



Interlaboratory comparison of $\delta^{13}\text{C}$ and δD measurements of atmospheric CH_4 for combined use of data sets from different laboratories

Taku Umezawa^{1,2}, Carl A. M. Brenninkmeijer¹, Thomas Röckmann³, Carina van der Veen³, Stanley C. Tyler^{4,5}, Ryo Fujita⁶, Shinji Morimoto^{6,7}, Shuji Aoki⁶, Todd Sowers⁸, Jochen Schmitt⁹, Michael Bock⁹, Jonas Beck⁹, Hubertus Fischer⁹, Sylvia E. Michel¹⁰, Bruce H. Vaughn¹⁰, John B. Miller¹⁰, James W. C. White¹⁰, Gordon Brailsford¹¹, Hinrich Schaefer¹¹, Peter Sperlich¹¹, Willi A. Brand¹², Michael Rothe¹², Thomas Blunier¹³, David Lowry¹⁴, Rebecca E. Fisher¹⁴, Euan G. Nisbet¹⁴, Andrew L. Rice¹⁵, Peter Bergamaschi¹⁶, Cordelia Veidt¹⁷, and Ingeborg Levin¹⁷

¹Max Planck Institute for Chemistry, Atmospheric Chemistry Department, 55128 Mainz, Germany

²Center for Environmental Measurement and Analysis, National Institute for Environmental Studies, Tsukuba 305-8506, Japan

³Institute for Marine and Atmospheric research Utrecht, Utrecht University, Utrecht, the Netherlands

⁴Earth System Science Department, University of California, Irvine, USA

⁵Chemistry Department, Norco College, Norco, CA 92860, USA

⁶Center for Atmospheric and Oceanic Studies, Graduate School of Science, Tohoku University, Sendai, Japan

⁷National Institute of Polar Research, Tokyo, Japan

⁸The Earth and Environmental Systems Institute, Pennsylvania State University, University Park, Pennsylvania 16802, USA

⁹Climate and Environmental Physics, Physics Institute and Oeschger Center for Climate Change Research, University of Bern, 3012 Bern, Switzerland

¹⁰Institute of Arctic and Alpine Research, University of Colorado Boulder, Boulder, Colorado, USA

¹¹National Institute of Water and Atmospheric Research, Wellington 6021, New Zealand

¹²Max Planck Institute for Biogeochemistry, 07745 Jena, Germany

¹³Centre for Ice and Climate, University of Copenhagen, Copenhagen, Denmark

¹⁴Department of Earth Sciences, Royal Holloway, University of London, Egham, UK

¹⁵Department of Physics, Portland State University, Portland, OR 97207, USA

¹⁶European Commission Joint Research Centre, Institute for Environment and Sustainability, Ispra (Va), Italy

¹⁷Institute for Environmental Physics, Heidelberg University, 69120 Heidelberg, Germany

Correspondence: Taku Umezawa (umezawa.taku@nies.go.jp)

Received: 31 July 2017 – Discussion started: 9 August 2017

Revised: 16 January 2018 – Accepted: 17 January 2018 – Published: 2 March 2018

Abstract. We report results from a worldwide interlaboratory comparison of samples among laboratories that measure (or measured) stable carbon and hydrogen isotope ratios of atmospheric CH_4 ($\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$). The offsets among the laboratories are larger than the measurement reproducibility of individual laboratories. To disentangle plausible measurement offsets, we evaluated and critically assessed a large number of intercomparison results, some of

which have been documented previously in the literature. The results indicate significant offsets of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements among data sets reported from different laboratories; the differences among laboratories at modern atmospheric CH_4 level spread over ranges of 0.5‰ for $\delta^{13}\text{C}\text{-CH}_4$ and 13‰ for $\delta\text{D}\text{-CH}_4$. The intercomparison results summarized in this study may be of help in future attempts to harmonize $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data sets from

different laboratories in order to jointly incorporate them into modelling studies. However, establishing a merged data set, which includes $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data from multiple laboratories with desirable compatibility, is still challenging due to differences among laboratories in instrument settings, correction methods, traceability to reference materials and long-term data management. Further efforts are needed to identify causes of the interlaboratory measurement offsets and to decrease those to move towards the best use of available $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data sets.

1 Introduction

Methane (CH_4) is an important anthropogenic and natural greenhouse gas, and it also has a large role in atmospheric chemistry through its reaction with the hydroxyl radical. Since individual CH_4 source types have characteristic isotope signatures and loss processes are associated with specific kinetic isotope effects, carbon and hydrogen isotope ratios of CH_4 ($\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$) have been useful in constraining the global CH_4 budget. Dictated by global mass balance, the average isotopic composition of CH_4 in the atmosphere ($\delta^{13}\text{C}\text{-CH}_4$ or $\delta\text{D}\text{-CH}_4$) equals the flux-weighted isotopic composition of the sources, corrected for the total kinetic isotope effects of removal processes (e.g. Stevens and Rust, 1982; Cicerone and Oremland, 1988; Quay et al., 1991, 1999; Miller et al., 2002; Turner et al., 2017; Rigby et al., 2017). It has been pointed out that assignment of representative isotopic signatures of various CH_4 sources remains uncertain due to their large spatial and temporal variability across the globe (e.g. Sherwood et al., 2017), which could result in large uncertainties of isotope-based estimates of the global CH_4 budget (Schwietzke et al., 2016). Nonetheless, the value of isotope measurements was amply demonstrated by recent studies which suggested shifts in the global CH_4 source over the last decades (Schaefer et al., 2016; Rice et al., 2016; Nisbet et al., 2016; Schwietzke et al., 2016); without isotopic analyses such conclusions would have been difficult to achieve. The isotopic ratios are commonly reported using the delta notation:

$$\delta = \frac{R_{\text{sample}}}{R_{\text{standard}}} - 1, \quad (1)$$

where R represents the atomic ratio of the less abundant isotope over the most abundant isotope in the sample or the standard. Conventionally, measured values are reported relative to the international isotope-scale VPDB (Vienna Pee Dee Belemnite) for $\delta^{13}\text{C}\text{-CH}_4$ and VSMOW (Vienna Standard Mean Ocean Water) for $\delta\text{D}\text{-CH}_4$ in per mil.

Given that the atmospheric lifetime of CH_4 is about a decade, its variation in background air is relatively small. For that reason, its mole fraction and isotopic measurements have to have high precision and accuracy. For $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$, researchers have achieved measurement reproducibil-

ity of <0.1 for $\delta^{13}\text{C}\text{-CH}_4$ and $<2\%$ for $\delta\text{D}\text{-CH}_4$. Incorporating $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data sets in chemistry transport models is useful for quantitatively separating different CH_4 source categories and attempts have been made to reduce uncertainties in the global CH_4 budget (e.g. Fung et al., 1991; Hein et al., 1997; Mikaloff Fletcher et al., 2004a, b; Monteil et al., 2011; Kirschke et al., 2013; Ghosh et al., 2015; Rice et al., 2016; Schaefer et al., 2016; Schwietzke et al., 2016; Röckmann et al., 2016; Turner et al., 2017; Rigby et al., 2017). However, although an increasing number of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data have been reported over the last decades, significant measurement offsets among laboratories have been found for both $\delta^{13}\text{C}\text{-CH}_4$ (e.g. Levin et al., 2012) and $\delta\text{D}\text{-CH}_4$ (Bock et al., 2014). It is clear that both traceability to the standard scales and interlaboratory comparisons (intercomparisons) are indispensable for combined use of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data from different laboratories. Many such intercomparisons have already been made, either on an ad hoc basis or on a more organized scale. However, a systematic evaluation of the underlying calibrations and related measurement offsets among laboratories has been lacking. It is also noted that some measurement programmes for $\delta^{13}\text{C}\text{-CH}_4$ and/or $\delta\text{D}\text{-CH}_4$ have been discontinued, and maintaining access to such data sets including well-established interlaboratory offsets is important. Here we combine and evaluate the existing comparison results to quantify interlaboratory measurement differences in order to facilitate the use of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data. This study therefore opens the possibility for merging historic CH_4 isotope data reported from multiple laboratories (i.e. synthesis analysis of the existing data sets) for a better understanding of the global CH_4 budget.

We first present a technical overview of atmospheric $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements and potential causes of measurement offsets among currently available data sets (Sect. 2), and then we summarize measurement methods by the laboratories that have conducted $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements for air and ice core samples (Sect. 3). In Sect. 4, we report new intercomparison exercises between some groups. We then link the intercomparison results through a survey of previously published intercomparisons and provide the current best estimates of measurement offsets among data sets from different laboratories (Sect. 5). Finally, we summarize the current status and briefly discuss possible causes of the measurement offsets as well as remaining issues that should be kept in mind when combining the use of currently existing data sets of isotopic composition of CH_4 (Sect. 6).

Table 1. List of laboratories that conduct measurements of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$. For each laboratory, measurement systems and relevant information that could have contributed, the interlaboratory measurement offsets are summarized. Brackets in the RM column indicate the laboratory from which the original standard scale was propagated. See Fig. 1 for overview of the past intercomparison exercises, Fig. 2 for intercomparison summary and the list of participating institution/project acronyms in the text for the laboratory names.

No.	Lab	System	IRMS	^{17}O correction ^a	Kr interference	Additional correction	RM		Reference ^c	$\delta^{13}\text{C}\text{-CH}_4$ offset with respect to NIWA	$\delta\text{D}\text{-CH}_4$ offset with respect to IMAU	Data availability ^d	
							$\delta^{13}\text{C}\text{-CH}_4$	$\delta\text{D}\text{-CH}_4$					
		Signifi- cance		Measure ^b		Intercomparison							
		Y		DI off- set/PCS		Drift cor- rection		Measur- ement					
1	NIWA	DI	Nuclide (1988–1996) MAT 252 (1996–2013) MAT 253 (2013–present) Isoprime	C1/C2	N	N	NBS-19, IAEA-CO-9, LSVEC	R1	R2, R3, R4, R5, R6, R7, R8, R9, R10, R11, R12, R13	Not measured	Not measured	WDCGG & NIWA website (R1, R2, R14)	
2	IMAU	CF	Delta plus XL	C2	Y	PCS	NBS-19 (MPIC)	R18	R10, R12, R18, R19, R20, R21, R22, R23	–	–	on request to Hinrich Schaefer (R17)	
3	MPIC	DI & TDLAS	MAT 252	C2	N	N	NBS-19	R27	R18, R19	-0.04 ± 0.07	$+0.3 \pm 1.1$	Utrecht University website (R25, R26)	
4	MPLBGC	CF	Delta V Plus	C4	Y	PCS	NBS-19 & LSVEC	R29	R21, R30	-0.07 ± 0.12	-4.2 ± 1.2	On request to Peter Bergam- aschi (R28)	
5	UCI	DI	MAT 252	C2	N	N	NBS-19 & IAEA-CO-9	R7, R31	R7, R12, R15, R22, R32, R33	-0.01 ± 0.09	-10.7 ± 0.7	On request to Heiko Moossen (R21, R29)	
6	TU	CF	Delta Plus XP	C3	Y	No	NBS-19	R37	R23, R37	$+0.36 \pm 0.08$	-13.1 ± 0.6	WDCGG (R35) & On request to Taku Umezawa (R36, R38)	
7	NIPR	CF	MAT 252	C3	N	N	NBS-19 (TU)	R8	R8, R23, R37	$+0.33 \pm 0.04$	Not measured	TU website (R8, R39)	
8	UW	DI	MAT 251	Information not available	N	–	NBS-19	R3	R2, R3, R4, R32	$+0.02 \pm 0.14$	No comparison available	CDIAC (R3)	
9	UHEI	DI	MAT 252	C2 with coefficients $a = 0.5$ & $K = 0.008335$	N	–	RM 8562, RM 8563 & RM 8564	R5	R4, R5, R6, R19, R30	-0.17 ± 0.03	-3.8 ± 1.3	UHEI website (R4)	
10	INSTAAR	CF	Optimal/isoprime	C3	Y	PCS	NBS-19 & IAEA-CO-9 (UCI)	R15	R4, R7, R9, R11, R12, R13, R15, R19, R41	$+0.13 \pm 0.02$	-13.2 ± 1.3	WDCGG & NOAA/ESRL/GMD website (R15)	

Table 1. Continued.

No	Lab	System	IRMS	^{17}O correction ^a	Kr interference	Additional correction	RM	Reference ^e	$\delta^{13}\text{C}\text{-CH}_4$ offset with respect to NIWA	$\delta\text{D}\text{-CH}_4$ offset with respect to IMAU	Data availability ^d
11	RHUL	CF DI	Isoprime Prism	C3	N	–	NBS-19 & IAEA-CO-9 Not mea- sured	R43 R9, R11, R13	+0.12 ± 0.03	Not mea- sured	CEDA & On request to Euan G. Nisbet (R11)
12	PDX	CF	Delta V	C2	Y	DI off- set/PCS	NBS-19 & IAEA-CO-9 (UCI)	R33, R34 R32, R33	+0.08 ± 0.09	–10.7 ± 1.5	PDX website (R32)
13	PSU	CF	MAT 252	C2	Y	Raw ion current correction/DI offset/PCS	RM 8563 & RM 8564 Oztech Gas (UCI)	R20, R44 R10, R12, R19, R20	+0.03 ± 0.16	–12.1 ± 1.5	NSIDC & On request to Todd Sowers
14	UB	CF	Isoprime	C2	Y	PCS	NBS-19 & IAEA-CO-9 (UCI via INSTAAR)	R10, R24, R41, R45 R10, R12, R20, R41 R12, R20, R24, R45	–0.18 ± 0.09	0.0 ± 1.6	PANGAEA (R46, R47)
15	AWI	CF	Isoprime	C3	Y	No	RM 8562, RM 8563 and RM 8564 (UHEH)	R6, R10, R19, R20	–0.09 ± 0.06	Not mea- sured	PANGAEA (R19)
16	CIC	DI ($\delta^{13}\text{C}$ - CH_4) TC/EA-IRMS & Pi-carro (δD - CH_4) CF	Delta V Plus Delta V Ad- vantage and Pi- carro	C2	N	–	RM 8563 VSMOW-2 & SLAP-2	R48 R21, R48 R21, R48	–0.03 ± 0.11	–2.1 ± 2.1	All data in papers (R16, R21, R48)

^a Cf. Allison et al. (1995), C2; Santrock et al. (1985), C3; Chang (1957), C4; Assonov and Bromberg (2003), b. Raw ion current correction; the Kr interference was corrected by subtracting the Kr-related anomalies in the raw ion current data. DI offset: the Kr interference was corrected by an offset relative to a DI IRMS measurement. PCS: Kr was separated by a post-combustion separation column or cryogenically. See Sect. 2.3. ^c R1: Lowe et al. (1991), R2: Lowe et al. (1994), R3: Quay et al. (1999), R4: Levin et al. (2012), R5: Paul (2003), R6: Behrens et al. (2008), R7: Tyler et al. (2007), R8: Morimoto et al. (2010), R9: Nisbet (2005), R10: Schmitt et al. (2013), R11: Nisbet et al. (2010), R12: this study (Sect. 4.4), R13: this study (Sect. 4.6), R14: Bergamaschi et al. (2001), R15: Miller et al. (2002), R16: Sperlich et al. (2005), R17: Ferrel et al. (2005), R18: Brass and Rockmann (2010), R19: Mollet et al. (2013), R20: Smart et al. (2011), R21: Sperlich et al. (2010), R22: this study (Sect. 4.1), R23: this study (Sect. 4.2), R24: Bock et al. (2010), R25: Rockmann et al. (2011), R26: Rockmann et al. (2009), R27: Bergamaschi et al. (1994), R28: Bergamaschi et al. (2003), R29: Brand et al. (2016), R30: this study (Sect. 4.3), R31: Tyler et al. (1999), R32: Rice et al. (2016), R33: Fauna (2013), R34: Rice et al. (2001), R35: Umezawa et al. (2012a), R36: Umezawa et al. (2012b), R37: Umezawa et al. (2009), R38: Umezawa et al. (2011), R39: Morimoto et al. (2017), R40: this study (Sect. 4.3), R41: Schmitt et al. (2014), R42: Lowry et al. (2001), R43: Fisher et al. (2006), R44: Sowers et al. (2005), R45: Bock et al. (2014), R46: Bock et al. (2010), R47: Bock et al. (2017), R48: Sperlich et al. (2012), d. WDCGG (World Data Centre for Greenhouse Gases): <http://ds.data.jma.go.jp/gm/ds/deg/wdgg.html>, NIWA website: www.niwa.co.nz, Utricht University website: <http://www.projects.science.uu.nl/atmosphere/imate/Data.php>, TU website: <http://enos.sakura.ne.jp/gf/datan/>, CDLAC (Carbon Dioxide Information Analysis Center): <http://cdiac.ess-dive.lbl.gov/pubs/ndb/1022db/1022.html> (UCI), <http://cdiac.ess-dive.lbl.gov/pubs/ndb/1022db/1022.html> (UCI), UHEH website: <http://www.projects.science.uu.nl/atmosphere/imate/Data.php>, NOAA/ES&R/GMD website: <http://www.esr.noaa.gov/gmd/valdat/PDX> website: <http://webpdx.edu/~arcedian>, CH4 hint, PANGAEA: <https://doi.org/10.1594/PANGAEA.873918>, CEDA (Centre for Environmental Data Analysis): <http://www.ceda.ac.uk>, Hinrich Schaefer: h.schaefer@niwa.co.nz, Peter Bergamaschi: peter.bergamaschi@ec.europa.eu, Heiko Moosser: heiko.moosser@bgc-jena.mpg.de, Taku Umezawa: umezawa.taku@nics.go.jp, Euan G. Nisbet: enisbet@hul.ac.uk, Todd Sowers: ts1@psu.edu.

