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Methodology and Information Content of the NOAA/NESDIS Operational Channel Selection for the Cross-Track Infrared Sounder (CrIS)

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Methodology and Information Content of the NOAA/NESDIS

Operational Channel Selection for the Cross-Track Infrared Sounder (CrIS)

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Methodology and information content of the NOAA NESDIS operational channel selection for the Cross-Track Infrared Sounder (CrIS)

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Abstract

We describe the NOAA NESDIS channel selection methodology for the Cross-track Infrared Sounder, scheduled to be launched in October 2011 and to become operational in 2012. We describe the main spectral characteristics and we discuss the information content of the final channel subset that will be distributed to the scientific community for operational data assimilation and retrieval applications. This channel selection is composed of 399 channels and its information content is shown to retain most of the total atmospheric variability contained in the original 1305 channel spectrum, up to instrumental noise.

1 Introduction

The Cross-track Infrared Sounder (CrIS) and the Advanced Technology Microwave Sounder (ATMS) are planned to be flown aboard the Joint Polar Satellite System (JPSS) and form the next generation of atmospheric operational sounders [4]. The CrIS instrument is a Fourier transform spectrometer with a total of 1305 infrared sounding channels spread over 3 bands covering the long-wave ($655\text{-}1095\text{ cm}^{-1}$), midwave ($1210\text{-}1750\text{ cm}^{-1}$), and shortwave ($2155\text{-}2550\text{ cm}^{-1}$) region, with a spectral resolution of 0.625 cm^{-1} , 1.25 cm^{-1} and 2.5 cm^{-1} respectively.

Given the large yield of the expected data flow, a channel selection is required that will expedite the near real time data distribution and assimilation. Although hyper spectral data are affected by a considerable redundancy, at-

attention must be paid aimed at minimizing the loss of information during the selection procedure, such that the final retrieval quality is not deteriorated.

The purpose of this paper is to describe the NOAA/NESDIS channel selection methodology applied to the CrIS instrument and to present the main spectral characteristics of the final channel subset that will be distributed to the scientific community for near real time data assimilation and retrieval applications. This paper is divided in two parts. Section 2 provides a detailed description of the channel selection methodology. Focus of this section is an assessment of the degree of spectral purity associated to each selected channel. Section 3 and 4 assess the information content present in the selection by performing a principal component analysis and quantifying the fraction of the full spectrum atmospheric variance explained by this selection.

2 Channel Selection Methodology

The finite spectral resolution of the instrument does not allow for *spectral purity*. That is, in the infrared domain, the signal associated to a channel of nominal frequency ν is the result of multiple molecular roto-vibrational transitions occurring in the same spectral range of the nominal spectral resolution, $\Delta\nu$, of that channel. For a given atmospheric species of interest then, the accuracy of the retrieval and its associated error characterization, rests on the knowledge of the degree of spectral purity present in the subset of channels selected to run the retrieval step of that given species.

We employ the UMBC SARTA radiative transfer model [14] to compute a spectral sensitivity analysis of the full CrIS spectrum. We use standard US atmospheric profiles to perform brute force perturbations of temperature, water vapor and trace gases (ozone, methane, carbon dioxide, carbon monoxide, HNO_3 , N_2O , and SO_2). We first analyze the radiance response to each perturbation, $\Delta R(\nu)$, represented by the difference between the perturbed and the unperturbed radiance spectra. In this methodology, derivatives of the radiance spectrum, defined as spectral sensitivity functions $S(\nu)$, are computed by finite differencing for each perturbation δX_j , according to the equations:

$$S(\nu)\delta X_j = \frac{\partial R(X_o, \nu)}{\partial X_j} \delta X_j \approx \Delta R_j(\nu) \quad (1)$$

$$\Delta R_j(\nu) = (R(X_o + \delta X_j) - R(X_o)) \quad (2)$$

Where R is the radiance signal, X_o represents the unperturbed atmospheric state, δX_j is the perturbation associated to species j . These radiance differences indicate the sensitivity of each channel to each specific atmospheric species of interest and are used in our study as indexes of spectral purity.

Figure 1, provides an illustrative example of this sensitivity analysis, applied to the case of a mid-latitude US standard atmospheric profile. Here the different curves represent differences in brightness temperature due to a perturbation in

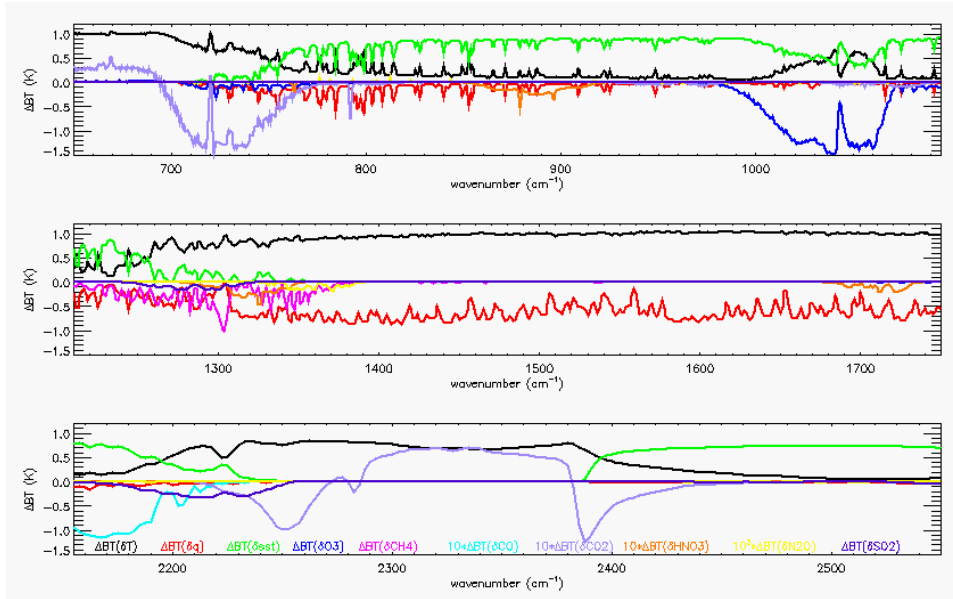


Figure 1: Sensitivity Analysis of CrIS channels. Black curve: temperature sensitivity analysis. Red curve: water vapor sensitivity analysis. Green curve: surface temperature sensitivity analysis. Blue curve: ozone sensitivity analysis. magenta curve: methane sensitivity analysis. Cyan curve: carbon monoxide sensitivity analysis. For display purposes a factor multiplier equal to 10 has been used. Light Purple: carbon dioxide sensitivity analysis. For display purposes a factor multiplier equal to 10 has been used. Orange curve: HNO_3 sensitivity analysis. For display purposes a factor multiplier, FM, equal to 10 has been used. Yellow curve: N_2O sensitivity analysis. For display purposes a factor multiplier equal to 10^3 has been used. Dark Purple: SO_2 sensitivity analysis.

temperature (1K degree perturbation, black curve), water vapor (10% perturbation, red curve), ozone (10% perturbation, blue curve), surface temperature (1K degree perturbation, green curve), methane (2% perturbation, magenta) and carbon monoxide (1% perturbation, cyan), carbon dioxide (1% perturbation, light purple), HNO_3 (1% perturbation, orange), N_2O (1% perturbation, yellow) and SO_2 (1% perturbation, dark purple).

A preliminary channel screening is computed on the basis of the sensitivity features highlighted in Figure 1. We section the spectrum into different channel subsets, each subset including channels highly sensitive to a specific atmospheric species of interest and with the lowest sensitivity to all remaining species. In doing so, for example, temperature sounding channels will be preferentially selected in the $670cm^{-1}$ and $2300cm^{-1}$ regions, where the sensitivity to temperature perturbation is high (black curve) and the interference of other species is minimal or is well known (see ahead discussion on noise covariance).

After this preliminary screening, we closely study the properties of the vertical sensitivity functions, $S(\nu)$, associated to the selected channels. Since the shape of the sensitivity functions is influenced by temperature and pressure as well as by the vertical variation of the atmospheric species mixing ratio, the selection of channels whose sensitivity functions peak at different altitudes enables vertical sounding of the atmosphere.

The finite spectral resolution of the instrument is responsible for broadening the channel sensitivity functions into coarse vertical ranges. Channels with sharper sensitivity functions are then preferentially selected over broadly structured ones, in the attempt of maximizing the vertical resolution of the retrieval product.

In summary, the NOAA/NESDIS methodology is a physically based procedure where channels are selected solely upon their spectral properties. Redundancy is removed by selecting channels with the highest spectral purity and whose sensitivity function structure maximizes vertical sounding coverage and resolution.

The example shown in figure 1 refers to a mid-latitude profile. We have performed this sensitivity study on multiple atmospheric test cases, representative of different atmospheric regimes (ocean/land, polar, mid-latitude and tropical regimes). The final channel selection proved uniform sensitivity features under different air mass conditions, a fundamental requirement to achieve global optimality and robustness in the retrieval applications.

