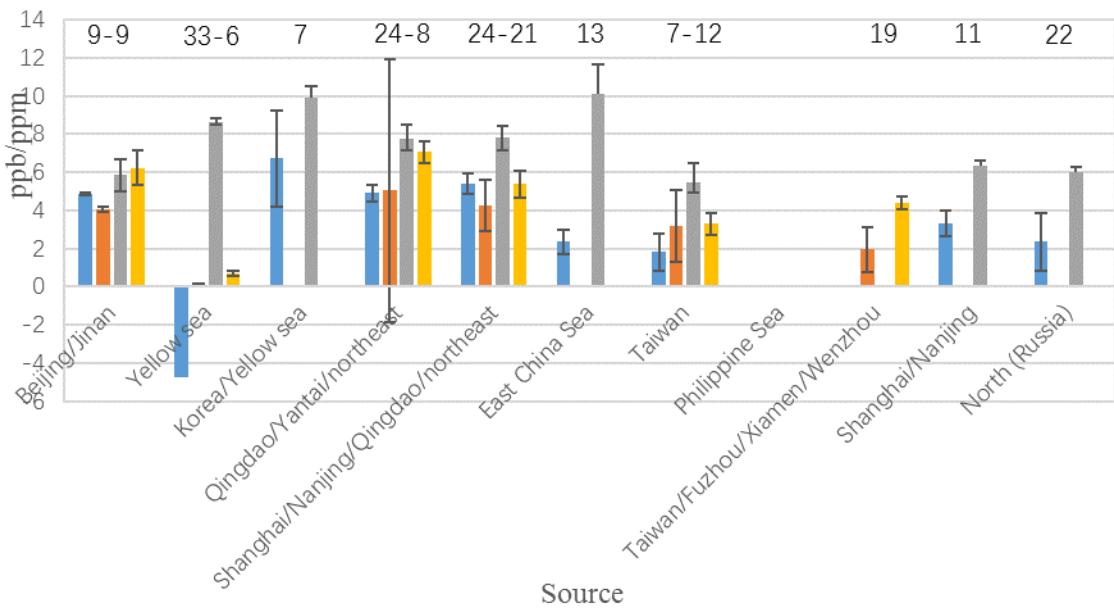
209 
$$[u(\Delta X_1)]^2 = \frac{\sum_{i=1}^n (\Delta x_{1i} - \bar{\Delta x}_1)^2}{n(n-1)} \text{ and } [u(\Delta X_2)]^2 = \frac{\sum_{i=1}^n (\Delta x_{2i} - \bar{\Delta x}_2)^2}{n(n-1)} \quad (3)$$

210 and where  $\Delta X_1, \Delta X_2$  can be one of  $\Delta \text{CO}_2, \Delta \text{CO}, \Delta \text{CH}_4, \Delta \text{N}_2\text{O}$ .



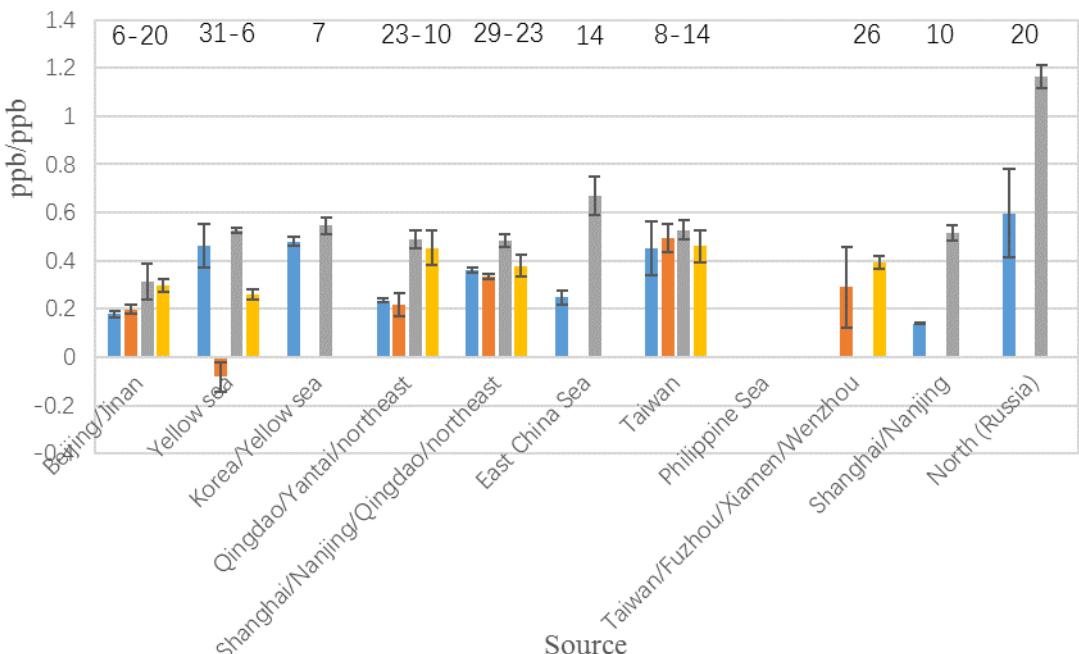
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212 **Fig. 10.** The enhancement ratios of  $\Delta \text{CO}/\Delta \text{CO}_2$ . For all plots 10-15, the numbers n1-n2 above  
213 the bars give the number of hourly averages used for each region, n1 for lower variability data  
214 (blue and gray bars), n2 for higher variability data (orange and yellow bars). We do not plot  
215 bars when they are based on 5 or fewer hourly averages.