2 Overview of atmospheric $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurement techniques

2.1 IRMS measurements for $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$

In the 1990s, atmospheric $\delta^{13}\text{C}\text{-CH}_4$ ($\delta\text{D}\text{-CH}_4$) was analysed using an offline technique in which CH_4 was separated from the sample air and converted to CO_2 (H_2) for subsequent offline $\delta^{13}\text{C}\text{-CH}_4$ ($\delta\text{D}\text{-CH}_4$) analyses by dual-inlet isotope ratio mass spectrometry (DI-IRMS; e.g. Stevens and Rust, 1982; Lowe et al., 1991; Quay et al., 1991, 1999; Sugawara et al., 1996; Poß, 2003). The original methodology was based on the combustion of CH_4 in sample air, with interfering compounds such as CO_2 , H_2O , N_2O , CO and non-methane hydrocarbons being removed cryogenically, chemically or by gas chromatography before CH_4 combustion. The number of measurements was limited not only because of laborious and time-consuming laboratory procedures but also because large volumes of air sample were required ($> 100\text{L}_{\text{STP}}$ for $\delta\text{D}\text{-CH}_4$). Later, a method based on a continuous-flow gas chromatography isotope ratio mass spectrometry (GC-IRMS) technique combined with combustion and pyrolysis furnaces became available (Merritt et al., 1995; Burgoyne and Hayes, 1998; Hilkert et al., 1999), which dramatically reduced time and effort in the laboratory and likewise the amount of sample air required (now typically $100\text{mL}_{\text{STP}}$). Such systems are now used in most laboratories worldwide to acquire $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data in the current and past atmosphere (Rice et al., 2001; Miller et al., 2002; Sowers et al., 2005; Ferretti et al., 2005; Morimoto et al., 2006; Fisher et al., 2006; Behrens et al., 2008; Umezawa et al., 2009; Brass and Röckmann, 2010; Sperlich et al., 2013; Schmitt et al., 2014; Bock et al., 2014; Brand et al., 2016; Röckmann et al., 2016). Although these systems use a similar measurement principle, they vary in the use of pre-concentration of CH_4 in sample air, GC separation and combustion/pyrolysis, data corrections and in the specific IRMS instrument among laboratories (see Schmitt et al., 2013, Sect. 3 and Table 1). Besides analysis by mass spectrometry, laser-based spectroscopy has also been developed for atmospheric $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements (Bergamaschi et al., 2000; Eyer et al., 2016), but detailed discussion on the technique is beyond the scope of this study.

2.2 Standard scales

VPDB and VSMOW are the standard scales for $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$, respectively. To make measurements traceable to these standard scales, each laboratory needs to calibrate its laboratory reference gases against reference materials (RMs) with known values on the standard scales. In this study, the term “calibration” means to measure a laboratory gas (for instance a laboratory working standard gas that is routinely compared with samples) against a standard at higher hierarchy level and to assign to that working standard a $\delta^{13}\text{C}\text{-CH}_4$

or $\delta\text{D}\text{-CH}_4$ value traceable to the standard scale. In principle, all measurements at individual laboratories intend to ultimately anchor their working standards and sample gases to the VPDB or VSMOW scale using the RMs provided by the International Atomic Energy Agency (IAEA) or National Institute of Standards and Technology (NIST; Coplen et al., 2006; Brand et al., 2014). However, since RMs and recommended calibration methods for measurements of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ in air have not yet been provided (Sperlich et al., 2012, 2016), individual groups have developed their own calibration strategies.

Since the $\delta^{13}\text{C}\text{-CH}_4$ measurement by IRMS is taken by $\delta^{13}\text{C}$ analysis in CO_2 oxidized from CH_4 in air, some laboratories use pure CO_2 gases as a working standard. In many laboratories, these internal CO_2 standard gases were calibrated against pure CO_2 produced from the primary anchor of the VPDB-scale NBS-19 or other RMs by using DI-IRMS (Table 1). Since the typical atmospheric $\delta^{13}\text{C}\text{-CH}_4$ value (about -47‰) differs considerably from the $\delta^{13}\text{C}$ value of NBS-19 ($+1.95\text{‰}$), some laboratories have used other RMs with VPDB values close to atmospheric $\delta^{13}\text{C}\text{-CH}_4$ such as LSVEC (lithium carbonate reference material prepared by Harry J. Svec), IAEA-CO-9 and RM 8563 as a second anchoring point of the VPDB scale (see Table 1). This minimizes the risk of significant errors in realization of the standard scale (due to scale contraction or ^{17}O correction, described in the following sections). A standard scale established this way at an individual laboratory was often propagated to laboratory-internal CO_2 standard gases at lower hierarchy levels, and they were used as the reference in DI-IRMS or GC-IRMS measurement of CO_2 processed from CH_4 in sample air. Ideally, this accurately links $\delta^{13}\text{C}\text{-CH}_4$ of the sample to the international isotope scale. In contrast, it has been recommended that a measured value of a sample is determined against a reference gas that undergoes the all preparation steps in the sample measurement line in order to cancel out possible isotopic fractionations due to different treatment between the sample and reference gases (principle of identical treatment; Werner and Brand, 2001). This concept has been taken into account in some laboratories; a working standard is calibrated for $\delta^{13}\text{C}\text{-CH}_4$ and sample measurements are referenced by comparison with measurements of that working standard processed in the same manner (e.g. Brand et al., 2016). Despite intentions of best traceability to RMs, the variety of calibrations has resulted in diverse realizations of the VPDB scale across $\delta^{13}\text{C}\text{-CH}_4$ measurement programmes. As in Table 1, the different RMs that have been applied to $\delta^{13}\text{C}\text{-CH}_4$ calibration include NBS-19 (limestone), IAEA-CO-9 (barium carbonate), LSVEC (lithium carbonate) and RM 8562–8564 (CO_2); see Coplen et al. (2006), Brand et al. (2014) and Sperlich et al. (2016). It is also noted that uncertainties of assigned values for these RMs range up to a few tenths per mille and the assigned values have been revised over time (Brand et al., 2014), which might have complicated the realization of the standard scale at each labora-

tory. Furthermore, most of these RMs are in different chemical forms, and different isotopic fractionations may have occurred during acid digestion to CO_2 , which could have biased calibrations at each laboratory. Lastly, the WMO (2016) has reported exhaustion of NBS-19 and instability of LSVEC, both of which are critical RMs for the VPDB scale. Associated possible revision of $\delta^{13}\text{C}$ values of RMs in the future will affect the consistency of the data sets from different laboratories.

For $\delta\text{D}-\text{CH}_4$, in the conventional offline measurements, CH_4 in sample air needs to be processed to H_2O followed by reduction to H_2 for a subsequent DI-IRMS measurement. GC-IRMS requires pyrolysis of CH_4 to H_2 . Therefore, individual laboratories have prepared internal standards of H_2O (liquid) or H_2 (gas), which were calibrated against primary RMs (water) or H_2 reference gases certified for δD (Table 1). Although the situation is less complicated compared to $\delta^{13}\text{C}-\text{CH}_4$ in terms of variety in chemical properties of RMs, the lack of RMs for $\delta\text{D}-\text{CH}_4$ forced laboratories to develop their calibration method standard scale individually. It is also noted that, similarly to $\delta^{13}\text{C}-\text{CH}_4$, this principle of identical treatment has not been followed strictly at the all laboratories. If not followed, sample measurements are subject to subtle changes in conditions of the all preparation steps (e.g. conversion of CH_4), while such changes do not affect the measured value of a reference gas injected directly into the IRMS.

2.3 Scale contraction

It has been found that cross contamination between sample and reference CO_2 gases shrinks the $\delta^{13}\text{C}$ distance measured on DI-IRMS (Meijer et al., 2000; Verkouteren et al., 2003a, b). This effect is known as the scale contraction or η effect, and the magnitude is specific to the IRMS instrument and its settings. Since the VPDB scale for $\delta^{13}\text{C}-\text{CH}_4$ has been realized and propagated via CO_2 calibrations by DI-IRMS at individual laboratories, the instrument-dependent scale contraction effect could have caused a significant difference in measurement values, especially at the low $\delta^{13}\text{C}$ values of atmospheric CH_4 of about -47% (Wendeborg et al., 2013).

2.4 ^{17}O correction

For measurement of $\delta^{13}\text{C}-\text{CH}_4$ by IRMS, CH_4 is first oxidized to CO_2 and the different isotopic variants of the produced CO_2 are then registered on Faraday cups with mass-to-charge ratios m/z of 44, 45 and 46. Since the raw ion beam intensity for $m/z = 45$ is the sum of $^{13}\text{C}^{16}\text{O}_2$ and $^{12}\text{C}^{17}\text{O}^{16}\text{O}$, the final $\delta^{13}\text{C}$ value is obtained by correcting for the contribution of the ^{17}O -containing isotopologue, known as ^{17}O correction (e.g. Assonov and Brenninkmeijer, 2003). Several algorithms such as Craig (1957) and Santrock et al. (1985) have been suggested (see Assonov and Brenninkmeijer, 2003 and references therein) and implemented

into software/programmes of the IRMS companies and individual laboratories. Assonov and Brenninkmeier (2003) showed that the bias caused by different ^{17}O -correction algorithms could exceed general repeatability achieved by IRMS measurements. The ^{17}O -correction method of each laboratory is listed in Table 1.

2.5 Krypton interference in GC-IRMS

The transition from DI-IRMS to GC-IRMS analyses reduced the analytical effort, but also introduced complications that were initially not recognized and taken into account. It was recently found that atmospheric krypton (Kr) interferes with the $\delta^{13}\text{C}-\text{CH}_4$ GC-IRMS analysis if Kr is present in the ion source during the data acquisition of the CO_2 peak generated from CH_4 oxidation (hereafter CH_4 -derived CO_2 peak) (Schmitt et al., 2013). Thus the $\delta^{13}\text{C}-\text{CH}_4$ measurements on a GC-IRMS system can be biased if Kr is not sufficiently separated either from CH_4 or from the CH_4 -derived CO_2 peak after the CH_4 combustion. Schmitt et al. (2013) demonstrated that the doubly charged krypton isotope $^{86}\text{Kr}^{2+}$, produced in the ion source of an IRMS, can cause lateral tailing extending into the Faraday cups used for $\delta^{13}\text{C}$ analysis (i.e. m/z of 44, 45 and 46), which compromises the measured signal of the CH_4 -derived CO_2 peak. This effect had not been recognized for more than a decade since the early years of GC-IRMS measurements (Merritt et al., 1995) and thus has not been taken into account in many data sets of atmospheric $\delta^{13}\text{C}-\text{CH}_4$ reported in the meantime (e.g. Miller et al., 2002; Morimoto et al., 2006; Fisher et al., 2011; Röckmann et al., 2011; Umezawa et al., 2012a, b). Furthermore, because the Kr effect is system dependent and variable with time (Schmitt et al., 2013), applying plausible corrections to past data may not be feasible. Likewise, several gas species including Kr can affect $\delta\text{D}-\text{CH}_4$ measurements, and this effect is also system dependent (Bock et al., 2014).

Several solutions have been suggested to eliminate or account for the Kr interference (Schmitt et al., 2013). Among them, three methods have been implemented at different laboratories (Table 1). Briefly, (1) after the CH_4 oxidation to CO_2 , Kr is separated from the CH_4 -derived CO_2 by using a post-combustion separation column (PCS) or cryogenically. (2) An offset due to the Kr interference is estimated by comparison with a DI-IRMS measurement (DI offset). (3) The Kr interference peak is subtracted from the raw ion current time series of the IRMS acquisition (raw ion current correction). A more detailed description has been presented in Schmitt et al. (2013).

3 Measurements of participating laboratories

In this section, we briefly document measurement systems of individual laboratories for ease of reference in the following intercomparisons (Sects. 4 and 5). For details, we refer to

more dedicated publications listed in Table 1. The table also visualizes differences among laboratories in terms of possible causes of the measurement offsets described in Sect. 2.

3.1 NIWA

The National Institute for Water and Atmospheric Research (NIWA, originally INS (Institute of Nuclear Sciences) and later INGS (Institute of Nuclear and Geological Sciences) until 1992) successfully initiated systematic measurements of atmospheric $\delta^{13}\text{C}\text{-CH}_4$ by means of offline CH_4 separation and conversion followed by a DI-IRMS measurement in 1988 (Lowe et al., 1988, 1991). A suite of CO_2 working gases with $\delta^{13}\text{C}\text{-CH}_4$ values around -47‰ referenced to IAEA materials were utilized to calibrate the measurements. An overall reproducibility of the $\delta^{13}\text{C}\text{-CH}_4$ measurement was evaluated to be 0.02‰ (Lowe et al., 1991). The $\delta^{13}\text{C}\text{-CH}_4$ measurements at NIWA are ultimately calibrated against CO_2 produced from NBS-19, IAEA-CO-9 and LSVEC. The long-term $\delta^{13}\text{C}\text{-CH}_4$ records have been presented since then (Lowe et al., 1994, 1997, 2004; Bergamaschi et al., 2001; Schaefer et al., 2016). Bromley et al. (2012) reported that repeated measurements of the two working reference gases and archived air indicated no detectable drift over 16 years since 1992. NIWA has also operated a GC-IRMS system since 2004 (Ferretti et al., 2005) with reproducibility of 0.1‰ . The Kr interference on the GC-IRMS $\delta^{13}\text{C}\text{-CH}_4$ measurement has been identified and is corrected by an offset relative to the conventional DI-IRMS measurement (see Sect. 4.1).

3.2 IMAU

The GC-IRMS system at the Institute for Marine and Atmospheric research Utrecht (IMAU) has been described by Brass and Röckmann (2010). The measurement reproducibility is estimated to be 0.07 and 2.3‰ for $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$, respectively. Sample air is measured against reference air that is processed in the GC-IRMS system in the same manner as a sample. The IMAU $\delta^{13}\text{C}\text{-CH}_4$ standard scale is based on a set of assigned values for 13 firn air samples measured at Max Planck Institute for Chemistry (MPIC; Bräunlich et al., 2001) and they are ultimately referenced to a CO_2 gas produced from NBS-19 (Röckmann, 1998; Bergamaschi et al., 2000). The $\delta\text{D}\text{-CH}_4$ standard scale is based on a set of reference gases originally produced at the MPIC (see Sect. 2.3). These calibration details have also been documented by Sperlich et al. (2016). The IMAU system was originally affected by Kr but later modified to remove this interference. A correction was applied for data obtained before the system modification (Schmitt et al., 2013).

3.3 MPIC

The MPIC has reported $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements at a baseline station (Bergamaschi et al., 2000) and

for firn air samples (Bräunlich et al., 2001) based on an offline DI-IRMS measurement for $\delta^{13}\text{C}\text{-CH}_4$ (Bergamaschi et al., 2000) and a tunable diode laser-based absorption spectrometer (TDLAS) for $\delta\text{D}\text{-CH}_4$ (Bergamaschi et al., 1994). Some firn air measurements by Bräunlich et al. (2001) were performed by using a GC-IRMS system at the Laboratory of Glaciology and Geophysics of the Environment. As described in Sect. 3.2, the $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ standard scales of MPIC are based on that of IMAU. For the $\delta^{13}\text{C}\text{-CH}_4$ DI-IRMS measurement, the CH_4 -derived CO_2 was measured against a working standard (pure CO_2) that was calibrated against NBS-19 on a DI-IRMS system (Röckmann, 1998; Bergamaschi et al., 2000). The MPIC $\delta\text{D}\text{-CH}_4$ scale is based on measurements of standard gases at the Bundesanstalt für Geowissenschaften und Rohstoffe, Hannover, Germany. CH_4 was combusted to CO_2 and H_2O , followed by reduction of H_2O to H_2 for subsequent DI-IRMS analysis on H_2 . They were calibrated against VSMOW and SLAP (Bergamaschi et al., 2000). The measurements of atmospheric $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ at the MPIC were discontinued.

3.4 MPI-BGC

The Max Planck Institute for Biogeochemistry (MPI-BGC) set up a GC-IRMS system for $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements, and it has been operated for air samples collected at baseline stations (Brand et al., 2016). The long-term (3 years) reproducibility was assessed to be 0.12 for $\delta^{13}\text{C}\text{-CH}_4$ and 1.0‰ for $\delta\text{D}\text{-CH}_4$. Initially, the GC-IRMS measurements had been anchored to a working standard air calibrated by IMAU. The Kr effect was eliminated by a PCS column, and the initial calibration has in the meantime been replaced by a new primary calibration, where measurements are ultimately anchored to NBS-19 and LSVEC for $\delta^{13}\text{C}\text{-CH}_4$ and VSMOW-2 and SLAP-2 for $\delta\text{D}\text{-CH}_4$ (Sperlich et al., 2016). This calibration, termed JRAS-M16, is the basis for the $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ values from MPI-BGC reported in this paper.

