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**Iterative Methods for Solving the Nonlinear Balance Equation
with Optimal Truncation**

Qin XU^{1*} and Jie CAO^{2,3}

¹*NOAA/National Severe Storms Laboratory, Norman, Oklahoma, USA*

²*Cooperative Institute for Mesoscale Meteorological Studies, University of Oklahoma,
Norman, Oklahoma, USA*

³*Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029*

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* Corresponding author: Qin Xu
E-mail: Qin.Xu@noaa.gov

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44 **Article Highlights:**

45 ● Two previous iterative methods for solving the NBE are re-derived by expanding the
46 solution asymptotically upon a small Rossby number Ro .

47 ● A criterion for optimal truncation of asymptotic expansion is proposed to obtain the super-
48 asymptotic approximation of the solution.

49 ● Using the integral-form Poisson solver for the boundary value problem in each iteration,
50 optimally truncated solutions can be obtained efficiently with improved accuracies.

51 ● Solution errors can be reduced effectively even when Ro increases to 0.4 for cyclonically
52 curved jet flows of meso- α scale.

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54

55 **1. Introduction**

56 For flows of synoptic and sub-synoptic scales in the middle and high latitudes, the
57 nonlinear balance equation (NBE) links the streamfunction field with the geopotential field
58 more accurately than the geostrophic balance (Charney, 1955; Bolin, 1955). However, solving
59 the streamfunction from the NBE for a given geopotential field can be very challenging due to
60 complicated issues on the existence of solution in conjunction with difficulties caused by
61 nonlinearity (Courant and Hilbert, 1962). It is well known mathematically that the NBE is a
62 special case of the Monge-Ampere differential equation for the streamfunction (Charney, 1955).
63 If the geostrophic vorticity (that is, the vorticity of geostrophic flow associated with the given
64 geopotential field) is larger than $-f/2$ for a constant f where f is the Coriolis parameter, then the
65 NBE is of the elliptic type and its associated boundary value problem can have no more than
66 two solutions (see Section 6.3 in Chapter 4 of Courant and Hilbert, 1962). If the geostrophic
67 vorticity is smaller than $-f/2$ in a local area, then the NBE becomes locally hyperbolic. In this
68 case, the boundary value problem becomes ill-posed and thus may have no solution although
69 the NBE can be integrated along the characteristic lines within the locally hyperbolic area (see
70 Section 3 of Appendix I in Chapter 5 of Courant and Hilbert, 1962).

71 To avoid the complication and difficulties caused by the local non-ellipticity in solving the
72 NBE, one can simply enforce the ellipticity condition to a certain extent by slightly smoothing
73 or adjusting the given geopotential field. This type of treatment has been commonly used in
74 previously developed iterative methods to solve the NBE as a boundary value problem (Bolin,
75 1955, 1956; Shuman, 1955, 1957; Miyakoda, 1956; Bushby and Huckle, 1956; Arnason, 1958;

76 Bring and Charasch, 1958; Liao and Chow, 1962; Asselin, 1967; Paegle and Tomlinson, 1975;
77 Bijlsma and Hoogendoorn, 1983). However, regardless of the above treatment, the convergence
78 properties of the previous iterative methods or any iterative methods can be not only scale-
79 dependent but also flow-dependent and thus very difficult to study theoretically and rigorously.

80 The above reviewed previous iterative methods can be classified into two types. In the first
81 type (originally proposed by Bolin, 1955), the NBE is transformed into a linearized equation
82 for a presumably small correction to the initial guess or to the subsequent updated solution
83 when this linearized equation is solved iteratively. In the second type (originally proposed by
84 Shuman, 1955, 1957; Miyakoda, 1956), the NBE is rearranged into a quadratic form of the
85 absolute vorticity and the positive root of this quadratic form is used in the form of Poisson
86 equation to solve for the streamfunction iteratively. The initial guess for both types is the
87 geostrophic streamfunction. Their convergence properties were analyzed theoretically, but the
88 analysis was lack of rigor and generality, because the coefficients of linearized differential
89 operator for the first type and the forcing terms on the right-hand side of the iterative form of
90 linearized equation for the second type were functions of space but treated as constants
91 (Arnason, 1958; Bijlsma and Hoogendoorn, 1983). Therefore the convergence properties of the
92 previously iterative methods were examined mainly through numerical experiments. Besides,
93 due to the very limited computer memories and speed in those early decades, the previous
94 iterative methods employed the memory-saving sequential relaxation scheme based on the
95 classical Liebmann-type iteration algorithm (Southwell, 1946) and applied to coarse resolution
96 grids for large-scale flows. The sequential relaxation and successive over-relaxation (SOR)

97 schemes have been used in the second type of iterative method (Shuman, 1955, 1957) to solve
98 the NBE for hurricane flows (Zhu et al., 2002). However, using the previous iterative methods
99 to solve the NBE still faces various difficulties especially when the spatial scale reduces to and
100 even below the sub-synoptic scale. In particular, there are unaddressed challenging issues
101 concerning whether and how the solutions can be obtained approximately and efficiently
102 through limited numbers of iterations, especially when the NBE becomes locally hyperbolic
103 (due mainly to reduced spatial scales) and thus the iterative methods fail to converge.

104 This paper aims to address the above concerned challenging issues. In particular, we will
105 re-derive the above two types of iterative methods formally and systematically by expanding
106 the solution asymptotically upon a small Rossby number and substituting it into the NBE. Since
107 the asymptotic expansion is not ensured to converge especially when the Rossby number is not
108 sufficiently small, the concept of optimal truncation of asymptotic expansion is employed and
109 a criterion is proposed for optimal truncation to obtain the super-asymptotic approximation of
110 the solution based on the heuristic theory of asymptotic analysis (Boyd, 1999). As will be seen
111 in this paper, by employing the optimal truncation, the issue on non-convergence of the iterative
112 methods caused by the increase of Rossby number can be addressed to a certain extent. Besides,
113 the recently developed Poisson solver based on integral formulas (Xu et al., 2011; Cao and Xu,
114 2011) will be used in comparison with the aforementioned classical SOR scheme to solve the
115 boundary value problem in each iterative step. In particular, for flows of sub-synoptic scale or
116 meso- α scale, the NBE can become locally hyperbolic and the solution will be checked in this
117 paper via the proposed optimal truncation under certain conditions.

118 The paper is organized as follows. The next section presents formal and systematical
119 derivations of the above reviewed two iterative methods. Section 2 formulates the criterion for
120 optimal truncation, and section 3 constructs four different iterative procedures with optimal
121 truncation and designs numerical experiments for testing the iterative procedures. Section 4
122 examines and compares the results of experiments performed with the four iterative procedures,
123 followed by conclusions in section 5.

124

125 **2. Derivations of two iterative methods**

126 ***2.1 Scaling and asymptotic expansion based on small Rossby number***

127 The NBE can be written into the following form (Charney, 1955):

128

$$129 \quad N(\psi) = \nabla^2 \phi, \tag{1a}$$

130

131 where $N(\psi) \equiv \nabla^2 \psi + (\nabla \psi) \cdot \nabla f + 2J_{xy}(\partial_x \psi, \partial_y \psi) = \nabla \cdot (f \nabla \psi) + 2J_{xy}(\partial_x \psi, \partial_y \psi) = \nabla^2(f\psi) - \nabla \cdot (\psi \nabla f)$
132 $+ 2J_{xy}(\partial_x \psi, \partial_y \psi)$, ψ is the streamfunction, ϕ is the geopotential, $\nabla \equiv (\partial_x, \partial_y)$, $\nabla^2 \equiv \nabla \cdot \nabla = \partial_x^2 +$
133 ∂_y^2 , and $J_{xy}(\partial_x \psi, \partial_y \psi) \equiv (\partial_x^2 \psi)(\partial_y^2 \psi) - (\partial_x \partial_y \psi)^2$. For large-scale and synoptic-scale flows, the
134 geostrophic approximation, $\nabla^2 \phi \approx \nabla^2(f\psi)$, is the leading-order balance in (1a) and thus $\nabla^2(f\psi)$
135 is the dominant term in $N(\psi)$. In this case, the boundary condition for solving ψ from (1a) over
136 a middle-latitude domain D can be given by

137

$$138 \quad \psi = \psi_g \quad \text{on } \partial D, \tag{1b}$$

139

140 where ∂D denotes the domain boundary, and $\psi_g \equiv \phi/f$ is the global geostrophic streamfunction

141 (Kuo, 1959; Charney and Stern, 1962; Schubert et al., 2009).

142 Formally, we can scale x and y by L , scale $f = f_0 + f'$ by f_0 , and scale ψ and ϕ by UL and
143 f_0UL , respectively, where U is the horizontal velocity scale, L is the horizontal length scale, f_0
144 is a constant reference value of f which can be the value of f at the domain center. The scaled
145 variables are still denoted by their respectively original symbols, so the NBE can have the
146 following non-dimensional form:

147

$$148 \quad \nabla^2(f\psi - \phi) = \nabla \cdot (\psi \nabla f) - R_0 2J_{xy}(\partial_x \psi, \partial_y \psi), \quad (2)$$

149

150 where $R_0 \equiv UL/f_0L$ is the Rossby number. For synoptic-scale and sub-synoptic-scale flows, the
151 above scaling can give $R_0 = \varepsilon < 1$. Substituting this into (2) gives

152

$$153 \quad \nabla^2(f\psi - \phi) = \varepsilon [\nabla \cdot (\psi \nabla F) - 2J_{xy}(\partial_x \psi, \partial_y \psi)], \quad (3)$$

154

155 where $F = f'/(f_0R_0) \leq O(1)$ and $O(\cdot)$ is the ‘order-of-magnitude’ symbol. Thus, ψ can have the
156 following asymptotic expansion:

157

$$158 \quad \psi = \psi_0 + \sum_{\perp} \varepsilon^k \delta \psi_k, \quad (4)$$

159

160 where $\psi_0 = \psi_g$ and \sum_{\perp} denotes the summation over k from 1 to ∞ . The k^{th} order truncation of
161 the asymptotic expansion of ψ in (4) is given by $\psi_k \equiv \psi_0 + \sum_{\perp}^k \varepsilon^{k'} \delta \psi_{k'}$, where \sum_{\perp}^k denotes the
162 summation over k' from 1 to k . Formally, $\psi = \psi_k + O(\varepsilon^{k+1})$, so ψ_k is accurate up to $O(\varepsilon^k)$ as an
163 approximation of ψ .

164 By substituting (4) into (3) and (1b), and then collecting terms of the same order of ε , we

165 obtain

166

167 $\nabla^2(f\delta\psi_1) = \nabla \cdot (\psi_0 \nabla F) - 2J_{xy}(\partial_x \psi_0, \partial_y \psi_0),$

168 $\nabla^2(f\delta\psi_2) = \nabla \cdot (\delta\psi_1 \nabla F) - 2[J_{xy}(\partial_x \psi_0, \partial_y \delta\psi_1) + J_{xy}(\partial_x \delta\psi_1, \partial_y \psi_0)],$

169 $\nabla^2(f\delta\psi_3) = \nabla \cdot (\delta\psi_2 \nabla F) - 2[J_{xy}(\partial_x \psi_0, \partial_y \delta\psi_2) + J_{xy}(\partial_x \delta\psi_2, \partial_y \psi_0) + J_{xy}(\partial_x \delta\psi_1, \partial_y \delta\psi_1)],$

170 $\nabla^2(f\delta\psi_4) = \nabla \cdot (\delta\psi_3 \nabla F) - 2[J_{xy}(\partial_x \psi_0, \partial_y \delta\psi_3) + J_{xy}(\partial_x \delta\psi_3, \partial_y \psi_0)$

171 $+ J_{xy}(\partial_x \delta\psi_1, \partial_y \delta\psi_2) + J_{xy}(\partial_x \delta\psi_2, \partial_y \delta\psi_1)],$

172 ...