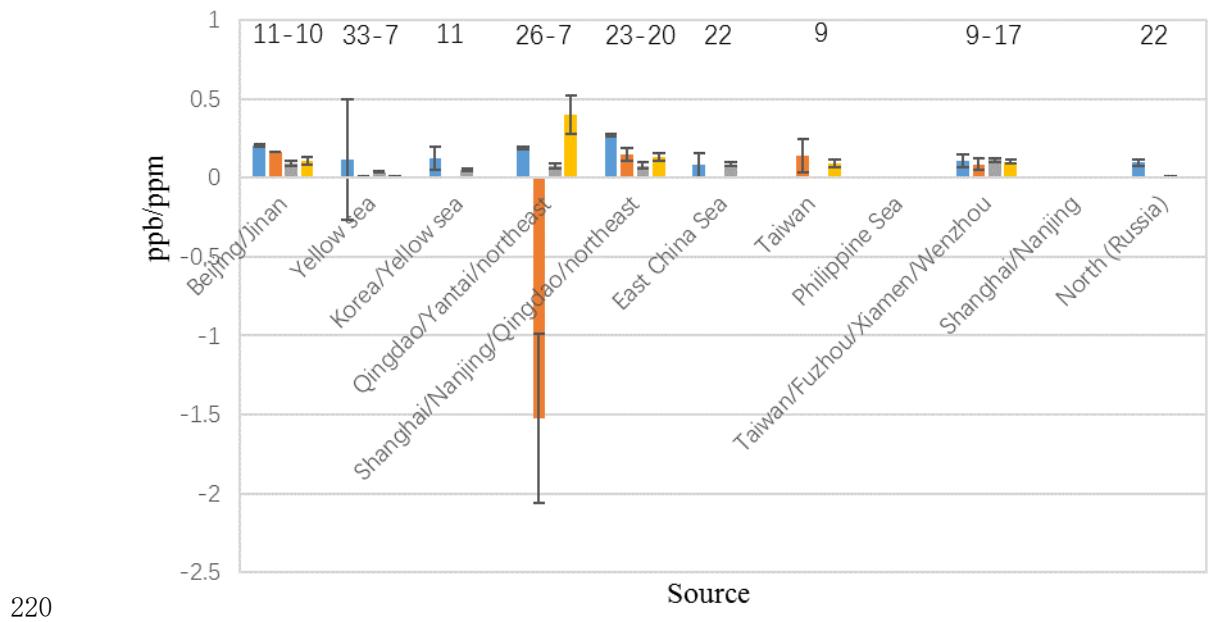
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217 **Fig. 11.** The enhancement ratios of  $\Delta\text{CH}_4/\Delta\text{CO}_2$ .