3.5 UCI

The University of California Irvine (UCI) measured atmospheric $\delta^{13}\text{C}\text{-CH}_4$ by offline DI-IRMS and $\delta\text{D}\text{-CH}_4$ by GC-IRMS (Tyler et al., 1999, 2007; Kai et al., 2011). The UCI GC-IRMS system for both $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ has been described in detail by Rice et al. (2001). The measurement reproducibility of the GC-IRMS system was estimated to be 0.05 and 1.5‰ for $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$, respectively, while that of the offline DI-IRMS $\delta^{13}\text{C}\text{-CH}_4$ measurement was 0.05‰ . Samples were measured against laboratory working standard gases of pure CO_2 for $\delta^{13}\text{C}\text{-CH}_4$ and pure H_2 for $\delta\text{D}\text{-CH}_4$. The $\delta^{13}\text{C}\text{-CH}_4$ calibration is based on a CO_2 reference gas provided by NIWA, which was compared with CO_2 produced from NBS-19 and IAEA-CO-9 (Lowe et al., 1999). The $\delta\text{D}\text{-CH}_4$ calibration is referenced to three H_2 gas

cylinders purchased from Oztech Gas Company (Rice et al., 2001). The possible Kr interference on the GC-IRMS system is unclear (the laboratory is now closed), but it appears that the Kr effect had been avoided using liquid nitrogen cooling of the GC column as surmised by Schmitt et al. (2013).

3.6 TU

The GC-IRMS system at Tohoku University (TU) has been described by Umezawa et al. (2009). The measurement reproducibility is estimated to be 0.08 for $\delta^{13}\text{C}\text{-CH}_4$ and 2.2‰ for $\delta\text{D}\text{-CH}_4$. Sample measurements are made against pure CO_2 and H_2 working standard gases for $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$, respectively. The $\delta^{13}\text{C}\text{-CH}_4$ calibration is based on a CO_2 primary gas produced from NBS-19. The H_2 working standard for the $\delta\text{D}\text{-CH}_4$ measurement is referenced to water laboratory standards that are calibrated against VSMOW and SLAP. Measured $\delta\text{D}\text{-CH}_4$ values are corrected so that the value of a laboratory test gas is kept constant over time to take into account fluctuations in the measured value due to the condition of the pyrolysis furnace (Umezawa et al., 2009, 2012a). The Kr interference in the $\delta^{13}\text{C}\text{-CH}_4$ measurement was identified, but modification or correction has not been implemented. It has been documented that the $\delta^{13}\text{C}\text{-CH}_4$ measurement at TU shifted by +0.27‰ after July 2008 (the cause of this sudden shift has yet to be identified) and measurements afterwards were corrected for this value to keep the data consistent (Umezawa et al., 2012a, b). Note that TU made a rigorous re-evaluation of the long-term measurements of their working standard gas recently, and the TU $\delta^{13}\text{C}\text{-CH}_4$ data sets will be revised accordingly. Therefore, the comparison numbers presented here are not comparable to those for earlier publications (Umezawa et al., 2009, 2011, 2012a, b).

3.7 NIPR

The National Institute of Polar Research (NIPR) reported $\delta^{13}\text{C}\text{-CH}_4$ measurements at an Arctic site using a GC-IRMS system (Morimoto et al., 2006, 2017). The measurement reproducibility was evaluated to be 0.06‰. The $\delta^{13}\text{C}\text{-CH}_4$ calibration follows same procedure as TU. By injecting different quantities of Kr, it was confirmed that ambient Kr does not significantly interfere with the $\delta^{13}\text{C}\text{-CH}_4$ measurements at NIPR.

3.8 UW

The University of Washington (UW) reported extensive $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements using an offline DI-IRMS system (Quay et al., 1991, 1999). The reproducibility was estimated to be 0.1 for $\delta^{13}\text{C}\text{-CH}_4$ and 3–4‰ for $\delta\text{D}\text{-CH}_4$. The $\delta^{13}\text{C}\text{-CH}_4$ calibration is based on measurements made against NBS-19 (Quay et al., 1999), while the earlier measurements were calibrated against NBS-20 and NBS-16 (Quay et al., 1991). The $\delta\text{D}\text{-CH}_4$ was anchored

to calibration by VSMOW and SLAP. Systematic measurements of air standards showed that no significant time shift ($+0.001 \pm 0.002\text{‰ yr}^{-1}$) affected their $\delta^{13}\text{C}\text{-CH}_4$ data set for 1988–1995 (Quay et al., 1999).

3.9 UHEI

The University of Heidelberg (UHEI) carried out $\delta^{13}\text{C}\text{-CH}_4$ measurements by DI-IRMS (Levin et al., 1999, 2012). The typical measurement reproducibility was evaluated to be 0.05‰ (Levin et al., 1999). The UHEI $\delta^{13}\text{C}\text{-CH}_4$ measurements are calibrated against CO_2 reference materials (RM 8562, RM 8563 and RM 8564; Behrens et al., 2008). Although reported previously only for signatures of source CH_4 (Levin et al., 1993), UHEI also took offline $\delta\text{D}\text{-CH}_4$ measurements of atmospheric samples by DI-IRMS and TDLAS (Poß, 2003). The $\delta\text{D}\text{-CH}_4$ measurements by DI-IRMS were taken for pure H_2 (H_2O from CH_4 oxidation converted to H_2 with zinc as catalyst) and were calibrated against VSMOW and SLAP. Note that UHEI recently re-evaluated all their atmospheric $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements rigorously, based on the history of laboratory standards used; therefore, comparison numbers published in earlier works are not comparable to the revised values presented here.

3.10 INSTAAR

The Institute of Arctic and Alpine Research (INSTAAR) of the University of Colorado, Boulder has measured $\delta^{13}\text{C}\text{-CH}_4$ and, intermittently, $\delta\text{D}\text{-CH}_4$ using a GC-IRMS system for flask air samples from the cooperative sampling network of National Oceanic and Atmospheric Administration (NOAA; Miller et al., 2002). Reproducibilities of the $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements are evaluated to be 0.08 and 2‰, respectively (Miller et al., 2002; White et al., 2016). The INSTAAR $\delta^{13}\text{C}\text{-CH}_4$ measurement currently follows the UCI calibration, while the $\delta\text{D}\text{-CH}_4$ measurement is not explicitly anchored to the VSMOW scale (White et al., 2016). The Kr interference in the $\delta^{13}\text{C}\text{-CH}_4$ measurement is significant, and a PCS column was therefore implemented into the system in May 2017. Correction of the data for the Kr interference (1998–present) is under evaluation. Of the data presented here, only the ice core intercomparison round robin (Sect. 3.4) and the INSTAAR–MPI-BGC comparison (Sect. 3.5) have not been interfered with by Kr.

3.11 RHUL

Royal Holloway University of London (RHUL) measured atmospheric $\delta^{13}\text{C}\text{-CH}_4$ using an offline DI-IRMS technique (Lowry et al., 2001) and a GC-IRMS system (Fisher et al., 2006, 2011; Nisbet et al., 2016). Reproducibility of the DI-IRMS measurement was evaluated to be 0.04‰ (Lowry et al., 2001) and that by the GC-IRMS is 0.05‰ (Fisher et al., 2006). They ultimately calibrated $\delta^{13}\text{C}\text{-CH}_4$ to IAEA carbonate materials NBS-19 and IAEA-CO-9 (Lowry et al.,

2001; Fisher et al., 2006). Note that RHUL applies an offset correction of -0.20‰ for the measured value by GC-IRMS (Sects. 4.6 and 5.11).

3.12 PDX

Portland State University (PDX) reported $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements for archive air samples (Rice et al., 2016). The PDX measurement system has been described in Teama (2013) with some updates since Rice et al. (2001). The $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ reproducibilities are 0.07 and 2.0‰, respectively, and PDX shares the standard scales with UCI for both $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ (Rice et al., 2016).

3.13 PSU

Pennsylvania State University (PSU) reported $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data from ice cores and firn air using a GC-IRMS system (e.g. Sowers et al., 2005; Sowers, 2010). The overall measurement reproducibility including every step for ice core measurements was evaluated to be 0.3 for $\delta^{13}\text{C}\text{-CH}_4$ and 3‰ for $\delta\text{D}\text{-CH}_4$ (Sowers, 2010). The PSU $\delta^{13}\text{C}\text{-CH}_4$ measurements are calibrated against CO_2 RMs (RM 8563 and RM 8564). The $\delta\text{D}\text{-CH}_4$ calibration is against H_2 gas bottles from Oztech Gas Company (Sowers, 2006).

3.14 UB

The University of Bern (UB) makes $\delta^{13}\text{C}\text{-CH}_4$ measurements from ice cores using a GC-IRMS system with an overall reproducibility of 0.15‰ (Schmitt et al., 2014; Bock et al., 2017). The UB measurements are referenced to a whole-air working standard with a CH_4 mole fraction of 1508.2 ppb and an assigned $\delta^{13}\text{C}\text{-CH}_4$ value of $-47.34 \pm 0.02\text{‰}$ (named “Boulder, CA08289” in Schmitt et al., 2014). This value is anchored to the standard scale used at INSTAAR (Sect. 3.10). UB also measures $\delta\text{D}\text{-CH}_4$ for ice core samples (Bock et al., 2010, 2014, 2017). The overall measurement precision for ice core sample (including extraction of air from an ice sample) was evaluated to be 2.3‰. The UB $\delta\text{D}\text{-CH}_4$ measurement is referenced by using an ambient air cylinder (named “Air Controlé”) with a $\delta\text{D}\text{-CH}_4$ value of $-93.6 \pm 2.8\text{‰}$, which was cross-referenced to a high-pressure cylinder filled at the Alert Station (“Alert 2002/11” with $\delta\text{D}\text{-CH}_4$ of $-82.2 \pm 1.0\text{‰}$) analysed on the scale maintained at UHEI (Bock et al., 2010, 2014). However, this value has to be corrected to $-85.2 \pm 1.0\text{‰}$ to account for the recent re-evaluation at UHEI (Sect. 3.9). All UB data published after 2011 are free of Kr interference.

3.15 AWI

The Alfred Wegener Institute Helmholtz Centre for Polar and Marine Research (AWI) reported $\delta^{13}\text{C}\text{-CH}_4$ measurements from ice cores using a GC-IRMS system (Behrens et al., 2008; Fischer et al., 2008; Möller et al., 2013). The measure-

ment reproducibility was estimated to be 0.2‰. The $\delta^{13}\text{C}\text{-CH}_4$ measurements employed the UHEI standard scale via comparison of measurements of an air sample collected at Neumayer Station, Antarctica (Möller et al., 2013).

3.16 CIC

The Centre for Ice and Climate (CIC) of the Niels Bohr Institute has reported $\delta^{13}\text{C}\text{-CH}_4$ measurements from ice cores (Sperlich et al., 2015) using a GC-IRMS system with measurement reproducibility of 0.09‰ (Sperlich et al., 2013). CIC also set up an offline combustion system for samples with a large amount of CH_4 , which is combined with DI-IRMS for $\delta^{13}\text{C}\text{-CH}_4$ and with either a high temperature conversion/elemental analyser (TC/EA) coupled to IRMS or laser spectroscopy for $\delta\text{D}\text{-CH}_4$ (Sperlich et al., 2012); the measurement reproducibility is 0.04 for $\delta^{13}\text{C}\text{-CH}_4$ and 0.7‰ for $\delta\text{D}\text{-CH}_4$. The CIC measurements are referenced to RM 8563 for $\delta^{13}\text{C}\text{-CH}_4$ and VSMOW-2 and SLAP-2 for $\delta\text{D}\text{-CH}_4$. The combined uncertainty of this analytical system including the uncertainty of the entire traceability chain was estimated at 0.07 for $\delta^{13}\text{C}\text{-CH}_4$ and 0.7‰ for $\delta\text{D}\text{-CH}_4$ (Sperlich et al., 2016).

4 Intercomparison exercises

4.1 Intercomparison between UCI and IMAU

An intercomparison between UCI and IMAU was made by analysing six air samples at both laboratories; the air samples were collected along a flight track of commercial aircraft in the upper troposphere in the early phase of the CARIBIC (Civil Aircraft for the Regular Investigation of the atmosphere Based on an Instrument Container) project (Brenninkmeijer et al., 1999). The original samples were collected into large stainless steel cylinders (21 L) and aliquots of them were transferred into smaller stainless steel canisters (~ 2.3 L) for storage after delivery to the MPIC laboratory. Different subsamples from identical original samples were sent to UCI and IMAU for analysis, and they were measured at UCI in 2008 and at IMAU in 2012 to 2013. The measurement results at both laboratories are summarized in Table 2. The result indicated significant differences of $+0.42 \pm 0.04\text{‰}$ for $\delta^{13}\text{C}\text{-CH}_4$ (UCI value is higher than at IMAU) and of $-10.7 \pm 0.7\text{‰}$ for $\delta\text{D}\text{-CH}_4$ (UCI value is lower than IMAU).

4.2 Intercomparison between TU/NIPR and IMAU

An intercomparison between TU/NIPR and IMAU was carried out during 2013–2015. The TU laboratory prepared four stainless steel canisters (~ 1 L) filled with dried ambient air (canisters MD1 and MD2) and CH_4 -in-synthetic air gas (canisters MD3 and MD4) with CH_4 mole fractions ranging from 899 to 2117 ppb on the TU CH_4 scale (Aoki et al., 1992;

Table 2. Result of intercomparison of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements between UCI and IMAU.

Sample ID	CH_4 (ppb) ^a	$\delta^{13}\text{C}\text{-CH}_4$ (‰) ^b			$\delta\text{D}\text{-CH}_4$ (‰) ^b		
		UCI	IMAU	UCI– IMAU	UCI	IMAU	UCI– IMAU
WAS-24-2	1784.7	-46.96 ± 0.07 ($N = 3$)	-47.33 ± 0.05 ($N = 3$)	+0.37	-91.6 (1.0, $N = 2$)	-78.9 ± 0.1 ($N = 4$)	–12.7
WAS-24-5	1825.8	-47.16 ($N = 1$)	-47.53 ± 0.02 ($N = 6$)	+0.37	-93.8 ($N = 1$)	-83.1 ± 0.2 ($N = 4$)	–10.7
WAS-24-6	1827.5	-47.08 (0.02, $N = 2$)	-47.55 ± 0.04 ($N = 6$)	+0.47	-92.1 (1.6, $N = 2$)	-83.6 ± 0.1 ($N = 4$)	–8.5
WAS-24-9	1799.8	-47.05 ($N = 1$)	-47.38 ± 0.02 ($N = 6$)	+0.33	-92.3 ± 1.8 ($N = 3$)	-79.8 ± 0.8 ($N = 4$)	–12.4
WAS-24-10	1789.8	-47.07 ($N = 1$)	-47.42 ± 0.02 ($N = 6$)	+0.35	-89.3 ($N = 1$)	-79.7 ± 0.8 ($N = 4$)	–9.6
WAS-24-11	1780.8	-46.77 ($N = 1$)	-47.37 ± 0.03 ($N = 6$)	+0.60	-89.0 (1.8, $N = 2$)	-78.7 ± 0.7 ($N = 4$)	–10.3
Average				$+0.42 \pm 0.04$ ^c			-10.7 ± 0.7 ^c

^a NOAA-2004 CH_4 scale (Dlugokencky et al., 2005). ^b Uncertainties are standard errors of the mean for measurements with $N \geq 3$. Difference in duplicate flask measurements ($N = 2$) is shown in parenthesis. ^c Uncertainties are standard errors of the mean for differences in the above lines.

Table 3. Result of intercomparison of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements between TU/NIPR and IMAU.

Sample ID	CH_4 (ppb) ^a	$\delta^{13}\text{C}\text{-CH}_4$ (‰) ^b				$\delta\text{D}\text{-CH}_4$ (‰) ^b		
		TU	NIPR	IMAU	Difference from IMAU	TU	IMAU	Difference from IMAU
MD1	1901.1	-47.04 ± 0.02 ($N = 16$)	-47.11 ± 0.02 ($N = 5$)	-47.40 ± 0.04 ($N = 9$)	$+0.36$ (TU) $+0.28$ (NIPR)	-97.2 ± 0.6 ($N = 10$)	-85.0 ± 0.1 ($N = 8$)	–12.2
MD2	2116.6	-46.81 ± 0.02 ($N = 16$)	-46.92 ± 0.03 ($N = 6$)	-47.26 ± 0.03 ($N = 9$)	$+0.45$ (TU) $+0.34$ (NIPR)	-118.5 ± 0.6 ($N = 10$)	-104.5 ± 0.3 ($N = 8$)	–14.0
MD3	899.1	-41.14 ± 0.04 ($N = 16$)	-41.05 ± 0.02 ($N = 5$)	-41.81 ± 0.03 ($N = 8$)	$+0.67$ (TU) $+0.76$ (NIPR)	-190.7 ± 0.6 ($N = 10$)	-175.8 ± 0.6 ($N = 8$)	–14.9
MD4	1700.5	-42.47 ± 0.03 ($N = 16$)	-42.43 ± 0.04 ($N = 5$)	-42.98 ± 0.02 ($N = 8$)	$+0.52$ (TU) $+0.56$ (NIPR)	-195.2 ± 0.6 ($N = 10$)	-180.6 ± 0.2 ($N = 8$)	–14.6
Average (ambient air)					$+0.40 \pm 0.04$ (TU) ^c $+0.31 \pm 0.03$ (NIPR) ^c			-13.1 ± 0.6 ^c
Average (all)					$+0.50 \pm 0.07$ (TU) ^c $+0.48 \pm 0.11$ (NIPR) ^c			-13.9 ± 0.9 ^c

^a Tohoku University CH_4 scale (Aoki et al., 1992; Umezawa et al., 2014). ^b Uncertainties are standard errors of the mean for the repetitive measurements. ^c Uncertainties are standard errors of the mean for differences in the above lines.