173 $\nabla^2(f\delta\psi_{k+1}) = \nabla \cdot (\delta\psi_k \nabla F) - 2[J_{xy}(\partial_x \psi_0, \partial_y \delta\psi_k) + J_{xy}(\partial_x \delta\psi_k, \partial_y \psi_0) + J_{xy}(\partial_x \delta\psi_1, \partial_y \delta\psi_{k-1})$

174 $+ J_{xy}(\partial_x \delta\psi_{k-1}, \partial_y \delta\psi_1) + J_{xy}(\partial_x \delta\psi_2, \partial_y \delta\psi_{k-2}) + J_{xy}(\partial_x \delta\psi_{k-2}, \partial_y \delta\psi_2) + \dots],$

175 ... (5a)

176 $\delta\psi_k = 0$ on ∂D for $k = 1, 2, 3, \dots$ (5b)

177

178 Here, (5) gives a formal series of linearized equations and boundary conditions for computing

179 $\delta\psi_k$ consecutively from $\delta\psi_1$ to increasingly higher-order term in the expansion of ψ in (5). The

180 equations in (5a), however, are inconvenient to use, because the equation at each given order

181 becomes increasingly complex as the order k increases. It is thus desirable to modify (5a) into

182 a recursive form, and this can be done non-uniquely by first combining the equations in (5a)

183 with $\nabla^2(f\psi_0) = \nabla^2\phi$ into a series of equations for ψ_k (instead of $\delta\psi_k$) and then adding properly

184 selected higher-order terms to the equation for ψ_k at each order without affecting the order of

185 accuracy of the equation. In particular, two different modifications will be made in the next

186 two subsections. From these two modifications, the two types of iterative methods reviewed in

187 the introduction for solving the NBE can be derived formally and systematically via the
 188 asymptotic expansion of ψ in (4).

189

190 **2.2 Derivation of method-1**

191 The equations in (5a) can be combined with $\nabla^2(f\psi_0) = \nabla^2\phi$ at $O(\varepsilon^0)$ into a series of equations
 192 for ψ_k defined in (4) as shown blow:

193

$$194 \quad \nabla^2(f\psi_1) = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_0\nabla F) - 2\varepsilon J_{xy}(\partial_x\psi_0, \partial_y\psi_0),$$

$$195 \quad \nabla^2(f\psi_2) = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_1\nabla F) - 2\varepsilon[J_{xy}(\partial_x\psi_0, \partial_y\psi_0) + \varepsilon J_{xy}(\partial_x\psi_0, \partial_y\delta\psi_1) + \varepsilon J_{xy}(\partial_x\delta\psi_1, \partial_y\psi_0)]$$

$$196 \quad = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_1\nabla F) - 2\varepsilon[J_{xy}(\partial_x\psi_1, \partial_y\psi_1) - \varepsilon^2 J_{xy}(\partial_x\delta\psi_1, \partial_y\delta\psi_1)]$$

$$197 \quad = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_1\nabla F) - 2\varepsilon J_{xy}(\partial_x\psi_1, \partial_y\psi_1) + O(\varepsilon^3),$$

$$198 \quad \nabla^2(f\psi_3) = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_2\nabla F) - 2\varepsilon[J_{xy}(\partial_x\psi_1, \partial_y\psi_1) + \varepsilon^2 J_{xy}(\partial_x\psi_0, \partial_y\delta\psi_2) + \varepsilon^2 J_{xy}(\partial_x\delta\psi_2, \partial_y\psi_0)]$$

$$199 \quad = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_2\nabla F) - 2\varepsilon[J_{xy}(\partial_x\psi_1, \partial_y\psi_1) + \varepsilon^2 J_{xy}(\partial_x\psi_1, \partial_y\delta\psi_2) + \varepsilon^2 J_{xy}(\partial_x\delta\psi_2, \partial_y\psi_1)$$

$$200 \quad \quad \quad - \varepsilon^3 J_{xy}(\partial_x\delta\psi_1, \partial_y\delta\psi_2) - \varepsilon^3 J_{xy}(\partial_x\delta\psi_2, \partial_y\delta\psi_1)]$$

$$201 \quad = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_2\nabla F) - 2\varepsilon[J_{xy}(\partial_x\psi_2, \partial_y\psi_2) - \varepsilon^3 J_{xy}(\partial_x\delta\psi_1, \partial_y\delta\psi_2) - \varepsilon^3 J_{xy}(\partial_x\delta\psi_2, \partial_y\delta\psi_1)$$

$$202 \quad \quad \quad - \varepsilon^4 J_{xy}(\partial_x\delta\psi_2, \partial_y\delta\psi_2)]$$

$$203 \quad = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_2\nabla F) - 2\varepsilon J_{xy}(\partial_x\psi_2, \partial_y\psi_2) + O(\varepsilon^4),$$

204 ...

$$205 \quad \nabla^2(f\psi_k) = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_{k-1}\nabla F) - 2\varepsilon J_{xy}(\partial_x\psi_{k-1}, \partial_y\psi_{k-1}) + O(\varepsilon^{k+1}),$$

206 ...

(6)

207

208 Formally ψ_k is accurate up to $O(\varepsilon^k)$ and so is $\nabla^2(f\psi_k)$ on the left-hand side of the above k^{th}

209 equation. This implies that the k^{th} equation is accurate only up to $O(\varepsilon^k)$, so the last term $O(\varepsilon^{k+1})$

210 (that represents all the high-order terms) on the right-hand side can be neglected without
 211 degrading the order of accuracy of the equation. This leads to the following recursive form of
 212 equation and boundary condition for solving the NBE iteratively:

$$214 \quad \nabla^2(f\psi_k) = \nabla^2\phi + \varepsilon\nabla\cdot(\psi_{k-1}\nabla F) - 2\varepsilon J_{xy}(\partial_x\psi_{k-1}, \partial_y\psi_{k-1}), \quad (7a)$$

$$215 \quad \psi_k = \psi_g \quad \text{on } \partial D \quad \text{for } k = 1, 2, 3, \dots \quad (7b)$$

216
 217 If ε is sufficiently small to ensure the convergence of the asymptotic expansion in (5), then ψ_k
 218 $\rightarrow \psi$ gives the solution of the NBE in the limit of $k \rightarrow \infty$.

219 Substituting $\varepsilon\nabla F = \nabla f f_0$ and $\varepsilon = R_0 \equiv U/f_0 L$ into (7) gives the dimensional form of (7):

$$221 \quad \nabla^2(f\psi_k) = \nabla^2\phi + \nabla\cdot(\psi_{k-1}\nabla f) - 2J_{xy}(\partial_x\psi_{k-1}, \partial_y\psi_{k-1}), \quad (8a)$$

$$222 \quad \psi_k = \psi_g \quad \text{on } \partial D. \quad (8b)$$

223
 224 For $f = \text{constant}$, (8a) recovers (5) of Bushby and Huckle (1956), but this recursive form of
 225 equation is derived here formally and systematically via the asymptotic expansion of the
 226 solution in (4). Substituting the dimensional form of $\psi_k = \psi_{k-1} + \varepsilon^k \delta\psi_k$, that is, $\psi_k = \psi_{k-1} + \delta\psi_k$
 227 into (8) gives

$$229 \quad \nabla^2(f\delta\psi_k) = \nabla^2\phi - N(\psi_{k-1}), \quad (9a)$$

$$230 \quad \delta\psi_k = 0 \quad \text{on } \partial D \quad \text{for } k = 1, 2, 3, \dots \quad (9b)$$

231
 232 where $N(\)$ is the nonlinear differential operator defined in (1a). Analytically, (9a) is identical
 233 to (8a) but expressed in an incremental form. Numerically, however, solving $\delta\psi_k$ from (9) and

234 updating ψ_{k-1} to $\psi_k = \psi_{k-1} + \delta\psi_k$ iteratively does not give exactly the same solution as that
 235 obtained by solving ψ_k from (8) iteratively. According to our additional numerical experiments
 236 (not shown), the solutions obtained from (8) are less accurate (by about an order of magnitude
 237 for the case of $R_o = 0.1$) than their counterpart solutions obtained from (9), so the non-
 238 incremental form of boundary value problem in (8) will not be considered in this paper.

239

240 **2.3 Derivation of method-2**

241 The equation for ψ_k in (7) can be multiplied by 2 and rewritten into

242

$$243 \quad 2f\nabla^2\psi_k = 2\nabla^2\phi - 2\varepsilon(\nabla F)\cdot\nabla\psi_{k-1} - 4\varepsilon J_{xy}(\partial_x\psi_{k-1}, \partial_y\psi_{k-1}) + O(\varepsilon^{k+1}), \quad (10)$$

244

245 where $\psi_k = \psi_{k-1} + \varepsilon^k\delta\psi_k = \psi_{k-1} + O(\varepsilon^k)$ and $\nabla^2(f\psi_k) = f\zeta_k + (\nabla f)\cdot\nabla\psi_k + \nabla\cdot(\psi_k\nabla f) = f\zeta_k +$
 246 $\varepsilon(\nabla F)\cdot\nabla\psi_k + \varepsilon\nabla\cdot(\psi_k\nabla F) = f\zeta_k + \varepsilon(\nabla F)\cdot(\nabla\psi_{k-1}) + \varepsilon\nabla\cdot(\psi_{k-1}\nabla F) + O(\varepsilon^{k+1})$ are used. One can verify
 247 that $-4\varepsilon J_{xy}(\partial_x\psi_k, \partial_y\psi_k) = \varepsilon(\zeta_k^2 - A_k^2 - B_k^2) = \varepsilon\zeta_k^2 - \varepsilon(A_{k-1}^2 + B_{k-1}^2) + O(\varepsilon^{k+1})$ where $\zeta_k = \nabla^2\psi_k$, A_k
 248 $\equiv (\partial_x^2 - \partial_y^2)\psi_k$, $B_k \equiv 2\partial_x\partial_y\psi_k$, and $A_k = (\partial_x^2 - \partial_y^2)(\psi_{k-1} + \varepsilon^k\delta\psi_k) = A_{k-1} + O(\varepsilon^k)$ and $B_k = 2\partial_x\partial_y(\psi_{k-1} + \varepsilon^k\delta\psi_k) = B_{k-1} + O(\varepsilon^k)$ are used. Substituting these into (10) gives

249

$$250 \quad \varepsilon\zeta_k^2 + 2f\zeta_k + 2\varepsilon(\nabla F)\cdot\nabla\psi_{k-1} - 2\nabla^2\phi - \varepsilon A_{k-1}^2 - \varepsilon B_{k-1}^2 = O(\varepsilon^{k+1}).$$

251

252 This leads to the following recursive form of equation that is accurate up to $O(\varepsilon^k)$:

253

$$254 \quad \varepsilon\zeta_k^2 + 2f\zeta_k + 2\varepsilon(\nabla F)\cdot\nabla\psi_{k-1} - 2\nabla^2\phi - \varepsilon A_{k-1}^2 - \varepsilon B_{k-1}^2 = 0. \quad (11a)$$

255

256 Substituting $\varepsilon\nabla F = \nabla f/f_o$ and $\varepsilon = R_o \equiv U/f_o L$ into (11a) gives its dimensional form which

258 can be rewritten into

259

260
$$(f + \zeta_k)^2 = M_{k-1} \equiv f^2 + 2\nabla^2\phi + A_{k-1}^2 + B_{k-1}^2 - 2(\nabla f) \cdot \nabla \psi_{k-1}. \quad (11b)$$

261

262 The non-negative condition of $(f + \zeta_k)^2 \geq 0$ requires $M_{k-1} \geq 0$ on the right-hand side of (11b).

263 Also, as a quadratic equation of $f + \zeta_k$ for given ϕ and ψ_{k-1} , (11b) has two roots, but only the

264 positive root, given by $f + \zeta_k = M_{k-1}^{1/2}$, is physically acceptable (because $f + \zeta_k \geq 0$ is required

265 for stably balanced flow). This leads to the following recursive form of equation and boundary

266 condition for solving the NBE iteratively:

267

268
$$\nabla^2 \psi_k = -f + M_{k-1}^{1/2}, \quad (12a)$$

269
$$\psi_k = \psi_g \quad \text{on } \partial D \quad \text{for } k = 1, 2, 3, \dots \quad (12b)$$

270

271 where $M_{k-1} \geq 0$ is ensured by setting $M_{k-1} = 0$ when the computed M_{k-1} from the previous step

272 becomes negative. Here, (12a) gives essentially the same recursive form of equation as that in

273 (8) of Shuman (1957) for solving the NBE iteratively, but this recursive form of equation is

274 derived here via the asymptotic expansion of the solution in (4).