The final channel selection is composed of a total of 399 channels distributed as 24 surface temperature and emissivity sounding channels, 87 temperature sounding channels, 62 water vapor, 53 ozone, 27 carbon monoxide, 54 methane, 53 carbon dioxide, 24 N_2O , 28 HNO_3 and 24 SO_2 sounding channels. Grey cross symbols on figures 2 indicate the location of all 1305 channels present in CrIS original spectra. Superimposed colored cross symbols indicate the 10 channel subsets forming our final channel selection. Here we have adopted the same color codes of figure 1 to indicate the different atmospheric species to which the selected channels are mainly sensitive. The full list of selected channels is given in the table 1 of the attached appendix. Here, the first column reports the channel index, the second column reports the actual channel frequency. In the third column, we have flagged each channel according to the atmospheric species of highest sensitivity, while the fourth column reports the total column spectral transmissivity of the main interfering atmospheric species (fixed composition gases, water vapor, ozone and liquid water) as an index of spectral purity.

The main spectral and noise characteristics of this channel selection are summarized in the paragraphs below.

2.1 Selection of temperature sounding channels

Temperature sounding requires channels sensitive to well mixed and slowly varying gases. Infrared sounders normally employ channels sensitive to carbon dioxide which has two strong fundamentals in the infrared spectrum, specifically in the $670cm^{-1}$ and $2300cm^{-1}$ regions (commonly expressed in microns as: $15\mu m$

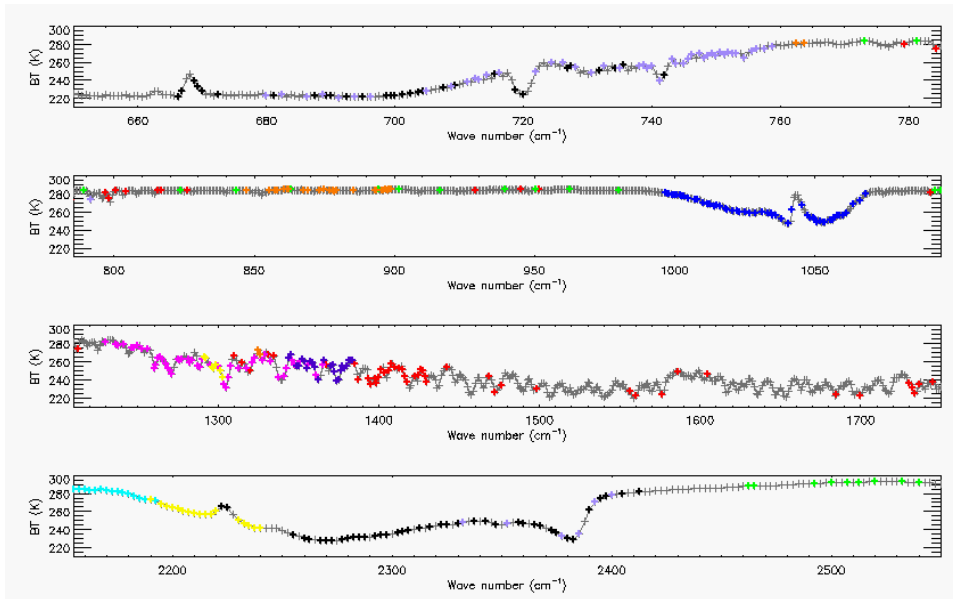


Figure 2: Final channel selection. Grey cross symbols indicate the location of all 1305 channels present in CrIS original spectra. Superimposed colored cross symbols indicate the 10 channel subsets forming our final channel selection. The final selection is composed of 24 surface temperature and emissivity sounding channels (green), 87 temperature sounding channels (black), 62 water vapor (red), 53 ozone (blue), 27 carbon monoxide (cyan), 54 methane (magenta), 53 carbon dioxide (light purple), 24 N_2O (yellow), 28 HNO_3 (orange) and 24 SO_2 (dark purple) sounding channels. The total number of channels is 399.

and $4.3 \mu m$ bands). We follow previous theoretical studies by kaplan *et al.* [10], and preferentially select channels along the wings of CO_2 transition lines, whose weighting functions were shown to have half the width at half peak as that of core channel weighting functions, due to the rapid increase of the absorption coefficient with pressure.

As shown in [9] the most pronounced sensitivity functions are found in the $2300 cm^{-1}$ band. This is due to the rapid increase of the absorption coefficient with tropospheric temperature. Besides the higher instrumental noise, the only complication of the shortwave band is the high solar reflectivity term affecting the surface and low troposphere sounding channels on one side, and the non local thermodynamic equilibrium affecting the upper and stratospheric sensitive channels, on the other. These factors, both occurring during daily conditions, are affected by large uncertainties during forward model calculations.

The complementary $670 cm^{-1}$ channels become fundamental for sounding under the conditions of partial cloudiness [2]. Long-wave channels have good sensitivity to temperature from the surface to the stratosphere, up to 1mb. The

only complication affecting this band rests in the interference of water vapor and ozone as opposed to the head band of the 2300cm^{-1} spectral region which instead shows low sensitivity to these two interfering species.

The temperature subset is also applicable for cloud filtering [15]. For cloud height retrievals based on CO_2 slicing techniques [2], only the long wave subset of this selection is used in order to avoid solar reflectance effects, as described earlier, and because of the uncertainty in the radiative properties of low cloud emissivity.

The selection of surface temperature channels requires high transmittance from all atmospheric components. A total of 24 channels have been selected in the $750 - 1100\text{ cm}^{-1}$ and $2150 - 2540\text{ cm}^{-1}$ window regions. Attention is required when using channels in the $1000 - 1275\text{ cm}^{-1}$ band, due to existing difficulties in the parametrization of desert emissivity.

2.2 Selection of atmospheric constituent sounding channels

For atmospheric constituents, core spectral line have been favored over wing line channels in order to maximize the signal to noise ratio, with equivalent attention to spectral purity and sensitivity function sharpness properties. Complementary wing line channels have been added to the selection in the attempt of improving the vertical coverage of the channel selection sensitivity, enabling vertical profile retrievals.

2.3 Instrumental and Apodization Noise effects

A large effort of the selection methodology also involves a thorough scrutiny of the channel noise properties. The selection generally tends to avoid channels with high instrumental noise such as those in the shortwave region, and channels affected by poor calibration.

Careful attention must also be paid to the noise correlation of adjacent channels resulting from apodization. The expansion of an apodization function into an infinite cosine series enables the determination of the apodization noise reduction factor and spectral correlation among adjacent channels.

Apodized radiances sampled at the Nyquist sampling can be written as a $2J - 1$ point running mean of the raw instrument radiances with weights, w_k , given by Eqn 4:

$$R_{apod}(v') = \sum_{k=-J+1}^{J-1} w_k R(v+k) \quad (3)$$

We can define the weights, w_k , in terms of the cosine expansion coefficients, a_k , as follows:

$$w_k = a_{|k|}, \text{ for } k = -(J-1), \dots, J-1 \quad (4)$$

The apodized noise is reduced by a factor, f , given by:

$$f = \left[\sum_{k=-J+1}^{J-1} w_k^2 \right]^{-1/2} \quad (5)$$

and the channel correlation for a channel separated by n Nyquist channels is given by:

$$CF_n = f^2 \sum_{k=-J+1}^{J-1-n} w_k w_{k+n} \quad (6)$$

where $n=1, \dots, 2J-2$.

CrIS radiances are apodized using a Hamming apodization function. Hamming apodization has $J = 2$ and $a_0 = 0.54$ and $a_1 = 0.23$, so that $w_1 = 0.23$, $w_0 = 0.54$, $w_1 = 0.23$ and $\sum(w_k) = 1$. Hamming apodization is able to reduce CrIS's noise by a factor $f = 1.5862$. In turn, adjacent channels are now correlated by a correlation factor $CF_1 = 62.5\%$ and alternate channels are correlated by $CF_2 = 13.3\%$. Channels separated by more than two indexes have zero correlation. In the attempt of maximizing the vertical sensitivity coverage, the proposed channel selection does contain, at times, few groups of adjacent channels. Users are advised then to pick every other third channel from the proposed selection, if in need of eliminating apodization correlation noise effects. In doing so though, attention must be paid to still retain uniform sensitivity coverage along the vertical atmospheric column.