Umezawa et al., 2014; Table 3). The canisters were analysed at TU and then sent to IMAU, after which they were sent back to TU and reanalysed to confirm the stability of the air samples in the canisters during the intercomparison exercise. The measurements at TU before and after the transport to IMAU from April 2013 to July 2015 indicated that possible drifts during canister storage and transportation are small (< 0.1 for $\delta^{13}\text{C}\text{-CH}_4$ and < 3.5 ‰ for $\delta\text{D}\text{-CH}_4$). NIPR also measured the canisters for $\delta^{13}\text{C}\text{-CH}_4$. The results indicate significant differences of $+0.50 \pm 0.07$ ‰ for $\delta^{13}\text{C}\text{-CH}_4$ (TU value is higher than IMAU) and of -13.9 ± 0.9 ‰ for $\delta\text{D}\text{-CH}_4$ (TU value is lower than IMAU; Table 3). The measurements of the four canisters at NIPR were $+0.48 \pm 0.11$ ‰ higher than IMAU. However, the differences of $\delta^{13}\text{C}\text{-CH}_4$ measurements are smaller for the ambient air samples (MD1 and MD2) than the CH_4 -in-synthetic air samples (MD3 and MD4). It is also noted that the $\delta^{13}\text{C}\text{-CH}_4$ difference between the laboratories is largest for the low CH_4 mole fraction (~ 900 ppb) sample (MD3). The cause is unclear, but might be related to (1) deviation in $\delta^{13}\text{C}\text{-CH}_4$ of the latter samples from the typical atmospheric value, i.e. scale contraction effect; (2) difference in air matrix, i.e. natural versus synthetic air; and (3) difference in linearity with respect to CH_4 mole fraction.

This result therefore indicates that the measurement offset is not constant for a wide range of $\delta^{13}\text{C}\text{-CH}_4$ values and CH_4 mole fractions or for differences in the air matrix. Since we focus in this study on comparison of atmospheric samples, the intercomparison results for the ambient air samples are considered as interlaboratory measurement offsets. The average differences for ambient air are $+0.40 \pm 0.04$ ‰ for TU and $+0.31 \pm 0.03$ ‰ for NIPR relative to IMAU. Likewise, the $\delta\text{D}\text{-CH}_4$ offset of TU versus IMAU is considered to be -13.1 ± 0.6 ‰.

4.3 Intercomparison between UHEI and MPI-BGC

An intercomparison between UHEI and MPI-BGC was conducted in 2013 on six archived air samples from Neumayer station, Antarctica. These samples, collected in the time period from 1988 to 2008, had been analysed by UHEI for $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ by DI-IRMS (two samples were analysed for $\delta\text{D}\text{-CH}_4$ additionally by TDLAS) during 2003–2010 and were stored in high-pressure cylinders. The typical reproducibility for the measurements is between 0.02 and 0.05‰ for $\delta^{13}\text{C}\text{-CH}_4$ and between 1.6 and 2.6‰ for $\delta\text{D}\text{-CH}_4$. In 2013, duplicate aliquots were sampled in 1 L glass

Table 4. Result of intercomparison of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements between UHEI and MPI-BGC. n.a.: not analysed.

Sample ID (Collection date)	Preparation Date UHEI	Analysis Date MPI-BGC	$\delta^{13}\text{C}\text{-CH}_4$ (‰)			$\delta\text{D}\text{-CH}_4$ (‰)		
			UHEI	MPI-BGC ^a	UHEI– MPI-BGC	UHEI ^a	MPI-BGC ^a	UHEI– MPI-BGC
GvN 88/20 (24 Jul 1988)	17 Dec 2003	9 Jul 2013	−47.54 (<i>N</i> = 1)	−47.66 (0.07, <i>N</i> = 2)	+0.13	−83.3 (<i>N</i> = 1)	−82.1 (0.8, <i>N</i> = 2)	−1.2
GvN 92/12 (11 May 1992)	11 Dec 2008	17 Jun 2013	−47.43 (<i>N</i> = 1)	−47.40 (0.04, <i>N</i> = 2)	−0.03	−79.1 (<i>N</i> = 1)	−81.2 (0.9, <i>N</i> = 2)	+2.1
GvN 96/03 (13 Feb 1996)	11 Nov 2003	17 Jun 2013	−47.27 (<i>N</i> = 1)	−47.18 (0.26, <i>N</i> = 2)	−0.08	−73.9 (<i>N</i> = 2)	−74.6 (0.9, <i>N</i> = 2)	+0.8
GvN 99/14 (29 Dec 1999)	3 Apr 2003	9 Jul 2013	−47.30 (<i>N</i> = 1)	−47.23 (0.16, <i>N</i> = 2)	−0.07	−75.2 (<i>N</i> = 2)	−74.6 (1.3, <i>N</i> = 2)	−0.5
GvN 06/14 (23 Sep 2006)	7 May 2003	9 Jul 2013	n.a.	−47.19 (0.09, <i>N</i> = 2)		−72.3 (<i>N</i> = 1)	−73.1 (0.0, <i>N</i> = 2)	+0.8
GvN 08/03 (6 Mar 2008)	28 Jul 2010	17 Jun 2013	−47.18 (<i>N</i> = 1)	−47.35 (0.05, <i>N</i> = 2)	+0.17	n.a.	−67.4 (2.9, <i>N</i> = 2)	
Average			+0.02 ± 0.05 ^b			+0.4 ± 0.6 ^b		

^a Difference in duplicate flask measurements is shown in parenthesis. ^b Uncertainties are standard errors of the mean for differences in the above lines.

flasks and analysed at MPI-BGC. The measurement results at both laboratories are summarized in Table 4. The results show insignificant measurement offsets of $+0.02 \pm 0.05$ for $\delta^{13}\text{C}\text{-CH}_4$ and of $+0.4 \pm 0.6$ ‰ for $\delta\text{D}\text{-CH}_4$ (with the MPI-BGC values being more negative than those from UHEI in both cases).

4.4 Round robin comparison of ice core analysis laboratories

A round robin cylinder exercise was initiated to facilitate intercomparison of laboratories that measure $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ in ice core and firn air samples. Part of this exercise has been presented previously (Table 2 in Schmitt et al., 2013). Three high-pressure Al cylinders were filled with varying trace gas compositions to mimic present day, pre-industrial and last-glacial air mole fractions. The CH_4 mole fractions of these cylinders were 1830.6 (CA 03560), 904.0 (CC 71560) and 372.2 ppb (CA 01179) on the NOAA-2004 CH_4 scale (Dlugokencky et al., 2005), respectively. The cylinders were shipped to the laboratories listed in Table 5 for analysis of all constituents that each lab was capable of measuring at that time. In Table 5, we list the $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ results from each laboratory. The Kr interfering artefact associated with GC-IRMS $\delta^{13}\text{C}\text{-CH}_4$ analyses was taken into account in many of the analyses (Schmitt et al., 2013). In some cases, aliquots from the tanks were measured using offline combustion to CO_2 followed by $\delta^{13}\text{C}\text{-CH}_4$ analyses via conventional DI-IRMS. The cylinders were remeasured at PSU at the end of the round robin to verify that the isotopic composition had not shifted over the 9 years during the transportation of the cylinders. The difference between the 2007 and 2016 $\delta^{13}\text{C}\text{-CH}_4$ measured at PSU was less than 0.14‰ for two of the three cylinders, indicating that the isotopic composition of the cylinder air was stable throughout

the intercomparison exercise. The third cylinder (CA 01179) was 0.58‰ off from the original measurement, which is just outside the analytical uncertainty associated with PSU measurements. There may have been a slight drift over the 9 years between measurements, although the cause has yet to be resolved. The results of the $\delta^{13}\text{C}\text{-CH}_4$ intercomparison showed agreement with the average standard deviation among all six participating laboratories better than 0.37‰ for the cylinders with high (CA 03560) and middle (CC 71560) mole fractions. $\delta\text{D}\text{-CH}_4$ results show more scatter due to the difficult nature of the measurements and the offset among the standard scales.

4.5 Intercomparison between INSTAAR and MPI-BGC

An intercomparison between INSTAAR and MPI-BGC was recently made by analysing three air cylinders at both laboratories. They were measured at MPI-BGC between April and July of 2016 and at INSTAAR between May and June of 2017. Two of the cylinders have ambient CH_4 mole fraction (~ 1900 ppb; HUEY-001 and DEWY-001) and the other has a lower value (~ 1500 ppb; LOUI-001; Table 6). In addition, air from another suite of cylinders was sampled into flasks at INSTAAR and sent to MPI-BGC. Measurements at MPI-BGC and INSTAAR were taken in January–February of 2017 and May–June of 2017, respectively. The four cylinders (CART-001, STAN-001, KENN-001 and KYLE-001) have different CH_4 mole fractions and $\delta^{13}\text{C}\text{-CH}_4$ values. The measurement results are summarized in Table 6. The INSTAAR data presented here were not interfered with by Kr by installing a PCS column into the system. The results show significant but consistent measurement offsets of $+0.28 \pm 0.01$ ‰ for the five cylinders with different CH_4 mole fractions and ambient $\delta^{13}\text{C}\text{-CH}_4$ values (with the IN-

Table 5. Results from the round robin comparison of ice core analysis laboratories conducted during 2007–2016.

Laboratory	CA 03560 (1830.6 ppb)		CA 71560 (904.0 ppb)		CA 01179 (372.2 ppb)		Kr corr.	Analysis Date $\delta^{13}\text{C-CH}_4$	Analysis Date $\delta\text{D-CH}_4$
	$\delta^{13}\text{C-CH}_4$ (‰) ^d	$\delta\text{D-CH}_4$ (‰) ^d	$\delta^{13}\text{C-CH}_4$ (‰) ^d	$\delta\text{D-CH}_4$ (‰) ^d	$\delta^{13}\text{C-CH}_4$ (‰) ^d	$\delta\text{D-CH}_4$ (‰) ^d			
PSU	-47.20 ± 0.16	-93.2 ± 0.9	-47.41 ± 0.10	-95.5 ± 2.3	-47.52 ± 0.06	-106.3 ± 2.4	Raw ion current correction ^a	Jul 2007	Jul 2007
	-46.96 ± 0.16		-47.20 ± 0.10		-47.41 ± 0.12		DI offset ^b	Jul 2007	
	-47.10 ± 0.05		-47.09 ± 0.06		-46.83 ± 0.12		PCS ^c	May 2016	
UCI (DI-IRMS)	-47.09 ± 0.12		-47.40 ± 0.08		-47.23 ± 0.06		–	Dec 2007 ^e	
INSTAAR	-47.08 ± 0.05		-47.20 ± 0.06		-46.78 ± 0.06		PCS ^c	Dec 2008	
NIWA (DI-IRMS)	-47.23 ± 0.02		-47.44 ± 0.02		-47.43 ± 0.02		–	Jun 2009	
NIWA (GC-IRMS)	-47.44 ± 0.21		-48.34 ± 0.28		-47.62 ± 0.11		DI offset ^b	Jun 2009	
UB	-47.41 ± 0.09	-80.4 ± 2.2	-47.37 ± 0.07	-81.0 ± 2.0	-47.31 ± 0.11	-86.2 ± 3.3	No ^d	Jan 2011	Dec 2010–Jan 2011
IMAU	-47.27 ± 0.07	-79.6 ± 1.2	-47.52 ± 0.11	-83.6 ± 3.8	-47.20 ± 0.20	-78.8 ± 12.4	PCS ^c	May & Aug 2012	May 2010

^a Raw ion current correction: the Kr interference was corrected by subtracting the Kr-caused anomalies in the raw ion current data (Sect. 5.4 of Schmitt et al., 2013). ^b DI offset: the Kr interference was corrected by an offset relative to a DI-IRMS measurement. ^c PCS: Kr was separated by a post-combustion separation column (Sect. 5.2 of Schmitt et al., 2013). ^d No: Measurements are affected by the Kr interference (old system without PCS) and raw ion current correction was not possible. ^e Estimated because no exact record on the analysis date at UCI is unfortunately available. ^d Uncertainties are standard deviations of multiple measurements at each laboratory.

Table 6. Results of a $\delta^{13}\text{C-CH}_4$ intercomparison between INSTAAR and MPI-BGC.

Sample ID	CH_4 (ppb) ^a	$\delta^{13}\text{C-CH}_4$ (‰) ^b		INSTAAR – MPI-BGC
		INSTAAR	MPI-BGC	
HUEY-001	1905.5	-47.37 ± 0.01 ($N = 22$)	-47.67 ± 0.01 ($N = 24$)	+0.29
DEWY-001	1879.9	-47.38 ± 0.01 ($N = 26$)	-47.67 ± 0.01 ($N = 22$)	+0.28
LOUI-001	1496.0	-47.26 ± 0.01 ($N = 17$)	-47.55 ± 0.02 ($N = 22$)	+0.29
CART-001	1848.1	-42.98 ± 0.01 ($N = 21$)	-43.30 ± 0.03 ($N = 7$)	+0.32
STAN-001	1696.4	-56.60 ± 0.01 ($N = 7$)	-57.20 ± 0.05 ($N = 8$)	+0.60
KENN-001	1847.6	-47.65 ± 0.01 ($N = 26$)	-47.94 ± 0.05 ($N = 7$)	+0.28
KYLE-001	1847.6	-47.27 ± 0.01 ($N = 29$)	-47.51 ± 0.07 ($N = 6$)	+0.24

^a NOAA-2004 CH_4 scale (Dlugokencky et al., 2005). ^b Uncertainties are standard errors of the mean for the repetitive measurements.

STAAR values being more positive than those from MPI-BGC). The measurements for the cylinder with low $\delta^{13}\text{C-CH}_4$ values were 0.60‰ off between both laboratories presumably due to the scale contraction effect. It is noted that the INSTAAR measurements without the Kr removal yielded a higher $\delta^{13}\text{C-CH}_4$ value ($+0.44 \pm 0.02$ ‰ relative to the MPI-BGC measurement) for one cylinder (LOUI-001), which presumably reflects the Kr interference pronounced at a lower CH_4 mole fraction.

4.6 Intercomparison based on co-located samples through the NOAA cooperative sampling network

The Cooperative Flask Sampling Network, operated by the NOAA Global Monitoring Division, collects air samples from numerous sites around the world, and INSTAAR has analysed those air samples for $\delta^{13}\text{C-CH}_4$ since 1998. There are several sites at which air samples have been concurrently collected by other laboratories. RHUL has analysed air samples at Alert (ALT), Canada and Ascension Island (ASC), and NIWA has done at Baring Head (BHD), New Zealand. Although the individual laboratories do not measure the same sample air in these cases, these co-located air samples provide an opportunity for assessment of possible measurement

offsets as examined previously (Levin et al., 2012). (1) For the RHUL–INSTAAR difference, the $\delta^{13}\text{C}\text{-CH}_4$ data at ALT during 2009–2014 and at ASC during 2010–2015 were compared to each other if both air samples were collected within a 10 h interval. The ALT and ASC comparisons indicated that the INSTAAR measurement is $+0.05 \pm 0.16$ ($N = 350$) and $0.00 \pm 0.17\%$ ($N = 80$) higher than RHUL, respectively. Note that, for this comparison, the RHUL GC-IRMS data were corrected by -0.20% ; the offset value was estimated from measurements of flasks filled from two different cylinders (CH_4 in air, both at ambient mole fraction level, one at ambient $\delta^{13}\text{C}\text{-CH}_4$ and the other at about -56% by spiking ^{13}C -depleted CH_4). (2) For the NIWA–INSTAAR comparison, the $\delta^{13}\text{C}\text{-CH}_4$ data at BHD during 2009–2014 from both laboratories were compared if both air samples were collected within a 15 h interval. The result indicates that the INSTAAR measurement is $+0.08 \pm 0.11\%$ ($N = 45$) higher than NIWA.

5 Measurement offsets among laboratories

Here we revisit intercomparisons published previously. Some laboratories employed a standard scale from another laboratory. Such intercomparisons and interlaboratory scale propagations reported in the literature are displayed in Fig. 1. In this section we review the previous and present intercomparison measurements and accordingly suggest plausible measurement offsets among different laboratories (Fig. 2). Relevant information is summarized in Table 1 and the subsections below correspond to those in Sect. 3. Since some laboratories focus on $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ measurements from ice core and firn air samples to elucidate changes of atmospheric CH_4 in the past, Fig. 2 also combines $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data both for the modern and past atmosphere. It is, however, noted that Fig. 2 suggests the measurement offsets at the modern CH_4 mole fraction and isotopic ratios and that such values could be different for the past atmosphere (see Sect. 4.2, 4.4 and 4.5).