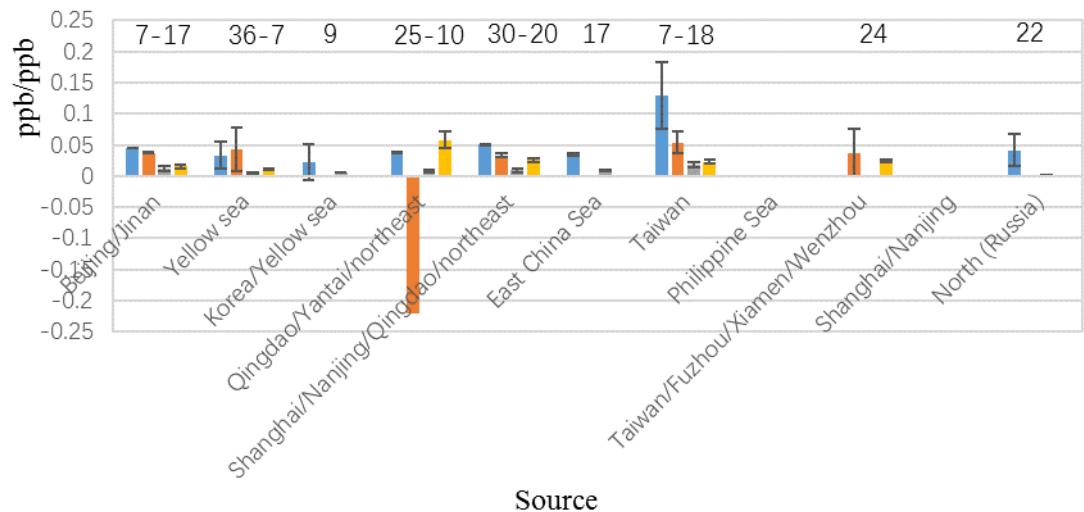


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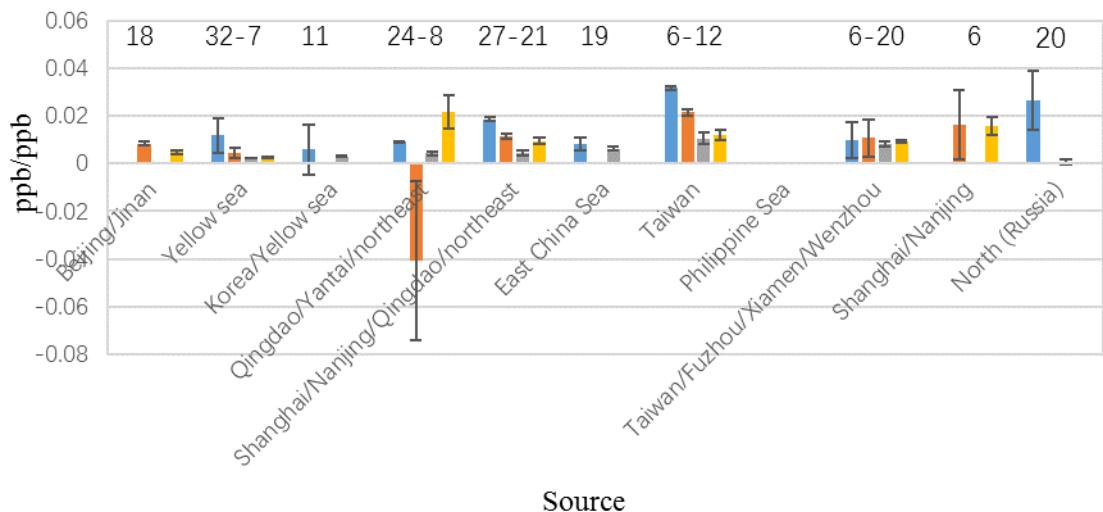
219 **Fig. 12.** The enhancement ratios of  $\Delta\text{CH}_4/\Delta\text{CO}$ .



221 **Fig. 13.** The enhancement ratios of  $\Delta\text{N}_2\text{O}/\Delta\text{CO}_2$ .



223 **Fig. 14.** The enhancement ratios of  $\Delta\text{N}_2\text{O}/\Delta\text{CH}_4$ .



224

225 **Fig. 15.** The enhancement ratios of  $\Delta N_2O/\Delta CO$ .

226

227 *3.3. Discussions*

228

229 We will briefly discuss some characteristics of the ratios of mole fraction enhancements as  
 230 shown in Figs. 10-15 for individual source regions because they are a measure of their relative  
 231 emissions during the time period of the cruise.