2.4 Applications

The issue of spectral impurity affecting the accuracy of the retrieval product can be circumvented by exploiting the distinctive non-linearity of the radiative interference sources shown in Figures 1. During the retrieval of a given species of interest, the interference signals of undesired species (held constant at that step) can be treated as radiance error sources that can be used as off-diagonal terms of the observed radiance error covariance matrix. In doing so, the least square minimization is made insensitive to the interfering species and the inversion problem can be solved by parts. In this scheme, the perturbation on species X_j , from which we compute the associated error source $\Delta R_j(\nu)$, is derived from an error estimate on the geophysical species, δX_j . The corresponding computational uncertainty $\Delta R_j(\nu)$ is a finite difference expressed by Eq. 7:

$$\Delta R_j^{s,i}(\nu) = (R(X_o^{s,i-1} + X_j * Q_j) - R(X_o^{s,i-1})) \quad (7)$$

where X_o is the current estimate of the geophysical state at a given step s and iteration i . δX_j is the perturbation on the X_j component.

Since δX_j is an RSS error estimate, it can be vertically cross-correlated and spectrally correlated with respect to other parameters (e.g., surface spectral emissivity error can be correlated with skin temperature). A scaling factor Q_j is needed to compensate for assumed correlations in these error estimates.

During the retrieval step s and iteration i of a given species of interest, the off diagonal terms of the error covariance matrix will be obtained from the summation of all radiance interference terms from undesired geophysical species (held constant during that retrieval step), as in Eq. 8:

$$\Delta R(\nu, \nu') = \sum_{j=0}^{N-1} \Delta R_j^{s,i}(\nu)(\Delta R_j^{s,i}(\nu'))^T \quad (8)$$

The accuracy of the solution will strongly depend on the accuracy of the radiative transfer model used to compute these off-diagonal noise terms.

The approach described above forms the basics of the current NASA operational hyper spectral inversion algorithm for the Atmospheric InfraRed Sounder (AIRS), [15], and the NOAA/NESDIS Infrared Atmospheric Sounding Instrument (IASI), [6], whose operational channel selections were performed following the same methodology described in this paper, [7]. The same inversion scheme will be employed for the NOAA/NESDIS operational processing of CrIS data.

On the other hand, this channel selection is also suitable for alternative 1-DVAR-like inversion techniques, which can resort to the interference of multiple species affecting a given channel, as described by the curves of Figure 1, to simultaneously solve for multiple species of interest. At present, the National Center for Environmental Prediction (NCEP) employs a 1-DVAR inversion technique to operationally process the aforementioned AIRS and IASI channel selections, along with the CrIS channel selection discussed in this paper.

3 Information content analysis

3.1 Singular value decomposition analysis (SVD)

We have used a sub-sample of a global CrIS proxy data from October 19 2007 [1] where we have considered only the first two scan lines of each granule, for a total of 4000 radiance spectra, uniformly including ocean/land/day/night cases. We use this data sample to perform a singular value decomposition (SVD) analysis of the noise weighted radiance covariance matrix, after mean centering of the data. We then presort the derived eigenvectors based on their associated eigenvalues, λ 's, from the highest to the lowest eigenvalue. The first five eigenvalues alone were found to account for more than 99% of the total variance contained in the full ensemble (not shown). We only focus our analysis then on the first five eigenfunctions of this analysis, $U_{n,k}$, as a function of frequency $\nu(n)$, where n is the frequency index and k is the eigenfunction index (Figure 3 for band1, 4 for band 2 and 5 for band 3).

The scope of this analysis is to analyze the contribution of each of the 399 selected channels to these most representative top eigenfunctions, as a qualitative measure of their effective representativeness of the total variance contained in the radiance sample.

While eigenfunctions are representative of the variation of combined atmospheric species, by comparison with figure 1, we can distinctly see the first

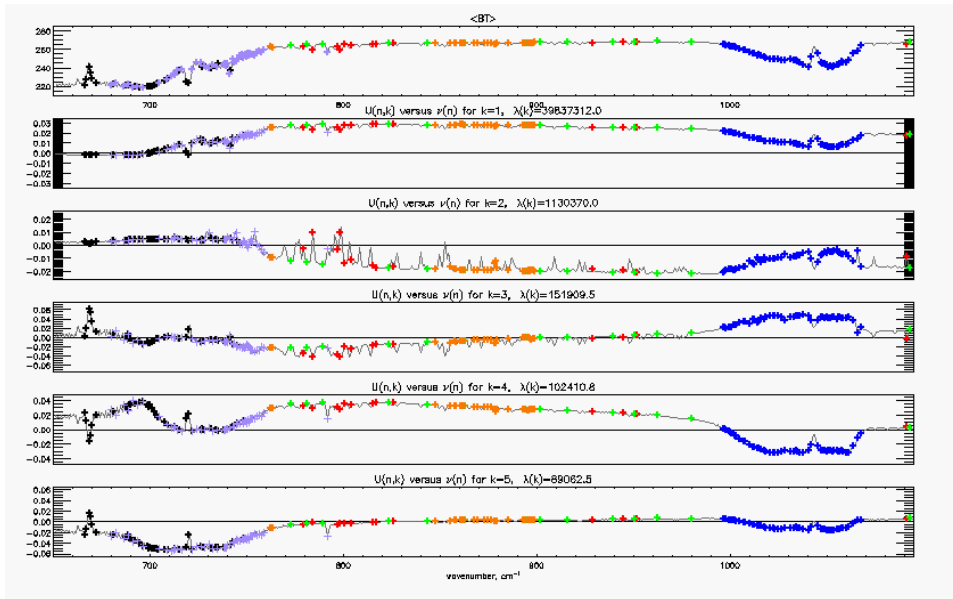


Figure 3: SVD analysis of a global sample of CrIS proxy data. The first 5 eigenfunction are seen to account for more than 99% of the total variance in the sample. Cross symbols indicated the location of the selected channels, color coded according to their main atmospheric sensitivity as in figure 1. In this figure only band 1 is depicted. Top: average brightness temperature.

eigenfunction being mainly representative of the surface temperature signal. We notice that the 24 selected surface temperature and emissivity channels are among the most important contributors to this first eigenvector (green cross symbols).

The highest components of the second eigenvector are found in the 1200-1500 cm^{-1} region where methane, water vapor, N_2O and SO_2 are normally active. We have indeed selected the majority of our methane (magenta crosses), water vapor (red crosses), N_2O (yellow crosses) and SO_2 (dark purple crosses) sounding channels among the highest peaking channels of this region. The highest components of third and fourth eigenvector fall in the 1000 - 1100 cm^{-1} range which is dominated by the ozone signal. We can confirm that our 53 ozone sounding channels (blue crosses) correspond to the highest peaking channels of both eigenvectors. Eigenvectors 4 and 5 are dominated by the carbon dioxide band signal. The selected temperature and carbon dioxide channels (black and light purple crosses) are among the highest components of these two eigenvectors. The selected carbon monoxide (cyan crosses) and N_2O channels (yellow crosses) are among the highest peaking channels of the first and third eigenvectors.

The analysis just shown indicates a high representativeness of the proposed

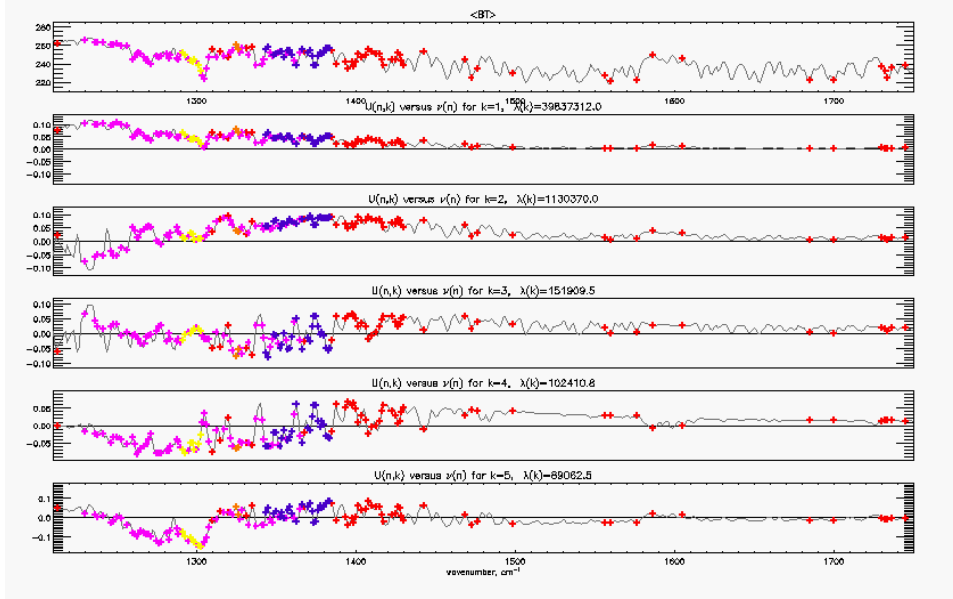


Figure 4: Same as for figure 3, but for band 2.

channel selection in the linear combinations of the orthonormal basis used to describe the main directions of variability contained in the full radiance spectra.