In this study, we report $\delta^{13}\text{C}\text{-CH}_4$ offsets with respect to the conventional DI-IRMS measurement at NIWA (Lowe et al., 1991) because NIWA's $\delta^{13}\text{C}\text{-CH}_4$ measurements have been compared with those from the most laboratories to date (Table 1 and Fig. 1). In contrast, $\delta\text{D}\text{-CH}_4$ measurements from different laboratories have been limited. We report $\delta\text{D}\text{-CH}_4$ offsets of different laboratories with respect to the IMAU measurement. The uncertainties presented in this study are generally standard errors of the mean, but numbers in the literature are cited as is. It should be therefore noted that the uncertainties, in particular those calculated by error propagation, are not rigorously consistent in all places in the paper.

5.1 NIWA

5.1.1 $\delta^{13}\text{C}\text{-CH}_4$

As listed in Table 1, the DI-IRMS measurement at NIWA has been repeatedly intercompared with other laboratories. Importantly for this comparison, Bromley et al. (2012) reported the long-term stability of the measurement over the years 1992–2007, and it is likewise confirmed until 2011. The NIWA GC-IRMS system, based on the methodology of Miller et al. (2002), has an offset relative to the DI-IRMS of $-0.19 \pm 0.26\%$. Measurements on the GC-IRMS informing this instrument comparison are subject to the Kr interference. A Kr-correction has since been derived in an empirical equation from the round robin intercomparison results (Schmitt et al., 2013 and Sect. 4.4), accounting for differences in the CH_4 mole fraction and an exponential fit to the GC-IRMS versus DI-IRMS results. The GC-IRMS system is currently equipped with a PCS column to eliminate the Kr interference.

5.2 IMAU

5.2.1 $\delta^{13}\text{C}\text{-CH}_4$

According to Schmitt et al. (2013), the IMAU measurement at the present CH_4 mole fraction level is in agreement with NIWA with an offset value of $-0.04 \pm 0.07\%$ (no. 2 in Fig. 2a). This corresponds to the round robin comparison for the cylinder with a CH_4 mole fraction of 1830.6 ppb (CA 03560) in Table 5 (Sect. 4.4). The difference is $-0.03 \pm 0.05\%$ for data analysed before the modification to remove the Kr interference (see Table 2 in Schmitt et al., 2013). The intercomparison in this study (Sect. 3.4) also shows that the IMAU offset is $-0.08 \pm 0.11\%$ for the cylinder with the CH_4 mole fraction of 904.0 ppb (CA 71560).

5.2.2 $\delta\text{D}\text{-CH}_4$

As listed in Table 1, IMAU has made the most intercomparisons with other laboratories so far. It is noted that the standard scale at IMAU was propagated from the MPIC (Bergamaschi et al., 2000; Sect. 2.2) and that it recently showed a reasonable agreement with the recent calibration at MPI-BGC (Sperlich et al., 2016).

5.3 MPIC

5.3.1 $\delta^{13}\text{C}\text{-CH}_4$

As written in Sect. 3.3, the standard scale at the MPIC was transferred to IMAU (Brass and Röckmann, 2010; Sperlich et al., 2016). Since no direct comparison with NIWA is available, the MPIC offset relative to NIWA is estimated to be $-0.04 \pm 0.07\%$, identical to the IMAU offset (no. 3 in Fig. 2a).

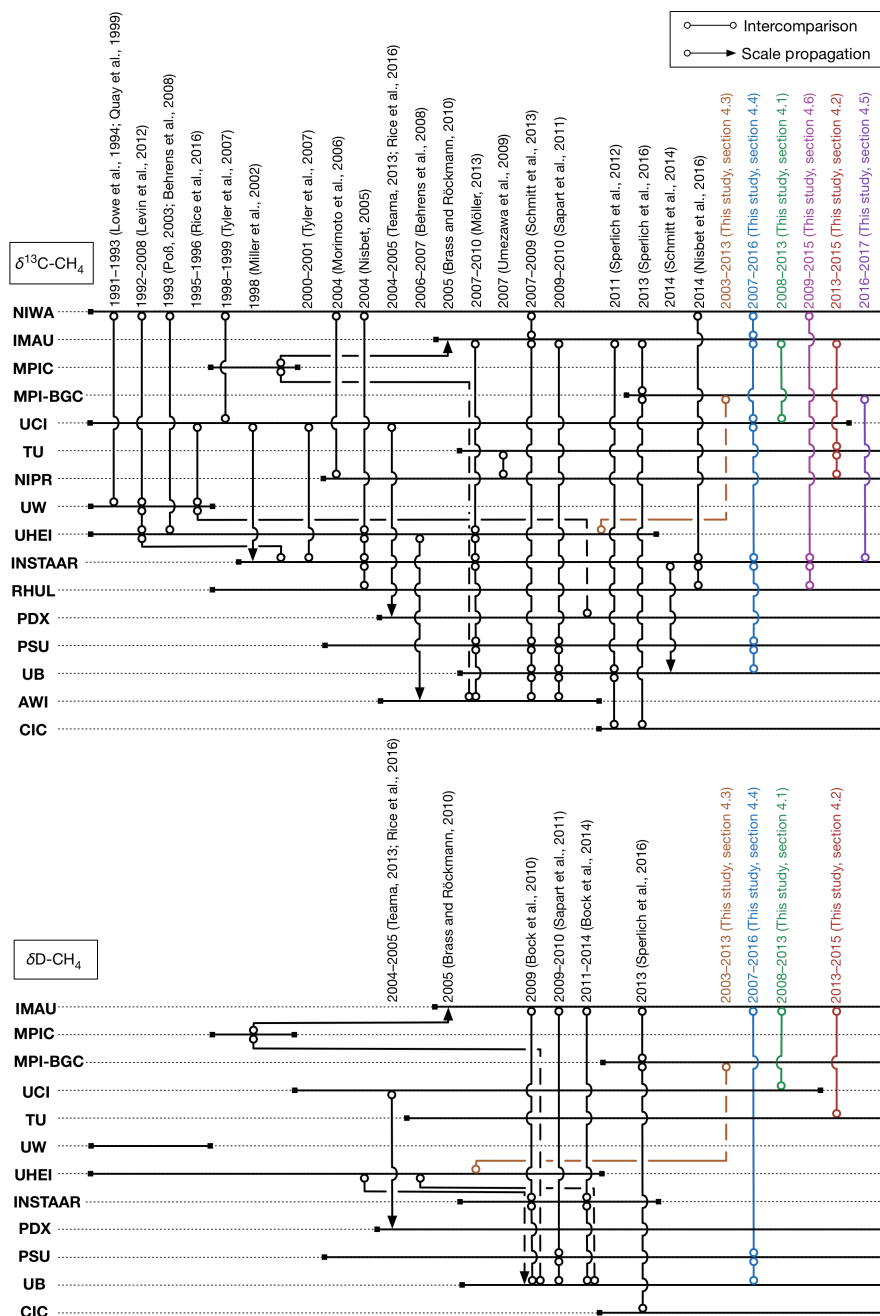


Figure 1. A schematic overview of the previous and present intercomparisons among laboratories for $\delta^{13}\text{C}\text{-CH}_4$ (top) and $\delta\text{D}\text{-CH}_4$ (bottom). Intercomparisons are marked by lines with open circles at both ends, and scale propagations are marked by lines with an arrow at one end.

5.3.2 $\delta\text{D}\text{-CH}_4$

Bock et al. (2010) reported an intercomparison using firm air samples between the UB and MPIC, which indicated that, combined with the UB $\delta\text{D}\text{-CH}_4$ offset (Sect. 5.14), the MPIC $\delta\text{D}\text{-CH}_4$ offset is $+0.3 \pm 1.1\text{‰}$ with respect to IMAU (no. 3 in Fig. 2b).

5.4 MPI-BGC

5.4.1 $\delta^{13}\text{C}\text{-CH}_4$

Sperlich et al. (2016) quantified the offset of the IMAU standard scale relative to the primary standard scale at MPI-BGC. It was indicated that the MPI-BGC measurement differs by $-0.03 \pm 0.10\text{‰}$ from the IMAU standard scale. Combined

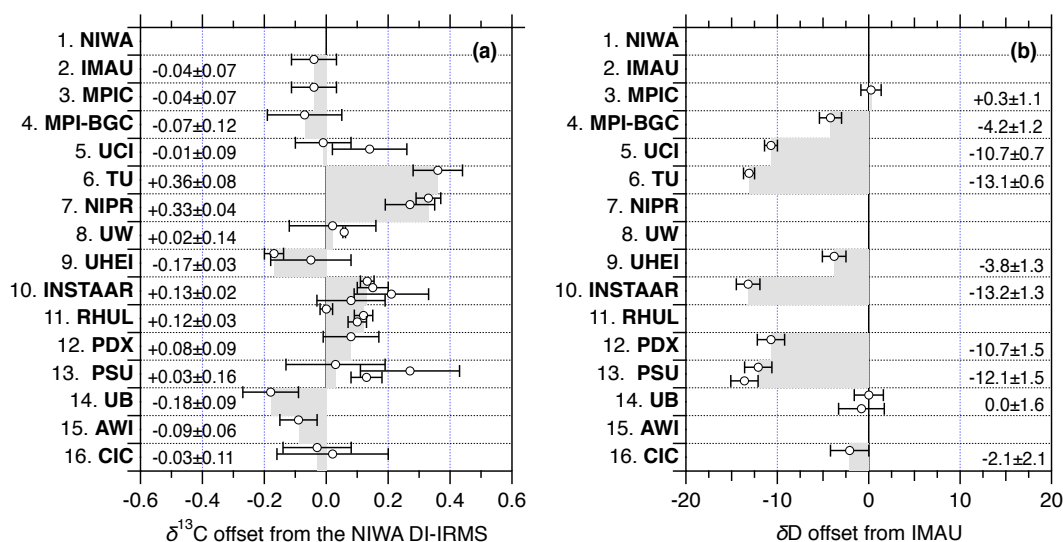


Figure 2. (a) $\delta^{13}\text{C}$ - CH_4 offsets of the different laboratories with respect to the NIWA DI-IRMS measurement with grey shades for ease of viewing. (b) δD - CH_4 offsets of the different laboratories with respect to the IMAU GC-IRMS measurement. See Table 1 and text for corresponding subsections in Sects. 3 and 5. Numbers shown in each laboratory column are the plausible measurement offsets estimated in this study. Note that this result represents intercomparisons for the modern atmospheric CH_4 .

with the IMAU offset relative to NIWA (Sect. 5.2), the MPI-BGC offset is estimated to be $-0.07 \pm 0.12\text{‰}$ (no. 4 in Fig. 2a).

5.4.2 δD - CH_4

According to Sperlich et al. (2016), the MPI-BGC measurement is $-4.2 \pm 1.2\text{‰}$ from IMAU (no. 4 in Fig. 2b).

5.5 UCI

5.5.1 $\delta^{13}\text{C}$ - CH_4

Intercomparison exercises of UCI with external laboratories have been made several times. The oldest intercomparison (Lowe et al., 1991) reported good agreement ($< 0.02\text{‰}$) between the former UCI laboratory (S. Tyler at NCAR) and NIWA (INS, IGNS at that time). Among the later measurements, there were two direct intercomparisons with NIWA. (1) Tyler et al. (2007) reported an intercomparison result of UCI to be $-0.01 \pm 0.09\text{‰}$ with respect to NIWA (top in no. 5, Fig. 2a). For this comparison, 16 air samples collected at Niwot Ridge, Colorado or Baring Head, New Zealand were exchanged between UCI and NIWA in 1998–1999. (2) This study (Sect. 4.4 and Table 5) shows that the UCI measurements are $+0.14 \pm 0.12$ (bottom of no. 5 in Fig. 2a) and $+0.04 \pm 0.08\text{‰}$ higher than NIWA for the cylinders with high (CA 03560) and middle (CC 71560) CH_4 mole fractions, respectively. (3) In contrast, the intercomparison in this study (Sect. 4.1 and Table 2) combined with the IMAU offset (Sect. 5.2) yields $+0.42 \pm 0.04\text{‰}$ relative to NIWA (not shown in Fig. 2a), but is inconsistent with the above inter-

comparison results. The determinate error has yet to be resolved.

5.5.2 δD - CH_4

According to the intercomparison in this study (Sect. 4.1), the UCI has a δD - CH_4 offset of $-10.7 \pm 0.7\text{‰}$ with respect to IMAU (no. 5 in Fig. 2b).

5.6 TU

5.6.1 $\delta^{13}\text{C}$ - CH_4

The intercomparison in this study (Sect. 3.2) and the IMAU offset (Sect. 5.2) give an offset of the TU measurements relative to NIWA to be $+0.36 \pm 0.08\text{‰}$ (no. 6 in Fig. 2b). Measurements at TU have been regularly compared with those at NIPR and they are in agreement within reproducibility of both systems (Umezawa et al., 2009 and additional measurements since then). This is consistent with the previous intercomparison between NIPR and NIWA (Sect. 5.7) and indicates long-term intra-laboratory consistency of TU and NIPR measurements. It is reasonable that TU shares the offset level with NIPR, because both institutions use the same standard scale. As described in Sect. 2.6, it should be noted that the above offset value is not for the data sets currently available to the research community (Umezawa et al., 2011, 2012a, b), for which $+0.32 \pm 0.08\text{‰}$ (not shown in Fig. 2) is recommended. Correction of the data sets from the earlier publications is under evaluation.

5.6.2 $\delta\text{D-CH}_4$

The intercomparison in this study (Sect. 4.2) gives an offset of $-13.1 \pm 0.6\text{‰}$ for the TU atmospheric $\delta\text{D-CH}_4$ measurement (no. 6 in Fig. 2b).

5.7 NIPR

5.7.1 $\delta^{13}\text{C-CH}_4$

An intercomparison between NIPR and NIWA was conducted in 2004 (Morimoto et al., 2006). After the recent update of the NIPR standard scale, the NIPR offset is evaluated to be $+0.33 \pm 0.04\text{‰}$ higher than NIWA (top in no. 7, Fig. 2a). The intercomparison in this study (Sect. 4.2) combined with the IMAU offset (Sect. 5.2) indicates that the NIPR measurement is $+0.27 \pm 0.08\text{‰}$ with respect to NIWA (bottom in no. 7, Fig. 2a), which is consistent with the above value.

5.8 UW

5.8.1 $\delta^{13}\text{C-CH}_4$

Quay et al. (1999) exchanged 30 air samples with NIWA. The average measurement offset was evaluated to be $+0.02 \pm 0.14\text{‰}$ (top in no. 8, Fig. 2a), although some individual samples disagreed by up to 0.5‰ (Lowe et al., 1994; Quay et al., 1999). Later, Levin et al. (2012) estimated that the UW offset is $+0.058 \pm 0.004\text{‰}$ with respect to NIWA based on co-located sampling at BHD (bottom in no. 8, Fig. 2a).

5.8.2 $\delta\text{D-CH}_4$

To our knowledge, no intercomparison exercises with UW have been reported.

5.9 UHEI

5.9.1 $\delta^{13}\text{C-CH}_4$

Levin et al. (2012) estimated the UHEI $\delta^{13}\text{C-CH}_4$ offset to be $-0.169 \pm 0.031\text{‰}$ relative to NIWA (top in no. 9, Fig. 2a). The intercomparison between UHEI and MPI-BGC in this study (Sect. 3.3), together with the MPI-BGC offset (Sect. 5.4), also infers the UHEI offset to be $-0.05 \pm 0.13\text{‰}$ (bottom in no. 9, Fig. 2a), which is consistent with the above value. Earlier measurements of three air samples at both UHEI and NIWA indicated that the UHEI offset is $-0.04 \pm 0.04\text{‰}$ relative to NIWA (Poß, 2003; Behrens et al., 2008). It is also noted that, in an intercomparison presented by Nisbet (2005), the UHEI measurement was $-0.07 \pm 0.04\text{‰}$ lower than NIWA. As these earlier comparison results have been published before the rigorous corrections of the UHEI measurements, these values are not included in Fig. 2a.

5.9.2 $\delta\text{D-CH}_4$

The intercomparison in this study (Sect. 4.3), combined with the MPI-BGC offset (Sect. 5.4), indicates that UHEI has an offset of $-3.8 \pm 1.3\text{‰}$ relative to IMAU.

5.10 INSTAAR

5.10.1 $\delta^{13}\text{C-CH}_4$

Levin et al. (2012) estimated that the INSTAAR measurements have an offset of $+0.132 \pm 0.022\text{‰}$ with respect to NIWA (top in no. 10, Fig. 2a). In an intercomparison exercise reported by Nisbet (2005), the INSTAAR measurement was $+0.14 \pm 0.06\text{‰}$ higher than NIWA (not shown in Fig. 2a), which is consistent with the above value. This study (Sect. 4.4) indicates that the INSTAAR measurement is $+0.15 \pm 0.05\text{‰}$ higher than NIWA for the cylinder with high CH_4 mole fraction (CA 03560; second top of no. 10 in Fig. 2a). The intercomparison between INSTAAR and MPI-BGC (Sect. 4.5) indicates that, combined with the MPI-BGC offset (Sect. 5.4), the INSTAAR offset is $+0.21 \pm 0.12\text{‰}$ relative to NIWA (second bottom in no. 10, Fig. 2a). Lastly, the co-located sample intercomparison (Sect. 4.6) indicates the INSTAAR offset to be $+0.08 \pm 0.11\text{‰}$ (bottom in no. 10, Fig. 2a). It is important to note again that only the round robin intercomparison measurements (Sect. 4.4 and second top of no. 10 in Fig. 2a) and the intercomparison with MPI-BGC (Sect. 4.5) were carried out with a PCS column to remove the Kr interference and that the data set currently available to the public from INSTAAR will be evaluated for future correction.