232

233 *3.3.1. Beijing Jinan*

234 In both lower and higher variability data the correlation between every pair of gases is  
 235 strong. The average mole fraction enhancements of the four gases CO<sub>2</sub>, CO, CH<sub>4</sub>, N<sub>2</sub>O in the  
 236 low-noise case were respectively 13.6 ppm, 244.7 ppb, 81.9 ppb, 1.1 ppb higher than the base  
 237 values and higher than for other source regions except for N<sub>2</sub>O and CH<sub>4</sub>. CO emissions are  
 238 very high, and likely caused by traffic. They were 3-4 times higher than CH<sub>4</sub> emissions. The

239 bars show ratios of emissions, so that if we have a trustworthy estimate for one of them, for  
240 example CO<sub>2</sub>, then we have estimates for the other gases. The higher variability data was  
241 similar to the lower variability data. CO<sub>2</sub> emissions from burning fossil fuels and factories, as  
242 well as CH<sub>4</sub> emissions from landfills, made up 97% GHGs of Beijing in 2014 (Li et al., 2017).  
243 Li et al. gathered and calculated the results of the inventory of Beijing GHG emissions  
244 including CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O for 2014. We use the Li et al.'s estimate of CO<sub>2</sub> emission from  
245 fossil fuels and the ratios to give us an estimate for CH<sub>4</sub> and N<sub>2</sub>O, about 355 kt/yr  
246 (1kt/yr=1000 ton/year) and 13 kt/year, which are lower than the values 789 kt/yr and 635 kt/yr  
247 respectively in Li et al.'s paper. The emissions of CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O increased significantly  
248 from 2000 to 2014 in Jinan, mainly because of vehicle emissions (Sun et al., 2016). We use  
249 the Sun et al.'s estimate of CO<sub>2</sub> and the ratio to estimate CO, about 1585 Gg/yr which is  
250 similar with the value 1723 Gg/yr in 2014. Anthropogenic emissions dominate during winter  
251 and spring. The temperature is very low in the winter and early spring, and the population  
252 density is high, so the total energy consumption, especially coal, is very large during winter  
253 and early spring (Zhao and Cui, 2014).

254

### 255 3.3.2. *Yellow Sea*

256 For the lower variability data the correlations are weak between every two gases. The  
257 highest R<sup>2</sup> is 0.68 between ΔCO and ΔCH<sub>4</sub> but compared to other source regions it is still  
258 weak. The average mole fractions enhancements of the four gases CO<sub>2</sub>, CO, CH<sub>4</sub>, N<sub>2</sub>O in the  
259 low-variability case were respectively 6.7 ppm, 110.4 ppb, 58.4 ppb, 0.28 ppb higher than the  
260 base values. The number of data points in this case is high. The air masses tended to stay over

261 the Yellow Sea, so that there were much fewer high variability cases. Mixing of air masses  
262 from different source regions, before our 72-hour back trajectories, may have weakened the  
263 correlations. As expected, the relationships between the average enhancements for CO<sub>2</sub>, CO,  
264 and CH<sub>4</sub> corresponded to the values above, 6.7 ppm, 110.4 ppb, 58.4 ppb respectively, but not  
265 for N<sub>2</sub>O. Were the ratios of the average enhancements caused by regional emissions in China,  
266 Korea, or do they came from larger areas as represented in MBL values?  $\overline{\Delta CO}/\overline{\Delta CO_2}=16.5$   
267 ppb/ppm,  $\overline{\Delta CH_4}/\overline{\Delta CO_2}= 8.7$  ppb/ppm, and  $\overline{\Delta CH_4}/\overline{\Delta CO}= 0.5$  ppb/ppb, whereas the MBL  
268 latitudinal gradients between ~54 N and 30 N in 2017 during the same time were 3.7 ppb/ppm,  
269 11.8 ppb/ppm, and 3.2 ppb/ppb respectively. Therefore the average enhancements are of  
270 regional origin.

271

### 272 3.3.3. *Korea/Yellow*

273 In the lower variability case we have significant correlations between the hourly data for  
274  $\Delta CO_2$ ,  $\Delta CO$ , and  $\Delta CH_4$ , but not for  $\Delta N_2O$ . The average mole fraction enhancements of the  
275 four gases CO<sub>2</sub>, CO, CH<sub>4</sub>, N<sub>2</sub>O are respectively 5.7 ppm, 105.4 ppb, 53.7 ppb, 0.28 ppb  
276 higher than the base values, similar to the Yellow Sea case. There was relatively less CO than  
277 in the Beijing case. Any influence from Korea appeared to be minor.