3.2 Variance Explained

As anticipated in the introduction, the scope of this channel selection is to optimize the system computational efficiency by selecting channels carrying unique pieces of information such that the internal variability contained in the original spectrum is still fully represented by the selection, up to instrumental noise. We now address the issue of quantitatively computing the actual fraction of the full spectrum variance explained by the proposed channel selection.

We use a CrIS proxy mid-latitude granule sample of full radiance spectra (1305 channels) and compute a polynomial linear fit to the correspondent channel selection spectra, as described by:

$$\Delta R_{N_f, N_s} = A_{N_f, N_{sub}} \Delta R_{N_{sub}, N_s} \quad (9)$$

where N_f is the size of the full spectrum (1305 channels), N_s is the number of spectra contained in one granule (5940), N_{sub} is the size of the channel selection, $\Delta R_{N_f, N_s}$ is the independent variable matrix of mean centered radiance values, $\Delta R_{N_{sub}, N_s}$ is the corresponding predictor matrix and $A_{N_f, N_{sub}}$ is the regression coefficient matrix, given by:

$$A_{N_f, N_{sub}} = \Delta R_{N_f, N_s} \Delta R_{N_s, N_{sub}}^T [\Delta R_{N_{sub}, N_s} \Delta R_{N_s, N_{sub}}^T]^{-1} \quad (10)$$

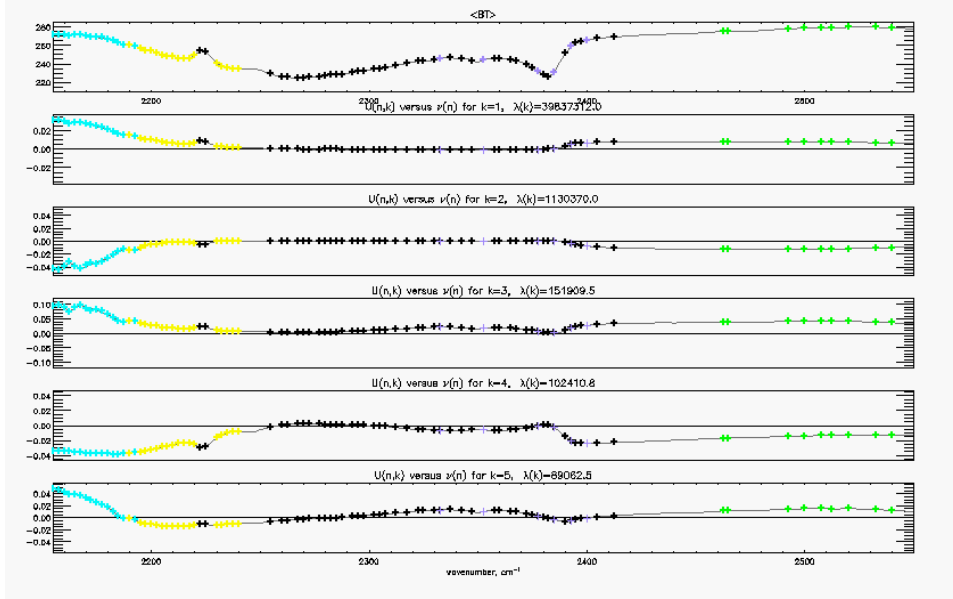


Figure 5: Same as for figure 3, but for band 3.

Once $A_{N_f, N_{sub}}$ is computed, we can derive the calculated $N_f x N_s$ matrix, as in:

$$\Delta R_{N_f, N_s}^{calc} = A_{N_f, N_{sub}} \Delta R_{N_{sub}, N_s} \quad (11)$$

We then compare the variance contained in the original spectra, Var_{tot} with the residual sum of squares, Var_{err} . Specifically, from the ratio of these two terms we derive what in some texts is referred to as the *coefficient of determination*, D , [5], defined as:

$$D = 1 - Var^{err} / Var^{tot} \quad (12)$$

In statistical methods, D is a metric used to determine the goodness of a fit to a model. In regression, D is a statistical measure of how close is the fit to the original data points. The closer D to unity, the better is the fit, that is to say the closer are the predictors (in our case, the actual selected channels) to explain the total variance of the original data sample.

The coefficient of determination is a useful tool for testing the impact of new channels that are either used as a replacement of old ones or that have been incrementally added to a pre-existing selection. Test channels are kept in the selection if they contribute at increasing the total variance explained. In least squares regressions though, like in our case, D increases with the number of predictors. As such, D is not meaningful for comparing selections differing by the actual number of channels. For a meaningful comparison then, we resort to an *adjusted* coefficient of determination, D^{adj} , defined as:

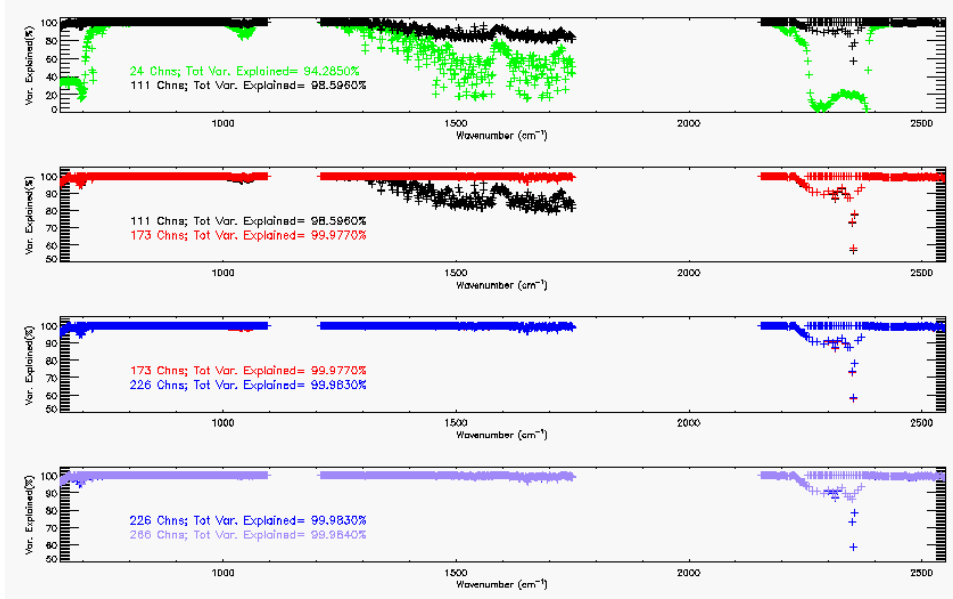


Figure 6: Percentage variance explained. Different colors indicate different channel subset used as predictors. Green curve: surface temperature channel selection only (24 channels). Black curve: previous subset plus 87 temperature sensitive channels. Red curve: like black curve, with 62 additional water vapor sounding channels. Blue: 53 additional ozone sounding channels. Light purple: 53 additional carbon dioxide sounding channels. Captions on figure report the total variance explained by the channel subset. See text for explanation.

$$D^{adj} = 1 - (Var^{err}/Var^{tot})(dof_{tot}/dof_{err}) \quad (13)$$

which represents a statistically unbiased version of D . Here dof_{tot} is the degree of freedom of the population variance of the dependent variable and is given by $N_s^{tot} - 1$, where N_s^{tot} is the sample size of the dependent variable population. dof_{err} is the degree of freedom of the predictor population variance, $N_s^{tot} - N_{sub} - 1$, where N_{sub} is again the number of predictors (selected channels).

We first compute the coefficient of determination for each channel. The variance contained in the original sample, $Var^{tot}(n)$ and the residual sum of squares, $Var^{err}(n)$ corresponding to channel ν of index n , are defined by equation 14 and equation 15, respectively:

$$Var^{tot}(n) = \sum_{s=1}^{N_s} [\Delta R_s(n) - (1/N_s)(\sum_{s=1}^{N_s} \Delta R_s(n))]^2 \quad (14)$$

$$Var^{err}(n) = \sum_{s=1}^{N_s} [\Delta R_s(n) - \Delta R_s^{calc}(n)]^2 \quad (15)$$