275

276 **3. Iterative procedures with optimal truncation and experiment design**

277 **3.1 Criterion for optimal truncation**

278 When the Rossby number is not sufficiently small to ensure the convergence of the

279 asymptotic expansion, the optimal truncation of the asymptotic expansion of ψ in (4) can be

280 determined (Boyd, 1999) by an empirical criterion in the following dimensional form:

281

282
$$E[N(\psi_K)] = \min E[N(\psi_k)] \text{ for } k = K, K \pm 1, \dots, K \pm m, \quad (13)$$

283

284 where $N(\cdot)$ is the function form defined in (1a), K is the number of optimal truncation, $E[N(\psi_k)]$

285 $\equiv \|\varepsilon[N(\psi_k)]\|$, $\|\cdot\|$ denotes the root-mean-square (RMS) of discretized field of the variable

286 inside $\|\cdot\|$ computed over all the interior grid points (excluding the boundary points) of domain

287 D , and $\varepsilon[N(\psi_k)] \equiv [N(\psi_k) - N(\psi_t)]/\|N(\psi_t)\| = [N(\psi_k) - \nabla^2 \phi]/\|\nabla^2 \phi\|$ is the relative error of $N(\psi_k)$

288 with respect to $N(\psi_t)$ which is also the normalized (by $\|\nabla^2 \phi\|$) residual error of the NBE caused

289 by the approximation of $\psi \approx \psi_k$, and ψ_t denotes the true solution. Here, $E[N(\psi_K)]$ is expected

290 to be the global minimum of $E[N(\psi_k)]$. If $E[N(\psi_k)]$ does not oscillate as k increases, then it is

291 sufficient to set $m = 1$ in (13). Otherwise, m should be sufficiently large to ensure $E[N(\psi_K)]$ be

292 the global minimum of $E[N(\psi_k)]$.

293

294 **3.2 Iterative procedures**

295 The iterative procedure for method-1 performs the following steps:

296 1. Start from $k = 0$ and set $\psi_0 = \psi_g \equiv \phi/f$ in D and ∂D .

297 2. Substitute ψ_{k-1} ($= \psi_0$ for $k = 1$) into $N(\psi_{k-1})$ to compute the right-hand-side of (9a), and then

298 solve the boundary value problem in (9) for $\delta\psi_k$.

299 3. Substitute $\psi_k = \psi_{k-1} + \alpha\delta\psi_k$ into $\|N(\psi_k) - \nabla^2 \phi\|$ and save the computed $\|N(\psi_k) - \nabla^2 \phi\|$ where

300 α is an adjustable parameter in the range of $0 < \alpha \leq 1$.

301 4. If $k \geq 2m$, then find $\min\|N(\psi_{k'}) - \nabla^2 \phi\|$, say at $k' = K'$, for $k' = k, k - 1, \dots, k - 2m$. If $K' < k -$

302 m , then $K = K'$ and ψ_K gives the optimally truncated solution – the final solution that ends the

303 iteration. Otherwise, go back to step 2.

304 When the Poisson solver (or SOR scheme) is used to solve boundary value problem in the
305 above step 2, the iterative procedure designed for method-1 is named M1a (or M1b). For the
306 Poisson solver used in this paper, the internally induced solution is obtained by using the
307 scheme S2 described in section 2.1 of Cao and Xu (2011) and the externally induced solution
308 obtained by using the Cauchy integral method described in section 4.1 of Cao and Xu (2011).
309 For M1a with $Ro < 0.4$ (or $Ro = 0.4$), it is sufficient to set $m = 1$ and $\alpha = 1$ (or $1/2$). For M1b,
310 it is sufficient to set $m = 3$ and $\alpha = 1$.

311 The iterative procedure for method-2 performs the following steps:

- 312 1. Start from $k = 0$ and set $\psi_0 = \psi_g \equiv \phi/f$ in D and ∂D .
- 313 2. Substitute ψ_{k-1} into M_{k-1} defined in (11b) to compute the right-hand-side of (12a), and then
314 solve the boundary value problem in (12) for ψ_k .
- 315 3. Compute and save $\|N(\psi_k) - \nabla^2 \phi\|$.
- 316 4. Perform this step as described above for step 4 of method-1.

317 When the Poisson solver (or SOR scheme) is used to solve boundary value problem in the
318 above step 2, the iterative procedure designed for method-2 is named M2a (or M2b). For M2a
319 and M2b, it is sufficient to set $m = 1$ and $\alpha = 1$.

320

321 **3.3 Experiment design**

322 To examine and compare the accuracies and computational efficiencies of the four iterative
323 procedures, the true streamfunction field is formulated for a wavering jet flow by

324

$$325 \quad \psi_t = -0.5UL \tanh[2y/L + 0.5 \cos(\pi x'/L)] \quad (14)$$

326

327 and the associated velocity components are given by

328

329
$$u_t \equiv -\partial_y \psi_t = U \operatorname{sech}^2[2y/L + 0.5\cos(\pi x'/L)] \quad (15a)$$

330 and
$$v_t \equiv \partial_x \psi_t = 0.25\pi U \sin(\pi x'/L) \operatorname{sech}^2[2y/L + 0.5\cos(\pi x'/L)], \quad (15b)$$

331

332 where $U = 20 \text{ ms}^{-1}$ is the maximum zonal speed of the wavering jet flow, $y = -0.25L\cos(\pi x'/L)$

333 is the longitudinal location (in y -coordinate) of the wavering jet axis as a function of $x' = x$

334 $-x_0$, and x_0 is the zonal location of wave ridge. By setting the half-wavelength L to 2000, 1000

335 and 500 km, the flow fields formulated in (14) and (15) resemble wavering westerly jet flows

336 on the synoptic, sub-synoptic and meso- α scales, respectively (as often observed on northern-

337 hemisphere mid-latitude 500 hPa weather maps).

338 Four sets of experiments are designed to test and compare the iterative procedures with ψ_t

339 given in (14) over a square domain of $D \equiv [-L \leq x \leq L, -L \leq y \leq L]$. The first set consists of four

340 experiments to test the four iterative procedures (that is, M1a, M1b, M2a and M2b) on the

341 synoptic scale by setting $L = 2000 \text{ km}$ and $x_0 = 0$ for ψ_t in (14). The second set also consists of

342 four experiments but to test the four iterative procedures on the sub-synoptic scale by setting L

343 $= 1000 \text{ km}$ and $x_0 = 0$ for ψ_t in (14). The third (or fourth) set still consists of four experiments

344 to test the four iterative procedures on the meso- α scale by setting $L = 500 \text{ km}$ and $x_0 = 0$ (or

345 L) for ψ_t in (14). Note that setting $x_0 = 0$ (or L) places the ridge (or trough) of the wavering jet

346 in the middle of domain D , so the nonlinearly balanced flow used for the tests in the third (or

347 fourth) set is curved anti-cyclonically (or cyclonically) in the middle of domain D . For

348 simplicity, the Coriolis parameter f is assumed to be constant and set to $f = f_0 = 10^{-4} \text{ s}^{-1}$ in all

349 the experiments. The Rossby number, defined by $Ro = U/f_oL$, is thus 0.1, 0.2 and 0.4 for $L =$
350 2000, 1000 and 500 km, respectively.

351 The true geopotential field, ϕ , is obtained by solving the Poisson equation, $\nabla^2\phi = N(\psi_t)$,
352 numerically on a 51×51 grid over domain D with the boundary condition given by $\phi = f\psi_t$. In
353 this case, ψ_t in (14) is also discretized on the same 51×51 grid over the same square domain,
354 and is used to compute the right-hand side of $\nabla^2\phi = N(\psi_t)$ via standard finite-differencing. Then,
355 ϕ is solved numerically by using the Poisson solver of Cao and Xu (2011). The SOR scheme
356 can be also used to solve for ϕ , but the solution is generally less accurate than that obtained by
357 using the Poisson solver. The NBE discretization error (scaled by $\|\nabla^2\phi\|'$) can be denoted and
358 defined by

$$359 \quad E(\nabla^2\phi) \equiv \|\nabla^2\phi - N(\psi_t)\|'/\|\nabla^2\phi\|'. \quad (16)$$

361
362 This error is 3.25×10^{-3} (or 4.33×10^{-3}) for ϕ obtained by using the Poisson solver with $L = 2000$
363 (or 1000) but increases to 5.58×10^{-3} (or 5.78×10^{-3}) for ϕ obtained by using the SOR scheme.
364 Thus, the solution obtained by using the Poisson solver is used as the input field of ϕ in the
365 NBE to test the iterative procedures in each set of experiments.

366

367 **4. Results of experiments**

368 **4.1 Results from first set of experiments**

369 For this set of experiments, ψ_t and (u_t, v_t) are plotted in Fig. 1a, ψ_g and $(u_g, v_g) \equiv (-\partial_y\psi_g,$
370 $\partial_x\psi_g)$ are plotted in Fig. 1b, the vorticity $\zeta_t \equiv \nabla^2\psi_t$ is plotted in Fig. 1c, and the geostrophic

371 vorticity $\zeta_g \equiv \nabla^2 \psi_g$ is plotted in Fig. 1d. Figure 1c shows that the absolute vorticity, defined by
372 $f + \zeta_i$, is positive everywhere so the nonlinearly balanced wavering jet flow is inertially stable
373 over the entire domain (see the proof in Appendix C of Xu, 1994). Figure 1c shows that the
374 geostrophic vorticity ζ_g is larger than $-f/2$ ($= -f_0/2$) everywhere, so the NBE is elliptic over the
375 entire domain and its associated boundary value problem in (1) is well posed.

376 The relative error of ψ_k with respect to ψ_t can be denoted and defined by

377

$$378 \quad \varepsilon(\psi_k) \equiv (\psi_k - \psi_t) / \|\psi_t\|, \quad (17)$$

379

380 where $\|\cdot\|$ denotes the RMS of discretized field of the variable inside $\|\cdot\|$ computed over all the
381 grid points (including the boundary points) of domain D . The accuracy of the solution ψ_k
382 obtained during the iterative process in each experiment can be evaluated by the RMS of $\varepsilon(\psi_k)$,
383 denoted and defined by

384

$$385 \quad E(\psi_k) \equiv \|\varepsilon(\psi_k)\|, \quad (18)$$

386

387 where $\|\cdot\|$ is defined in (17). The accuracy to which the NBE is satisfied by ψ_k can be measured
388 by $E[N(\psi_k)]$ defined in (13).

389 Table 1 lists the values of $E(\psi_k)$ and $E[N(\psi_k)]$ for the initial guess ψ_0 ($= \psi_g$) in row 1 and
390 the optimally truncated solutions ψ_K from the four experiments in rows 2-5. As shown in row
391 2 versus row 1 of Table 1, M1a reaches the optimal truncation at $k = K = 6$ where $E[N(\psi_k)]$ is
392 reduced (from 0.120 at $k = 0$) to its minimum [$= 2.411 \times 10^{-3} < E(\nabla^2 \phi) = 3.25 \times 10^{-3}$ – the NBE
393 discretization error defined in (16)] with $E(\psi_k)$ reduced (from 2.43×10^{-2} at $k = 0$) to 4.87×10^{-4} .

394 Figure 2a shows that $E(\psi_k)$ reaches its minimum ($= 4.79 \times 10^{-4}$) at $k = 10$. This minimum is
395 slightly below $E(\psi_k) = 4.87 \times 10^{-4}$ but undetectable in real-case applications.