As described in Sect. 2.10, INSTAAR follows the standard scale of UCI. Tyler et al. (2007) reported that measurements of 10 air cylinders filled at Niwot Ridge, Colorado in 2000–2001 were analysed at both laboratories and that the result indicated an offset of INSTAAR to be $+0.04 \pm 0.12\text{‰}$ relative to UCI. The collection of air samples at Niwot Ridge for the UCI–INSTAAR comparison continued until 2003. A revisit to the measurement record showed that the INSTAAR offset relative to UCI had shifted over the years; the average differences are $+0.02 \pm 0.08$ for 2000 ($N = 7$), $+0.12 \pm 0.07\text{‰}$ for 2001 ($N = 2$) and $+0.26 \pm 0.03\text{‰}$ for 2002 ($N = 12$). This fact may suggest excursions of the internal calibration of either laboratory for these years, but the cause has yet to be resolved; this problem will be addressed in a subsequent paper from either group. It is noted that the offsets relative to NIWA for both laboratories inferred from the different intercomparison pathways are consistent with each other within the uncertainties (Fig. 2a).

5.10.2 $\delta\text{D-CH}_4$

Bock et al. (2010) reported an intercomparison between UB and INSTAAR. This indicates that the INSTAAR measure-

ment offset is $-13.2 \pm 1.3\text{‰}$ with respect to IMAU (no. 10 in Fig. 2b).

5.11 RHUL

5.11.1 $\delta^{13}\text{C}\text{-CH}_4$

Nisbet (2005) reported that the RHUL DI-IRMS measurements agreed well with NIWA with an offset of $0.00 \pm 0.02\text{‰}$ (top in no. 11, Fig. 2a). At the same time, they indicated that the RHUL GC-IRMS measurement has an offset of $+0.11 \pm 0.13\text{‰}$ with respect to NIWA, and later Nisbet et al. (2016) reported that the GC-IRMS system has an offset of about $+0.3\text{‰}$ relative to NIWA (not shown in Fig. 2a). Based on measurements of air in two cylinders exchanged between RHUL and NIWA in 2011 and 2014, RHUL applied an offset correction (-0.20‰) to all data (see Sect. 4.6), by which the RHUL offset has now been evaluated to be $+0.12 \pm 0.03\text{‰}$ (middle of no. 11 in Fig. 2a). The intercomparisons based on the co-located air samples via INSTAAR (Sect. 4.6), combined with the INSTAAR offset (Sect. 5.10), infer that the RHUL offset is $+0.10 \pm 0.03\text{‰}$ relative to NIWA (bottom in no. 11, Fig. 2a).

5.12 PDX

5.12.1 $\delta^{13}\text{C}\text{-CH}_4$

Rice et al. (2016) presented an offset of $+0.024 \pm 0.088\text{‰}$ of the PDX measurements relative to UW by comparing coinciding measurements of archive air samples at PDX and $\delta^{13}\text{C}\text{-CH}_4$ records from Quay et al. (1999) from stations Mauna Loa, Hawaii and Tutuila, American Samoa (1995–1996). With the UW offset with respect to NIWA (Sect. 5.8), it is indicated that the PDX measurement is $+0.08 \pm 0.09\text{‰}$ higher than NIWA (no. 12 in Fig. 2a). This offset is consistent with the UCI offset with respect to NIWA within the uncertainties (note that PDX follows the UCI standard scale).

5.12.2 $\delta\text{D}\text{-CH}_4$

Since PDX follows the UCI standard scale (Teama, 2013; Rice et al., 2016), the likely offset is the same as that of UCI (no. 12 in Fig. 2b).

5.13 PSU

5.13.1 $\delta^{13}\text{C}\text{-CH}_4$

According to Schmitt et al. (2013), the PSU measurement has an offset of $+0.03 \pm 0.16\text{‰}$ relative to NIWA after being corrected for the Kr interference. The measurements of the cylinder with a high CH_4 mole fraction (CA 03560) at PSU are $+0.03 \pm 0.16$, $+0.27 \pm 0.16$ and $+0.13 \pm 0.05\text{‰}$ (no. 13 top, middle and bottom, respectively in Fig. 2a) higher than NIWA for different Kr corrections at different measurement

times, these values being consistent with each other within the uncertainties.

5.13.2 $\delta\text{D}\text{-CH}_4$

An intercomparison result using three firm air samples gives the PSU an offset of $-12.1 \pm 1.5\text{‰}$ relative to the IMAU measurement (Sapart et al., 2011; top in no. 13, Fig. 2b). The intercomparison in this study (Sect. 4.4) gives $-13.6 \pm 1.5\text{‰}$ relative to IMAU for the cylinder with a high CH_4 mole fraction (CA 03560; bottom in no. 13, Fig. 2b).

5.14 UB

5.14.1 $\delta^{13}\text{C}\text{-CH}_4$

The UB measurement has an offset of $-0.18 \pm 0.09\text{‰}$ relative to NIWA (Schmitt et al., 2013; no. 14 in Fig. 2a). This was determined by the round robin intercomparison (Sect. 4.4 and Table 5).

5.14.2 $\delta\text{D}\text{-CH}_4$

Sapart et al. (2011) gives an intercomparison result between UB and IMAU, indicating the UB offset of $0.0 \pm 1.6\text{‰}$ relative to IMAU (top in no. 14, Fig. 2b). This value is consistent with the intercomparisons between UB and IMAU reported by Bock et al. (2010). Later UB modified the measurement set-up, but the measurements of same air samples before and after all modifications were in good agreement as presented by Bock et al. (2014). The intercomparison in this study (Sect. 3.4) shows that the UB measurement differs insignificantly by $-0.8 \pm 2.5\text{‰}$ with respect to IMAU for the cylinder with high CH_4 mole fraction (CA 03560; bottom in no. 14, Fig. 2b).

5.15 AWI

5.15.1 $\delta^{13}\text{C}\text{-CH}_4$

The AWI offset is reported to be $-0.09 \pm 0.06\text{‰}$ with respect to NIWA (Schmitt et al., 2013; no. 15 in Fig. 2a).

5.16 CIC

5.16.1 $\delta^{13}\text{C}\text{-CH}_4$

Sperlich et al. (2012) reported measurements of an air cylinder at CIC, IMAU and UB. The CIC measurement insignificantly different by $+0.01 \pm 0.09\text{‰}$ from IMAU, and the CIC offset with respect to NIWA is estimated to be -0.03 ± 0.11 (top in no. 16, Fig. 2a). They have also reported that the CIC measurement is in agreement with UB with difference of $+0.00 \pm 0.14\text{‰}$. It is noted that, although the UB offset relative to NIWA is estimated to be significant (Sect. 5.14), the difference is still within uncertainties of the intercomparison exercises. Two pure CH_4 gases prepared by Sperlich

et al. (2012) constitute crucial components of the reference gas series developed at MPI-BGC (Sperlich et al., 2016). This has provided a direct intercomparison between CIC and MPI-BGC. The CIC measurement is $+0.09 \pm 0.14\text{‰}$ higher than MPI-BGC. Combined with the MPI-BGC offset (Sect. 5.4), the CIC offset with respect to NIWA is estimated to be $+0.02 \pm 0.18\text{‰}$ (bottom in no. 16, Fig. 2a), which is consistent with the aforementioned value.

5.16.2 $\delta\text{D-CH}_4$

Sperlich et al. (2016) reported $\delta\text{D-CH}_4$ measurement results of the two reference gases prepared by Sperlich et al. (2012) at CIC and MPI-BGC. The results indicated that the CIC measurement differs by $+2.1 \pm 1.8\text{‰}$ from MPI-BGC. Combined with the MPI-BGC offset (Sect. 4.4), the CIC offset relative to IMAU is estimated to be $-2.1 \pm 2.1\text{‰}$ (no. 16 in Fig. 2b).

6 Summary and discussion

We carried out interlaboratory comparison exercises for atmospheric $\delta^{13}\text{C-CH}_4$ and $\delta\text{D-CH}_4$ covering many laboratories around the world. In addition, we reviewed previously published intercomparison results. The results indicated measurement offsets among laboratories, which range from -0.2 to $+0.3\text{‰}$ with respect to the NIWA DI-IRMS measurement for $\delta^{13}\text{C-CH}_4$ and up to -13‰ with respect to the IMAU measurement for $\delta\text{D-CH}_4$. These offset values are larger than the measurement uncertainties from individual laboratories.

The significant $\delta^{13}\text{C-CH}_4$ measurement offsets among laboratories are obvious even though all laboratories ultimately refer to the VPDB scale. We have presented potential causes of the measurement offsets in individual laboratories (Sect. 2), with possible further causes being hidden in all preparation and measurement steps of standard materials. (1) The scale contraction effect for DI-IRMS CO_2 analysis, which is instrument dependent, could be responsible for a considerable part of the observed offsets, given the fact that the atmospheric $\delta^{13}\text{C-CH}_4$ value (about -47‰) differs considerably from the primary anchor of the VPDB scale (NBS-19). (2) Individual laboratories have carried out calibrations against different RMs with different uncertainties of assigned values; such diverse calibration trajectories have also definitely contributed to the interlaboratory measurement offsets. Such RMs have different chemical properties and are processed to CO_2 at individual laboratories, at which different fractionation is possible. (3) Different algorithms for ^{17}O correction have been used for $\delta^{13}\text{C}$ measurements at different laboratories, which could have caused biases among available data sets. (4) The Kr interference on a GC-IRMS system is in several cases a probable cause of the offsets, and unfortunately, this effect is system dependent and can vary

with time and the instrument settings. Lastly, it is important to note that we summarized $\delta^{13}\text{C-CH}_4$ measurement offsets at the modern atmospheric CH_4 mole fraction level, but the offset may vary with the amount of CH_4 analysed (e.g. lower mole fractions in ice core analyses, see Tables 3, 5 and 6), because of a non-linear response of IRMS (Umezawa et al., 2009) and because the Kr interference is directly dependent on the Kr-to- CH_4 ratio (Schmitt et al., 2013). Furthermore, the intercomparisons presented here focus on modern atmospheric CH_4 of typically -47‰ and such comparisons for high and low $\delta^{13}\text{C-CH}_4$ values (e.g. CH_4 from ice cores or enriched/depleted source signatures) are to date very limited (Tables 3 and 6 in this study).

Concerning $\delta\text{D-CH}_4$ measurement offsets among laboratories, it is interesting that the listed laboratories can be roughly split into two groups whose $\delta\text{D-CH}_4$ measurements differ by $\sim 10\text{‰}$. Some laboratories with higher $\delta\text{D-CH}_4$ values refer to an identical set of standards produced at the MPIC (MPIC and IMAU) or to the UHEI calibration (UHEI and UB), and measurements of these groups have been cross-referenced (see Sects. 2 and 4), thereby showing the reasonable agreements. The original calibrations were carried out using an offline CH_4 processing technique (cryogenic separation and conversion of CH_4 to CO_2 and H_2O followed by H_2O reduction to H_2) with subsequent analysis by DI-IRMS. The other laboratories with higher $\delta\text{D-CH}_4$ values recently developed their own primary calibrations independently (CIC and MPI-BGC). CIC used an offline CH_4 processing technique combined with DI-IRMS, whereas MPI-BGC adopted TC/EA coupled to continuous-flow IRMS. For the lower $\delta\text{D-CH}_4$ group, some laboratories carried out calibrations against Oztech H_2 gases (UCI, PDX and PSU) or have other calibration pathways (TU and INSTAAR; see Sect. 2). These laboratories used local H_2 working gas standards for GC-IRMS, which were calibrated with a separate DI-IRMS procedure. As is the case for $\delta^{13}\text{C-CH}_4$, possible causes of the observed $\delta\text{D-CH}_4$ discrepancies could have arisen in all preparation and measurement steps. (1) The classical technique for DI-IRMS involves processing of H_2O , and the associated steps in experimental lines are prone to surface adhesion and contamination of H_2O ; thereby considerable memory effect is possible (Bergmaschi et al., 2000). (2) Similarly to $\delta^{13}\text{C-CH}_4$, calibration for $\delta\text{D-CH}_4$ involves measurements of standards with different chemical properties (H_2O and H_2), and such calibrations at different laboratories could contribute to the offset. (3) Difficulties in maintaining stable pyrolysis conditions for GC-IRMS (Bock et al. 2010) might have affected measurements against local H_2 working standards in the cases where the principle of identical treatment (Werner and Brand, 2001) was not strictly followed. Lastly, it is noted that the non-linearity of the IRMS in $\delta\text{D-CH}_4$ measurements (Brass and Röckmann, 2010) may also play a role for samples with low mole fractions such as ice core analyses.

The measurement offsets summarized in this study should be thoroughly taken into account when data from different laboratories are combined, and this study will be of help when incorporating merged $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data sets into a state-of-the-art chemistry transport model. However, it is recommended that data users contact the data providers directly for the latest information whenever possible. The Kr interference is under evaluation at some laboratories and it will possibly involve an update of the data sets that are currently available. More importantly, it is imperative to have common reference gases with transparent and reproducible traceability (for instance, Sperlich et al. 2016) and to carry out a systematic intercomparison programme (flask or cylinder round robin) in the research community to attain the necessary but ambitious high-compatibility goals of 0.02 ‰ for $\delta^{13}\text{C}\text{-CH}_4$ and 1 ‰ for $\delta\text{D}\text{-CH}_4$ (WMO, 2016). Such thorough efforts will facilitate optimized use of $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ data sets in a combined way and maximize the number of isotope data sets (and thus their spatial and temporal coverage) usable for enhancing our understanding of the global CH_4 cycle.

We welcome collaborative works that analyse the multiple data sets from laboratories that participated in this study (see data availability listed in Table 1). Data users can examine the offset numbers (Table 1 and Fig. 2) and adjust the data sets at least for data points with values close to the modern atmosphere in $\delta^{13}\text{C}\text{-CH}_4$ and $\delta\text{D}\text{-CH}_4$ as well as the CH_4 mole fraction. For data with CH_4 mole fractions and isotopic ratios that are far from modern background values (e.g. sample air from ice core and stratosphere and those influenced by sources), more intercomparisons are needed to establish correction factors among data sets.

Data availability. All the interlaboratory comparison data presented in this study are included in the tables of this paper.

Appendix A: List of participating institution/project acronyms

AWI:	Alfred Wegener Institute Helmholtz Centre for Polar and Marine Research, Bremerhaven, Germany
CARIBIC:	Civil Aircraft for the Regular Investigation of the atmosphere Based on an Instrument Container
CIC:	Centre for Ice and Climate, Niels Bohr Institute, University of Copenhagen, Copenhagen, Denmark
IMAU:	Institute for Marine and Atmospheric research Utrecht, Utrecht University, Utrecht, the Netherlands
INSTAAR:	Institute of Arctic and Alpine Research University of Colorado Boulder, Boulder, USA
MPI-BGC:	Max Planck Institute for Biogeochemistry, Jena, Germany
MPIC:	Max Planck Institute for Chemistry, Mainz, Germany
NCAR:	National Center for Atmospheric Research, Boulder, USA
NIPR:	National Institute of Polar Research, Tokyo, Japan
NIWA:	National Institute for Water and Atmospheric Research, Wellington, New Zealand
NOAA:	National Oceanic and Atmospheric Administration, USA
PDX:	Portland State University, Portland, USA
PSU:	Pennsylvania State University, Pennsylvania, USA
RHUL:	Royal Holloway, University of London, Egham, UK
TU:	Tohoku University, Sendai, Japan
UB:	University of Bern, Bern, Switzerland
UCI:	University of California Irvine, Irvine, USA
UHEI:	University of Heidelberg, Heidelberg, Germany
UW:	University of Washington, Seattle, USA

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. We thank Morimasa Sato for measurements at TU. We thank Owen Sherwood for his contribution in setting up the INSTAAR measurement system and Ken Masarie and John Mund, who developed software enabling data comparisons from co-located samples measured by INSTAAR, RHUL and NIWA. We thank Paul Quay (UW) for isotopic CH_4 data and intercalibration information and Doaa Teama for isotopic measurements at PDX. We thank Célia Sapart for helpful information about the past intercomparison. Part of this work at NIES was supported by the Environment Research and Technology Development Fund (2-1710) of the Ministry of the Environment, Japan and Environmental Restoration and Conservation Agency. Funding for work at PDX was provided by the US National Science Foundation (Atmospheric and Geospace Sciences Grant 0952307). CH_4 isotope ice core work at UB received funding from the European Research Council under the European Union's Seventh Framework Programme (FP7/2007-2013) ERC grant agreement no. 226172 and the Swiss National Science Foundation (grant no. 200020_159563 & 20020_172506). NIWA's isotope measurements are funded under the Climate and Atmosphere Research Programme CAAC1804 (2017/18 SCI). We are grateful to the referee Sergey Assonov and one anonymous referee for their insightful and constructive comments for improving this paper.

The article processing charges for this open-access publication were covered by the Max Planck Society.