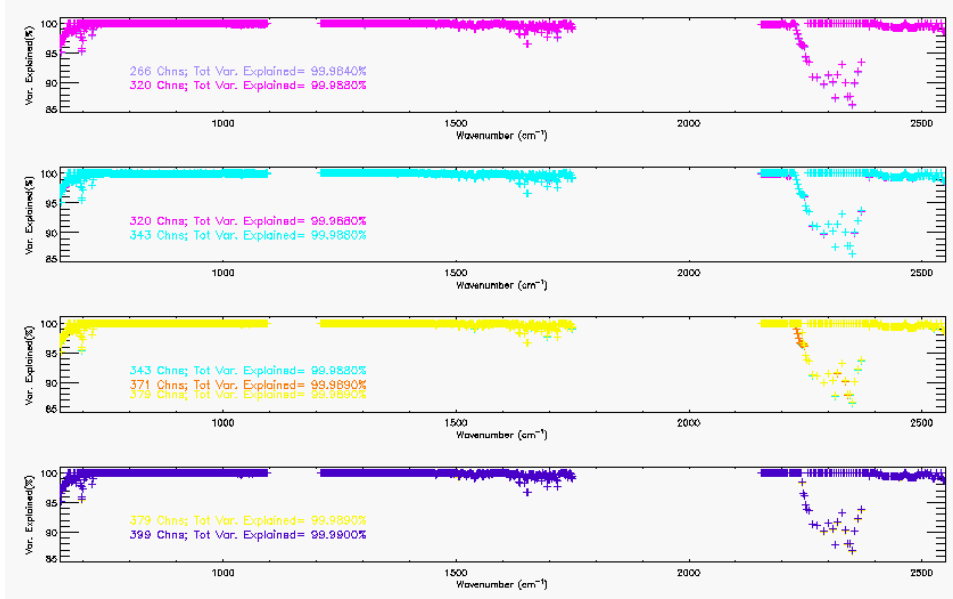


Figure 7: Same as Figure 6. The magenta curve is due to the addition of 54 methane sounding channels. The cyan curve represents the impact of 27 additional carbon monoxide sounding channels. The orange curve refers to the addition of 28 HNO_3 sounding channels, the yellow to 24 N_2O and the dark purple to 24 SO_2 additional sounding channels. For comparison, the results of a channel subset addition are plotted on top of the results of the previous partial selection. Captions on figure report the total population variance explained by the channel subset. See text for explanation.

We then have, for each channel:

$$D^{adj}(n) = 1 - [Var^{err}(n)/Var^{tot}(n)][(N_s - 1)/(N_s - N_{sub} - 1)] \quad (16)$$

For completeness, we have plotted the full sequence of improvements obtained by incrementally adding each of the ten channel subsets described in the previous sessions, each including channels mainly sensitive to a given atmospheric species of interest. Results are plotted in figure 6 and 7. For comparison, the results of a channel subset addition are plotted on top of the results of the previous partial selection. We use the same color code of previous figures.

We start by considering the sole 24 selected surface temperature sensitive channels, and we perform the regression as in equations 9, 10 and 11. We then compute the coefficient of determination for the full spectrum, according to equations 14, 15 and 16. Here $N_{sub}=24$. Green symbols on figure 6 indicate the percentage variance explained by the selected 24 surface temperature and emissivity channels. Regions of the spectrum from where these 24 channels

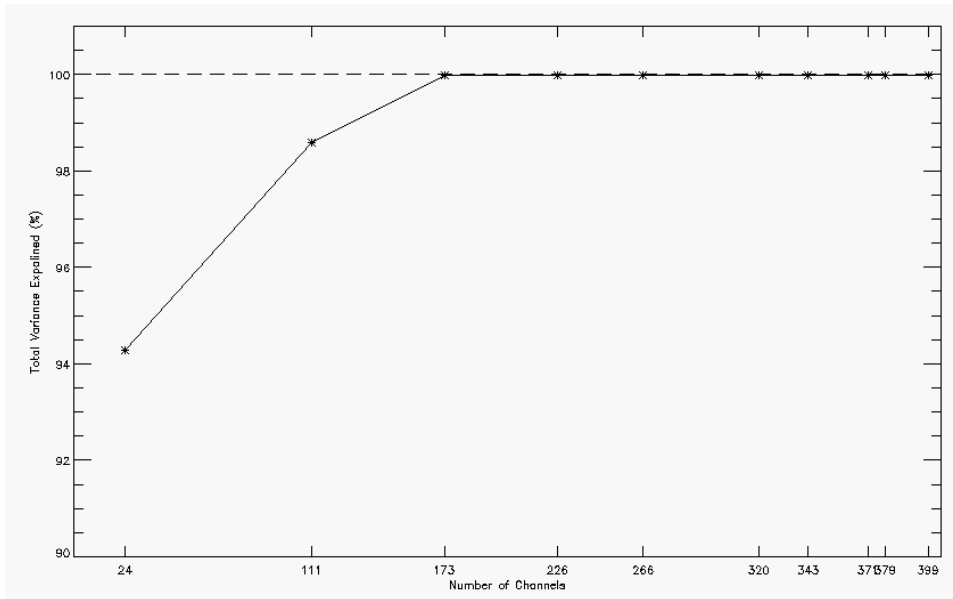


Figure 8: Total population variance explained by each channel subset.

were selected are almost completely explained. This was expected due to the redundancy present across the spectrum.

When we add the 87 temperature channels (670cm^{-1} and 2300cm^{-1} region) we considerably improve the variance explained across most of the longwave and shortwave regions. The addition of the 62 selected water vapor channels (which mainly fall in the $1200 - 1700\text{cm}^{-1}$ region) brings the explained variance in the full mid wave band to more than 99% (red cross symbols). A few percentage more is gained in the $1000 - 1100\text{cm}^{-1}$ region with the addition of the 53 ozone sounding channels (blue cross symbols) and a few percentage improvement is obtained with the addition of the selected 53 CO_2 channels in the shortwave region. The addition of the remaining trace gas subsets brings an improvement of the order of 10^{-3} across the spectrum, as shown in figure 7. The reader should notice the change of scale in the reported values of the y axis. The magenta curve is due to the addition of 54 methane sounding channels. The cyan curve represents the impact of 27 additional carbon monoxide sounding channels. The orange curve refers to the addition of 28 HNO_3 sounding channels, the yellow to 24 N_2O and the dark purple to 24 SO_2 additional sounding channels.

In conclusion, the full 399 channel selection is seen to account for more than 99% of the total variance across the whole spectrum, except for the $600 - 700\text{cm}^{-1}$ and the 1700cm^{-1} regions, where the variance explained is about 95% and for the $2200 - 2300\text{cm}^{-1}$ region, where it oscillates between 85% and 99%. This result was expected and originates from the lower representativeness of these spectral bands in our channel selection. This choice rested in the high instru-

mental noise of the 600 - 700 cm^{-1} region, in the lower spectral purity of the 1700 cm^{-1} with respect to the 1400 cm^{-1} band (from where we have preferentially selected most of the water vapor channels), and in the high instrumental and calibration noise of the shortwave band. For the reasons just explained, retrieval users generally tend to avoid these low signal to noise spectral regions. Accordingly, we do not expect the lower representativeness of these regions to be penalizing in terms of retrieval application performance.

We now summarize the previous discussion by investigating the variance of the full population in the sample. We define:

$$VAR^{TOT} = \sum_{n=1}^{N_f} Var^{tot}(n) \quad (17)$$

$$VAR^{ERR} = \sum_{n=1}^{N_f} Var^{tot}(n) \quad (18)$$

We then have, for the total population variance explained by each subset:

$$D^{adjTOT} = 1 - [VAR^{ERR}/VAR^{TOT}] [(N_f * N_s - 1)/(N_f * N_s - N_{sub} - 1)] \quad (19)$$

Results for the total variance explained, D^{adjTOT} , upon the incremental addition of each subset are listed as captions on figure 6 and 7 and are summarized in figure 8. From this analysis we can see that the first 173 channels including surface and profile temperature and water vapor sensitive channels, are able to already account for more than 99% of the total variability contained in the full sample, while the total variance explained by the complete 399 channel selection is equal to about 99.99%. The case under exam is a mid-latitude case, but tests have been conducted globally and showed equivalent results.

The results just shown confirm that the internal variability contained in the original spectrum is fully explained by the proposed channel selection up to instrumental noise, implying that no significant penalty will be paid in terms of information performance and retrieval application.

4 Conclusions and Future Work

We have described the NOAA/NESDIS physically-based methodology for the near real time channel selection of the Cross-track Infrared Sounder. This channel selection is composed of 24 surface temperature and emissivity sounding channels, 87 temperature sounding channels, 62 water vapor, 53 ozone, 27 carbon monoxide, 54 methane, 53 carbon dioxide, 24 N_2O , 28 HNO_3 and 24 SO_2 sounding channels. The described methodology takes in great care spectral purity and vertical sensitivity features, as well as instrumental and apodization noise properties of the channels. A detailed PCA analysis was performed to

identify the information content carried by this selection. The proposed channel selection is found to carry a high representativeness of the atmospheric variability contained in the full radiance spectra. Furthermore, from a regression analysis, we found that the total internal variability contained in the original spectrum is fully explained by the proposed channel selection up to instrumental noise, implying that no significant penalty will be paid in terms of information content and retrieval performance.

Test applications have shown this channel selection to be significantly advantageous in terms of computational efficiency. For example, the computational time of routine operations, such as the computation of radiance covariance matrices, have been observed to improve by a factor of 10 when going from assimilating full spectra down to the 399 channel selection.

The methodology described in this paper has been previously adopted for the operational channel selection of the AIRS and IASI sensors, [15], [7]. National Weather Prediction users have also found great advantage in employing these AIRS and IASI near real time channel selections to expedite data assimilation processing, while guaranteeing good quality in the retrieval results [8].