396 On the contrary, as shown in row 3 of Table 1 and Fig. 2b, M1b reaches the optimal
397 truncation very slowly at $k = K = 38493$ where $E[N(\psi_k)]$ is reduced to its global minimum ($=$
398 1.81×10^{-2}) with $E(\psi_k)$ reduced to 1.68×10^{-3} . Here, $E[N(\psi_k)]$ has three extremely shallow and
399 small local minima (at $k = 32408, 38490$ and 38497) not visible in Fig. 2b. These local minima
400 are detected and passed by setting $m = 3$ in (13) for M1b. Clearly M1b is less accurate and
401 much less efficient than M1a.

402 Figure 2c (or 2d) shows that M2a (or M2b) reaches the optimal truncation at $k = K = 19$ (or
403 26) where $E[N(\psi_k)]$ is reduced to its global minimum [$= 3.55 \times 10^{-2}$ (or 2.66×10^{-2})] with $E(\psi_k)$
404 reduced to 4.55×10^{-3} (or 2.69×10^{-3}), and $E(\psi_k)$ decreases continuously toward its minimum [$=$
405 2.45×10^{-3} (or 1.62×10^{-3})] as k increases beyond K . Thus, M2a and M2b are less efficient and
406 much less accurate than M1a for $Ro = 0.1$.

407

408 **4.2 Results from second set of experiments**

409 For this set of experiments, ψ_t and (u_t, v_t) have the same patterns as those in Fig. 1a, and
410 ψ_g and (u_g, v_g) are similar to those in Fig. 1b, but the contour intervals of ψ_t and ψ_g are reduced
411 by 50% as L is reduced from 2000 to 1000 km with Ro increased to 0.2, so the wavering jet
412 flow is on the sub-synoptic scale. In this case, the nonlinearly balanced jet flow is still inertially
413 stable over the entire domain since $\zeta_t > -f$ everywhere as shown in Fig. 3a, but $\zeta_g < -f/2$ in the
414 two small yellow colored areas as shown in Fig. 3b, so the NBE becomes hyperbolic locally in
415 this small area and the boundary value problem in (1) is not fully well posed.

416 In this case, as shown in row 2 versus row 1 of Table 2, M1a reaches the optimal truncation
417 at $k = K = 13$ where $E[N(\psi_k)]$ is reduced (from 0.243 at $k = 0$) to its minimum [$= 5.23 \times 10^{-3}$
418 close to $E(\nabla^2 \phi) = 4.33 \times 10^{-3}$] with $E(\psi_k)$ reduced (from 4.86×10^{-2} at $k = 0$) to 1.24×10^{-3} . The
419 rapid descending processes of $E(\psi_k)$ and $E[N(\psi_k)]$ (not shown) are similar to those in Fig. 2a
420 for M1a in the first set of experiments.

421 As shown in row 3 of Table 2, M1b takes $K = 48057$ iterations to reach the optimal
422 truncation and the values of $E[N(\psi_k)]$ and $E(\psi_k)$ at $k = K$ are about four times larger than those
423 from M1a. The extremely slow descending processes of $E(\psi_k)$ and $E[N(\psi_k)]$ (not shown) are
424 similar to those in Fig. 2b for M1b in the first set of experiments. As shown in row 4 (or 5) of
425 Table 2, M2a (or M2b) reaches the optimal truncation at $k = K = 26$ (or 35) and the values of
426 $E[N(\psi_k)]$ and $E(\psi_k)$ are more than (or about) 4 times of those from M1a. Thus, M1a is still
427 more accurate and much more efficient than M1b and is more efficient and much more accurate
428 than M2a and M2b for $Ro = 0.2$, although the boundary value problem in (1) in this case is not
429 fully (but nearly) well posed.

430

431 **4.3 Results from third set of experiments**

432 For this set of experiments, ψ_t and (u_t, v_t) have the same patterns as those in Fig. 1a but the
433 contour interval of ψ_t is reduced 4 times as L is reduced from 2000 to 500 km with Ro increased
434 to 0.4, so the wavering jet flow is on the meso- α scale. Figure 4a shows the fields of ψ_g and
435 (u_g, v_g) for the nonlinearly balanced jet flow. This nonlinearly balanced jet flow is inertially
436 unstable in the yellow colored area south of the ridge of wavering jet axis in the middle of
437 domain D where $\zeta_t < -f$ as shown in Fig. 4c. Figure 4d shows that $\zeta_g < -f/2$ in the long and broad

438 yellow colored area along and around the wavering jet, so the NBE is hyperbolic in this area
439 and the boundary value problem in (1) becomes seriously ill-posed.

440 In this case, as shown in row 2 of Table 3 and Fig. 5a, M1a reaches the optimal truncation
441 at $k = K = 2$ where $E[N(\psi_k)]$ is decreased (from 0.57 at $k = 0$) to its minimum ($= 0.13$), while
442 $E(\psi_k)$ decreases from 9.72×10^{-2} at $k = 0$ to 8.20×10^{-2} at $k = K = 2$ and then to its minimum ($=$
443 7.38×10^{-2}) at $k = 6$. As k increases beyond 6, M1a diverges. Its optimally truncated solution ψ_K
444 is merely slightly more accurate than the initial guess ψ_0 . As shown in row 3 of Table 3 and
445 Fig. 5b, M1b reaches the optimal truncation at $k = K = 10325$ where $E[N(\psi_k)]$ is decreased to
446 its global minimum ($= 0.15$), while $E(\psi_k)$ decreases to 8.31×10^{-2} at $k = K$ and then to its
447 minimum ($= 7.68 \times 10^{-2}$) at $k = 23515$. Thus, M1b is still less accurate and much efficient than
448 M1a.

449 Figure 5c (or 5d) shows that M2a (or M2b) reaches the optimal truncation at $k = K = 26$ (or
450 29) where $E[N(\psi_k)]$ is reduced to its minimum [$= 0.11$ (or 0.10)], while $E(\psi_k)$ is reduced to its
451 minimum [$= 8.24 \times 10^{-2}$ (or 8.24×10^{-2})] at $k = 25$ (or 26) and then increases slightly to 8.25×10^{-2}
452 (or 8.26×10^{-2}) at $k = K = 26$ (or 29). As shown in row 4 (or 5) versus row 2 of Table 3, $E(\psi_K)$
453 from M2a (or M2b) is larger than that from M1a, so M2a (or M2b) is still less accurate than
454 M1a in this case.

455 Figure 6a (or 6b) shows that $\varepsilon(\psi_K)$ from M1a (or M1b) peaks positively and negatively in
456 the middle of domain D as $\varepsilon(\psi_0)$ does in Fig. 4b but with slightly reduced amplitudes. Figure
457 6c (or 6d) shows that $\varepsilon(\psi_K)$ from M2a (or M2b) has a broad negative peak south of the ridge
458 of wavering jet axis similar to that of $\varepsilon(\psi_0)$ in Fig. 4b but with a slightly enhanced amplitude.

459 In this case, M1a is still slightly more accurate than other three iterative procedures but it cannot
460 effectively reduce the solution error in the central part of the domain where not only the NBE
461 is hyperbolic (with $\zeta_g < -f/2$ as shown in Fig. 4d) but also the jet flow is strongly anti-
462 cyclonically curved and subject to inertial instability (with $\zeta_t < -f$ as shown in Fig. 4c).

463

464 **4.4 Results from fourth set of experiments**

465 For this set of experiments, ψ_t and (u_t, v_t) are plotted in Fig. 7a. These fields represent the
466 same nonlinearly balanced wavering westerly jet flow as that in the third set of experiments
467 except that the wave fields are shifted by a half of wavelength so the jet flow is curved
468 cyclonically in the middle of domain D . In this case, ψ_g and (u_g, v_g) are nearly the same as the
469 half-wavelength shifted fields (not shown) from Fig. 4a but with small differences mainly along
470 and around the trough and ridge lines due to the boundary condition, $\phi \equiv f\psi_g = f\psi_t$, used here
471 along the two trough lines (instead of the two ridge lines in Fig. 4a) for solving ϕ from $\nabla^2\phi =$
472 $N(\psi_t)$. Figure 7c shows the jet flow becomes inertially unstable in the two yellow colored areas
473 (where $\zeta_t < -f$) around the west and east boundaries of domain D . Figure 7d shows that the NBE
474 becomes hyperbolic in the long and broad yellow colored area (where $\zeta_g < -f/2$) that is nearly
475 the same as the yellow colored area in Fig. 4d but half-wavelength shifted, so the area of $\zeta_g <$
476 $-f$ (that is, the area of $\zeta_0 + f < 0$ in which the initial guess field is inertially unstable) in Fig. 4d
477 is moved with the ridge line to the west and east boundaries in Fig. 7d. As the area of $\zeta_g < -f$
478 and area of $\zeta_t < -f$ are moved away from the domain center to the domain boundaries where ψ_t
479 is known and given by ϕ/f , solving the NBE becomes less difficult in this fourth set of
480 experiments than in the third set.

481 In this case, as shown in row 2 of Table 4 and Fig. 8a, M1a reaches the optimal truncation
482 at $k = K = 7$ where $E[N(\psi_k)]$ is decreased (from 0.76 at $k = 0$) to its minimum ($= 3.81 \times 10^{-2}$),
483 while $E(\psi_k)$ decreases from 9.71×10^{-2} at $k = 0$ to 2.29×10^{-2} at $k = K = 7$ and then to its flat
484 minimum ($= 2.25 \times 10^{-2}$) at $k = 12$, so ψ_K is significantly more accurate than ψ_0 and slightly less
485 accurate than ψ_k at $k = 12$ (which is undetectable in real-case applications). As shown in row 3
486 of Table 4 and Fig. 8b, M1b reaches the optimal truncation at $k = K = 31830$ where $E[N(\psi_k)]$
487 is decreased to its global minimum ($= 4.54 \times 10^{-2}$), while $E(\psi_k)$ decreases to 2.37×10^{-2} at $k = K$
488 and then to its minimum ($= 2.21 \times 10^{-2}$) at $k = 57586$. Thus, M1b is still much less efficient and
489 less accurate than M1a.

490 Figure 8c (or 8d) shows that M2a (or M2b) reaches the optimal truncation at $k = K = 27$ (or
491 32) where $E[N(\psi_k)]$ is reduced to its minimum [$= 5.42 \times 10^{-2}$ (or 4.66×10^{-2})], $E(\psi_k)$ reduces to
492 3.03×10^{-2} (or 2.64×10^{-2}) at $k = K$ and then to its minimum [$= 2.72 \times 10^{-2}$ (or 2.43×10^{-2})] at $k =$
493 36 (or 44), so M2a (or M2b) is still less efficient and less accurate than M1a in this case.

494 Figure 7b shows that $\varepsilon(\psi_0)$ has a broad positive (or negative) peak south (or north) of the
495 trough of wavering jet axis in the middle of domain D . These broad peaks are mostly reduced
496 by M1a as shown by $\varepsilon(\psi_K)$ in Fig. 9a but slightly less reduced by M1b as shown in Fig. 9b and
497 less reduced by M2a (or M2b) as shown in Fig. 9c (or 9d). However, the small secondary
498 negative peak of $\varepsilon(\psi_g)$ near the west (or east) boundary in Fig. 7b is reduced only about 30%
499 by M1a (or M1b) as shown by $\varepsilon(\psi_K)$ in Fig. 9a (or 9b) and even less reduced by M2a (or M2b)
500 as shown in Fig. 9c (or 9d). Thus, all the four iterative procedures have difficulties to reduce
501 the errors of their optimally truncated solutions near the west and east boundaries where not

502 only the NBE is hyperbolic (with $\zeta_g < -f/2$ as shown in Fig. 7d) but also the jet flow is subject
503 to inertial instability (with $\zeta_t < -f$ as shown in Fig. 7c). Nevertheless, since the area of $\zeta_t < -f$ is
504 moved with the ridge of wavering jet axis to the domain boundaries in Fig. 7c, all the four
505 iterative procedures perform significantly better in this set of experiments than in the previous
506 third set, as shown in Fig. 9 and Table 4 versus Fig. 6 and Table 3. In this case, M1a is still
507 most accurate and M1b is still least efficient among the four iterative procedures.