Edited by: Frank Keppler

Reviewed by: Sergey Assonov and one anonymous referee

References

- Allison, C. E., Francey, R. J., and Meijer, H. A. J.: Recommendations for the reporting of stable isotope measurements of carbon and oxygen in CO_2 gas, TECDOC-825 IAEA, Vienna, 155–162, 1995.
- Aoki, S., Nakazawa, T., Murayama, S., and Kawaguchi, S.: Measurements of atmospheric methane at the Japanese Antarctic Station, Syowa, *Tellus*, 44B, 273–281, 1992.
- Assonov, S. S. and Brenninkmeijer, C. A. M.: On the ^{17}O correction for CO_2 mass spectrometric isotopic analysis, *Rapid Commun. Mass Sp.*, 17, 1007–1016, <https://doi.org/10.1002/rcm.1012>, 2003.
- Behrens, M., Schmitt, J., Richter, K.-U., Bock, M., Richter, U., Levin, I., and Fischer, H.: A gas chromatography/combustion/isotope ratio mass spectrometry system for high-precision $\delta^{13}\text{C}$ measurements of atmospheric methane extracted from ice core samples, *Rapid Commun. Mass Sp.*, 22, 3261–3269, <https://doi.org/10.1002/rcm.3720>, 2008.
- Bergamaschi, P., Schupp, M., and Harris, G. W.: High-precision direct measurements of $^{13}\text{CH}_4/^{12}\text{CH}_4$ and $^{12}\text{CH}_3\text{D}/^{12}\text{CH}_4$ ratios in atmospheric methane sources by means of a long-path tunable diode laser absorption spectrometer, *Appl. Opt.*, 33, 7704–7716, 1994.
- Bergamaschi, P., Bräunlich, M., Marik, T., and Brenninkmeijer, C. A. M.: Measurements of the carbon and hydrogen isotopes of atmospheric methane at Izana, Tenerife: Seasonal cycles and synoptic-scale variations, *J. Geophys. Res.-Atmos.*, 105, 14531–14546, <https://doi.org/10.1029/1999JD901176>, 2000.
- Bergamaschi, P., Lowe, D. C., Manning, M. R., Moss, R., Bromley, T., and Clarkson, T. S.: Transects of atmospheric CO , CH_4 , and their isotopic composition across the Pacific: Shipboard measurements and validation of inverse models, *J. Geophys. Res.*, 106, 7993–8011, <https://doi.org/10.1029/2000JD900576>, 2001.
- Bock, M., Schmitt, J., Behrens, M., Möller, L., Schneider, R., Sapart, C., and Fischer, H.: A gas chromatography/pyrolysis/isotope ratio mass spectrometry system for high-precision δD measurements of atmospheric methane extracted from ice cores, *Rapid Commun. Mass Sp.*, 24, 621–633, <https://doi.org/10.1002/rcm.4429>, 2010.
- Bock, M., Schmitt, J., Beck, J., Schneider, R., and Fischer, H.: Improving accuracy and precision of ice core $\delta\text{D}(\text{CH}_4)$ analyses using methane pre-pyrolysis and hydrogen post-pyrolysis trapping and subsequent chromatographic separation, *Atmos. Meas. Tech.*, 7, 1999–2012, <https://doi.org/10.5194/amt-7-1999-2014>, 2014.
- Bock, M., Schmitt, J., Beck, J., Seth, B., Chappellaz, J., and Fischer, H.: Glacial/interglacial wetland, biomass burning, and geologic methane emissions constrained by dual stable isotopic CH_4 ice core records, *P. Natl. Acad. Sci. USA*, E5778–E5786, <https://doi.org/10.1073/pnas.1613883114>, 2017.
- Brand, W., Coplen, T. B., Vogl, J., Rosner, M., and Prohaska, T.: Assessment of international reference materials for isotope-ratio analysis (IUPAC Technical Report), *Pure Appl. Chem.*, 86, 425–467, <https://doi.org/10.1515/pac-2013-1023>, 2014.
- Brand, W., Rothe, M., Sperlich, P., Strube, M., and Wendeberg M.: Automated simultaneous measurement of the $\delta^{13}\text{C}$ and $\delta^2\text{H}$ values of methane and the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of carbon dioxide in flask samples using a new multi cryo-trap/gas chromatography/isotope ratio mass spectrometry system, *Rapid Commun. Mass Sp.*, 30, 1523–1539, <https://doi.org/10.1002/rcm.7587>, 2016.
- Brass, M. and Röckmann, T.: Continuous-flow isotope ratio mass spectrometry method for carbon and hydrogen isotope measurements on atmospheric methane, *Atmos. Meas. Tech.*, 3, 1707–1721, <https://doi.org/10.5194/amt-3-1707-2010>, 2010.
- Bräunlich, M., Aballanin, O., Marik, T., Jockel, P., Brenninkmeijer, C. A. M., Chappellaz, J., Barnola, J. M., Mulvaney, R., and Sturges, W. T.: Changes in the global atmospheric methane budget over the last decades inferred from ^{13}C and D isotopic analysis of Antarctic firn air, *J. Geophys. Res.-Atmos.*, 106, 20465–20481, <https://doi.org/10.1029/2001JD900190>, 2001.
- Brenninkmeijer, C. A. M., Crutzen, P. J., Fischer, H., Gusten, H., Hans, W., Heinrich, G., Heintzenberg, J., Hermann, M., Immelmann, T., Kersting, D., Maiss, M., Nolle, M., Pitscheider, A., Pohlkamp, H., Scharffe, D., Specht, K., and Wiedensohler, A.: CARIBIC – Civil aircraft for global measurement of trace gases and aerosols in the tropopause region, *J. Atmos. Ocean. Tech.*, 16, 1373–1383, [https://doi.org/10.1175/1520-0426\(1999\)016<1373:Ccafgm>2.0.Co;2](https://doi.org/10.1175/1520-0426(1999)016<1373:Ccafgm>2.0.Co;2), 1999.
- Bromley, T., Allan, W., Martin, R., Mikaloff Fletcher, S. E., Lowe, D. C., Struthers, H., and Moss, R.: Shipboard measurements and modeling of the distribution of CH_4 and

- $^{13}\text{CH}_4$ in the western Pacific, *J. Geophys. Res.*, 117, D04307, <https://doi.org/10.1029/2011JD016494>, 2012.
- Burgoyne, T. W. and Hayes J. M.: Quantitative production of H_2 by Pyrolysis of gas chromatographic effluents, *Anal. Chem.*, 70, 24, 5136–5141, <https://doi.org/10.1021/ac980248v>, 1998.
- Cicerone, R. J. and Oremland, R. S.: Biogeochemical aspects of atmospheric methane, *Global Biogeochem. Cy.*, 2, 299–327, <https://doi.org/10.1029/GB002i004p00299>, 1988.
- Coplen, T. B., Brand, W. A., Gehre, M., Gröning, M., Meijer, H. A. J., Toman, B., and Verkouteren, R. M.: New guidelines for $\delta^{13}\text{C}$ measurements, *Anal. Chem.*, 78, 2439–2441, <https://doi.org/10.1021/ac052027c>, 2006.
- Craig, H.: Isotopic standards for carbon and oxygen and correction factors for mass-spectrometric analysis of carbon dioxide, *Geochim. Cosmochim. Ac.*, 12, 133–149, [https://doi.org/10.1016/0016-7037\(57\)90024-8](https://doi.org/10.1016/0016-7037(57)90024-8), 1957.
- Dlugokencky, E. J., Myers, R. C., Lang, P. M., Masarie, K. A., Crotwell, A. M., Thoning, K. W., Hall, B. D., Elkins, J. W., and Steele, L. P.: Conversion of NOAA atmospheric dry air CH_4 mole fractions to a gravimetrically prepared standard scale, *J. Geophys. Res.*, 110, D18306, <https://doi.org/10.1029/2005JD006035>, 2005.
- Eyer, S., Tuzson, B., Popa, M. E., van der Veen, C., Röckmann, T., Rothe, M., Brand, W. A., Fisher, R., Lowry, D., Nisbet, E. G., Brennwald, M. S., Harris, E., Zellweger, C., Emmenegger, L., Fischer, H., and Mohn, J.: Real-time analysis of $\delta^{13}\text{C}$ - and δD - CH_4 in ambient air with laser spectroscopy: method development and first intercomparison results, *Atmos. Meas. Tech.*, 9, 263–280, <https://doi.org/10.5194/amt-9-263-2016>, 2016.
- Ferretti, D. F., Miller, J. B., White, J. W. C., Etheridge, D. M., Lassey, K. R., Lowe, D. C., MacFarling Meure, C. M., Dreier, M. F., C. M. Trudinger, C. M., van Ommen, T. D., and Langenfelds, R. L.: Unexpected changes to the global methane budget over the past 2000 years, *Science*, 309, 1714–1717, <https://doi.org/10.1126/science.1115193>, 2005.
- Fischer, H., Behrens, M., Bock, M., Richter, U., Schmitt, J., Loulergue, L., Chappellaz, J., Spahni, R., Blunier, T., Leuenberger, M., and Stocker, T. F.: Changing boreal methane sources and constant biomass burning during the last termination, *Nature*, 452, 864–867, <https://doi.org/10.1038/nature06825>, 2008.
- Fisher, R., Lowry, D., Wilkin, O., Sriskantharajah, S., and Nisbet, E. G.: High-precision, automated stable isotope analysis of atmospheric methane and carbon dioxide using continuous-flow isotope-ratio mass spectrometry, *Rapid Commun. Mass Sp.*, 20, 200–208, <https://doi.org/10.1002/rcm.2300>, 2006.
- Fisher, R. E., Sriskantharajah, S., Lowry, D., Lanouisellé, M., Fowler, C. M. R., James, R. H., Hermansen, O., Lund Myhre, C., Stohl, A., Greinert, J., Nisbet-Jones, P. B. R., Mienert, J., and Nisbet, E. G.: Arctic methane sources: Isotopic evidence for atmospheric inputs, *Geophys. Res. Lett.*, 38, L21803, <https://doi.org/10.1029/2011GL049319>, 2011.
- Fung, I., John, J., Lerner, J., Matthews, E., Prather, M., Steele, L. P., and Fraser, P. J.: Three-dimensional model synthesis of the global methane cycle, *J. Geophys. Res.*, 96, 13033–13065, <https://doi.org/10.1029/91JD01247>, 1991.
- Ghosh, A., Patra, P. K., Ishijima, K., Umezawa, T., Ito, A., Etheridge, D. M., Sugawara, S., Kawamura, K., Miller, J. B., Dlugokencky, E. J., Krummel, P. B., Fraser, P. J., Steele, L. P., Langenfelds, R. L., Trudinger, C. M., White, J. W. C., Vaughn, B., Saeki, T., Aoki, S., and Nakazawa, T.: Variations in global methane sources and sinks during 1910–2010, *Atmos. Chem. Phys.*, 15, 2595–2612, <https://doi.org/10.5194/acp-15-2595-2015>, 2015.
- Hein, R., Crutzen, P. J., and Heimann, M.: An inverse modeling approach to investigate the global atmospheric methane cycle, *Global Biogeochem. Cy.*, 11, 43–76, <https://doi.org/10.1029/96GB03043>, 1997.
- Hilkert, A. W., Douthitt, C. B., Schlüter, H. J., and Brand W. A.: Isotope ratio monitoring gas chromatography/mass Spectrometry of D/H by high temperature conversion isotope ratio mass spectrometry, *Rapid Comm. Mass Spec.*, 13, 1226–1230, [https://doi.org/10.1002/\(SICI\)1097-0231\(19990715\)13:13<1226::AID-RCM575>3.0.CO;2-9](https://doi.org/10.1002/(SICI)1097-0231(19990715)13:13<1226::AID-RCM575>3.0.CO;2-9), 1999.
- Kai, F. M., Tyler, S. C., Randerson, J. T., and Blake, D. R.: Reduced methane growth rate explained by decreased Northern Hemisphere microbial sources, *Nature*, 476, 194–197, <https://doi.org/10.1038/nature10259>, 2011.
- Kirschke, S., Bousquet, P., Ciais, P., Saunoy, M., Canadell, J. G., Dlugokencky, E. J., Bergamaschi, P., Bergmann, D., Blake, D. R., Bruhwiler, L., Cameron-Smith, P., Castaldi, S., Chevallier, F., Feng, L., Fraser, A., Heimann, M., Hodson, E. L., Houweling, S., Josse, B., Fraser, P. J., Krummel, P. B., Lamarque, J.-F., Langenfelds, R. L., Le Quééré, C., Naik, V., O'Doherty, S., Palmer, P. I., Pison, I., Plummer, D., Poulter, B., Prinn, R. G., Rigby, M., Ringeval, B., Santini, M., Schmidt, M., Shindell, D. T., Simpson, I. J., Spahni, R., Steele, L. P., Strode, S. A., Sudo, K., Szopa, S., van der Werf, G. R., Voulgarakis, A., van Weele, M., Weiss, R. F., Williams, J. E., and Zeng, G.: Three decades of global methane sources and sinks, *Nat. Geosci.*, 6, 813–823, <https://doi.org/10.1038/ngeo1955>, 2013.
- Levin, I., Bergamaschi, P., Dörr, H., and Trapp, D.: Stable isotopic signature of methane from major sources in Germany, *Chemosphere*, 26, 1–4, 161–177, [https://doi.org/10.1016/0045-6535\(93\)90419-6](https://doi.org/10.1016/0045-6535(93)90419-6), 1993.
- Levin, I., Glatzel-Mattheier, H., Marik, T., Cuntz, M., and Schmidt, M.: Verification of German methane emission inventories and their recent changes based on atmospheric observations, *J. Geophys. Res.*, 104, 3447–3456, <https://doi.org/10.1029/1998JD100064>, 1999.
- Levin, I., Veidt, C., Vaughn, B. H., Brailsford, G., Bromley, T., Heinz, R., Lowe, D., Miller, J. B., Poss, C., and White, J. W. C.: No inter-hemispheric $\delta^{13}\text{C}$ trend observed, *Nature*, 486, E3–E4, <https://doi.org/10.1038/nature11175>, 2012.
- Lowe, D. C., Brenninkmeijer, C. A. M., Manning M. R., Sparks, R., and Wallace, G.: Radiocarbon determination of atmospheric methane at Baring Head, New Zealand, *Nature*, 332, 522–525, <https://doi.org/10.1038/332522a0>, 1988.
- Lowe, D. C., Brenninkmeijer, C. A. M., Tyler, S. C., and Dlugokencky, E. J.: Determination of the Isotopic Composition of Atmospheric Methane and its Application in the Antarctic, *J. Geophys. Res.-Atmos.*, 96, 15455–15467, <https://doi.org/10.1029/91JD01119>, 1991.
- Lowe, D. C., Brenninkmeijer, C. A. M., Brailsford, G. W., Lassey, K. R., and Gomez, A. J.: Concentration and ^{13}C records of atmospheric methane in New Zealand and Antarctica: Evidence for changes in methane sources, *J. Geophys. Res.*, 99, 16913–16925, <https://doi.org/10.1029/94JD00908>, 1994.