Existing methodologies, as described in [12], [11] and [3], follow Rodgers statistical iterative methodology [13] where successive tests are performed on the incremental addition of each channels. At each addition, the channel is retained if the information content (based on the entropy of the system) is observed to increase. For near real time operational applications, a *constant* channel selection is normally used, which is derived as an average from multiple optimal selections computed over the globe.

Differently from Rodgers's statistical technique [13], the NOAA/NESDIS methodology described in this paper employs a physically-based procedure where channels are selected solely upon their spectral properties: high priority is given to spectral purity, avoidance of redundancy and vertical sensitivity properties, along with low instrumental noise and global optimality.

Fourrie and Thepaut (2002) have tested the global optimality of the ECMWF *constant* channel selection and compared it to the AIRS NOAA/NESDIS channel selection. They found the performance of the two selections to be comparable, with the NOAA/NESDIS selection providing global optimality with respect to different background error specifications.

We can expect the proposed near real time CrIS channel selection to carry on AIRS and IASI heritage with similar robustness and global optimality. A cross-comparison of the information content contained in the AIRS, IASI and CrIS operational channel selections is under exam. Vertical resolution and degree of freedom of the signal will be also part of a separate study, once the CrIS NOAA/NESDIS retrieval algorithm is fully optimized and retrieval products become operationally available.

5 Appendix 1

T.1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
27	666.250	T, CF, CH	0.00	0.68	0.98
28	666.875	T, CF, CH	0.00	0.68	0.89
31	668.750	T, CF, CH	0.00	0.69	1.00
32	669.375	T, CF, CH	0.00	0.69	0.90
33	670.000	T, CF, CH	0.00	0.69	0.99
37	672.500	T, CF, CH	0.00	0.70	0.97
49	680.000	CO_2	0.00	0.71	0.97
51	681.250	T, CF, CH	0.00	0.71	0.98
53	682.500	CO_2	0.00	0.72	0.97
59	686.250	CO_2	0.00	0.72	0.95
61	687.500	T, CF, CH	0.00	0.71	0.94
63	688.750	T, CF, CH, CO_2	0.00	0.73	0.96
64	689.375	T, CF, CH	0.00	0.73	0.87
65	690.000	T, CF, CH	0.00	0.73	0.95
67	691.250	CO_2	0.00	0.73	0.93
69	692.500	T, CF, CH	0.00	0.74	0.88
71	693.750	CO_2	0.00	0.73	0.97
73	695.000	T, CF, CH	0.00	0.74	0.97
75	696.250	T, CF, CH, CO_2	0.00	0.74	0.95
79	698.750	T, CF, CH	0.00	0.75	0.97
80	699.375	T, CF, CH	0.00	0.25	0.85
81	700.000	T, CF, CH	0.00	0.75	0.97
83	701.250	T, CF, CH	0.00	0.75	0.98
85	702.500	T, CF, CH	0.00	0.75	0.99
87	703.750	T, CF, CH	0.00	0.70	0.98
88	704.375	T, CF, CH	0.00	0.05	0.80
89	705.000	CO_2	0.00	0.62	0.95
93	707.500	T, CF, CH	0.00	0.68	0.95
95	708.750	CO_2	0.00	0.76	0.89
96	709.375	T, CF, CH	0.00	0.40	0.88
99	711.250	CO_2	0.00	0.76	0.96
101	712.500	CO_2	0.01	0.75	0.93
102	713.125	T, CF, CH, CO_2	0.01	0.49	0.92
104	714.375	CO_2	0.02	0.50	0.92
106	715.625	T, CF, CH	0.04	0.74	0.91
107	716.250	CO_2	0.03	0.76	0.92
111	718.750	T, CF, CH	0.00	0.78	0.93
113	720.000	T, CF, CH	0.00	0.78	0.95
116	721.875	T, CF, CH, CO_2	0.04	0.74	0.93
120	724.375	CO_2	0.14	0.77	0.96
123	726.250	T, CF, CH, CO_2	0.14	0.78	0.97

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
124	726.875	T, CF, CH	0.09	0.77	0.97
125	727.500	T, CF, CH	0.11	0.76	0.97
126	728.125	CO_2	0.10	0.72	0.96
130	730.625	CO_2	0.01	0.08	0.96
132	731.875	T, CF, CH	0.05	0.59	0.98
133	732.500	CO_2	0.09	0.75	0.97
136	734.375	T, CF, CH	0.09	0.78	0.97
137	735.000	CO_2	0.08	0.80	0.96
138	735.625	T, CF, CH	0.10	0.79	0.97
142	738.125	T, CF, CH	0.09	0.78	0.97
143	738.750	CO_2	0.10	0.79	0.98
144	739.375	CO_2	0.07	0.68	0.98
145	740.000	CO_2	0.06	0.53	0.98
147	741.250	CO_2	0.01	0.65	0.97
148	741.875	T, CF, CH	0.02	0.53	0.96
150	743.125	T, CF, CH, CO_2	0.18	0.64	0.94
151	743.750	CO_2	0.11	0.47	0.96
153	745.000	CO_2	0.06	0.22	0.97
154	745.625	T, CF, CH, CO_2	0.20	0.65	0.98
155	746.250	CO_2	0.28	0.77	0.98
157	747.500	CO_2	0.29	0.70	0.98
158	748.125	CO_2	0.21	0.61	0.98
159	748.750	T, CF, CH, CO_2	0.25	0.64	0.98
160	749.375	CO_2	0.30	0.72	0.98
161	750.000	CO_2	0.29	0.78	0.96
162	750.625	CO_2	0.37	0.79	0.95
163	751.250	CO_2	0.29	0.74	0.96
164	751.875	CO_2	0.27	0.55	0.98
165	752.500	CO_2	0.24	0.50	0.98
166	753.125	CO_2	0.27	0.58	0.98
168	754.375	CO_2	0.13	0.27	0.98
170	755.625	CO_2	0.34	0.62	0.98
171	756.250	CO_2	0.41	0.72	0.98
173	757.500	CO_2	0.45	0.81	0.97
175	758.750	CO_2	0.54	0.82	0.98
181	762.500	HNO_3	0.68	0.83	0.99
183	763.750	HNO_3	0.68	0.83	0.98
198	773.125	Srf	0.77	0.83	0.99
208	779.375	H_2O	0.55	0.58	0.99
211	781.250	Srf	0.80	0.83	0.99
216	784.375	H_2O	0.30	0.31	0.99
224	789.375	Srf	0.83	0.85	1.00
228	791.875	CO_2	0.42	0.81	1.00

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
236	796.875	H_2O	0.55	0.57	1.00
238	798.125	H_2O	0.29	0.31	1.00
242	800.625	H_2O	0.80	0.82	1.00
248	804.375	H_2O	0.73	0.76	1.00
266	815.625	H_2O	0.82	0.84	1.00
268	816.875	H_2O	0.85	0.87	1.00
279	823.750	Srf	0.87	0.88	1.00
283	826.250	H_2O	0.85	0.85	1.00
311	843.750	Srf	0.87	0.89	1.00
317	847.500	HNO_3	0.87	0.89	1.00
330	855.625	HNO_3	0.85	0.85	1.00
333	857.500	HNO_3	0.89	0.89	1.00
334	858.125	HNO_3	0.87	0.88	1.00
338	860.625	HNO_3	0.89	0.90	1.00
340	861.875	HNO_3	0.90	0.90	1.00
341	862.500	HNO_3	0.89	0.90	1.00
342	863.125	Srf	0.90	0.90	1.00
349	867.500	HNO_3	0.90	0.90	1.00
352	869.375	HNO_3	0.90	0.90	1.00
358	873.125	HNO_3	0.90	0.90	1.00
361	875.000	HNO_3	0.90	0.91	1.00
364	876.875	HNO_3	0.90	0.90	1.00
366	878.125	HNO_3	0.77	0.78	1.00
367	878.750	HNO_3	0.74	0.75	1.00
368	879.375	HNO_3	0.85	0.87	1.00
378	885.625	HNO_3	0.89	0.90	1.00
390	893.125	HNO_3	0.91	0.92	1.00
391	893.750	HNO_3	0.91	0.91	1.00
392	894.375	Srf	0.91	0.92	1.00
394	895.625	HNO_3	0.90	0.91	1.00
395	896.250	HNO_3	0.90	0.91	1.00
396	896.875	HNO_3	0.90	0.91	1.00
397	897.500	HNO_3	0.89	0.89	1.00
398	898.125	HNO_3	0.89	0.90	1.00
399	898.750	HNO_3	0.91	0.91	1.00
404	901.875	Srf	0.91	0.92	1.00
427	916.250	Srf	0.90	0.92	1.00
447	928.750	H_2O	0.87	0.91	1.00
464	939.375	Srf	0.91	0.93	1.00
473	945.000	H_2O	0.88	0.90	1.00
482	950.625	Srf	0.91	0.94	1.00
484	951.875	H_2O	0.91	0.94	1.00
501	962.500	Srf	0.93	0.94	1.00