508

509 **5. Conclusions**

510 In this paper, two types of previous iterative methods for solving the NBE are reviewed and
511 revisited. The first type was originally proposed by Bolin (1955), in which the NBE is
512 transformed into a linearized equation for a presumably small correction to the initial guess or
513 the subsequently updated solution. The second type was originally proposed by Shuman (1955,
514 1957) and Miyakoda (1956), in which the NBE is rearranged into a quadratic form of the
515 absolute vorticity and the positive root of this quadratic form is used in the form of Poisson
516 equation to obtain the solution iteratively. These two types of methods are re-derived formally
517 by expanding the solution asymptotically upon a small Rossby number (see section 2), and the
518 re-derived methods are called method-1 and method-2, respectively.

519 Since the rearranged asymptotic expansion is not ensured to converge especially when the
520 Rossby number is not sufficiently small, a criterion for optimal truncation of asymptotic
521 expansion is proposed [see (13)] to obtain the super-asymptotic approximation of the solution
522 based on the heuristic theory of asymptotic analysis (Boyd, 1999). In addition, the Poisson

523 solver based on the integral formulas (Xu et al., 2011; Cao and Xu, 2011) is used versus the
524 SOR scheme to solve the boundary value problem in each iterative step.

525 The four iterative procedures are tested with analytically formulated wavering jet flows on
526 different spatial scales in four sets of experiments. The computational domain covers one full
527 wavelength and is centered at the ridge of the wavering jet in the first three sets of experiments
528 but centered at the trough in the last set. In the first set of experiments, the wavering jet flow is
529 formulated on the synoptic scale [with the half wavelength $L = 2000$ km and the associated
530 Rossby number $Ro = 0.1$]. In this case, the NBE is of the elliptic type over the entire domain
531 and therefore its boundary value problem is well posed. In the second set of experiments, the
532 wavering jet flow is formulated on the sub-synoptic scale [with $L = 1000$ km and $Ro = 0.2$]. In
533 this case, the NBE is of the elliptic type nearly over the entire domain so that its boundary value
534 problem is nearly well posed. In the third (or fourth) set of experiments, the wavering jet flow
535 is formulated on the meso- α scale with $Ro = 0.4$, the wavering jet flow is curved anti-
536 cyclonically (or cyclonically) in the middle of the domain where the absolute vorticity is locally
537 negative (or strongly positive), and the NBE becomes hyperbolic broadly along and around the
538 wavering jet so that its boundary value problem is seriously ill-posed.

539 The test results can be summarized as follows: For wavering jet flows on the synoptic and
540 sub-synoptic scales, all the four iterative procedures can reach their respective optimal
541 truncations and the solution error (originally from the initial guess – the geostrophic
542 streamfunction) can be reduced at the optimal truncation by an order of magnitude or nearly so
543 even when the NBE is not entirely elliptic. Among the four iterative procedures, M1a is most

544 accurate and efficient while M1b is least efficient. The results for wavering jet flows on the
545 synoptic and sub-synoptic scales are insensitive to the location of wavering jet in the
546 computational domain. In particular, according to our additional experiments (not shown in
547 this paper), when the wavering jet is shifted zonally by a half of wavelength (with the trough
548 moved to the domain center), the solution errors become slightly smaller and the optimal
549 truncation numbers for M1a and M1b (or M2a and M2b) become slightly smaller (or larger)
550 than those listed in Tables 1 and 2. For wavering jet flows on the meso- α scale in which the
551 NBE's boundary value problem is seriously ill-posed, the four iterative procedures still can
552 reach their respective optimal truncations with the solution error reduced effectively for
553 cyclonically curved part of the wavering jet flow but not for the anti-cyclonically curved part.
554 In this case, M1a is still most accurate and efficient while M1b is least efficient.

555 In comparison with M1b, the high accuracy and efficiency of M1a can be explained by the
556 fact that the solution obtained by the Poisson solver based on the integral formulas is not only
557 more accurate but also smoother than the solution obtained by the SOR scheme in each step of
558 nonlinear iteration. Consequently, in each next step, the nonlinear differential term on the right-
559 hand side of the incremental-form iteration equation [see (9a)] is computed more accurately in
560 M1a than in M1b and so is the entire right-hand side. This is especially true and important
561 when the entire right-hand side becomes very small (toward zero) in the late stage of iterations,
562 as it also explains why M1b reaches the optimal truncation much slower than M1a (see Tables
563 1–4). In comparison with M2a and M2b, the high accuracy and efficiency of M1a can be
564 explained by the fact that the solution in M1a is updated incrementally and the increment is

565 small relative to the entire solution and so is the error of the increment computed in each step
566 of nonlinear iteration. On the other hand, the solution in M2a or M2b is updated entirely and
567 the entire solution is large relative to the increment and so is the error of the entire solution
568 computed in each step of nonlinear iteration. Moreover, the recursive form of equation [see
569 (12)] used by M2a and M2b contains a square root term on its right-hand side, so it cannot be
570 converted into an incremental form. Furthermore, this square root term must set to zero when
571 the term inside the square root becomes negative, although the term inside the square root
572 corresponds to the squared absolute vorticity. This problem is caused by the non-negative
573 absolute vorticity assumed in the derivation of the recursive form of equation for M2a and M2b.

574 Cyclonically curved meso- α scale jet flows in the middle and upper troposphere are often
575 precursors of severe weather especially when the curved jet flow evolves into a cut-off cyclone
576 atop a meso- α scale low pressure system in the lower troposphere. In this case, M1a can be
577 potentially and particularly useful for severe weather analyses in the context of semi-balanced
578 dynamics (Xu, 1994; Xu and Cao, 2012). In addition, since the mass fields can be estimated
579 from Advanced Microwave Sounding Unit (AMSU) observations, using the NBE to retrieve
580 the horizontal winds in and around tropical cyclones (TC) from the estimated mass fields have
581 potentially important applications for TC warnings and improving TC initial conditions in
582 numerical predictions (Velden and Smith, 1983; Bessho et al, 2006). Applications of M1a in
583 the aforementioned directions deserve continued studies. In particular, the gradient wind can
584 be easily computed for the axi-symmetric part of a cut-off cyclone (or TC) and used to improve
585 the initial guess for the iterative procedure. This use of gradient wind can be somewhat similar

586 to the use of gradient wind associated with the axisymmetric part of a hurricane to improve the
587 basic-state potential vorticity (PV) construction for hurricane PV diagnoses (Wang and Zhang,
588 2003; Kieu and Zhang, 2010). Furthermore, either the gradient wind or the optimal truncated
589 solution from M1a can be used as a new improved initial guess. In this case, the asymptotic
590 expansion can be reformulated upon a new small parameter associated with the reduced error
591 of the new initial guess and this new small parameter can be smaller or much smaller than the
592 Rossby number used for the asymptotic expansion in this paper. The reformulated asymptotic
593 expansion may be truncated to yield a more accurate ‘hyerasymptotic’ approximation of the
594 solution according to the heuristic theory of asymptotic analysis (see section 5 of Boyd, 1999).
595 This approach deserves further explorations.

596

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662 Table 1. Values of $E(\psi_k)$ and $E[N(\psi_k)]$ listed in row 1 for the initial guess $\psi_0 (= \psi_g)$ with $k = 0$
663 and in rows 2-5 for ψ_K from the four iterative procedures in the first set of experiments (with
664 $Ro = 0.1$). Here, $E(\psi_k)$ is defined in (18), $E[N(\psi_k)]$ is defined in (13), k is the iteration number,
665 and ψ_K is the optimally truncated solution at $k = K$.

	$E(\psi_k)$	$E[N(\psi_k)]$	k
ψ_0	2.43×10^{-2}	0.120	$k = 0$
M1a	4.87×10^{-4}	2.41×10^{-3}	$k = K = 6$
M1b	1.68×10^{-3}	1.81×10^{-2}	$k = K = 38493$
M2a	4.55×10^{-3}	3.55×10^{-2}	$k = K = 19$
M2b	2.69×10^{-3}	2.66×10^{-2}	$k = K = 26$

666
667 Table 2. As in Table 1 but for the second set of experiments (with $Ro = 0.2$).

	$E(\psi_k)$	$E[N(\psi_k)]$	k
ψ_0	4.86×10^{-2}	0.243	$k = 0$
M1a	1.24×10^{-3}	5.23×10^{-3}	$k = K = 13$
M1b	5.14×10^{-3}	2.20×10^{-2}	$k = K = 48057$
M2a	6.31×10^{-3}	4.17×10^{-2}	$k = K = 26$
M2b	3.96×10^{-3}	2.94×10^{-2}	$k = K = 35$

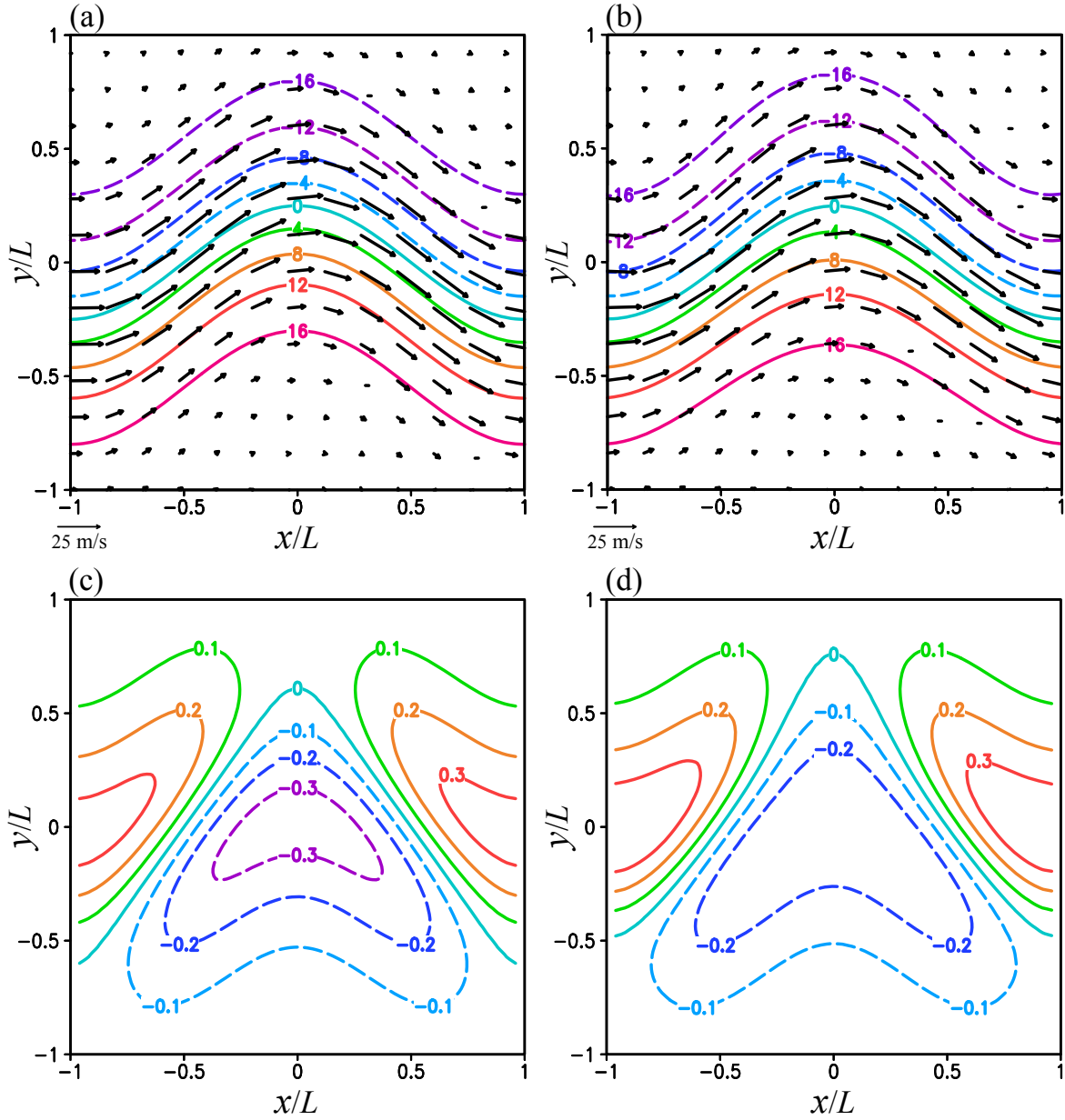
668
669 Table 3. As in Table 1 but for the third set of experiments (with $Ro = 0.4$ and $x_0 = 0$).