- Lowe, D. C., Manning, M. R., Brailsford, G., and Bromley, A. M.: The 1991–1992 atmospheric methane anomaly: Southern Hemisphere ^{13}C decrease and growth rate fluctuations, *Geophys. Res. Lett.*, 24, 8, 857–860, <https://doi.org/10.1029/97GL00830>, 1997.
- Lowe, D. C., Allan, W., Manning, M. R., Bromley, T., Brailsford, G., Ferretti, D., Gomez, A., Knobben, R., Martin, R., Mei, Z., Moss, R., Koshy, K., and Maata, M.: Shipboard determinations of the distribution of C-13 in atmospheric methane in the Pacific, *J. Geophys. Res.-Atmos.*, 104, 26125–26135, <https://doi.org/10.1029/1999jd900452>, 1999.
- Lowe, D. C., Koshy, K., Bromley, T., Allan, W., Struthers, H., Mani, F., and Maata, M.: Seasonal cycles of mixing ratio and ^{13}C in atmospheric methane at Suva, Fiji, *J. Geophys. Res.*, 109, D23308, <https://doi.org/10.1029/2004JD005166>, 2004.
- Lowry, D., Holmes, C. W., Rata, N. D., O'Brien, P., and Nisbet, E. G.: London methane emissions: Use of diurnal changes in concentration and $\delta^{13}\text{C}$ to identify urban sources and verify inventories, *J. Geophys. Res.*, 106, 7427–7448, <https://doi.org/10.1029/2000JD900601>, 2001.
- Meijer, H. A. J., Neubert, R. E. M., and Visser, G. H.: Cross contamination in dual inlet isotope ratio mass spectrometers, *Int. J. Mass Spectrom.*, 198, 45–61, 2000.
- Merritt, D. A., Hayes, J. M., and Des Marais, D. J.: Carbon isotopic analysis of atmospheric methane by isotope-ratio-monitoring gas-chromatography mass-spectrometry, *J. Geophys. Res.*, 100, 1317–1326, <https://doi.org/10.1029/94JD02689>, 1995.
- Mikaloff Fletcher, S. E., Tans, P. P., Bruhwiler, L. M., Miller, J. B., and Heimann, M.: CH_4 sources estimated from atmospheric observations of CH_4 and its $^{13}\text{C}/^{12}\text{C}$ isotopic ratios: 1. Inverse modeling of source processes, *Global Biogeochem. Cy.*, 18, GB4004, <https://doi.org/10.1029/2004GB002223>, 2004a.
- Mikaloff Fletcher, S. E., Tans, P. P., Bruhwiler, L. M., Miller, J. B., and Heimann, M.: CH_4 sources estimated from atmospheric observations of CH_4 and its $^{13}\text{C}/^{12}\text{C}$ isotopic ratios: 2. Inverse modeling of CH_4 fluxes from geographical regions, *Global Biogeochem. Cy.*, 18, GB4005, <https://doi.org/10.1029/2004GB002224>, 2004b.
- Miller, J. B., Mack, K. A., Dissly, R., White, J. W. C., Dlugokencky, E. J., and Tans, P. P.: Development of analytical methods and measurements of $\delta^{13}\text{C}$ in atmospheric CH_4 from the NOAA/CMDL global air sampling network, *J. Geophys. Res.*, 107, 4178, <https://doi.org/10.1029/2001JD000630>, 2002.
- Möller, L., Sowers, T., Bock, M., Spahn R., Behrens, M., Schmitt, J., Miller, H., and Fischer, H.: Independent variations of CH_4 emissions and isotopic composition over the past 160000 years, *Nat. Geosci.*, 6, 885–890, <https://doi.org/10.1038/NCEO1922>, 2013.
- Monteil, G., Houweling, S., Dlugokencky, E. J., Maenhout, G., Vaughn, B. H., White, J. W. C., and Rockmann, T.: Interpreting methane variations in the past two decades using measurements of CH_4 mixing ratio and isotopic composition, *Atmos. Chem. Phys.*, 11, 9141–9153, <https://doi.org/10.5194/acp-11-9141-2011>, 2011.
- Morimoto, S., Aoki, S., Nakazawa, T., and Yamanouchi, T.: Temporal variations of the carbon isotopic ratio of atmospheric methane observed at Ny Ålesund, Svalbard from 1996 to 2004, *Geophys. Res. Lett.*, 33, L01807, <https://doi.org/10.1029/2005GL024648>, 2006.
- Morimoto, S., Fujita, R., Aoki, S., Goto, D., and Nakazawa, T.: Long-term variations of the mole fraction and carbon isotope ratio of atmospheric methane observed at Ny-Ålesund, Svalbard from 1996 to 2013, *Tellus B*, 69, 1380497, <https://doi.org/10.1080/16000889.2017.1380497>, 2017.
- Nisbet, E. (Ed.): Meth-MonitEU: Methane Monitoring in the European Union and Russia, *Eur. Comm.*, Brussels, 33–37, 2005.
- Nisbet, E. G., Dlugokencky, E. J., Manning, M. R., Lowry, D., Fisher, R. E., France, J. L., Michel, S. E., Miller, J. B., White, J. W. C., Vaughn, B., Bousquet, P., Pyle, J. A., Warwick, N. J., Cain, M., Brownlow, R., Zazzeri, G., Lanoisellé, M., Manning, A. C., Gloor, E., Worthy, D. E. J., Brunke, E.-G., Labuschagne, C., Wolff, E. W., and Ganesan, A. L.: Rising atmospheric methane: 2007–2014 growth and isotopic shift, *Global Biogeochem. Cy.*, 30, 1356–1370, <https://doi.org/10.1002/2016GB005406>, 2016.
- Poß, C.: Investigation of the variability of atmospheric methane in polar regions based on trajectory analysis and the measurement of stable isotopes (in German), PhD thesis, Heidelberg University, 17–32, 2003.
- Quay, P., Stutsman, J., Wilbur, D., Snover, A., Dlugokencky, E., and Brown, T.: The isotopic composition of atmospheric methane, *Global Biogeochem. Cy.*, 13, 445–461, 1999.
- Quay, P. D., King, S. L., Stutsman, J., Wilbur, D. O., Steele, L. P., Fung, I., Gammon, R. H., Brown, T. A., Farwell, G. W., Grootes, P. M., and Schmidt, F. H.: Carbon isotopic composition of atmospheric CH_4 : Fossil and biomass burning strengths, *Global Biogeochem. Cy.*, 5, 25–47, 1991.
- Rice, A. L., Gotoh, A. A., Ajie, H. O., and Tyler, S. C.: High-precision continuous-flow measurement of $\delta^{13}\text{C}$ and δD of atmospheric CH_4 , *Anal. Chem.*, 73, 4104–4110, <https://doi.org/10.1021/ac0155106>, 2001.
- Rice, A. L., Butenhoff, C. L., Teama, D. G., Röger, F. H., Khalil, M. A. K., and Rasmussen, R. A.: Atmospheric methane isotopic record favors fossil sources flat in 1980s and 1990s with recent increase, *P. Natl. Acad. Sci. USA*, 113, 10791–10796, <https://doi.org/10.1073/pnas.1522923113>, 2016.
- Rigby, M., Montzka, S. A., Prinn, R. G., White, J. W. C., Young, D., O'Doherty, S., Lunta, M. F., Ganesan, A. L., Manning, A. J., Simmonds, P. G., Salameh, P. K., Harth, C. M., Mühle, J., Weiss, R. F., Fraser, P. J., L. Steele, P. Krummel, P. B., McCulloch, A., and Park, S.: Role of atmospheric oxidation in recent methane growth, *P. Natl. Acad. Sci. USA*, 114, 5373–5377, <https://doi.org/10.1073/pnas.1616426114>, 2017.
- Röckmann, T.: Measurement and interpretation of ^{13}C , ^{14}C , ^{17}O and ^{18}O variations in atmospheric carbon monoxide, PhD thesis, Univ. Heidelberg, Heidelberg, Germany, 56–57, 1998.
- Röckmann, T., Brass, M., Borchers, R., and Engel, A.: The isotopic composition of methane in the stratosphere: high-altitude balloon sample measurements, *Atmos. Chem. Phys.*, 11, 13287–13304, <https://doi.org/10.5194/acp-11-13287-2011>, 2011.
- Röckmann, T., Eyer, S., van der Veen, C., Popa, M. E., Tuzson, B., Monteil, G., Houweling, S., Harris, E., Brunner, D., Fischer, H., Zazzeri, G., Lowry, D., Nisbet, E. G., Brand, W. A., Necki, J. M., Emmenegger, L., and Mohn, J.: In situ observations of the isotopic composition of methane at the Cabauw tall tower site, *Atmos. Chem. Phys.*, 16, 10469–10487, <https://doi.org/10.5194/acp-16-10469-2016>, 2016.

- Santrock, J., Studley, S. A., and Hayes, J. M.: Isotopic analyses based on the mass spectra of carbon dioxide, *Anal. Chem.*, 57, 1444, <https://doi.org/10.1021/ac00284a060>, 1985.
- Sapart, C. J., van der Veen, C., Vigano, I., Brass, M., van de Wal, R. S. W., Bock, M., Fischer, H., Sowers, T., Buizert, C., Sperlich, P., Blunier, T., Behrens, M., Schmitt, J., Seth, B., and Röckmann, T.: Simultaneous stable isotope analysis of methane and nitrous oxide on ice core samples, *Atmos. Meas. Tech.*, 4, 2607–2618, <https://doi.org/10.5194/amt-4-2607-2011>, 2011.
- Schaefer, H., Mikaloff Fletcher, S. E., Veidt, C., Lassey, K. R., Brailsford, G. W., Bromley, T. M., Dlugokencky, E. J., Michel, S. E., Miller, J. B., Levin, I., Lowe, D. C., Martin, R. J., Vaughn, B. H., and White, J. W. C.: A 21st-century shift from fossil-fuel to biogenic methane emissions indicated by $^{13}\text{CH}_4$, *Science*, 352, 80–84, <https://doi.org/10.1126/science.aad2705>, 2016.
- Schmitt, J., Seth, B., Bock, M., van der Veen, C., Möller, L., Sapart, C. J., Prokopiou, M., Sowers, T., Röckmann, T., and Fischer, H.: On the interference of Kr during carbon isotope analysis of methane using continuous-flow combustion-isotope ratio mass spectrometry, *Atmos. Meas. Tech.*, 6, 1425–1445, <https://doi.org/10.5194/amt-6-1425-2013>, 2013.
- Schmitt, J., Seth, B., Bock, M., and Fischer, H.: Online technique for isotope and mixing ratios of CH_4 , N_2O , Xe and mixing ratios of organic trace gases on a single ice core sample, *Atmos. Meas. Tech.*, 7, 2645–2665, <https://doi.org/10.5194/amt-7-2645-2014>, 2014.
- Schwietzke, S., Sherwood, O. A., Bruhwiler, L. M. P., Miller, J. B., Etiope, G., Dlugokencky, E. J., Michel, S. E., Arling, V. A., Vaughn, B. H., White, J. W. C., and Tans, P. P.: Upward revision of global fossil fuel methane emissions based on isotope database, *Nature*, 538, 88–91, <https://doi.org/10.1038/nature19797>, 2016.
- Sherwood, O. A., Schwietzke, S., Arling, V. A., and Etiope, G.: Global Inventory of Gas Geochemistry Data from Fossil Fuel, Microbial and Burning Sources, version 2017, *Earth Syst. Sci. Data*, 9, 639–656, <https://doi.org/10.5194/essd-9-639-2017>, 2017.
- Sowers, T.: Late Quaternary atmospheric CH_4 isotope record suggests marine clathrates are stable, *Science*, 311, 838–840, <https://doi.org/10.1126/science.1121235>, 2006.
- Sowers, T.: Atmospheric methane isotope records covering the Holocene period, *Quaternary Sci. Rev.*, 29, 213–221, <https://doi.org/10.1016/j.quascirev.2009.05.023>, 2010.
- Sowers, T., Bernard, S., Aballain, O., Chappellaz, J., Barnola, J. M., and Marik, T.: Records of the $\delta^{13}\text{C}$ of atmospheric CH_4 over the last 2 centuries as recorded in Antarctic snow and ice, *Global Biogeochem. Cy.*, 19, GB2002, <https://doi.org/10.1029/2004GB002408>, 2005.
- Sperlich, P., Guillevic, M., Buizert, C., Jenk, T. M., Sapart, C. J., Schaefer, H., Popp, T. J., and Blunier, T.: A combustion setup to precisely reference $\delta^{13}\text{C}$ and $\delta^2\text{H}$ isotope ratios of pure CH_4 to produce isotope reference gases of $\delta^{13}\text{C}\text{-CH}_4$ in synthetic air, *Atmos. Meas. Tech.*, 5, 2227–2236, <https://doi.org/10.5194/amt-5-2227-2012>, 2012.
- Sperlich, P., Buizert, C., Jenk, T. M., Sapart, C. J., Prokopiou, M., Röckmann, T., and Blunier, T.: An automated GC-C-GC-IRMS setup to measure palaeoatmospheric $\delta^{13}\text{C}\text{-CH}_4$, $\delta^{15}\text{N}\text{-N}_2\text{O}$ and $\delta^{18}\text{O}\text{-N}_2\text{O}$ in one ice core sample, *Atmos. Meas. Tech.*, 6, 2027–2041, <https://doi.org/10.5194/amt-6-2027-2013>, 2013.
- Sperlich, P., Schaefer, H., Mikaloff Fletcher, S. E., Guillevic, M., Lassey, K., Sapart, C. J., Röckmann, T., and Blunier, T.: Carbon isotope ratios suggest no additional methane from boreal wetlands during the rapid Greenland Interstadial 21.2, *Global Biogeochem. Cy.*, 29, 1962–1976, <https://doi.org/10.1002/2014GB005007>, 2015.
- Sperlich, P., Uitslag, N. A. M., Richter, J. M., Rothe, M., Geilmann, H., van der Veen, C., Röckmann, T., Blunier, T., and Brand, W. A.: Development and evaluation of a suite of isotope reference gases for methane in air, *Atmos. Meas. Tech.*, 9, 3717–3737, <https://doi.org/10.5194/amt-9-3717-2016>, 2016.
- Stevens, C. M. and Rust, F. E.: The Carbon Isotopic Composition of Atmospheric Methane, *J. Geophys. Res.-Ocean. Atmos.*, 87, 4879–4882, <https://doi.org/10.1029/JC087iC07p04879>, 1982.
- Sugawara, S., Nakazawa, T., Inoue, G., Machida, T., Mukai, H., Vinnichenko, N. K., and Khattatov, V. U.: Aircraft measurements of the stable carbon isotopic ratio of atmospheric methane over Siberia, *Global Biogeochem. Cy.*, 10, 223–231, <https://doi.org/10.1029/96GB00038>, 1996.
- Teama, D. G. M.: A 30-Year Record of the Isotopic Composition of Atmospheric Methane, *Dissertations and Theses, Paper 642*, <https://doi.org/10.15760/etd.642>, 2013.
- Turner, A. J., C. Frankenberg, C., Wennberg, P. O., and Jacob, D. J.: Ambiguity in the causes for decadal trends in atmospheric methane and hydroxyl, *P. Natl. Acad. Sci. USA*, 114, 5367–5372, <https://doi.org/10.1073/pnas.1616020114>, 2017.
- Tyler, S. C., Ajie, H. O., Gupta, M. L., Cicerone, R. J., Blake, D. R., and Dlugokencky, E. J.: Stable carbon isotopic composition of atmospheric methane: A comparison of surface level and free tropospheric air, *J. Geophys. Res.*, 104, 13895–13910, <https://doi.org/10.1029/1999JD900029>, 1999.
- Tyler, S. C., Rice, A. L., and Ajie, H. O.: Stable isotope ratios in atmospheric CH_4 : Implications for seasonal sources and sinks, *J. Geophys. Res.*, 112, D03303, <https://doi.org/10.1029/2006JD007231>, 2007.
- Umezawa, T., Aoki, S., Nakazawa, T., and Morimoto, S.: A High-precision Measurement System for Carbon and Hydrogen Isotopic Ratios of Atmospheric Methane and Its Application to Air Samples Collected in the Western Pacific Region, *J. Meteorol. Soc. Jpn.*, 87, 365–379, <https://doi.org/10.2151/jmsj.87.365>, 2009.
- Umezawa, T., Aoki, S., Kim, Y., Morimoto, S., and Nakazawa, T.: Carbon and hydrogen stable isotopic ratios of methane emitted from wetlands and wildfires in Alaska: Aircraft observations and bonfire experiments, *J. Geophys. Res.*, 116, D15305, <https://doi.org/10.1029/2010JD015545>, 2011.
- Umezawa, T., Machida, T., Ishijima, K., Matsueda, H., Sawa, Y., Patra, P. K., Aoki, S., and Nakazawa, T.: Carbon and hydrogen isotopic ratios of atmospheric methane in the upper troposphere over the Western Pacific, *Atmos. Chem. Phys.*, 12, 8095–8113, <https://doi.org/10.5194/acp-12-8095-2012>, 2012a.
- Umezawa, T., Machida, T., Aoki, S., and Nakazawa, T.: Contributions of natural and anthropogenic sources to atmospheric methane variations over western Siberia estimated from its carbon and hydrogen stable isotopes, *Global Biogeochem. Cy.*, 26, GB4009, <https://doi.org/10.1029/2011GB004232>, 2012b.
- Umezawa, T., Goto, D., Aoki, S., Ishijima, K., Patra, P. K., Sugawara, S., Morimoto, S., and Nakazawa, T.: Variations of tro-

- ospheric methane over Japan during 1988–2010, *Tellus B*, 66, 23837, <https://doi.org/10.3402/tellusb.v66.23837>, 2014.
- Wendeberg, M., Richter, J. M., Rothe, M., and Brand, W. A.: Jena Reference Air Set (JRAS): a multi-point scale anchor for isotope measurements of CO_2 in air, *Atmos. Meas. Tech.*, 6, 817–822, <https://doi.org/10.5194/amt-6-817-2013>, 2013.
- Verkouteren, R. M., Allison, C. E., Studley, S. A., and Leckrone, K. J.: Isotopic metrology of carbon dioxide. I. Interlaboratory comparison and empirical modeling of inlet equilibration time, inlet pressure, and ion source conductance, *Rapid Commun. Mass Sp.*, 17, 771–776, <https://doi.org/10.1002/rcm.905>, 2003a.
- Verkouteren, R. M., Assonov, S., Klinedinst, D. B., and Brand, W. A.: Isotopic metrology of carbon dioxide. II. Effects of ion source materials, conductance, emission, and accelerating voltage on dual-inlet cross contamination, *Rapid Commun. Mass Sp.*, 17, 777–782, <https://doi.org/10.1002/rcm.906>, 2003b.
- Werner R. A. and Brand, W. A.: Referencing strategies and techniques in stable isotope ratio analysis, *Rapid Commun. Mass Sp.*, 15, 501–519, <https://doi.org/10.1002/rcm.258>, 2001.
- White, J. W. C., Vaughn, B. H., and Michel, S. E.: University of Colorado, Institute of Arctic and Alpine Research (INSTAAR), Stable Isotopic Composition of Atmospheric Methane (^2H) from the NOAA ESRL Carbon Cycle Cooperative Global Air Sampling Network, 2005–2009, Version: 2016-04-26, available at: ftp://aftp.cmdl.noaa.gov/data/trace_gases/ch4h2/flask/, last access: 21 November 2016.
- World Meteorological Organization: 18th WMO/IAEA Meeting on carbon dioxide, other greenhouse gases, and related measurement techniques (GGMT-2015), GAW report No. 229, 2016.