278

### 279 3.3.4. *Qingdao/Yantai/northeast*

280 The pair-wise correlations for the lower variability data look very similar to Beijing except  
281 for  $\Delta N_2O$ . All the R<sup>2</sup> values were above 0.8. The average mole fractions of the four gases CO<sub>2</sub>,  
282 CO, CH<sub>4</sub>, N<sub>2</sub>O are respectively 7.8 ppm, 127.5 ppb, 57.9 ppb, 0.61 ppb higher than the base

283 values. The air mass was mainly from northeast of China and then through Qingdao and  
284 Yantai. The back-trajectories fall in two groups. One group, the majority, stays mostly above 2  
285 km and a small group stays below 500 m. In this case we may have had fairly clean air that  
286 picked up pollution at the end. The atmospheric CO<sub>2</sub> and CH<sub>4</sub> content observed at  
287 Longfengshan regional station (Fig.9) located in the northeast of China is influenced all year  
288 by anthropogenic emissions, and the mole fractions of CO<sub>2</sub> and CH<sub>4</sub> are higher in winter and  
289 spring. (Fang et al., 2017). The higher variability data show the same pair-wise correlations as  
290 the lower variability data for  $\Delta$ CO<sub>2</sub>,  $\Delta$ CO, and  $\Delta$ CH<sub>4</sub>. The average enhancement of N<sub>2</sub>O  
291 appears to be the highest, 4.4 ppb. The emissions of anthropogenic sources, including  
292 municipal solid waste, increased much in the northern provinces. Moreover, the landfill sites  
293 have contributed more significantly to the N<sub>2</sub>O emissions in recent years in China (Du et al.,  
294 2017; Wang et al., 2017; Long et al., 2018).

295

#### 296 3.3.5. *Shanghai/Nanjing/Qingdao/northeast*

297 The correlation relationships were relatively strong, all of them have R<sup>2</sup> above 0.8. Again  
298 back-trajectories stay mostly above 2 km, but near the end they went through several big  
299 cities of China. The air mass mainly went through the Yangtze River Delta (YRD) economic  
300 development zone where the economy is developed and the population is dense. Moreover,  
301 industrial emissions are mainly in Jilin and Liaoning provinces located in the northeast of  
302 China. Therefore the anthropogenic source contributed most in this case.(Pu et al., 2012; Cai  
303 et al., 2018). Moreover, Lixin Liu et al. found that the CO value at Shangdianzi (SDZ) station  
304 which is located near Beijing was considerably higher than CO at Linan (LAN) station which

305 is located in the center of the YRD region in spring (Liu et al., 2018), which is consistent with  
306 the observation in this study. The average mole fraction enhancements of the four gases CO<sub>2</sub>,  
307 CO, CH<sub>4</sub>, N<sub>2</sub>O in the lower variability case are respectively 8.6 ppm, 113.6 ppb, 51.2 ppb,  
308 0.77 ppb higher than the base values. For the higher variability data CH<sub>4</sub> is especially higher  
309 than the base value in this case, about 111.3 ppb. N<sub>2</sub>O is also high, about 2.5 ppb higher than  
310 the base line. The correlation relationships are also relatively significant in the high variability  
311 case.

312

### 313 *3.3.6. East China Sea*

314 There are only lower variability data in this case. The R<sup>2</sup> values of the correlations between  
315 the hourly data of ΔCO<sub>2</sub>, ΔCH<sub>4</sub> and ΔCO were not as high, like in the Yellow Sea case. The  
316 average mole fraction enhancements of the four gases CO<sub>2</sub>, CO, CH<sub>4</sub>, N<sub>2</sub>O are respectively  
317 4.9 ppm, 63.7 ppb, 42.8 ppb, 0.41 ppb higher than their corresponding base values. The ratios  
318 between these average enhancements are similar to the Yellow Sea case, and are thus a  
319 regional signature, not strongly polluted by any single source region.