	$E(\psi_k)$	$E[N(\psi_k)]$	k
ψ_0	9.72×10^{-2}	0.57	$k = 0$
M1a	8.20×10^{-2}	0.13	$k = K = 2$
M1b	8.31×10^{-2}	0.15	$k = K = 10325$
M2a	8.25×10^{-2}	0.11	$k = K = 26$
M2b	8.26×10^{-2}	0.10	$k = K = 29$

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671 Table 4. As in Table 1 but for the fourth set of experiments (with $Ro = 0.4$ and $x_0 = L$).

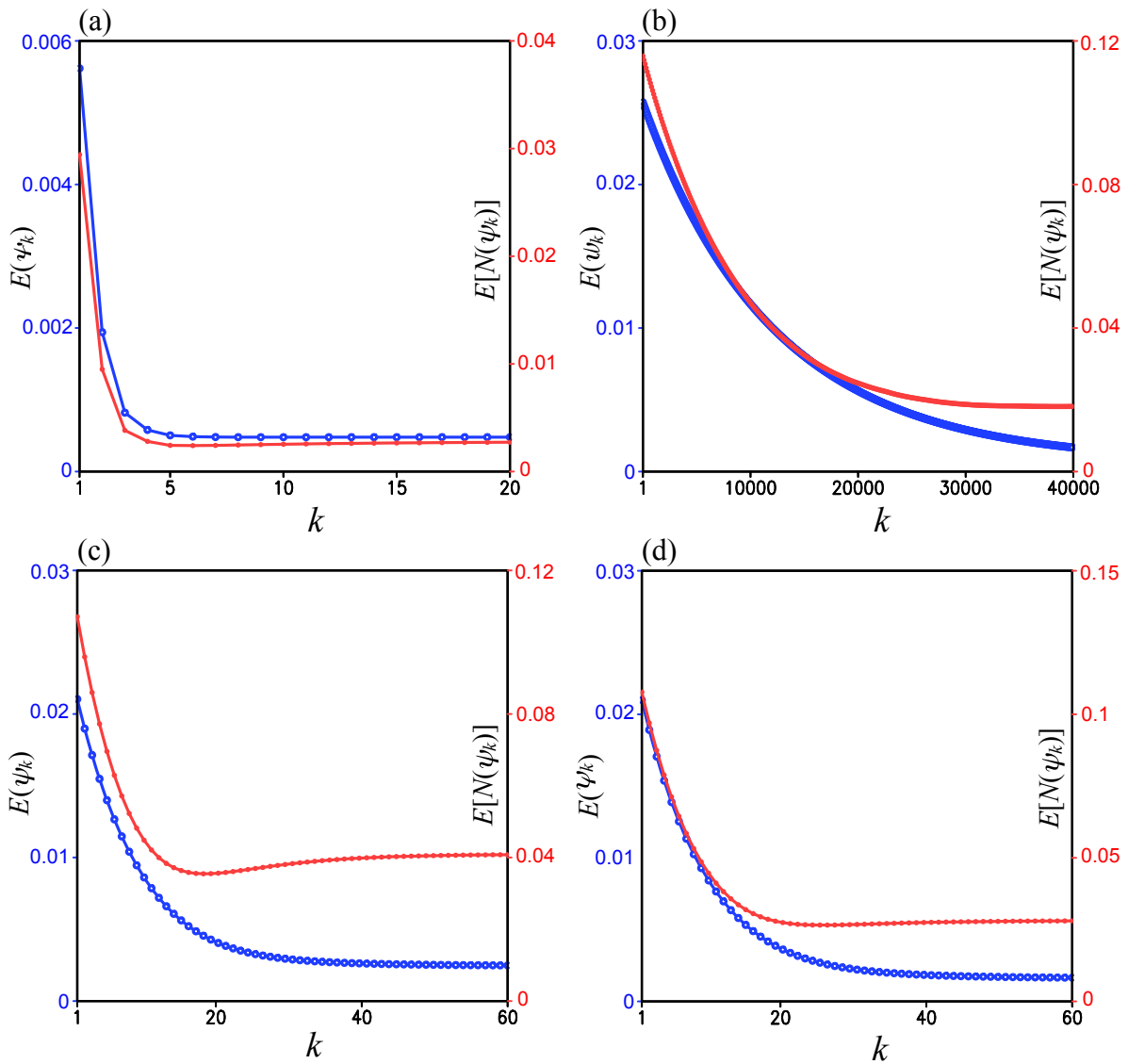
	$E(\psi_k)$	$E[N(\psi_k)]$	k
ψ_0	9.71×10^{-2}	0.76	$k = 0$
M1a	2.29×10^{-2}	3.81×10^{-2}	$k = K = 7$
M1b	2.37×10^{-2}	4.54×10^{-2}	$k = K = 31830$
M2a	3.03×10^{-2}	5.42×10^{-2}	$k = K = 27$
M2b	2.64×10^{-2}	4.66×10^{-2}	$k = K = 32$

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675 **Fig. 1.** (a) ψ_t plotted by color contours every 4.0 in the unit of $10^6 \text{ m}^2\text{s}^{-1}$ and (u_t, v_t) plotted by
676 black arrows over domain $D \equiv [-L \leq x \leq L, -L \leq y \leq L]$ with $L = 2000 \text{ km}$ for the first set of
677 experiments. (b) As in (a) but for ψ_g and (u_g, v_g) with $\psi_g \equiv \phi/f$ and ϕ computed from ψ_t by
678 setting $f = f_o = 10^{-4} \text{ s}^{-1}$ as described in section 3.3. (c) Vorticity $\zeta_t \equiv \nabla^2 \psi_t$ plotted by color
679 contours every 0.1 in the unit of 10^{-4} s^{-1} over domain D . (d) As in (c) but for geostrophic
680 vorticity $\zeta_g \equiv \nabla^2 \psi_g$. The wavering jet axis is along the green contour of $\psi_t = 0$ in (a) with its
681 ridge at $x = 0$ and two troughs at $x = \pm L$ on the west and east boundaries of domain D .

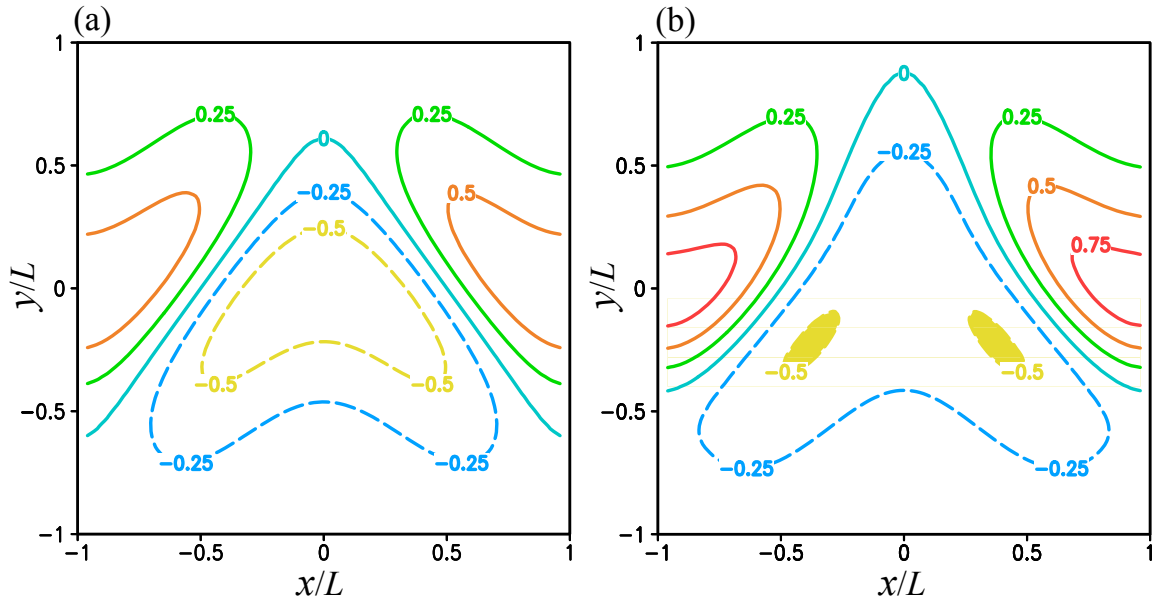


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685 **Fig. 2.** (a) $E[N(\psi_k)]$ and $E(\psi_k)$ from M1a in the first set of experiments plotted by red and blue686 curves, respectively, as functions of k over the range of $1 \leq k \leq 20$. (b) As in (a) but from M1b687 plotted over the range of $1 \leq k \leq 4 \times 10^4$. (c) As in (a) but from M2a plotted over the range of 1688 $\leq k \leq 60$. (d) As in (c) but from M2b. In each panel, the ordinate of $E[N(\psi_k)]$ is on the left side689 labeled in red and the ordinate of $E(\psi_k)$ is on the right side labeled in blue.

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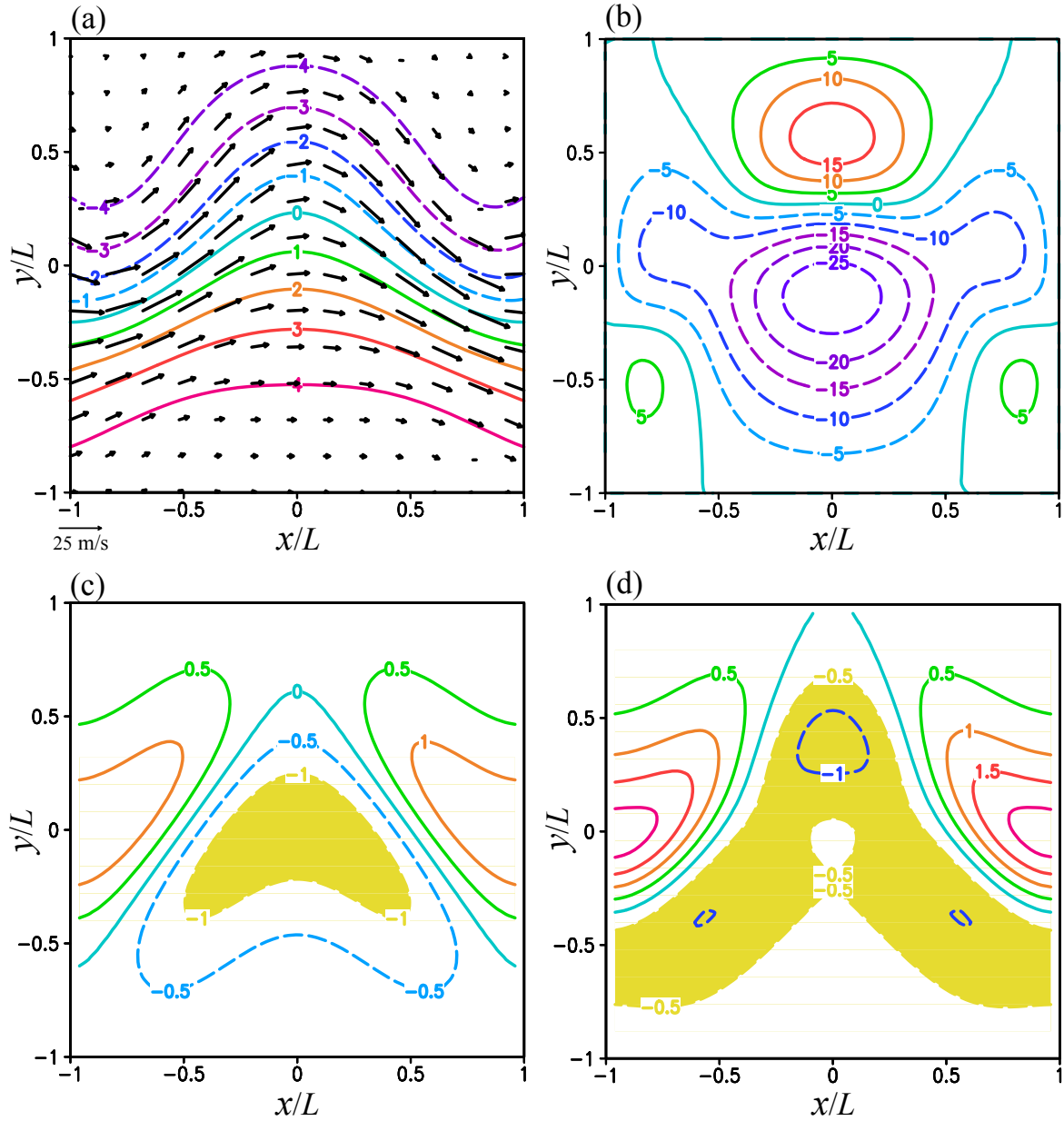
693 **Fig. 3.** (a) ζ_t plotted by color contours every 0.25 in the unit of 10^{-4} s^{-1} in domain D with $L =$