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
529	980.000	Srf	0.90	0.94	0.99
556	996.875	O_3	0.85	0.95	0.89
557	997.500	O_3	0.83	0.95	0.88
558	998.125	O_3	0.81	0.93	0.87
560	999.375	O_3	0.78	0.92	0.85
561	1000.000	O_3	0.77	0.91	0.84
562	1000.625	O_3	0.77	0.91	0.84
564	1001.875	O_3	0.78	0.95	0.82
565	1002.500	O_3	0.76	0.95	0.81
566	1003.125	O_3	0.75	0.93	0.80
569	1005.000	O_3	0.71	0.94	0.76
573	1007.500	O_3	0.67	0.95	0.71
574	1008.125	O_3	0.68	0.95	0.72
577	1010.000	O_3	0.55	0.83	0.66
580	1011.875	O_3	0.58	0.94	0.62
581	1012.500	O_3	0.58	0.95	0.62
584	1014.375	O_3	0.45	0.79	0.57
585	1015.000	O_3	0.50	0.85	0.59
587	1016.250	O_3	0.52	0.95	0.55
590	1018.125	O_3	0.43	0.86	0.51
591	1018.750	O_3	0.45	0.93	0.49
594	1020.625	O_3	0.43	0.95	0.45
597	1022.500	O_3	0.43	0.95	0.45
598	1023.125	O_3	0.42	0.96	0.44
601	1025.000	O_3	0.40	0.96	0.42
604	1026.875	O_3	0.40	0.95	0.42
607	1028.750	O_3	0.33	0.79	0.43
611	1031.250	O_3	0.39	0.93	0.43
614	1033.125	O_3	0.38	0.94	0.41
616	1034.375	O_3	0.38	0.96	0.40
617	1035.000	O_3	0.34	0.95	0.36
619	1036.250	O_3	0.34	0.96	0.36
622	1038.125	O_3	0.28	0.94	0.30
626	1040.625	O_3	0.21	0.96	0.22
628	1041.875	O_3	0.46	0.93	0.51
634	1045.625	O_3	0.55	0.96	0.60
637	1047.500	O_3	0.33	0.96	0.36
638	1048.125	O_3	0.30	0.96	0.32
640	1049.375	O_3	0.27	0.87	0.32
641	1050.000	O_3	0.26	0.92	0.31
642	1050.625	O_3	0.20	0.90	0.25
644	1051.875	O_3	0.23	0.90	0.28
646	1053.125	O_3	0.22	0.96	0.24

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
647	1053.750	O_3	0.22	0.96	0.24
650	1055.625	O_3	0.21	0.82	0.28
651	1056.250	O_3	0.23	0.91	0.27
652	1056.875	O_3	0.28	0.96	0.31
654	1058.125	O_3	0.32	0.94	0.35
655	1058.750	O_3	0.31	0.93	0.35
657	1060.000	O_3	0.30	0.87	0.35
659	1061.250	O_3	0.38	0.95	0.41
663	1063.750	O_3	0.56	0.95	0.59
667	1066.250	O_3	0.43	0.58	0.78
670	1068.125	O_3	0.80	0.96	0.87
707	1091.250	H_2O	0.65	0.69	0.98
710	1093.125	Srf	0.91	0.96	0.97
713	1095.000	Srf	0.91	0.96	0.97
716	1212.500	H_2O	0.29	0.30	1.00
730	1230.000	CH_4	0.71	0.88	1.00
735	1236.250	CH_4	0.59	0.79	1.00
736	1237.500	CH_4	0.54	0.74	1.00
739	1241.250	CH_4	0.55	0.71	1.00
743	1246.250	CH_4	0.45	0.71	1.00
744	1247.500	CH_4	0.50	0.77	1.00
746	1250.000	CH_4	0.53	0.73	1.00
748	1252.500	CH_4	0.40	0.58	1.00
751	1256.250	CH_4	0.41	0.67	1.00
754	1260.000	CH_4	0.03	0.08	1.00
755	1261.250	CH_4	0.07	0.20	1.00
756	1262.500	CH_4	0.16	0.43	1.00
757	1263.750	CH_4	0.18	0.42	1.00
758	1265.000	CH_4	0.11	0.25	1.00
760	1267.500	CH_4	0.05	0.14	1.00
761	1268.750	CH_4	0.01	0.05	1.00
762	1270.000	CH_4	0.00	0.01	1.00
763	1271.250	CH_4	0.00	0.01	1.00
766	1275.000	CH_4	0.19	0.59	1.00
767	1276.250	CH_4	0.19	0.71	1.00
768	1277.500	CH_4	0.23	0.74	1.00
771	1281.250	CH_4	0.13	0.31	1.00
772	1282.500	CH_4	0.13	0.43	1.00
773	1283.750	CH_4	0.11	0.26	1.00
776	1287.500	CH_4	0.03	0.09	1.00
777	1288.750	CH_4	0.05	0.16	1.00
778	1290.000	CH_4	0.11	0.25	1.00
779	1291.250	N_2O	0.17	0.36	1.00
780	1292.500	CH_4, N_2O	0.17	0.47	1.00

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
782	1295.00	CH_4	0.10	0.41	1.0
783	1296.25	CH_4, N_2O	0.05	0.21	1.0
784	1297.50	N_2O	0.07	0.30	1.0
785	1298.75	N_2O	0.13	0.63	1.0
786	1300.00	CH_4	0.12	0.67	1.0
787	1301.25	CH_4, N_2O	0.07	0.70	1.0
788	1302.50	CH_4, N_2O	0.03	0.65	1.0
789	1303.75	CH_4	0.01	0.78	1.0
790	1305.00	CH_4	0.00	0.28	1.0
791	1306.25	CH_4	0.02	0.12	1.0
792	1307.50	CH_4	0.03	0.07	1.0
794	1310.00	H_2O	0.10	0.17	1.0
796	1312.50	CH_4	0.01	0.02	1.0
798	1315.00	H_2O	0.00	0.01	1.0
800	1317.50	CH_4	0.00	0.00	1.0
802	1320.00	H_2O	0.00	0.00	1.0
803	1321.25	CH_4	0.02	0.02	1.0
804	1322.50	CH_4	0.04	0.05	1.0
806	1325.00	HNO_3	0.16	0.18	1.0
807	1326.25	HNO_3	0.13	0.20	1.0
808	1327.50	CH_4	0.08	0.15	1.0
809	1328.75	CH_4	0.10	0.13	1.0
811	1331.25	H_2O	0.15	0.22	1.0
812	1332.50	CH_4	0.08	0.18	1.0
814	1335.00	H_2O	0.04	0.04	1.0
816	1337.50	CH_4	0.00	0.00	1.0
819	1341.25	CH_4	0.00	0.00	1.0
820	1342.50	CH_4	0.01	0.03	1.0
821	1343.75	SO_2	0.03	0.05	1.0
822	1345.00	H_2O, SO_2	0.05	0.07	1.0
823	1346.25	CH_4	0.03	0.07	1.0
824	1347.50	CH_4	0.02	0.04	1.0
825	1348.75	SO_2	0.01	0.01	1.0
826	1350.00	SO_2	0.01	0.01	1.0
827	1351.25	CH_4	0.02	0.03	1.0
828	1352.50	SO_2	0.03	0.04	1.0
829	1353.75	SO_2	0.02	0.02	1.0
830	1355.00	SO_2	0.00	0.01	1.0
831	1356.25	CH_4	0.01	0.01	1.0
832	1357.50	H_2O, SO_2	0.01	0.01	1.0
833	1358.75	CH_4, SO_2	0.00	0.00	1.0