320

### 321 *3.3.7. Taiwan*

322 For the lower variability data some of the correlation relationships are good, the R<sup>2</sup> are  
323 above 0.8, but not for ΔCH<sub>4</sub>/ΔCO<sub>2</sub>, ΔCH<sub>4</sub>/ΔCO and ΔN<sub>2</sub>O/ΔCH<sub>4</sub>, with R<sup>2</sup> values of about 0.36,  
324 0.62 and 0.44 respectively. The average mole fraction enhancements of the four gases CO<sub>2</sub>,  
325 CO, CH<sub>4</sub>, N<sub>2</sub>O are respectively 8.6 ppm, 85.5 ppb, 44.6 ppb, 0.95 ppb higher than the base  
326 values. The air came from Taiwan. The emissions of N<sub>2</sub>O and CO<sub>2</sub> appear to be relatively

327 higher than the other two gas species. For the higher variability data the relationships of  
328  $\Delta N_2O/\Delta CO$  and  $\Delta CO/\Delta CH_4$  are well-defined. Compared to other source regions, the pollution  
329 in Taiwan appears to be less than in several big cities in China, such as Beijing, Shanghai. The  
330 research of Chang and Lee (Chang and Lee, 2007) has shown that the air quality in Taipei  
331 City improved since 1994 according to the observation data from the monitoring stations  
332 established by the Taiwan Environmental Protection Administration.

333

### 334 *3.3.8. Taiwan/Fuzhou/Wenzhou*

335 There are much fewer lower variability data than higher variability data. The correlations of  
336 the data in this case are weak whether they are low or high variability data. The pollution  
337 came from several big cities along the coast. The contaminated air mass was not mixed evenly  
338 when it reached the measured point. The enhancement of  $CO_2$  in this case is just below that  
339 from the source region of Beijing/Jinan.

340

### 341 *3.3.9. Shanghai/Nanjing*

342 There are only lower variability data in pair-wise comparisons of  $\Delta CO_2$ ,  $\Delta CH_4$ ,  $\Delta CO$  and  
343 there are not enough data for  $\Delta N_2O$ . The average mole fraction enhancements of the gases  
344  $CO_2$ ,  $CO$ ,  $CH_4$  are respectively 10.2 ppm, 105.4 ppb, 55.0 ppb higher than the base values.  
345 The correlations of  $\Delta CO_2/\Delta CO$ ,  $\Delta CO/\Delta CH_4$  are much better. The higher variability data only  
346 appear in the  $\Delta N_2O/\Delta CO$  case, but there are only 6 data points. The correlation of  $\Delta N_2O/\Delta CO$   
347 is weak. There are high spikes of  $CO$  without high  $N_2O$ . In this case, the pollution was more  
348 concentrated.

349

350 *3.3.10. North (Russia)*

351 The correlations are not strong for the lower variability case. The air mass came from the  
352 North (Russia) and it also came from altitudes about 2 km. Near the end situation the altitude  
353 became low, and the air mass may have been influenced by some cities in China. The average  
354 mole fraction enhancements of the four gases CO<sub>2</sub>, CO, CH<sub>4</sub>, N<sub>2</sub>O were respectively 8.6 ppm,  
355 45.0 ppb, 52.0 ppb, 0.035 ppb higher than the base values. The ratios of the averages are on  
356 average closer to north-south gradients observed in the MBL, compared to the Yellow Sea and  
357 East China Sea. Compared with other source regions, the mole fractions of CO and N<sub>2</sub>O in  
358 this case were the lowest.

359

360 **4. Conclusions**

361

362 Using the NOAA HYSPLIT model, the data were assigned to ten different source regions,  
363 which were mainly in North China and coastal cities of East China. We used two different  
364 methods to calculate the enhancement ratios of every pair of the gases and found that the  
365 enhancement ratio of  $\Delta\text{CH}_4/\Delta\text{CO}$  is lower, and the enhancement ratio of  $\Delta\text{CO}/\Delta\text{CO}_2$  is much  
366 higher in big cities like Beijing and Shanghai mainly because of the high CO emissions in  
367 these source regions in winter and early spring.

368 Compared with the MBL values of CO<sub>2</sub>, CO, CH<sub>4</sub> and N<sub>2</sub>O during the same time, the ratios  
369 of the average enhancements when the air had stayed over the Yellow Sea and the East China  
370 Sea for several days were caused by the emissions from the general region of Eastern China in

371 the spring. In this case the correlations of pairs of gases in every individual sample are weaker  
372 than for the urban source regions. For the source region called North (Russia) the  
373 enhancements ratios of the averages were on average closer to the MBL.