694 1000 km and $Ro = 0.2$ for the second set of experiments . (b) As in (a) but for ζ_g . As shown in

695 (b), $\zeta_g < -f/2$ ($= -f_0/2$) in the two small yellow colored areas where the NBE becomes locally

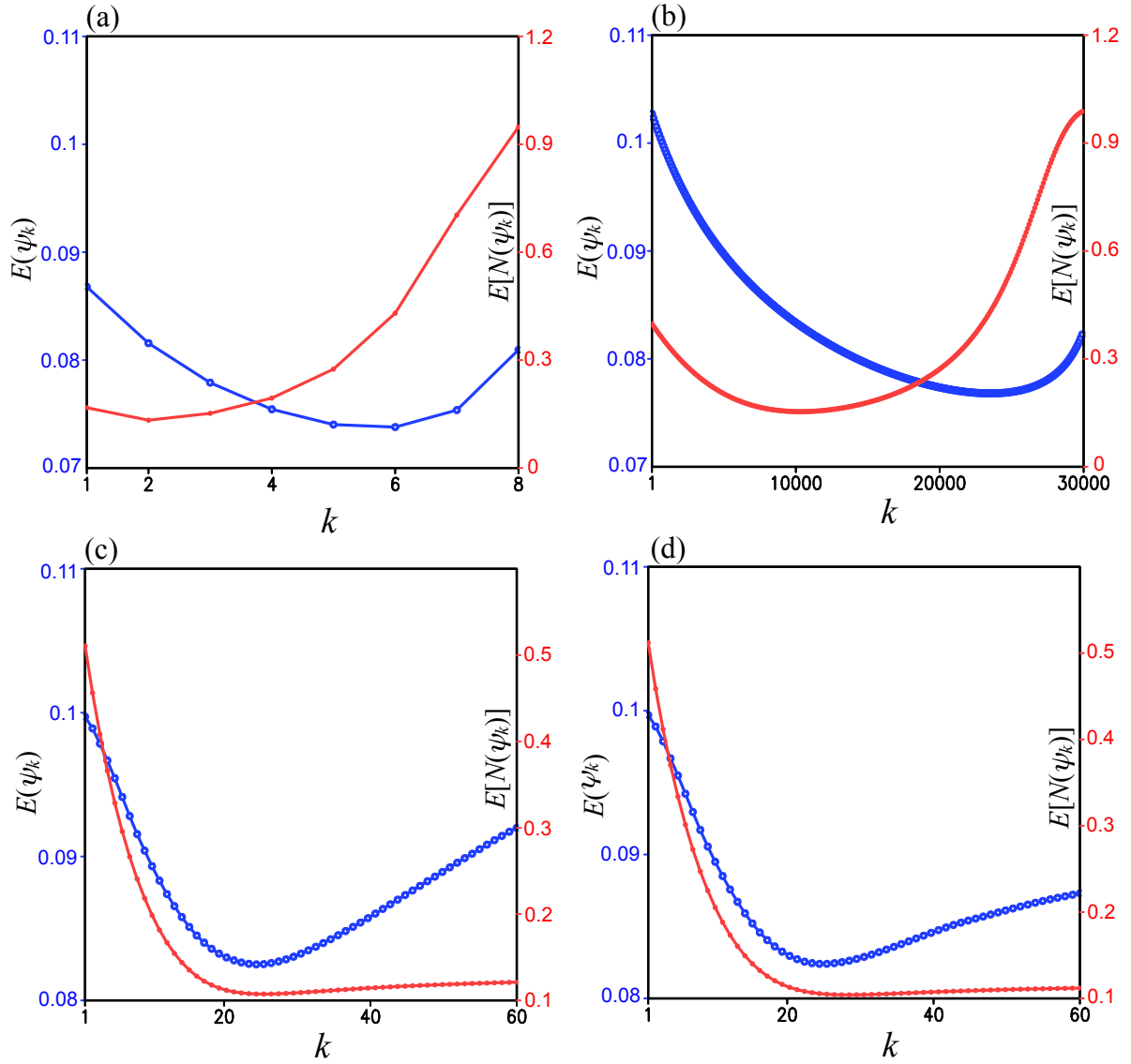
696 hyperbolic.

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699 **Fig. 4.** (a) ψ_g plotted by color contours every 1.0 in the unit of $10^6 \text{ m}^2\text{s}^{-1}$ and (u_g, v_g) plotted by
700 black arrows over domain D with $L = 500$ km and $Ro = 0.4$ for the third set of experiments. (b)
701 As in (a) but for $\varepsilon(\psi_0) = \varepsilon(\psi_g)$ plotted by color contours every 5.0 in the unit of 10^{-2} . (c) As in
702 (a) but for ζ_t plotted by color contours every 0.5 in the unit of 10^{-4} s^{-1} in domain D . (d) As in
703 (c) but for ζ_g . As shown in (c), $\zeta_t < -f$ in the yellow colored area south of the ridge of wavering
704 jet axis where the jet flow becomes inertially unstable. As shown in (c), $\zeta_g < -f/2$ ($= -f_0/2$) in the
705 long and broad yellow colored area (along and around the wavering jet) where the NBE
706 becomes hyperbolic.



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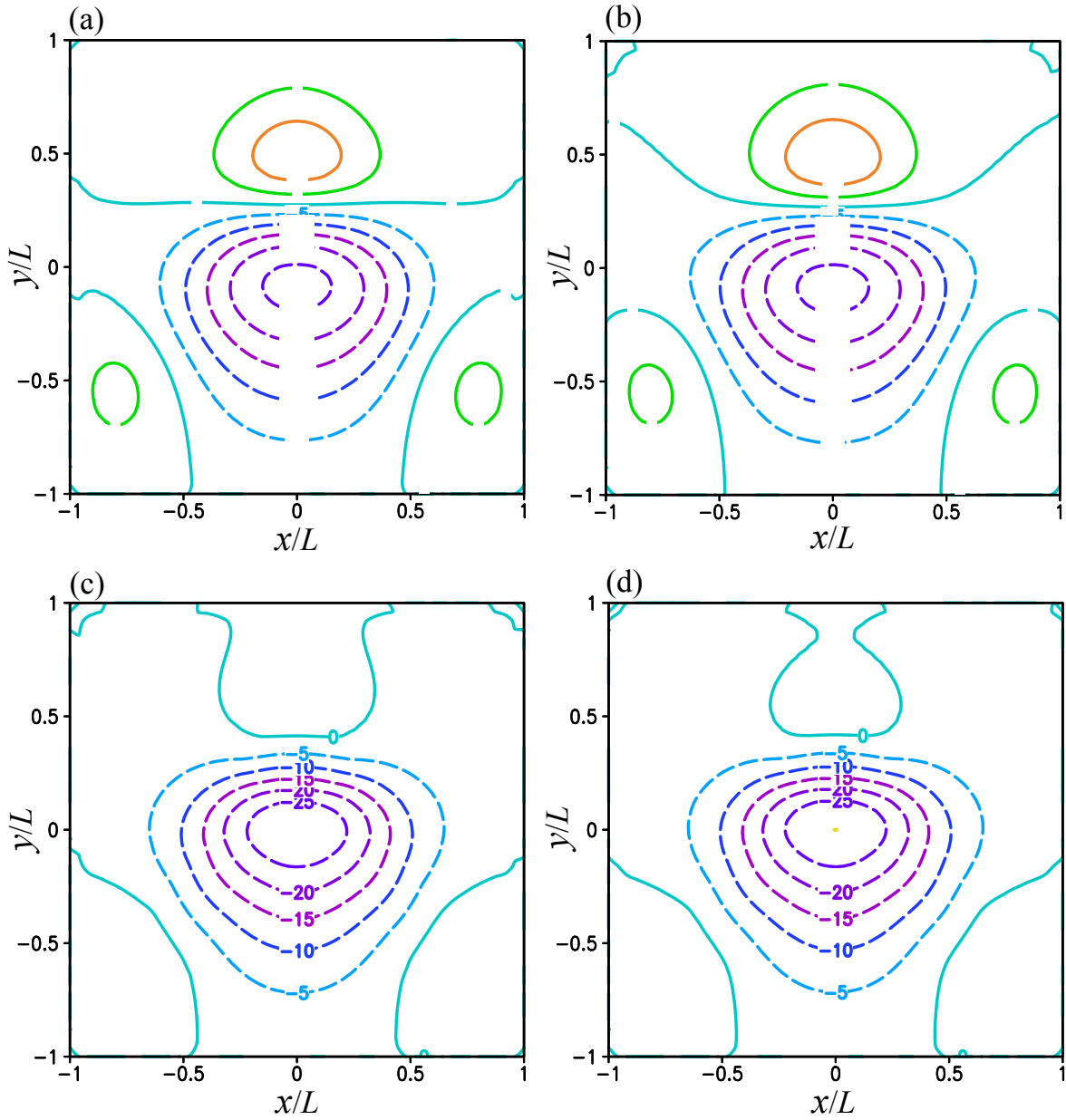
709 **Fig. 5.** (a) $E[N(\psi_k)]$ and $E(\psi_k)$ from M1a in the third set of experiments plotted by red and blue

710 curves, respectively, as functions of k over the range of $1 \leq k \leq 8$, (b) As in (a) but from M1b

711 plotted over the range of $1 \leq k \leq 3 \times 10^4$. (c) As in (a) but from M2a plotted over the range of 1

712 $\leq k \leq 60$. (d) As in (a) but from M2b. In each panel, the ordinates of $E[N(\psi_k)]$ and $E(\psi_k)$ are

713 placed and labeled as in Fig. 2.