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
834	1360.00	SO_2	0.0	0.0	1.0
835	1361.25	CH_4	0.0	0.0	1.0
836	1362.50	SO_2	0.0	0.0	1.0
838	1365.000	CH_4	0.0	0.0	1.0
839	1366.250	SO_2	0.0	0.0	1.0
840	1367.500	H_2O	0.0	0.0	1.0
842	1370.000	SO_2	0.0	0.0	1.0
843	1371.250	SO_2	0.0	0.0	1.0
844	1372.500	SO_2	0.0	0.0	1.0
845	1373.750	SO_2	0.0	0.0	1.0
846	1375.000	SO_2	0.0	0.0	1.0
847	1376.250	SO_2	0.0	0.0	1.0
848	1377.500	SO_2	0.0	0.0	1.0
849	1378.750	SO_2	0.0	0.0	1.0
850	1380.000	SO_2	0.0	0.0	1.0
851	1381.250	H_2O, SO_2	0.0	0.0	1.0
852	1382.500	SO_2	0.0	0.0	1.0
853	1383.750	SO_2	0.0	0.0	1.0
854	1385.000	H_2O	0.0	0.0	1.0
856	1387.500	H_2O	0.0	0.0	1.0
861	1393.750	H_2O	0.0	0.0	1.0
862	1395.000	H_2O	0.0	0.0	1.0
864	1397.500	H_2O	0.0	0.0	1.0
865	1398.750	H_2O	0.0	0.0	1.0
866	1400.000	H_2O	0.0	0.0	1.0
867	1401.250	H_2O	0.0	0.0	1.0
869	1403.750	H_2O	0.0	0.0	1.0
871	1406.250	H_2O	0.0	0.0	1.0
872	1407.500	H_2O	0.0	0.0	1.0
874	1410.000	H_2O	0.0	0.0	1.0
876	1412.500	H_2O	0.0	0.0	1.0
878	1415.000	H_2O	0.0	0.0	1.0
879	1416.250	H_2O	0.0	0.0	1.0
880	1417.500	H_2O	0.0	0.0	1.0
884	1422.500	H_2O	0.0	0.0	1.0
886	1425.000	H_2O	0.0	0.0	1.0
887	1426.250	H_2O	0.0	0.0	1.0
888	1427.500	H_2O	0.0	0.0	1.0
889	1428.750	H_2O	0.0	0.0	1.0
890	1430.000	H_2O	0.0	0.0	1.0
900	1442.500	H_2O	0.0	0.0	1.0
921	1468.750	H_2O	0.0	0.0	1.0
924	1472.500	H_2O	0.0	0.0	1.0

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
927	1476.25	H_2O	0.00	0.00	1.00
945	1498.75	H_2O	0.00	0.00	1.00
991	1556.25	H_2O	0.00	0.00	1.00
994	1560.00	H_2O	0.00	0.00	1.00
1007	1576.25	H_2O	0.00	0.00	1.00
1015	1586.25	H_2O	0.00	0.00	1.00
1030	1605.00	H_2O	0.00	0.00	1.00
1094	1685.00	H_2O	0.00	0.00	1.00
1106	1700.00	H_2O	0.00	0.00	1.00
1130	1730.00	H_2O	0.00	0.00	1.00
1132	1732.50	H_2O	0.00	0.00	1.00
1133	1733.75	H_2O	0.00	0.00	1.00
1135	1736.25	H_2O	0.00	0.00	0.98
1142	1745.00	H_2O	0.00	0.00	1.00
1147	2155.00	CO	0.73	0.85	1.00
1148	2157.50	CO	0.75	0.89	1.00
1149	2160.00	CO	0.70	0.86	1.00
1150	2162.50	CO	0.64	0.79	1.00
1151	2165.00	CO	0.69	0.88	1.00
1152	2167.50	CO	0.72	0.95	1.00
1153	2170.00	CO	0.66	0.90	1.00
1154	2172.50	CO	0.63	0.89	1.00
1155	2175.00	CO	0.63	0.95	1.00
1156	2177.50	CO	0.60	0.96	1.00
1157	2180.00	CO	0.52	0.92	1.00
1158	2182.50	CO	0.44	0.89	1.00
1159	2185.00	CO	0.36	0.86	1.00
1160	2187.50	CO	0.32	0.86	1.00
1161	2190.00	CO, N_2O	0.33	0.94	1.00
1162	2192.50	CO	0.31	0.94	1.00
1163	2195.00	CO, N_2O	0.24	0.96	1.00
1164	2197.50	CO, N_2O	0.19	0.93	1.00
1165	2200.00	T, CF, CO, N_2O	0.17	0.90	1.00
1166	2202.50	T, CF, CO, N_2O	0.15	0.89	1.00
1167	2205.00	CO, N_2O	0.11	0.89	1.00
1168	2207.50	T, CF, CO, N_2O	0.10	0.86	1.00
1169	2210.00	CO, N_2O	0.10	0.89	1.00
1170	2212.50	CO, N_2O	0.08	0.89	1.00
1171	2215.00	CO, N_2O	0.08	0.94	1.00
1172	2217.50	CO, N_2O	0.08	0.94	1.00
1173	2220.00	T, CF, CO, N_2O	0.12	0.96	1.00
1174	2222.50	T, CF	0.18	0.96	1.00
1175	2225.00	T, CF	0.16	0.93	1.00
1177	2230.00	N_2O	0.03	0.94	1.00

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
1178	2232.5	N_2O	0.02	0.93	1.00
1179	2235.0	N_2O	0.01	0.98	1.00
1180	2237.5	N_2O	0.01	0.89	1.00
1181	2240.0	N_2O	0.01	0.98	1.00
1187	2255.0	T, CF	0.00	0.97	0.95
1189	2260.0	T, CF	0.00	0.98	0.94
1190	2262.5	T, CF	0.00	0.72	1.00
1192	2267.5	T, CF	0.00	0.84	1.00
1193	2270.0	T, CF	0.00	0.98	0.93
1194	2272.5	T, CF	0.00	0.84	1.00
1196	2277.5	T, CF	0.00	0.85	1.00
1197	2280.0	T, CF	0.00	0.98	0.94
1198	2282.5	T, CF	0.00	0.73	1.00
1199	2285.0	T, CF	0.00	0.98	0.95
1200	2287.5	T, CF	0.00	0.88	1.00
1202	2292.5	T, CF	0.00	0.98	1.00
1203	2295.0	T, CF	0.00	0.98	0.92
1204	2297.5	T, CF	0.00	0.98	1.00
1206	2302.5	T, CF	0.00	0.98	1.00
1207	2305.0	T, CF	0.00	0.98	0.91
1208	2307.5	T, CF	0.00	0.98	1.00
1210	2312.5	T, CF	0.00	0.98	1.00
1212	2317.5	T, CF	0.00	0.99	1.00
1214	2322.5	T, CF	0.00	0.99	1.00
1215	2325.0	T, CF	0.00	0.99	0.93
1217	2330.0	T, CF	0.00	0.99	0.94
1218	2332.5	T, CF, CO_2	0.00	0.99	1.00
1220	2337.5	T, CF	0.00	0.99	1.00
1222	2342.5	T, CF	0.00	0.99	1.00
1224	2347.5	T, CF	0.00	0.99	1.00
1226	2352.5	CO_2	0.00	0.99	1.00
1228	2357.5	T, CF	0.00	0.99	1.00
1229	2360.0	T, CF	0.00	0.99	0.98
1231	2365.0	T, CF	0.00	0.99	0.97
1232	2367.5	T, CF	0.00	0.99	1.00
1234	2372.5	T, CF	0.00	0.99	1.00
1235	2375.0	T, CF	0.00	0.99	0.94
1236	2377.5	T, CF, CO_2	0.00	0.99	1.00
1237	2380.0	T, CF	0.00	0.99	0.96
1238	2382.5	T, CF	0.00	0.90	1.00
1239	2385.0	T, CF, CO_2	0.00	0.99	0.99
1241	2390.0	T, CF	0.08	0.99	1.00

T1: The NOAA/NESDIS CrIS Channel Selection (See legend at the bottom)

I	cm^{-1}	Chn Sensitivity	τ_{fix}	τ_{HO_2}	τ_{O_3}
1242	2392.5	T, CF, CO_2	0.20	0.99	1.0
1243	2395.0	T, CF	0.31	0.99	1.0
1244	2397.5	T, CF	0.37	0.99	1.0
1245	2400.0	CO_2	0.41	0.99	1.0
1247	2405.0	T, CF	0.47	0.99	1.0
1250	2412.5	T, CF	0.52	0.99	1.0
1270	2462.5	Srf	0.76	0.99	1.0
1271	2465.0	Srf	0.77	0.99	1.0
1282	2492.5	Srf	0.86	0.99	1.0
1285	2500.0	Srf	0.89	0.99	1.0
1288	2507.5	Srf	0.90	0.99	1.0
1290	2512.5	Srf	0.90	0.99	1.0
1293	2520.0	Srf	0.92	0.99	1.0
1298	2532.5	Srf	0.90	0.98	1.0
1301	2540.0	Srf	0.87	0.99	1.0

5.1 Table legend

- column 1: frequency index
- column 2: wave number
- column 3: channel sensitivity flags:
 - T: temperature sensitive channels
 - Srf: Surface temperature and emissivity sensitive channels
 - H_2O : water vapor sensitive channels
 - O_3 : ozone sensitive channels
 - CF: cloud filtering channels
 - CH: cloud height channels
 - CO_2 : carbon dioxide sensitive channels
 - CH_4 : methane sensitive channels
 - CO : carbon monoxide sensitive channels
 - HNO_3 : HNO_3 sensitive channels
 - N_2O : N_2O sensitive channels
 - SO_2 : SO_2 sensitive channels
- column 4: Mid-latitude atmosphere total transmittance of fixed concentration atmospheric gases
- column 5: Mid-latitude atmosphere water vapor total transmittance
- column 6: Mid-latitude atmosphere ozone and liquid water total transmittance

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