374 The methods used to calculate the enhancement ratios and the uncertainties of the ratios of  
375 the enhancements of every pair of gases in this study can be used to compare with emissions  
376 inventories as a completely independent check. Moreover, the observation data and the results  
377 can be used for multi-species inverse estimates of the sources and sinks of greenhouse gases  
378 (Pison et al., 2009; Wada et al., 2011). However, we want to emphasize that our observations  
379 of enhancement ratios do not depend on inverse modeling. Carefully calibrated data stand on  
380 their own, and will be “forever” if data management (archival) is handled correctly. We  
381 recommend that our approach be carried out in other seasons, to provide more comprehensive  
382 comparisons with emissions inventories.

383

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385

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395

396

397 **References**

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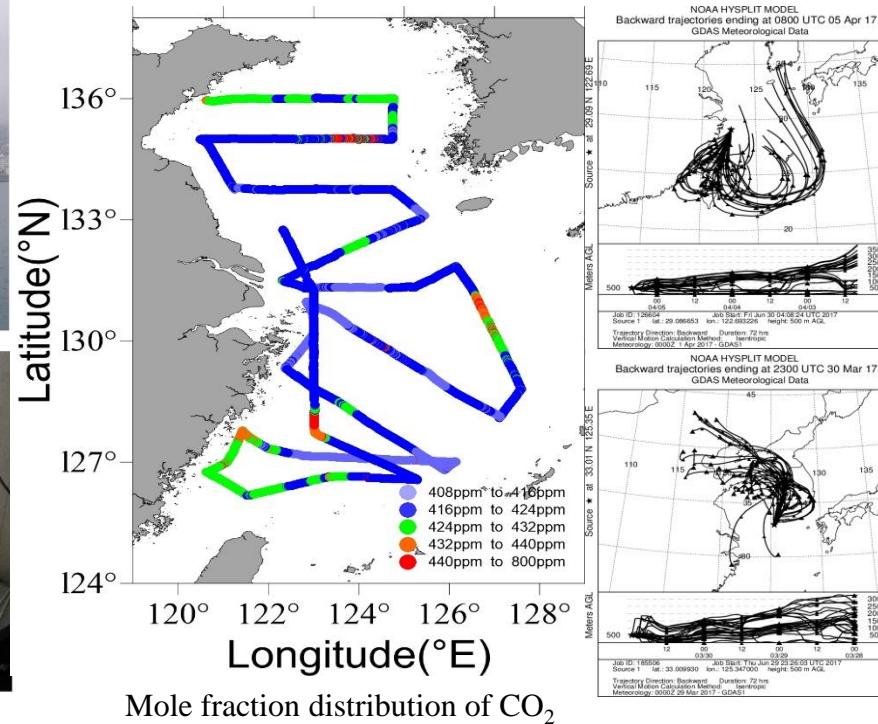
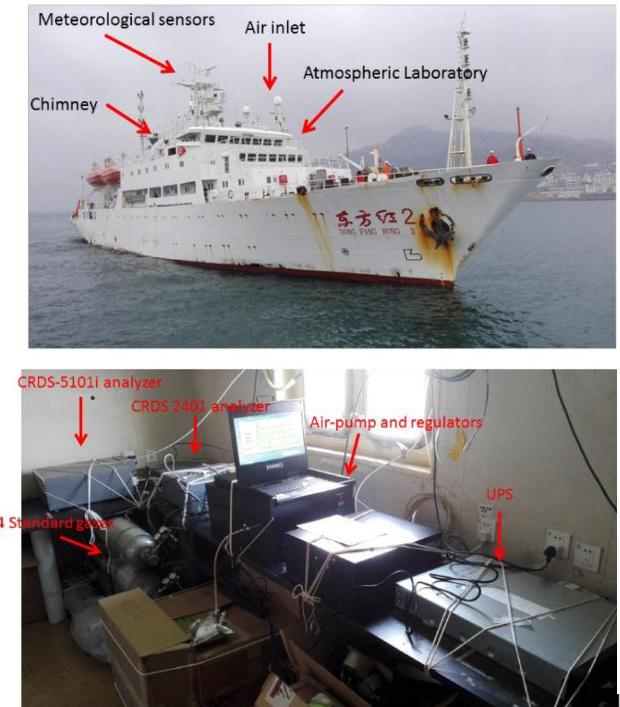
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520



A ship-board continuously observation system was applied to observe the atmospheric distributions of CO<sub>2</sub>, CO, CH<sub>4</sub> and N<sub>2</sub>O over the Yellow Sea and the East China Sea and we obtained the high-precision data to analyze the source regions through the NOAA HYSPLIT Model and study the emission characteristics of these areas.