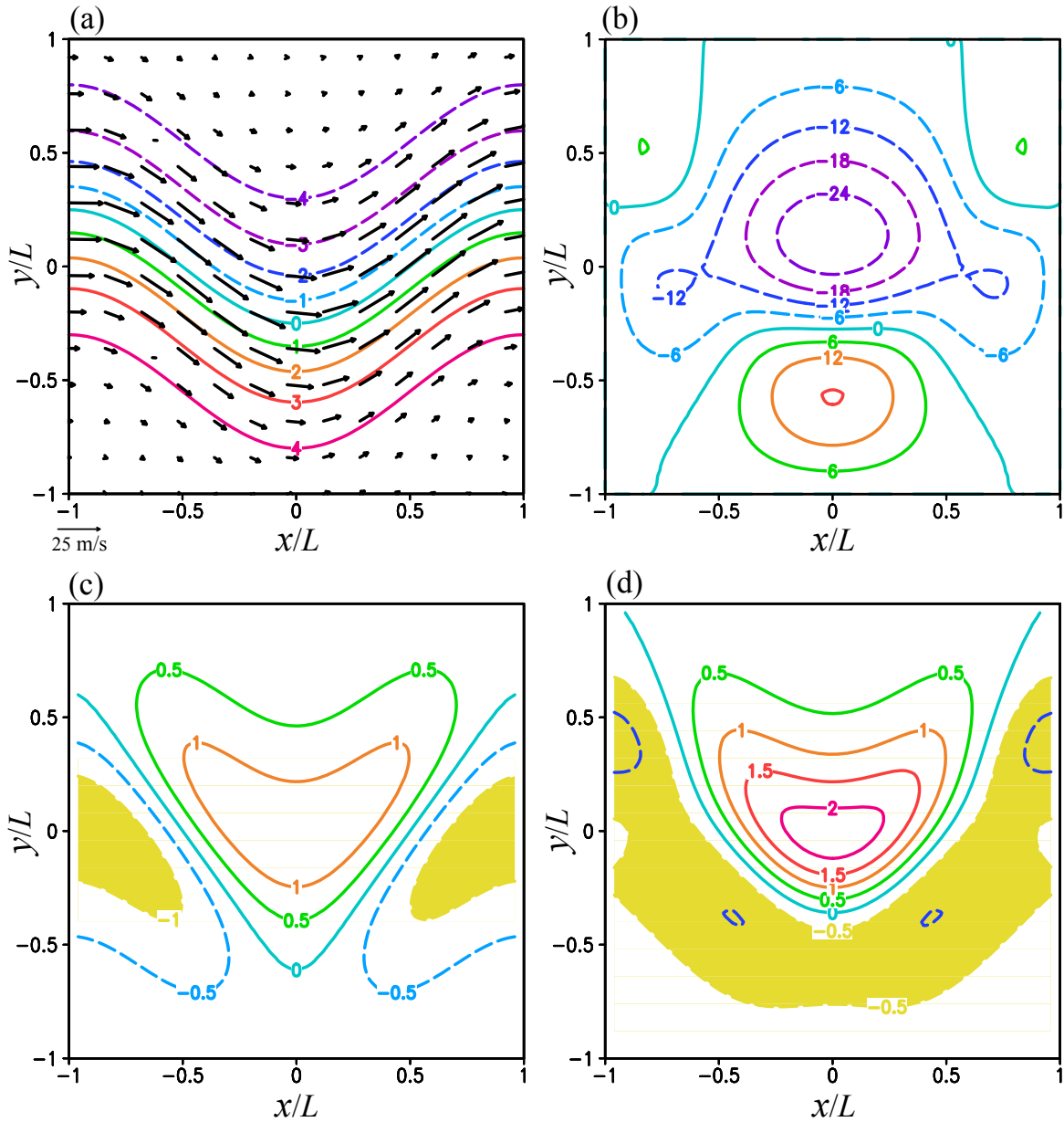
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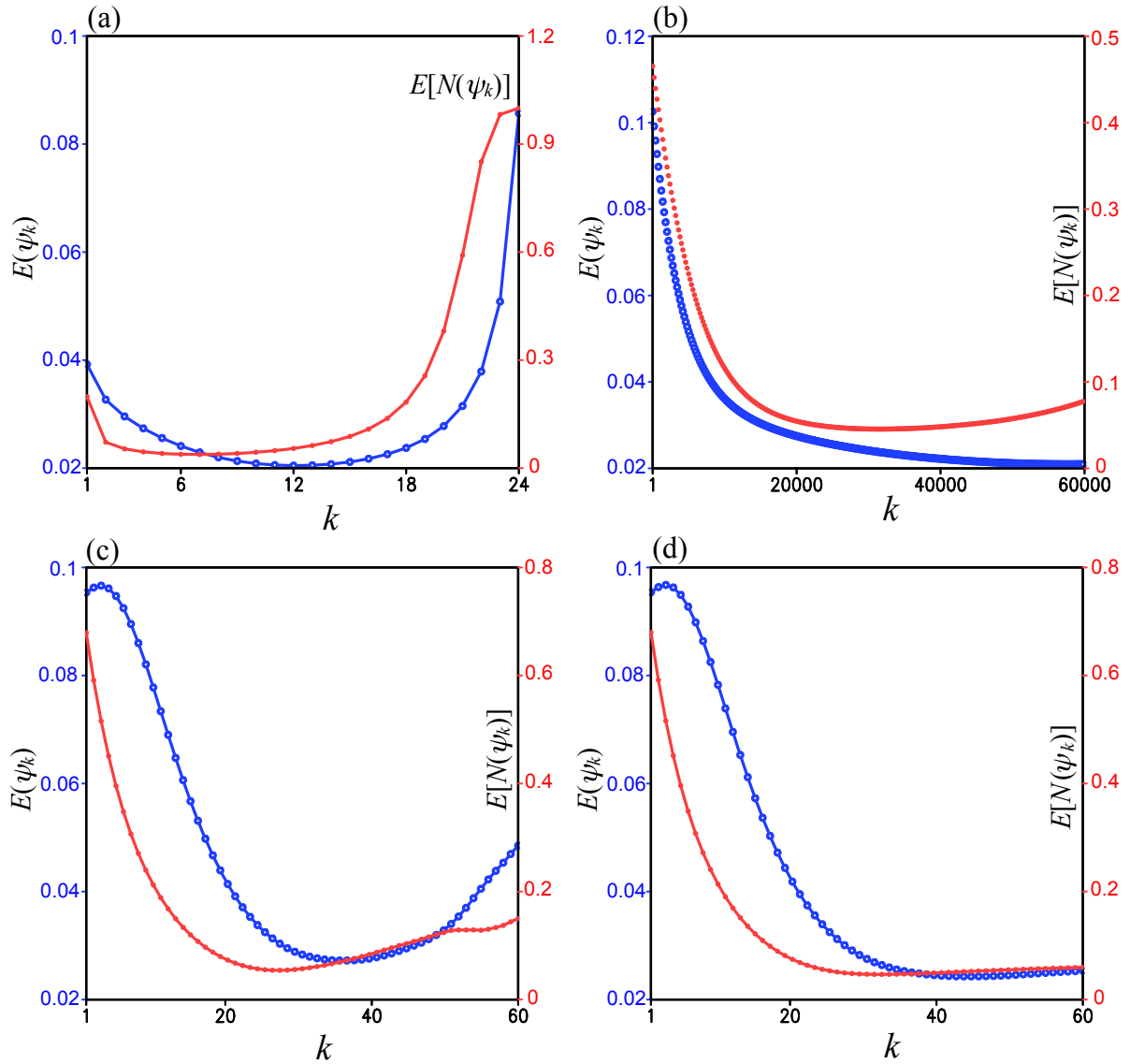
717 **Fig. 6.** $\varepsilon(\psi_K)$ plotted by color contours every 0.5 in the unit of 10^{-2} for ψ_K from (a) M1a, (b)

718 M1b, (c) M2a and (d) M2b in the third set of experiments.



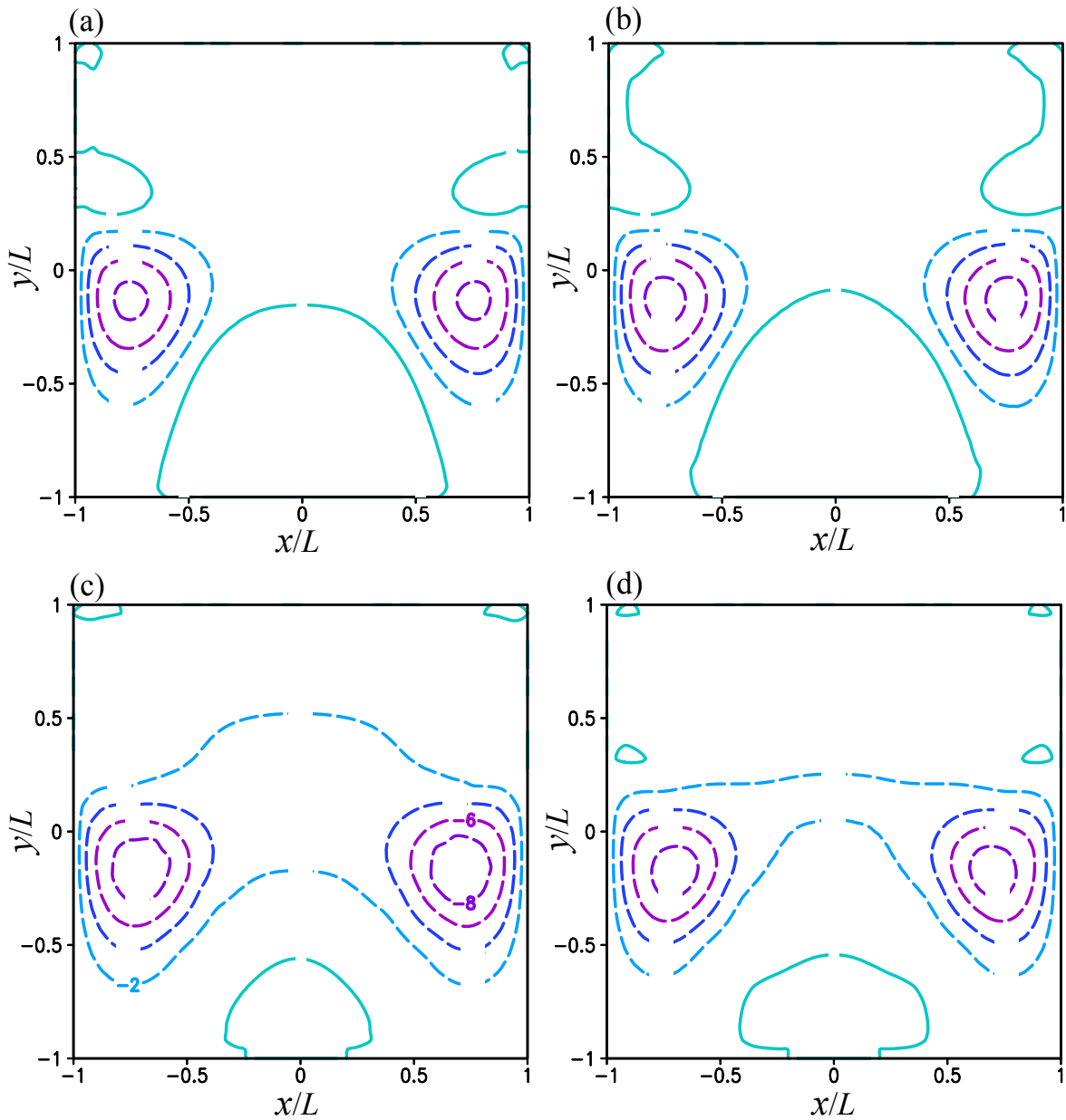
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722 **Fig. 7.** (a) As in Fig. 4a but for ψ_t and (u_t, v_t) in the fourth set of experiments with $L = 500$ km
723 and $x_0 = L$ (instead of $x_0 = 0$). (b) As in (a) but for $\varepsilon(\psi_0) = \varepsilon(\psi_g)$ plotted by color contours every
724 6.0 in the unit of 10^{-2} . (c) As in (a) but for ζ_t plotted by color contours every 0.5 in the unit of
725 10^{-4} s^{-1} in domain D . (d) As in (c) but for ζ_g .



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Fig. 8. (a) $E[N(\psi_k)]$ and $E(\psi_k)$ from M1a in the fourth set of experiments plotted by red and blue curves, respectively, as functions of k over the range of $1 \leq k \leq 24$, (b) As in (a) but from M1b plotted over the range of $1 \leq k \leq 6 \times 10^4$. (c) As in (a) but from M2a plotted over the range of $0 \leq k \leq 60$. (d) As in (a) but from M2b. In each panel, the ordinates of $E[N(\psi_k)]$ and $E(\psi_k)$ are placed and labeled as in Fig. 2.



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738 **Fig. 9.** $\varepsilon(\psi_K)$ plotted by color contours every 2.0 in the unit of 10^{-2} for ψ_K from (a) M1a, (b)

739 M1b, (c) M2a and (d) M2b in the fourth set of experiments.

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